

Review

Not peer-reviewed version

---

# Evidence for a New Oxidation Mechanism for Sulfur Dioxide from Laboratory Measurements

---

[William R. Stockwell](#)<sup>\*</sup> and [Rosa M. Fitzgerald](#)

Posted Date: 7 August 2025

doi: 10.20944/preprints202508.0554.v1

Keywords: particulate matter; sulfur dioxide; sulfate particles; aerosol; gas-oxidation; heterogeneous processes



Preprints.org is a free multidisciplinary platform providing preprint service that is dedicated to making early versions of research outputs permanently available and citable. Preprints posted at Preprints.org appear in Web of Science, Crossref, Google Scholar, Scilit, Europe PMC.

Copyright: This open access article is published under a Creative Commons CC BY 4.0 license, which permit the free download, distribution, and reuse, provided that the author and preprint are cited in any reuse.

Disclaimer/Publisher's Note: The statements, opinions, and data contained in all publications are solely those of the individual author(s) and contributor(s) and not of MDPI and/or the editor(s). MDPI and/or the editor(s) disclaim responsibility for any injury to people or property resulting from any ideas, methods, instructions, or products referred to in the content.

Review

# Evidence for a New Oxidation Mechanism for Sulfur Dioxide from Laboratory Measurements

William R. Stockwell \* and Rosa M. Fitzgerald

Physics Department, The University of Texas at El Paso, El Paso, TX 79902, USA

\* Correspondence: william.r.stockwell@gmail.com

## Abstract

The oxidization of sulfur dioxide ( $\text{SO}_2$ ) occurs in the gas and liquid-phase and this oxidation contributes to particulate matter and acid precipitation. The production of sulfate particles is significant because of their impacts on climate, precipitation acidification, and human health. In this paper the focus is on the oxidation of  $\text{SO}_2$  and on the possibility of unknown heterogeneous reactions that may occur on sulfate aerosol surfaces. These results are based on a reanalysis of a foundational set of  $\text{SO}_2$  laboratory oxidation measurements. The experiments involved two sets of photochemical studies of nitrous acid (HONO), nitrogen oxides ( $\text{NO}_x = \text{NO} + \text{NO}_2$ ),  $\text{SO}_2$ , carbon monoxide (CO) and water vapor ( $\text{H}_2\text{O}$ ) mixtures made in molecular nitrogen ( $\text{N}_2$ ) with trace of molecular oxygen or in synthetic air. The reanalysis strongly suggests that there are uncharacterized processes for the oxidation of  $\text{SO}_2$  that are nearly three times faster than the known gas-phase reactions. The uncharacterized processes may involve sulfate aerosol surface reactions in the presence of nitrogen oxides. If these processes can be included in current atmospheric chemistry models, greater conversion rates of  $\text{SO}_2$  to sulfate aerosol will be calculated and this may reduce modelling bias.

**Keywords:** particulate matter; sulfur dioxide; sulfate particles; aerosol; gas-oxidation; heterogeneous processes

## 1. Introduction

The oxidation of sulfur dioxide ( $\text{SO}_2$ ) produces products that are significant sources of atmospheric particulate matter and acid precipitation (acid rain) [1–4]. The products, small droplets of liquid sulfuric acid ( $\text{H}_2\text{SO}_4$ ), bisulfate ( $\text{HSO}_4^-$ ) and sulfate ( $\text{SO}_4^{2-}$ ) have consequences for health, agriculture, climate, etc.  $\text{SO}_2$  is emitted from many natural and anthropogenic sources [4,5]. The greatest source of sulfur emissions into the atmosphere comes from biological processes in the Earth's oceans and this means that these sources are stronger in the southern hemisphere. Dimethyl sulfide (DMS,  $\text{CH}_3\text{SCH}_3$ ) constitutes a significant fraction of the sulfur emissions from the oceans [5–10]. Hydrogen sulfide ( $\text{H}_2\text{S}$ ) is another biogenically emitted compound.  $\text{H}_2\text{S}$  and DMS are converted to  $\text{SO}_2$  by a complicated chemical reaction mechanism [11].

Natural sources of  $\text{SO}_2$  include volcanoes [5] and its photochemical production from biogenic emissions such as hydrogen sulfide ( $\text{H}_2\text{S}$ ) and dimethyl sulfide ( $\text{CH}_3\text{SCH}_3$ ). Biomass burning, i.e. wildfires, is an underappreciated source of sulfur compound emissions [12]. The wildfire sources of sulfur compound emissions is increasing because of the increasing number and intensity of wildfires [13]. The major anthropogenic sources include the production of metals from sulfur containing ores (smelting) and the combustion of fossil fuels [14]. Coal is an example of a fossil fuel with sources that have large differences in sulfur content [15].

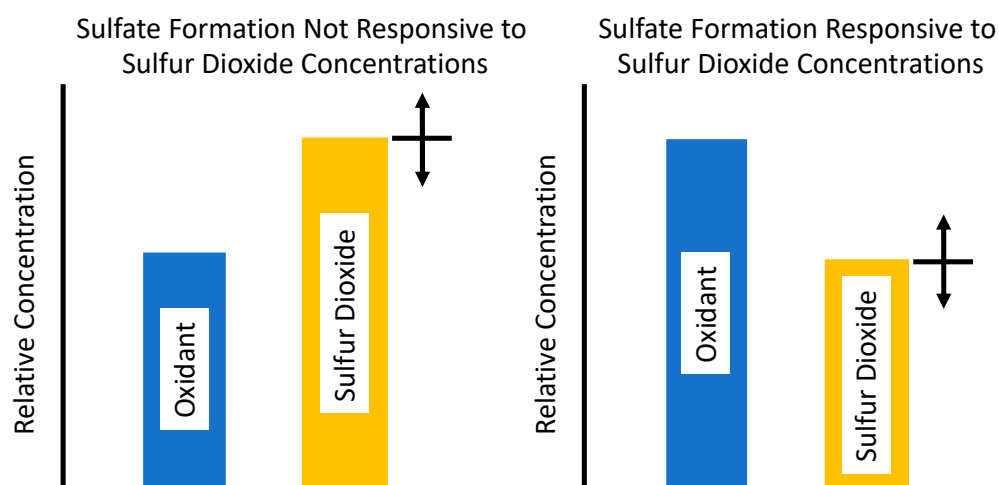
Sulfur dioxide is oxidized in the atmosphere to produce sulfate particles, i.e.,  $\text{H}_2\text{SO}_4$ ,  $\text{HSO}_4^-$  or sulfate  $\text{SO}_4^{2-}$ , that are chemical components of particulate matter or aerosols [3]. This oxidation of sulfur dioxide to produce particulate matter is very important because of their effects on climate, precipitation patterns, amounts and acidification, and their health effects [4]. The chemistry may have possible applications to exoplanetary atmospheres too [16].

Sulfur containing particles have very significant effects on climate because they strongly scatter solar radiation which effects the Earth's radiation budget [4]. In fact, the emission of sulfate containing particles has been suggested as a method of cooling the Earth's atmosphere to counteract global warming [17].

Sulfate containing particles, along with other forms of particulate matter, contribute to the secondary aerosol effect [4]. Fine aerosol particles provide nucleation sites for the formation of cloud water droplets. If there is an increase in the concentration of fine aerosol particles, then there will be more, but smaller droplets formed for a given level of water vapor in the atmosphere. Depending upon conditions the secondary aerosol effect can affect the amount and geographical distribution of precipitation.

Sulfate compounds are acidic in liquid water, and they contribute strongly to the acidification of precipitation along with nitrates ( $\text{NO}_3^-$ ) and organic acids such as formic ( $\text{HCOOH}$ ) and acetic acid ( $\text{CH}_3\text{COOH}$ ) [18]. For this reason and direct health effects, emissions of  $\text{SO}_2$  have regulated in the United States by the Clean Air Act Amendments since 1970 [19]. Acid precipitation was recognized as a major problem in Europe, Canada and the United States [20,21] during the late 1970s and 1980s. One of the major driving political and scientific questions for public policy regarding acid precipitation was to determine if the production of sulfate was limited by the available sulfur dioxide ( $\text{SO}_2$ ) or by the available oxidant; this uncertainty was known as the oxidant limitation question. This was a major issue of several international programs including the American National Acid Precipitation Program (NAPAP) and the U.S. Department of Energy's Processing of Emissions by Clouds Program (PRECP) [22–26].

Figure 1 shows the difference between a situation where the oxidation of  $\text{SO}_2$  is limited by the available oxidant or not [27]. If there is a limited amount of oxidant relative to the amount of  $\text{SO}_2$ , left plot, then moderate reductions in  $\text{SO}_2$  will not result in less sulfate produced; this condition was known as oxidant limited. The right panel shows that if there is more oxidant available than  $\text{SO}_2$  then reductions in  $\text{SO}_2$  will result in reductions in the amount of sulfate produced. Finding the chemical mechanisms for the oxidation of  $\text{SO}_2$  is important for determining the emission reduction policies needed to reduce acid precipitation.



**Figure 1.** The plot to the left shows an oxidant limited condition. In that case reductions or increases in  $\text{SO}_2$  concentrations do not affect the formation of sulfate. The plot to the right shows the case where there is sufficient oxidant to convert  $\text{SO}_2$  to sulfate and decreases in  $\text{SO}_2$  concentrations lead to decreases in sulfate production.

The oxidation of  $\text{SO}_2$  may be important on local, urban scales as well as regional and global. In general, the particulate matter produced through atmospheric chemical reactions are fine particles with an aerodynamic diameter of  $2.5 \mu\text{m}$  and these particles are known as  $\text{PM}_{2.5}$ . Larger particles with an aerodynamic diameter of  $10 \mu\text{m}$  are known as  $\text{PM}_{10}$  [19]. As an example, the city of El Paso

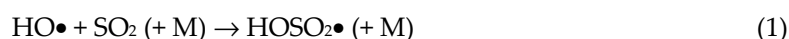
Texas in the Paso Del Norte Border Region of the United States and Mexico has had problems with high concentrations of PM<sub>2.5</sub> and PM<sub>10</sub> [28,29]. High emissions of SO<sub>2</sub>, NO<sub>x</sub> and volatile organic compounds and wind-swept mineral dust have occurred in El Paso which contributes to particulate pollution. Emissions of nitrogen oxides are significant because the region contains the American – Mexican border. Truck and auto traffic are queued near the border until they are cleared to pass the border into the United States.

### 1.1. SO<sub>2</sub> Oxidation Mechanisms

Many atmospheric chemical reactions produce products that condense to form particulate matter, PM<sub>2.5</sub> or PM<sub>10</sub>. In general, SO<sub>2</sub> is oxidized in either gas or aqueous phases or it may be deposited directly to the ground. In this paper we present evidence that there may be chemical reactions that may involve nitrogen oxides and/or heterogeneous processes involving sulfate aerosols.

#### 1.1.1. The Gas-Phase Oxidation of Sulfur Dioxide

There are two gas-phase mechanisms known for the gas-phase oxidation of sulfur dioxide. The hydroxyl radical (HO•) is the oxidant [30]. As shown by the reactions given below the HO• reacts with SO<sub>2</sub> to produce the adduct HOSO<sub>2</sub>•, note that M is any molecule of air that stabilizes the formation of the adduct by absorbing excess collision energy.

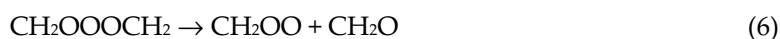
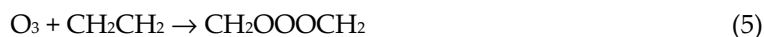


The HOSO<sub>2</sub>• adduct reacts with molecular oxygen to produce the hydroperoxyl radical (HO<sub>2</sub>•) and sulfur trioxide (SO<sub>3</sub>) [30]. SO<sub>3</sub> reacts with water vapor to produce sulfuric acid. In a nitrogen oxide polluted atmosphere, HO<sub>2</sub>• reacts with NO to reproduce HO• as shown below.



The hydroxyl radical produced by Reaction 4 can react with SO<sub>2</sub> making this mechanism a chain mechanism. When the concentration of nitric oxide (NO) is low, HO<sub>2</sub>• reacts with another HO<sub>2</sub>• or with organic peroxy radicals to produce hydrogen peroxide (H<sub>2</sub>O<sub>2</sub>) or an organic peroxide.

Another gas-phase process for the oxidation of SO<sub>2</sub> reaction with Criegee intermediates [31,32]. Criegee intermediates are produced by reactions of ozone (O<sub>3</sub>) with alkenes. An example of the mechanism for the production of Criegee intermediates from ethene (CH<sub>2</sub>CH<sub>2</sub>) and its oxidation of SO<sub>2</sub> is given below, Reactions 5–6.



The average total rates of the HO• radical and the Criegee intermediate mechanisms are a few percent per hour for a range of realistic conditions with a maximum rate of 6.13 % hr<sup>-1</sup> [33].

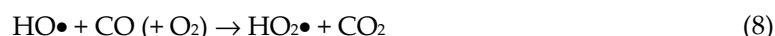
In cloud water or water coated aerosols the H<sub>2</sub>O<sub>2</sub> may react with SO<sub>2</sub> to produce sulfate [1,4,34] and in the gas-phase SO<sub>2</sub> does not affect the concentration of the HO•. Either way there is little or no oxidant limitation to sulfate production from SO<sub>2</sub> in the lower troposphere.

However, there is reason to suspect that there could be additional oxidation processes. Relative to observations, air quality models underestimate sulfate production across the Eastern U.S and the models underestimate the effect of temperature on its production [35,36]. These newer modeling studies for United States, Alaska and Beijing, China indicate that estimates of particulate sulfur may

be improved by the addition of additional heterogeneous chemical reactions [36]. In this paper we present an analysis of experimental data that suggests a process for the oxidation of SO<sub>2</sub> that is not included in models.

## 2. Materials and Methods

The experiments that are reanalyzed here were made with the objective to measure the rate coefficient of the HO• radical with SO<sub>2</sub> (HO+SO<sub>2</sub>) reaction) relative to the rate coefficient of the HO• radical with carbon monoxide (HO+CO) reaction [30,37]. The reaction of HO• with SO<sub>2</sub> (Reaction 1) is given above and the reaction of with HO• with CO is given below.



The original analysis determined that HO• radical concentrations were not affected by SO<sub>2</sub> concentrations, and this observation led to the proposal of the mechanism given by Reactions 1 through 4 [30]. This study was cited as a landmark in sulfur oxide chemistry [38]. The proposed mechanism was verified by several independent studies [39,40].

Mixtures of SO<sub>2</sub>, carbon monoxide (CO), nitrous acid (HONO), nitric oxide (NO) and nitrogen dioxide (NO<sub>2</sub>) were photolyzed in molecular nitrogen (N<sub>2</sub>) or synthetic air as a background gas, Table 1 [37]. The initial reactive nitrogen is the total molecular concentration of NO, NO<sub>2</sub> and HONO, Table 1 gives the initial conditions for the experiments made in N<sub>2</sub>; note that there was a small trace of molecular (O<sub>2</sub>) introduced during the filling of the chamber. Table 2 gives the initial conditions for the experiments made in synthetic air. Photolysis of nitrous acid produced HO• radicals.



The photolysis experiments were made using a long-path glass chamber that was 6m long and the concentration changes were measured with a Fourier Transform Spectrometer [30,37]. The concentrations of CO were sufficiently high so that they did not appreciably change during the experiments so the relative extent of the HO•+CO reaction was taken as the total CO<sub>2</sub> produced during each experiment. The rates of the HO•+SO<sub>2</sub> and HO•+CO reactions are given below where k<sub>SO<sub>2</sub></sub> and k<sub>CO</sub> are rate coefficients for these reactions.

$$\frac{[\text{SO}_2]}{dt} = k_{\text{HO}+\text{SO}_2} [\text{HO}\bullet] [\text{SO}_2] \quad (9)$$

$$\frac{[\text{CO}]}{dt} = k_{\text{HO}+\text{CO}} [\text{HO}\bullet] [\text{CO}] \quad (10)$$

Equations 9 and 10 may be rearranged to derive the ratio, k<sub>HO+SO<sub>2</sub></sub>/k<sub>HO+CO</sub>. To derive Equation 11 for the experiments, note that the HO• concentrations cancel out, that the concentrations of SO<sub>2</sub> and CO may be averaged over the photolysis period ([SO<sub>2</sub>]<sub>AVG</sub>, [CO]<sub>AVG</sub>) and that the change in CO is equal to the amount of CO<sub>2</sub> produced (Δ[CO<sub>2</sub>]). The change in the SO<sub>2</sub> concentration is given by Δ[SO<sub>2</sub>].

$$\frac{k_{\text{HO}+\text{SO}_2}}{k_{\text{HO}+\text{CO}}} = \frac{[\text{CO}]_{\text{AVG}} \Delta[\text{SO}_2]}{[\text{SO}_2]_{\text{AVG}} \Delta[\text{CO}_2]} \quad (11)$$

The gas-phase rate coefficients for the Reactions 1 and 8 are well known now. These were used to calculate the literature k<sub>HO+SO<sub>2</sub></sub>/k<sub>HO+CO</sub> ratio. The rate coefficients for the reaction of HO• with SO<sub>2</sub> is pressure and temperature dependent [41]. This reaction involves two small molecules forming a reactive intermediate that requires its collision energy to be dissipated through collisions with a third bodies. The rate coefficient n units of cm<sup>6</sup>molecule<sup>-1</sup>s<sup>-1</sup> is calculated through the following equations.

$$k_0(T) = k_{0,298} \left( \frac{T}{298} \right)^{-n} \quad (12)$$

$$k_\infty(T) = k_{\infty,298} \left( \frac{T}{298} \right)^{-m} \quad (13)$$

$$k_{HO+SO_2}(T, [M]) = \left\{ \frac{k_{\infty}(T)k_0(T)[M]}{k_{\infty}(T) + k_0(T)[M]} \right\} 0.6 \left\{ 1 + \left[ \log_{10} \left( \frac{k_0(T)[M]}{k_{\infty}(T)} \right) \right]^2 \right\}^{-1} \quad (14)$$

where:  $k_{0,298} = 2.90 \times 10^{-31}$ ,  $n = 4.1$ ,  $k_{\infty,298} = 1.70 \times 10^{-12}$  and  $m = -0.2$ ; and  $[M]$  is the concentration of the total background gas in molecules  $\text{cm}^{-3}$  [41].

The rate coefficient for the reaction of  $\text{HO}\bullet$  with CO in units of  $\text{cm}^6 \text{molecule}^{-1} \text{s}^{-1}$  is given is given by equation 15 [42,43]:

$$k_{HO+CO} = 1.44 \times 10^{-13} \times \left( 1.0 + \frac{0.8 \times [N_2]}{4 \times 10^{+19}} \right) \quad (15)$$

These were used to calculate literature ratio of the coefficients given in the Results section of this work.

**Table 1.** Initial conditions for sulfur dioxide oxidation experiments that were made using molecular nitrogen as the background gas. All experiments contained a trace of molecular oxygen with a pressure assumed to be  $4.64 \pm 0.06 \times 10^{-5}$  atm [37]. The initial reactive nitrogen is the total initial sum of the HONO, NO and  $\text{NO}_2$  concentrations.

Experiment Number	$\text{N}_2$ (atm)	Initial Reactive Nitrogen (Molec $\text{cm}^{-3} \times 10^{-15}$ )	Initial CO (Molec $\text{cm}^{-3} \times 10^{-16}$ )	Initial $\text{SO}_2$ (Molec $\text{cm}^{-3} \times 10^{-15}$ )
1	0.921	0.817	2.31	2.96
2	0.954	0.882	3.43	3.88
3	0.925	0.917	2.31	2.23
4	0.924	1.09	1.60	2.07
5	0.922	1.23	1.62	1.69
6	0.925	2.17	1.10	1.70
7	0.933	2.38	1.03	1.28
8	0.921	1.91	1.07	2.01

**Table 2.** Initial conditions for sulfur dioxide oxidation experiments that were made using synthetic air as the background gas [37].

Experiment Number	$\text{N}_2$ (atm)	$\text{O}_2$ (atm)	Initial Reactive Nitrogen (Molec $\text{cm}^{-3} \times 10^{-15}$ )	Initial CO (Molec $\text{cm}^{-3} \times 10^{-16}$ )	Initial $\text{SO}_2$ (Molec $\text{cm}^{-3} \times 10^{-15}$ )
O1	0.761	0.176	0.739	2.29	3.20
O2	0.748	0.171	1.43	2.32	2.26
O3	0.749	0.172	1.34	1.86	3.42
O4	0.746	0.175	0.828	1.83	2.31
O5	0.752	0.169	0.908	1.38	3.14
O6	0.742	0.176	1.02	1.39	2.36
O7	0.750	0.171	1.20	0.935	3.56

### 3. Results

Table 3 shows experimental results for the photolysis experiments.

**Table 3.** Experimental Results for sulfur dioxide oxidation [37].

Experiment Number	Reaction Time (min)	$\Delta$ Reactive Nitrogen (Molec cm <sup>-3</sup> ×10 <sup>-14</sup> )	$\Delta$ CO <sub>2</sub> (Molec cm <sup>-3</sup> ×10 <sup>-14</sup> )	$\Delta$ SO <sub>2</sub> (Molec cm <sup>-3</sup> ×10 <sup>-14</sup> )	Experimental $\frac{k_{HO+SO_2}}{k_{HO+CO}}$
1	29.12	0.088	1.60	2.53	12.9
2	25.37	0.52	1.01	1.03	9.1
3	16.98	0.060	1.58	1.38	9.4
4	25.57	-0.004	1.07	1.61	12.1
5	25.72	1.20	0.979	1.26	12.8
6	25.78	3.90	0.816	1.72	14.5
7	25.82	2.57	1.00	1.72	14.9
8	28.47	19.10	0.383	1.03	14.8
O1	28.20	0.100	1.50	2.34	11.6
O2	28.87	0.917	1.27	1.75	14.7
O3	28.37	-0.024	0.832	2.45	16.7
O4	28.62	-0.062	1.28	2.57	16.8
O5	28.87	-0.609	0.916	1.64	8.0
O6	28.32	0.382	0.962	2.45	15.9
O7	37.20	0.803	0.763	2.92	10.5

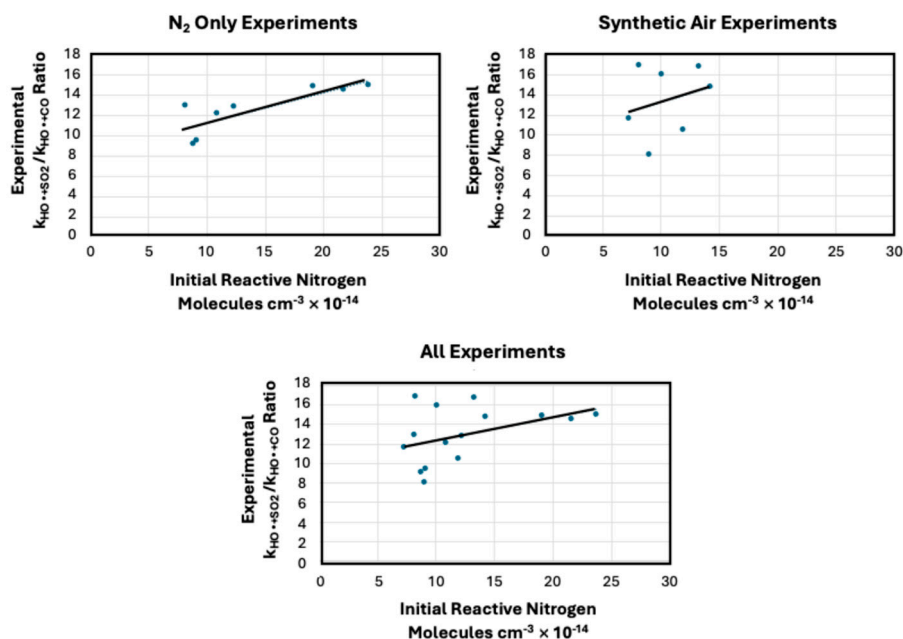
Table 3 shows the experimental  $k_{HO+SO_2}/k_{HO+CO}$  ratios calculated using Equation 11 from the initial concentrations. The observed changes in SO<sub>2</sub> and CO<sub>2</sub> concentrations The ratios ranges from 9.1 to 14.9 for the experiments made with background N<sub>2</sub>. The ratio ranges from 0.8 to 16.7 for the experiments made with background synthetic air.

Table 4 shows calculated  $k_{HO+SO_2}/k_{HO+CO}$  ratios from the literature rate coefficients. The expected values of the ratios are between 4 and 4.5. The experimental values of the  $k_{HO+SO_2}/k_{HO+CO}$  ratios are between 1.80 and 3.75 times greater than the literature calculated values.

**Table 4.** Comparison of experimental  $k_{HO+SO_2}/k_{HO+CO}$  ratio with literature values. The values of  $k_{HO+SO_2}$  and  $k_{HO+CO}$  were calculated for a temperature of 300K and the N<sub>2</sub> pressures given in Tables 1 and 2.

Experiment Number	$k_{HO+SO_2}$ (cm <sup>3</sup> Molec <sup>-1</sup> ×10 <sup>13</sup> )	$k_{HO+CO}$ (cm <sup>3</sup> Molec <sup>-1</sup> ×10 <sup>13</sup> )	Literature $\frac{k_{HO+SO_2}}{k_{HO+CO}}$	Experimental/Literature $\frac{k_{HO+SO_2}}{k_{HO+CO}}$
1	9.22	2.21	4.17	3.10
2	9.33	2.24	4.17	2.18
3	9.23	2.22	4.17	2.26
4	9.23	2.21	4.17	2.90
5	9.22	2.21	4.17	3.07
6	9.23	2.22	4.17	3.48
7	9.26	2.22	4.17	3.58
8	9.22	2.21	4.17	3.55
O1	9.27	2.08	4.46	2.60
O2	9.21	2.07	4.46	3.30
O3	9.22	2.07	4.46	3.75
O4	9.22	2.07	4.46	3.76
O5	9.22	2.07	4.45	1.80
O6	9.21	2.06	4.47	3.56
O7	9.22	2.07	4.46	2.36

Plots were made to examine the possibility that there was an effect of reactive nitrogen on the  $k_{\text{HO}+\text{SO}_2}/k_{\text{HO}+\text{CO}}$  ratios. Figure 2 shows plots of the experimental  $k_{\text{HO}+\text{SO}_2}/k_{\text{HO}+\text{CO}}$  ratios as functions of the initial reactive nitrogen oxides (HONO+NO+NO<sub>2</sub>). Table 5 shows the average reactive nitrogen in the three sets of experiments, the average  $k_{\text{HO}+\text{SO}_2}/k_{\text{HO}+\text{CO}}$  ratio and the shared variance between them.



**Figure 2.** Figure shows the relationship between the experimental  $k_{\text{HO}+\text{SO}_2}/k_{\text{HO}+\text{CO}}$  ratios and the initial reactive nitrogen used in the experiments. The initial reactive nitrogen is the total initial sum of the HONO, NO and NO<sub>2</sub> concentrations. The plot at the top left is for the experiments made with N<sub>2</sub> as the background gas. The plot at the top right is for the experiments made with synthetic air as the background gas. The plot at the lower center shows all experiments plotted together. The line in each plot is the regression line.

**Table 5.** Comparison of average experimental  $k_{\text{HO}+\text{SO}_2}/k_{\text{HO}+\text{CO}}$  ratio to the average reactive nitrogen concentrations and the shared variance between them as determined from the plots shown in Figure 2.

Experimental Series	Average Reactive Nitrogen Conc. (Molec $\text{cm}^{-3} \times 10^{-14}$ )	Average $k_{\text{HO}+\text{SO}_2}/k_{\text{HO}+\text{CO}}$ Ratio	Shared Variance (R <sup>2</sup> )
N <sub>2</sub> Only	14.2	12.6	0.65
Synthetic Air	10.7	13.5	0.076
All	12.6	13.0	0.18

#### 4. Discussion

This new analysis of photolysis experiments made with relatively high concentrations of HONO, NO, NO<sub>2</sub>, SO<sub>2</sub> and CO showed that there may be an uncharacterized oxidation process for SO<sub>2</sub>. The literature value of the  $k_{\text{HO}+\text{SO}_2}/k_{\text{HO}+\text{CO}}$  ratio for the experimental conditions is near 4 to 4.5 while the measured ratio is around three times greater for the N<sub>2</sub> only, synthetic air background gas experiments and both sets of experiments considered together.

The experimental ratio for the experiments with N<sub>2</sub> only as the background gas has a shared variance of 0.65 with the initial total reactive nitrogen concentration. However, the shared variance between the experimental ratio and the initial total reactive nitrogen concentration is much lower for the synthetic air case. Although there is much uncertainty, this could be due to an oxygen effect that is not taken into account in this analysis.

If there is a new oxidation process for SO<sub>2</sub> that is three times greater than the known hydroxyl radical reaction that occurs in NO<sub>x</sub> polluted atmospheres then conversion rates in the range of 10 to



20% hr<sup>-1</sup> may occur in urban regions such as El Paso, Texas with sources of SO<sub>2</sub> and NO<sub>x</sub> (using reference 33 as the baseline).

Sulfuric acid has been produced on a commercial scale from the direct reaction of SO<sub>2</sub> and NO<sub>x</sub> and H<sub>2</sub>O. These were mixed at very high concentrations in large chambers, made from lead, where reactions produced H<sub>2</sub>SO<sub>4</sub>. [44] Based on our experiments we show that similar reactions may occur at lower concentrations. These reactions are probably heterogeneous and appear to be faster than the gas-phase reactions. The reactions may be important for converting SO<sub>2</sub> to sulfate on urban scales and therefore contribute to local PM<sub>2.5</sub> formation. Improved representation of these processes may improve the agreement between measurements and air quality modeling results [35,36] for the United States, China and elsewhere. New research involving laboratory studies and field measurements are needed to better characterize the oxidation of SO<sub>2</sub> because it is possible that heterogeneous reactions occurred on sulfate particles as soon as they were formed and/or other reactive nitrogen catalyzed processes occur. Therefore, better measurements of formation and properties of the sulfate particles and their surface reactions are needed.

**Author Contributions:** Conceptualization, W.S. and R.F.; methodology, W.S.; validation, W.S., and R.F.; formal analysis, W.S.; investigation, W.S.; resources, W.S. and R.F.; data curation, W.S.; writing—original draft preparation, W.S.; writing—review and editing, R.F.; visualization, W.S.; supervision, W.R. and R.F.; project administration, R.F.; funding acquisition, R.F. All authors have read and agreed to the published version of the manuscript.

**Funding:** This research was funded by the NOAA Center for Atmospheric Science—Meteorology (NCASM II), which is funded by the U.S. Department of Commerce, National Oceanic and Atmospheric Administration, Educational Partnership Program under Agreement No. NA22SEC4810015.

**Data Availability Statement:** Original data is available in: Stockwell, William. "The chemistry of nitrogen oxides/sulphur oxides/hydrogen oxides systems." Doctoral dissertation, Ohio State University, 1981. [http://rave.ohiolink.edu/etdc/view?acc\\_num=osu1487171566434414](http://rave.ohiolink.edu/etdc/view?acc_num=osu1487171566434414).

**Conflicts of Interest:** The authors declare no conflicts of interest.

## References

1. Middleton, P.M.; Klang, C.S.; Mohnen, V.A. The relative importance of various urban sulfate aerosol production mechanisms – A Theoretical comparison. In *Heterogeneous Atmospheric Chemistry*; Schryer D.R.; American Geophysical Union; Washington D.C., United States, 1982, pp. 221–230.
2. Walcek, C.J.; Taylor, G.R. A Theoretical Method for Computing Vertical Distributions of Acidity and Sulfate Production within Cumulus Clouds. *J. Atmos. Sci.* **1986**, *43*, 339–55. [https://doi.org/10.1175/1520-0469\(1986\)043<0339:atmfcv>2.0.co;2](https://doi.org/10.1175/1520-0469(1986)043<0339:atmfcv>2.0.co;2).
3. Seinfeld, J.H.; Pandis, S.N. *Atmospheric Chemistry and Physics: From Air Pollution to Climate Change* 3rd. 2016, John Wiley & Sons, New York.
4. Kolb, C.F.; Worsnop, D.R. Chemistry and Composition of Atmospheric Aerosol Particles. *Annu. Rev. Phys. Chem.* **2012**, *63*, 471–491. <https://doi.org/10.1146/annurev-physchem-032511-143706>
5. Warneck, P., Chapter 10. Sulphur compounds in the atmosphere. In *Chemistry of the Natural Atmosphere*, 2ed; Academic Press, San Diego, California, United States, 2000, pp. 587–655.
6. Brasseur, G.P.; Orlando, J.J.; Tyndall, G.S. *Atmospheric Chemistry and Global Change*, Oxford University Press, United Kingdom, 1999.
7. Rotstain, L.D.; Lohmann, U. Simulation of the Tropospheric Sulfur Cycle in a Global Model with a Physically Based Cloud Scheme. *J. Geophys. Res.* **2002**, *107*, 4592. <https://doi.org/10.1029/2002JD002128>.
8. Gondwe, M.; Krol, M.; Gieskes, W.; Klaassen, W.; de Baar, H. The Contribution of Ocean-Leaving DMS to the Global Atmospheric Burdens of DMS, MSA, SO<sub>2</sub>, and NSS SO<sub>4</sub><sup>2-</sup>. *Global Biogeochem. Cycles* **2003**, *17*, 1056. <https://doi.org/10.1029/2002GB001937>.
9. Gondwe, M.; Krol, M.; Gieskes, W.; Klaassen, W.; de Baar, H. Correction to "The Contribution of Ocean-Leaving DMS to the Global Atmospheric Burdens of DMS, MSA, SO<sub>2</sub>, and NSS SO<sub>4</sub><sup>2-</sup>". *Global Biogeochem. Cycles* **2003**, *17*, 1106. <https://doi.org/10.1029/2003GB002153>.

10. Lana, A.; Bell, T.G.; Simó, R.; Vallina, S.M.; Ballabrera-Poy, J.; Kettle, A.J.; Dachs, J.; Bopp, L.; Saltzman, E.S.; Stefels, J.; Johnson, J.E.; Liss, P.S. An Updated Climatology of Surface Dimethylsulfide Concentrations and Emission Fluxes in the Global Ocean. *Global Biogeochem. Cycles* **2011**, *25*, GB 1004. <https://doi.org/10.1029/2010GB003850>.
11. Goss, M.B.; Kroll, J.H. Chamber Studies of OH + Dimethyl Sulfoxide and Dimethyl Disulfide: Insights into the Dimethyl Sulfide Oxidation Mechanism. *Atmos. Chem. Phys.* **2024**, *24*, 1299–1314. <https://doi.org/10.5194/acp-24-1299-2024>.
12. Rickly, P.S.; Guo, H.; Campuzano-Jost, P.; Jimenez, J. L.; Wolfe, G.M.; Bennett, R.; Bourgeois, I.; Crouse, J.D.; Dibb, J.E.; DiGangi, J.P.; Diskin, G.S.; Dollner, M.; Gargulinski, E.M.; Hall, S.R.; Halliday, H.S.; Hanisco, T.F.; Hannun, R.A.; Liao, J.; Moore, R.; Nault, B.A.; Nowak, J.B.; Peischl, J.; Robinson, C.E.; Ryerson, T.; Sanchez, K.J.; Schöberl, M.; Soja, A.J.; St. Clair, J.M.; Thornhill, K.L.; Ullmann, K.; Wennberg, P.O.; Weinzierl, B.; Wiggins, E.B.; Winstead, E.L.; Rollins, A.W. Emission factors and evolution of SO<sub>2</sub> measured from biomass burning in wildfires and agricultural fires. *Atmos. Chem. Phys.*, **2022**, *22*, 15603–15620. <https://doi.org/10.5194/acp-22-15603-2022>
13. Pausas, J.G.; Keeley, J.E. Wildfires and Global Change. *Frontiers in Ecology and the Environment*, **2021**, *19*, 387–395. <https://doi.org/10.1002/fee.2359>
14. Smith, S.J.; van Aardenne, J.; Klimont, Z.; Andres, R.J.; Volke, A.; Delgado Arias, S. Anthropogenic Sulfur Dioxide Emissions: 1850–2005. *Atmos. Chem. Phys.* **2011**, *11*, 1101–1116. <https://doi.org/10.5194/acp-11-1101-2011>.
15. Calkins, W.H. The Chemical Forms of Sulfur in Coal: A Review. *Fuel* **1994**, *73*, 475–484. [https://doi.org/10.1016/0016-2361\(94\)90028-0](https://doi.org/10.1016/0016-2361(94)90028-0)
16. Hu, R.; Seager S.; Bains W. Photochemistry in Terrestrial Exoplanet Atmospheres. II. H<sub>2</sub>S and SO<sub>2</sub> Photochemistry in Anoxic Atmospheres. *The Astrophysical Journal*. **2013**, *769*, 1–14. <https://doi.org/10.1088/0004-637X/769/1/6>
17. Visioni, D.; Pitari, G.; Aquila, V. Sulfate Geoengineering: A Review of the Factors Controlling the Needed Injection of Sulfur Dioxide. *Atmos. Chem. Phys.* **2017**, *17*, 3879–3889. <https://doi.org/10.5194/acp-17-3879-2017>
18. Tilgner, A.; Schaefer, T.; Alexander, B.; Barth, M.; Collett Jr., J.L.; Fahey, K.M.; Nenes, A.; Pye, H.O.T.; Herrmann H.; McNeill, V.F. Acidity and the Multiphase Chemistry of Atmospheric Aqueous Particles and Clouds. *Atmos. Chem. Phys.* **2021**, *21*, 13483–536. doi.org:10.5194/acp-21-13483-2021
19. Jacobson, M.Z. Atmospheric Pollution, History, Science and Regulation. Cambridge University Press, Cambridge, United Kingdom, **2002**.
20. Bolin, B.; Persson, C. Regional Dispersion and Deposition of Atmospheric Pollutants with Particular Application to Sulfur Pollution over Western Europe. *Tellus* **1975**, *27*, 281–310. <https://doi.org/10.1111/j.2153-3490.1975.tb01679.x>
21. Cowling, E.B. Acid Precipitation in Historical Perspective. *Environ. Sci. Technol.* **1982**, *16*, 110A-123A. <https://doi.org/10.1021/es00096a002>
22. Acid Deposition: State of Science and Technology, Volume 1 – Emissions, Atmospheric Processes and Deposition; National Acid Assessment Program; Superintendent of Documents; Washington D.C. United States, 1990.
23. Acid Deposition: State of Science and Technology, Volume 2 – Aquatic Processes and Effects; National Acid Assessment Program; Superintendent of Documents; Washington D.C. United States, 1990.
24. Acid Deposition: State of Science and Technology, Volume 3 – Terrestrial, Materials, Health and Visibility Effects; National Acid Assessment Program; Superintendent of Documents; Washington D.C. United States, 1990
25. Acid Deposition: State of Science and Technology, Volume 4 – Control Technologies, Future Emissions and Effects Valuation; National Acid Assessment Program; Superintendent of Documents; Washington D.C. United States, 1990.
26. Tanner, R.L.; Tichler, J.; Brown, R.; Davis, W.; Johnson, S.; Patrinos, A.A.; Sisterson, D.; Slinn, W.G. PRECP: the Department of Energy's Program on the Nonlinearity of Acid Precipitation Processes (No. BNL-38776). Brookhaven National Lab., Upton, NY (USA); Pacific Northwest Lab., Richland, WA (USA); Argonne National Lab., IL (USA) 1986.
27. Rhode, H.; Crutzen, R.; Vanderpol, A. Formation of Sulfuric and Nitric Acid in the Atmosphere During Long-Range Transport. *Tellus*, **1981**, *33*, 132 – 141.

28. Chen, L.-W.A.; Tropp, R.; Li, W.-W.; Zhu, D.; Chow, J.C.; Watson, J.C.; Zielinska, B. Aerosol and Air Toxics Exposure in El Paso, Texas: A Pilot Study. *Aerosol and Air Quality Research*, 12: 2012, 169–179. <https://doi.org/10.4209/aaqr.2011.10.0169>
29. Karle, N.N.; Mahmud, S.; Sakai, R.K.; Fitzgerald, R.M.; Morris, V.R.; Stockwell, W.R. Investigation of the Successive Ozone Episodes in the El Paso–Juarez Region in the Summer of 2017. *Atmosphere* 2020, 11, 532. <https://doi.org/10.3390/atmos11050532>.
30. Stockwell, W.R.; Calvert, J.G. The Mechanism of the HO–SO<sub>2</sub> Reaction. *Atmos. Environ.* 1983, 17, 2231–2235.
31. Sarwar, G.; Fahey, F.; Kwok, R.; Gilliam, R.; Xue, J.; Jianzhen, Y.; Carter, W.P.L. Potential impacts of two SO<sub>2</sub> oxidation pathways on regional sulfate concentrations: aqueous-phase oxidation by NO<sub>2</sub> and gas-phase oxidation by Stabilized Criegee Intermediates. *Atmos. Environ.* 2013, 68, 186–197.
32. Sarwar, G.; Simon, H.; Fahey, K.; Mathur, R.; Goliff, W.S.; Stockwell, W.R. Impact of Sulfur Dioxide Oxidation by Stabilized Criegee Intermediate on Sulfate, *Atmos. Environ.*, 2014, 85, 204–214.
33. Calvert, J.G.; Stockwell, W.R. Acid Generation in the Troposphere by Gas Phase Chemistry. *Envir. Sci. Technol.*, 1983, 17, 428A–443A.
34. Graedel, T.E.; Weschler, C.J. Chemistry Within Aqueous Atmospheric Aerosols and Raindrops. *Rev. Geophys. and Space Phys.* 1981, 19, 505–539.
35. Vannucci, P. F.; Foley, K.; Murphy, B. N.; Hogrefe, C.; Cohen, R. C.; and Pye, H. O. T.; Temperature-Dependent Composition of Summertime PM<sub>2.5</sub> in Observations and Model Predictions across the Eastern U.S., *ACS Earth Space Chem.*, 8 (2), 381–392, 2024. <https://doi.org/10.1021/acsearthspacechem.3c00333>.
36. Farrell, S. L.; Pye, H. O. T.; Gilliam, R.; Pouliot, G.; Huff, D.; Sarwar, G.; Vizuete, W.; Briggs, N.; Duan, F.; Ma, T.; Zhang, S.; Fahey, K. Predicted Impacts of Heterogeneous Chemical Pathways on Particulate Sulfur over Fairbanks (Alaska), the Northern Hemisphere and the Contiguous United States, *Atmos. Chem. Phys.* 2025, 25, 3287–3312. <https://doi.org/10.5194/acp-25-3287-2025>.
37. Stockwell, W.R. The Chemistry of Nitrogen Oxides/Sulphur Oxides/Hydrogen Oxides Systems. Ph.D. Thesis, The Ohio State University, Columbus, OH, USA, 1981. [http://rave.ohiolink.edu/etdc/view?acc\\_num=osu1487171566434414](http://rave.ohiolink.edu/etdc/view?acc_num=osu1487171566434414)
38. Crutzen, P.J. My Life with O<sub>3</sub>, NO<sub>x</sub> and Other YZO<sub>x</sub>s, Nobel Lecture. December 8, 1995. Available online: <https://www.nobelprize.org/prizes/chemistry/1995/crutzen/lecture/> (Accessed on 4 August 2025).
39. Gleason, J. F.; Sinha, A.; Howard, C. J. Kinetics of the Gas-Phase Reaction HOSO<sub>2</sub> + O<sub>2</sub> → HO<sub>2</sub> + SO<sub>3</sub>. *J. Phys. Chem.* 1987, 91, 719–724. <https://doi.org/10.1021/j100287a045>.
40. Egsgaard, H.; Carlson, L.; Florencio, H.; Drewello, T.; Schwarz, H. Experimental Evidence for the Gaseous HSO<sub>3</sub> Radical. The Key Intermediate in the Oxidation of SO<sub>2</sub> in the Atmosphere. *Chem. Phys. Lett.* 1988, 148, 537–540. [https://doi.org/10.1016/0009-2614\(88\)80327-0](https://doi.org/10.1016/0009-2614(88)80327-0).
41. Burkholder, J.B.; Sander, S.P.; Abbatt, J.; Barker, J.R.; Cappa, C.; Crouse, J.D.; Dibble, T.S.; Huie, R.E.; Kolb, C.E.; Kurylo, M.J.; Orkin, V.L.; Percival, C.J.; Wilmouth, D.M.; Wine, P.H. Chemical Kinetics and Photochemical Data for Use in Atmospheric Studies, Evaluation No. 19," JPL Publication 19-5, Jet Propulsion Laboratory, Pasadena, 2019 <http://jpldataeval.jpl.nasa.gov>.
42. Atkinson, R.; Baulch, D.L.; Cox, R.A.; Crowley, J.N.; Hampson, R.F.; Hynes, R.G.; Jenkin, M.E.; Rossi, M.J.; Troe, J. Evaluated Kinetic and Photochemical Data for Atmospheric Chemistry: Volume I – Gas Phase Reactions of O<sub>x</sub>, HO<sub>x</sub>, NO<sub>x</sub> and SO<sub>x</sub> Species. *Atmos. Chem. Phys.* 2004, 4, 1461–1738. <https://doi.org/10.5194/acp-4-1461-2004>, 2004.
43. Calvert, J.G.; Orlando, J.J.; Stockwell, W.R.; Wallington, T.J. The Mechanisms of Reactions Influencing Atmospheric Ozone, Oxford University Press: Oxford, United Kingdom, 2015; pp. 412–424.
44. Schroeder, W.H.; Urone, P. (1978). Isolation and Identification of Nitrosonium Hydrogen Sulfate (NOHSO<sub>4</sub>) as a Photochemical Reaction Product in Air Containing Sulfur Dioxide and Nitrogen Dioxide. *Envir. Sci. Technol.* 1978, 12, 545–550.

**Disclaimer/Publisher's Note:** The statements, opinions and data contained in all publications are solely those of the individual author(s) and contributor(s) and not of MDPI and/or the editor(s). MDPI and/or the editor(s) disclaim responsibility for any injury to people or property resulting from any ideas, methods, instructions or products referred to in the content.