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Article

Increased Extreme Precipitation in Western North America from Cut-Off Lows Under a Warming Climate

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Abstract: Cut-off Low (COL) pressure systems significantly influence local weather in regions with high COL frequency, such as western North America. However, future changes in COL frequency, intensity, and precipitation patterns remain uncertain. This study examines projected COL changes and their drivers in western North America under a high-emission scenario (SSP585) using a multi-model ensemble from CMIP6 and a feature tracking algorithm. We compare historical simulations (1980–2009) and future projections (2070–2099), revealing a marked increase in COL track density during summer in the northeast Pacific and western United States, while a strong decrease is projected for winter, associated with shifts in jet streams. Climate models project an increase in COL-related precipitation in future climate, with winter and spring experiencing more intense and localized precipitation, while autumn showing a more widespread precipitation pattern. Additionally, there is an increased frequency of extreme precipitation events, though accompanied by large uncertainties. The projected increase in extreme precipitation highlights the need to understand COL dynamics for effective climate adaptation in affected areas. Further research should aim to refine projections and reduce uncertainties, supporting better-informed policy and decision-making.

Keywords: cut-off lows; CMIP6; multi-model ensemble; climate projections; extreme precipitation; North America

1. Introduction

The scientific community has increasingly focused on understanding the potential impacts of climate change on various meteorological phenomena, including cut-off low (COL) pressure systems. COLs are synoptic-scale low-pressure systems displaced equatorward from the main westerly flow and are typically associated with Rossby wave breaking [1]. These systems are slow-moving or stationary, often leading to prolonged periods of extreme precipitation. Persistent storms associated with COLs have caused severe flooding in Europe and Asia [2–6], while in the Americas, COLs are a major driver of extreme rainfall along the western slopes of the Andes [7,8] and the Rocky Mountains [9–11], contributing more than 50% of annual precipitation. Despite historical analyses, there remains a lack of comprehensive research on projected changes in COL frequency, intensity, and associated precipitation patterns under high-emission future climate scenarios. Understanding these changes is crucial for predicting regional weather patterns and developing effective climate adaptation strategies.

Previous studies have extensively documented the climatology and impacts of COLs. Early work by [12] provided a comprehensive climatology of COLs at 200 hPa in the Northern Hemisphere, detailing their frequency and seasonal distribution. [13] expanded on this by examining the vertical structure and precipitation patterns of COLs in the Mediterranean region. Further comprehensive

climatologies in both hemispheres have been developed with reanalysis data and advanced methodologies for identifying and tracking COLs [14–19]. However, uncertainties remain due to variations in detection criteria, atmospheric levels, and datasets used [16,17,20,21].

Recent studies have enhanced our understanding of the dynamical processes and life cycles of COLs [22–24]. Additionally, evaluations of climate models indicate significant advancements in simulating COLs, with the sixth generation of Coupled Model Intercomparison Project (CMIP6) models outperforming the earlier CMIP5 simulations [25]. Furthermore, investigations into the effects of El Niño and La Niña on global COL predictability have provided valuable insights [26].

Despite extensive historical analyses of COLs, a substantial gap remains in understanding their response to future climate change. Although some studies have examined future changes in COLs in specific regions, they underscore the need for comprehensive research in other regions. For instance, high-resolution models have projected significant increases in frequency and intensity of mid-tropospheric closed-lows and extreme rainfall events over southern Africa [27]. Similarly, future climate conditions in western Europe could lead to more frequent and severe precipitation events associated with COLs [28,29]. In addition, high-intensity COLs with longer lifetimes are expected to become more frequent during the spring over land regions of the Northern Hemisphere [30]. However, these studies are either restricted to specific regions or focus solely on changes in COL frequency, leaving a critical gap in understanding how COLs will affect regional precipitation patterns under high-emission scenarios. Addressing this gap is crucial for developing effective climate adaptation and mitigation strategies.

Predicting the future behavior of COLs is challenging due to the numerous interacting factors that influence their regional patterns, including shifts in jet streams, wave breaking, changes in sea surface temperature patterns, land-sea temperature contrasts, and diabatic feedback directly affecting COL dynamics. The strength and position of jet streams, influenced by amplitude waves and split-jet structures, play a crucial role in COL variability. In particular, a weaker subtropical jet increases the COL frequency, while a stronger upstream polar front jet tends to deepen the COL vertical circulation [31]. However, these dynamics are generally not well represented in state-of-the-art climate models, including higher-resolution models, where an equatorward bias in the Northern Hemisphere jet streams leads to underestimations of COL activity in regions with high observed occurrence [25].

This study addresses the critical gap in the current understanding of future COL behavior in North America by providing projections under a high-emission scenario (SSP585). Using a multi-model ensemble from CMIP6 and a feature tracking algorithm, we examine seasonal changes in COL frequency, intensity, and associated precipitation patterns. Employing a COL-centered methodology, we directly measure the impact of climate warming on precipitation associated with COLs in climate model projections. This research offers detailed insights into how COL dynamics and precipitation may change under future climate conditions, particularly in North America's most active COL region, contributing valuable knowledge to climate science and aiding in the development of mitigation and adaptation strategies under a warming climate.

Key science questions this study seeks to answer include:

1. How will COL frequency and intensity change under a high-emission scenario (SSP585)?
2. How will COL-associated precipitation patterns, including extreme precipitation events, vary seasonally in future climate projections?
3. What dynamical processes drive COL changes, and what uncertainties do models present in these projections?

2. Data and Methods

We use climate model outputs from the CMIP6 [32] archive, specifically from eight high-performing models: ACCESS-CM2, BCC-CSM2-MR, CNRM-CM6-1-HR, MIROC-ES2L, MPI-ESM1-2-HR, MRI-ESM2-0, NESM3, and NorESM2-LM (Table 1). These models, which provide 6-hourly data at horizontal resolutions ranging from 250 km to 50 km, have shown to have improved performance in simulating COLs compared to CMIP5 models [25]. Analyses are conducted for each

season over a historical period (1980-2009) and a future period (2070-2099) under the SSP585 high-emission scenario.

Table 1. List of ERA5 reanalysis and CMIP6 models with their horizontal resolution of the atmospheric component of the model (in grid points and nominal resolution in parentheses) and the corresponding identified number of 250-hPa COL tracks per season within the region 20°-50°N and 140°-60°W for the historical period (1980-2009) and future model projections (2070-2099).

Model	Institute	Resolution	Tracks per season	
			1980-2009	2070-2099
ACCESS-CM2	CSIRO, Australia	192 × 145 (250 km)	10.1	9.8
BCC-CSM2-MR	BCC, China	320 × 160 (100 km)	11.8	11.8
CNRM-CM6-1-HR	CNRM, France	720 × 360 (50 km)	15.6	13.2
MIROC-ES2L	MIROC, Japan	128 × 64 (500 km)	10.1	9.0
MPI-ESM1-2-HR	MPI, Germany	384 × 192 (100 km)	10.5	10.7
MRI-ESM2-0	MRI, Japan	320 × 160 (100 km)	10.3	9.8
NESM3	NUIST, China	192 × 96 (250 km)	9.8	9.1
NorESM2-LM	NCC, Norway	144 × 96 (250 km)	6.9	7.4
CMIP6 mean			10.6	10.1
ERA5	ECMWF, Europe	1440 × 721 (31 km)	12.3	-

The model performance is evaluated using data from the European Centre for Medium Range Weather Forecasts (ECMWF) 5th generation reanalysis (ERA5) [33] for the period 1980-2009. Biases in COL locations and associated features (e.g., maximum intensity and precipitation) are assessed by comparing model outputs with the reanalysis data for the historical period. These biases are calculated for each model and ensemble individually as well as for the ensemble multi-model mean. Projected changes in COL features are determined by comparing future projections (2070-2099) with historical simulations (1980-2009).

We applied an objective feature tracking algorithm, the TRACK algorithm [34,35], to identify COLs in both reanalysis data and climate model simulations. This algorithm has been used extensively in previous studies by the authors [18,20,21,23,25,26,31,36]. COLs were identified as features exceeding $1 \times 10^{-5} \text{ s}^{-1}$ in 6-hourly 250 hPa relative vorticity. Prior to tracking, data were spectrally filtered to T42 resolution to remove small-scale noise (total wavenumbers >42) and planetary scales (total wavenumbers <5), emphasizing the synoptic scale while excluding mesoscale storm systems. The tracking is performed on the sphere by initializing tracks using nearest-neighboring feature points and refining them with a cost function for smoothness with adaptive constraints appropriate for the typical COL motion [20]. A post-tracking filter is applied to ensure the presence of a cut-off circulation by analyzing the 250-hPa horizontal wind components at four offset points around the COL center [20]. Tracks with at least 60% of their points within the region of interest (20°-50°N and 140°-60°W) and a minimum lifetime of one day were included in the analysis. Seasonal and yearly spatial statistics are computed using spherical kernel estimators to determine track density and mean intensity based on T42 250-hPa vorticity [37].

To identify COL-associated precipitation, a radius of influence of 5 geodesic degrees from the vorticity center is used, capturing precipitation within this radius at each 6-hourly timestep for both reanalysis and model data. We calculated precipitation from different radii (5°, 10° and 15°) and found that the rate increases with reduced radius. Two approaches are applied to compute the precipitation associated with COLs: 1) maximum precipitation (mm/h), which is calculated at the time of maximum precipitation accumulated over a 6-hour period; and 2) cumulative precipitation (mm/event), obtained by summing all precipitation values between the first and last track points of each COL system. Extreme precipitation events, defined as those exceeding the 95th percentile of the

historical period (1980–2009) for each ensemble model and reanalysis, are analyzed and categorized as extreme COLs.

The composite structure of COLs is examined using all storms using a system-centered composite method described by Bengtsson et al. [38,39] and [36]. Atmospheric fields such as horizontal winds, vorticity, and precipitation associated with each COL are sampled onto a 15° longitude-latitude rectangular grid with a 0.5° resolution, initially setup on the equator and then rotated to be centered on the COL's maximum 250-hPa relative vorticity. This grid setup ensures a fair spatial comparison across different latitudes. Lifecycle composites of precipitation are produced by averaging precipitation around the COL center at various lifecycle stages, centered on the time of maximum intensity. Model biases are calculated by subtracting the values corresponding to the reanalysis from those corresponding to the models across the 15° longitude-latitude domain for each atmospheric field.

3. Results

3.1. Northern Hemisphere Perspective of Cut-Off Lows

Building confidence in future climate projections requires that models accurately represent synoptic-scale features with sufficient fidelity. To address this, we begin by summarizing the findings of [25], who assessed the performance of the CMIP5 and CMIP6 historical simulations in simulating COLs. Recent work by [25] shows that CMIP6 models generally simulate COL characteristics more accurately than CMIP5 models when contrasted with reanalyses. Figure 1 exhibits this comparison, showing the annual COL track density of the ERA5 reanalysis (Figure 1a) compared to the multi-model means for CMIP5 (Figure 1b) and CMIP6 (Figure 1c). These comparisons use the same period (1979–2014) and model set (46 models for each CMIP phase) as in [25]. Stippling indicates systematic biases, where all models in each CMIP phase exhibit biases of the same sign.

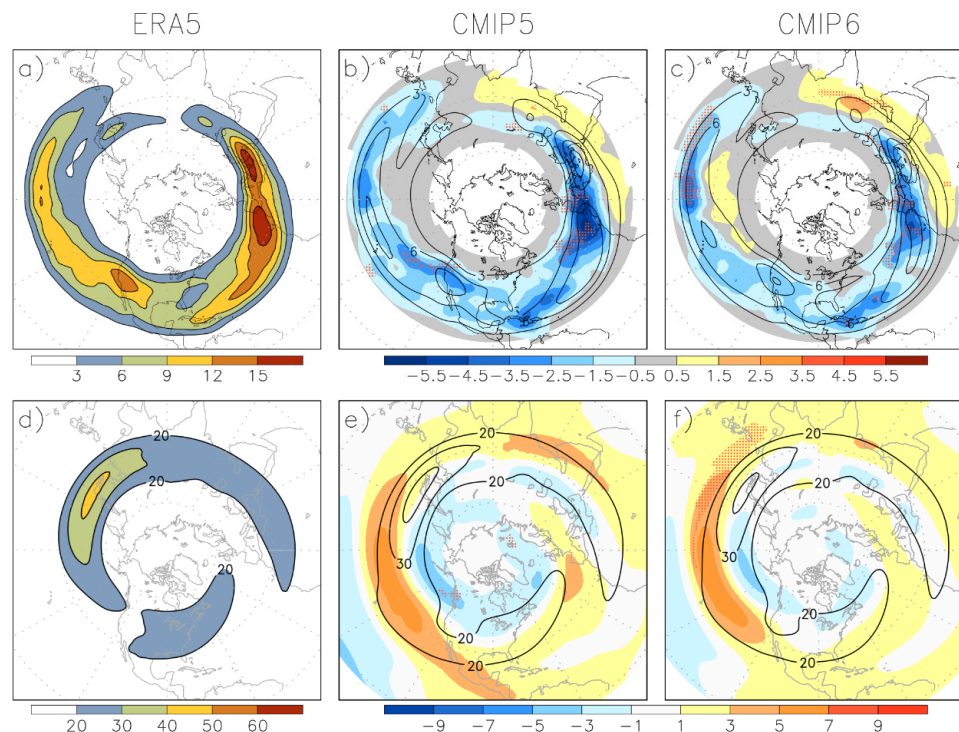


Figure 1. Top panels: Track density of 250-hPa relative vorticity averaged across all seasons: (a) ERA5 and multi-model mean (contours) with bias (shaded) for (b) CMIP5 and (c) CMIP6. Bottom panels: Zonal winds at 250 hPa averaged across all seasons: (d) ERA5 and multi-model mean (contours) with bias (shaded) for (e) CMIP5 and (f) CMIP6. Biases are relative to ERA5 for the period 1980–2009. Units:

m s^{-1} for zonal winds and number per season per unit area (5° spherical cap) for track density. Adapted from [25].

The Northern Hemisphere COL track density shows similar bias patterns in both CMIP5 and CMIP6 models. However, a significant improvement is evident in CMIP6, where the magnitude of mean biases is considerably lower than in CMIP5, indicating enhanced model performance. The largest track density biases occur in regions of high COL activity, where both CMIP5 and CMIP6 tend to underestimate COL frequency. This underestimation is more significant in CMIP5, whereas CMIP6 models show a reduced bias, suggesting that the CMIP6 better captures COL frequency in these critical regions.

Positive biases, indicating an overestimation of COLs, are more prevalent at lower latitudes, particularly from North Africa across the Middle East and into India. This reflects a persistent equatorward displacement of COL tracks in climate models — a systematic bias that persists in CMIP6. The underestimation of COL track density from the North Pacific into western North America, and from the North Atlantic into central Europe is likely related to a dipole bias characterized by overly strong winds on the equatorward side and weaker winds on the poleward side of the climatological jet stream. This pattern, shown in Figures 1e and 1f and demonstrated in [25], leads to a significant underestimation of COL track density in these regions. While the bias in 250-hPa zonal winds is a persistent challenge across climate models, CMIP6 models show a notable reduction in this bias compared to CMIP5, reflecting an improved representation of the jet stream and the associated COL distribution in the latest model generation.

3.2. Cut-OFF low in Western North America

3.2.1. Historical Climate Simulations

To provide a more detailed analysis of the regional climate impacts linked to changes in COLs, this study investigates the ability of CMIP6 models to capture their spatial distribution and intensity over western North America (20° - 50°N and 140° - 60°W). The annual track density and mean intensity of COLs (measured by the T42 vorticity at 250 hPa) in CMIP6 historical simulations (Figure 2b) are compared against ERA5 reanalysis (Figure 2a) for the period 1980–2009, using eight CMIP6 models (Table 1). The smaller number of models compared to [25] is due the limited availability of models for future projection scenarios.

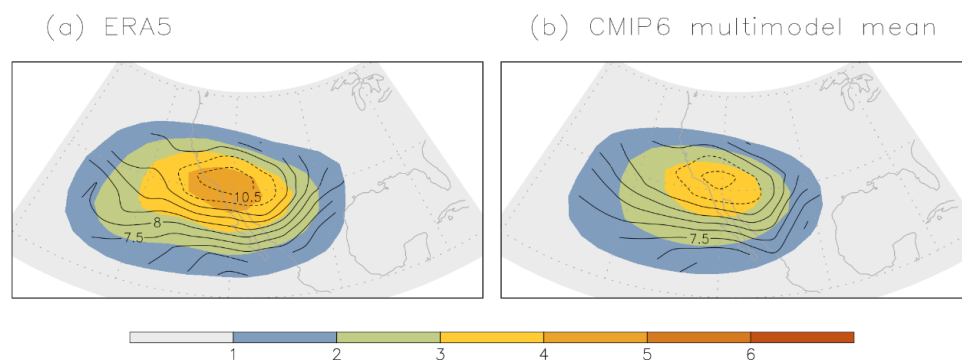


Figure 2. Track density (shaded) and mean intensity (contour) of T42 250-hPa relative vorticity, averaged across all seasons for (a) ERA5 and (b) CMIP6 multi-model mean. Tracks are identified in the region within 20° - 50°N and 140° - 60°W . Units are number per season per unit area (5° spherical cap) for track density and 10^{-5} s^{-1} for mean intensity, for contour intervals 0.5 with dotted line at 10.0.

Figure 2a shows that the COL activity is concentrated along the western edge of North America, specifically over northwestern Mexico and the southwestern United States, including California, Arizona, and Nevada. The region of maximum COL intensity closely aligns with areas of highest

track density, consistent with previous studies [12,16,19,40]. While the spatial distribution of track density in the CMIP6 multi-model mean is similar to ERA5, there is a noticeable underestimation in the CMIP6 models, likely due to their coarser horizontal resolution compared to reanalysis data. As shown in Table 1, most CMIP6 models underestimate the number of COL tracks in this region, with the exception of the higher-resolution CNRM-CM6-1-HR model. Despite this underestimation, the mean bias in intensity based on T42 250-hPa relative vorticity is relatively small in the multi-model mean, which enhances confidence in the future projections of COLs.

3.2.2. Projected Response of the Cut-Off Low Activity and Jet Streams Under a Warming Scenario

In this section, we examine the projected response of western North American COLs to climate change based on CMIP6 model projections. As shown in Table 1, the overall number of COL tracks is expected to decrease slightly by the end of the 21st century, with the multi-model mean reducing from 10.6 to 10.1 tracks per season. This decline is observed in six of the eight models, with particularly notable reductions in CNRM-CM6-1-HR and MIROC-ES2L, while MPI-ESM1-2-HR and NorESM2-LM project slight increases.

Seasonal analysis provides further insight into these projected changes, highlighting distinct patterns in COL characteristics under the SSP585 scenario. Here, it is important to distinguish between track density (Figures 3a-d) – a spatial measure of track distribution – and the overall number of identified tracks. While changes in track number generally reflect shifts in track density, the two metrics are not always correlated, as shown in Tables S1 and S2.

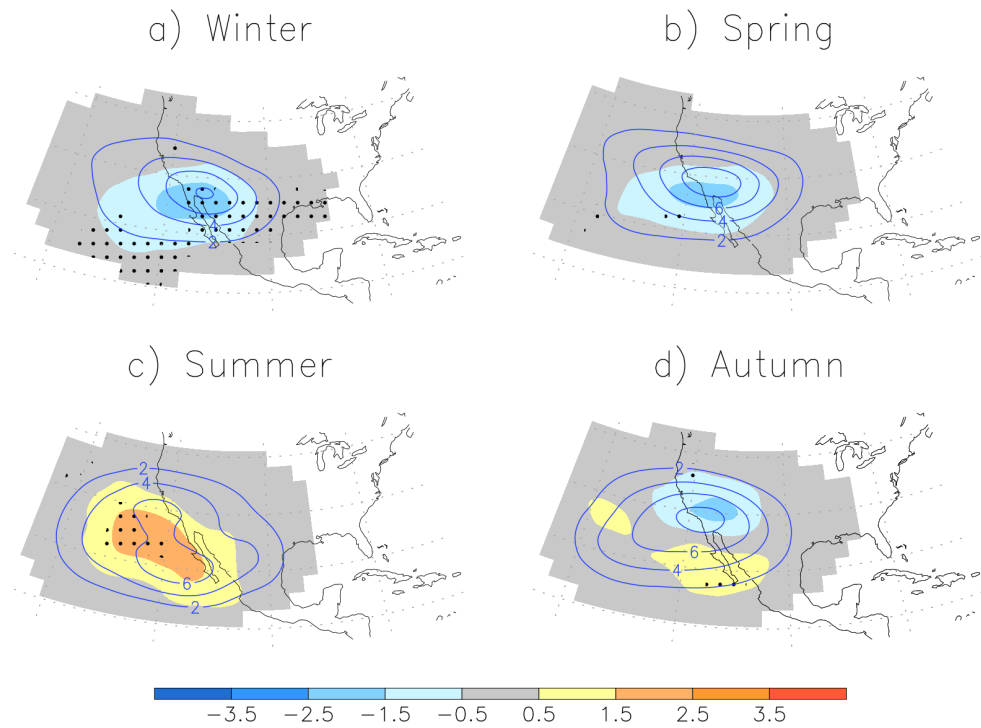


Figure 3. Seasonal projected changes in COL track density by CMIP6 multi-model mean calculated from future (SSP585) and historical climate simulations (2099-2070 minus 1980-2009). Stippling indicates where at least 80% of the models agree on the sign of change.

In summer, both track number and track density (Figure 3c) increase markedly, indicating heightened COL activity over the northeast Pacific and western United States. The multi-model mean shows a 32.2% increase in track number and an 18.4% increase in track density. This trend is robust across most models, with MPI-ESM1-2-HR showing the most substantial increases in both metrics

(69.9% in track number, 44.5% in track density). NESM3, however, presents a notable exception, with only a slight rise in track number (5.5%) and a decrease in density (-3.8%).

In winter (Figure 3a), the trend reverses, with both metrics indicating declines in COL activity. The average change in track number is -11.5%, with a corresponding reduction in track density of -16.3%. Track density is a smooth measure, sensitive to track length and location, which can sometimes lead to discrepancies between track number and density changes. For example, while MPI-ESM1-2-HR shows large declines in both track number (-24.1%) and density (-36.2%), CNRM-CM6-1-HR shows a significant decline in track number (-44.0%) but an increase in track density (10.9%) due to projected longer tracks in future climates (not shown).

Spring (Figure 3b) exhibits a modest decline in track number (-0.2%) and a more pronounced decrease in track density (-12.2%), though model agreement is inconsistent. NorESM2-LM, for example, projects an increase in track number (20.6%) and density (7.8%), contrasting with models like MIROC-ES2L and MRI-ESM2-0, which show substantial reductions in density (-21.1% and -25.0%, respectively).

Autumn (Figure 3d) presents a more complex picture, with average increases of 12.6% in track number and 2.1% in track density, though model projections vary significantly. While NorESM2-LM and MPI-ESM1-2-HR suggest increases in both metrics, models like MIROC-ES2L and CNRM-CM6-1-HR project reductions. An equatorward shift in COL track density in autumn, extending activity into the Baja California Peninsula, is noted in six models, marking a unique seasonal feature not observed in other periods.

Overall, CMIP6 models generally align on seasonal COL patterns, particularly regarding increased summer and decreased winter activity by the end of the 21st century. However, the inconsistencies between track number and density changes in autumn and spring highlight the variability in regional responses across models and seasons. As illustrated in the CNRM-CM6-1-HR winter projection, longer tracks can result in higher track densities, and vice versa. This model, for instance, projects a decrease in track number but a positive change in track density due to longer tracks in future climates (mean life time of 4.0 days) compared to present-day simulations (mean life time of 3.3 days).

Changes in the location and intensity of jet streams have important implications for COL characteristics. Previous studies have examined the representation of the northern and southern hemisphere jet streams and their impact on biases in simulated COLs using CMIP5 and CMIP6 historical simulations [25,26]. Notably, positive (negative) mean COL track density biases tend to correspond with negative (positive) upper-tropospheric zonal wind biases, highlighting the role of jet streams in modulating COL behavior.

Figure 4 shows the seasonal response of 250-hPa zonal wind speeds in the SSP585 scenario. CMIP6 models project a notable intensification in the exit region of the wintertime North Pacific jet (Figure 4a), a consistent feature across most models. The upper-tropospheric zonal mean flow accelerates over both western and eastern North America, with the eastern jet showing an equatorward shift, resembling an El Niño-like pattern [41]. The winter jet stream strengthening is primarily associated with a reduced COL track density (see Figure 3a), supported by a significant negative correlation (-0.63) between the two variables. The CMIP6 multi-model projections exhibit a comparable response in spring (Figure 4b), although with weaker wind speed changes than in winter. The spatial response of jet streams across the North Pacific and North Atlantic is similar to that projected by CMIP6 for the SSP245 scenario and by CMIP5 for RCP4.5 [42].

In contrast, the summertime jet stream (Figure 4c) demonstrates a robust weakening in maximum wind speeds, coupled with strengthening to the south and north. In autumn (Figure 4d), the North America jet exhibits a spatial response pattern similar to summer, though the southern jet flank shows a less pronounced strengthening. Unlike winter and spring, the summer weakening of the North America jet correlates with increased COL activity in western North America, though without a statistically significant correlation. Notably, the equatorward shift in the main COL track density in autumn (see Figure 3d) seems to result from weakening along the southern flank of the

North America jet. The peak magnitude of jet stream changes reaches approximately 25% (40%) of the mean fields in winter (summer) in regions with the highest COL activity.

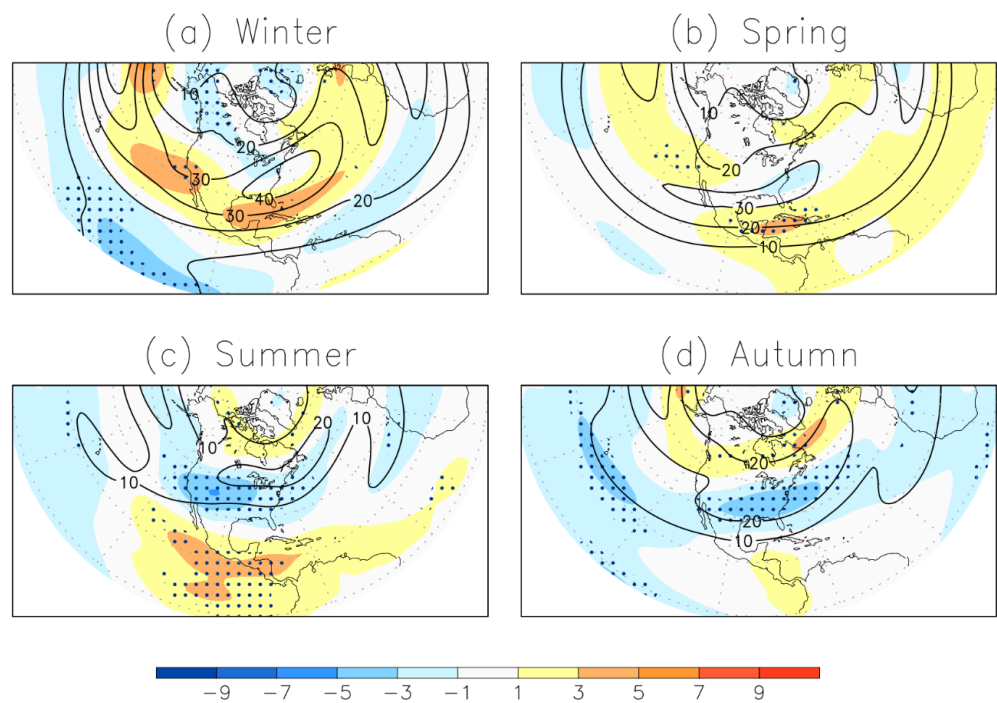


Figure 4. Seasonal projected changes in 250-hPa zonal winds in CMIP6 multi-model mean calculated from future (SSP585) and historical climate simulations (2099-2070 minus 1980-2009). Contours show the multi-model mean historical simulations for contour intervals 10 ms-1. Stippling indicates where at least 80% of the models agree on the sign of change.

3.2.3. Cut-Off Low Related Precipitation Features

Given that the frequency and location of COLs are likely to change under future climate conditions, understanding the precipitation associated with these systems is critical for accurately projecting their future impacts. We begin by assessing how well CMIP6 models reproduce the observed COL-related precipitation, which is essential for evaluating their reliability in future projections. Figure 5 shows the distribution of COL-associated precipitation in western North America, measured by both maximum precipitation rate and accumulated precipitation (see Data and Methods), with results from historical and future simulations compared against ERA5 reanalysis.

In general, the CMIP6 models capture the maximum precipitation distribution in the historical simulations, though they tend to include a relatively large number of strongly precipitation systems (i.e., the upper end of the distribution). The reason for this overestimation is unclear but may be linked to biases in the representation of latent heating, as noted in previous studies of climate models in extratropical regions [43,44]. However, the accumulated precipitation matches closely between the reanalysis and the models, indicating that these models adequately simulate the overall precipitation lifecycle of COLs. This consistency provides confidence in the model simulations, even though their horizontal resolution is coarser than that of the reanalysis data.

In future projections, the overall distribution of COL-associated precipitation remains largely similar to the historical period. However, significant changes are expected in the frequency of extreme precipitation events, which are projected to increase under warmer climate conditions. The most intense precipitation events—often the most damaging—show marked increases, despite the uncertainty indicated by the model spread (red dashed lines). Specifically, precipitation events exceeding the 95th percentile of the historical period (hereafter referred to as extreme COLs), as

shown in the inset of Figure 5a, nearly double by the end of the century, with six out of eight models projecting increases of at least 50%.

A similar trend is observed for cumulative precipitation, which accounts for rainfall accumulated over the entire COL lifecycle. Projections suggest an increase in extreme COLs by the end of the 21st century, with all models, except for ACCESS-CM2, showing this trend. On average, there is a 45% increase in the number of extreme precipitation systems, despite a reduction in the total number of COLs (see Table 1). This intensification is possibly influenced by both dynamic and thermodynamic factors, which will be explored with more detail in the following sections of this paper.

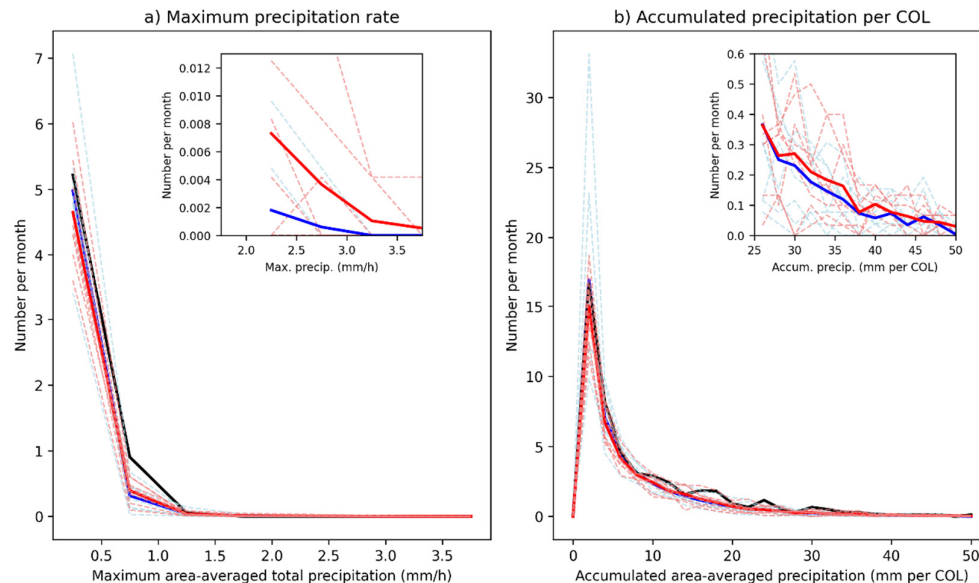


Figure 5. Monthly distribution of COL-associated precipitation in western North America for a) maximum precipitation rate (mm/h) and accumulated precipitation (mm per COL event). Blue and red solid lines indicate the ensemble model means of historical simulations (1980-2009) and future projections (2070-2099), respectively; dashed lines represent each model; and black line represents ERA5 reanalysis. The inset shows the tails of distribution for the values above the 95th percentile.

To assess the temporal evolution of extreme precipitation events associated with COLs, we analyze trends in maximum precipitation rates exceeding the 95th percentile of the historical period. This approach focuses on the most impactful events, providing a clearer picture of regional changes in their frequency. Climate model projections from the CMIP6 consistently indicate a substantial and increasing trend in the frequency of COL-related extreme precipitation events. By the end of the century, such events are projected to nearly double compared to historical levels, according to the CMIP6 mean. This trend is evident in six out of eight climate models, with BCC-CSM2-MR and CNRM-CM6-1-HR being exceptions. The pronounced increase in extreme precipitation under a high-emission pathway highlights the growing influence of warming on regional rainfall extremes.

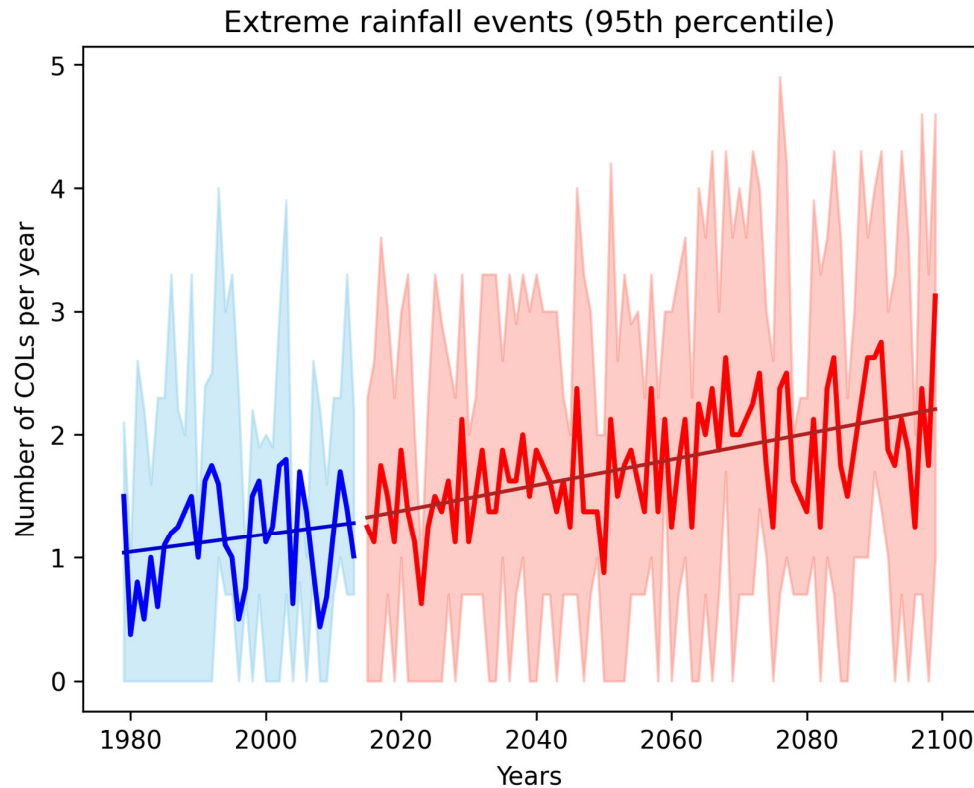


Figure 6. Annual count of western North America COLs with maximum precipitation rate exceeding the 95th percentile corresponding from the historical period. Solid lines represent multi-model means, with shading indicating model spread. Dark lines depict linear regression trends. Blue corresponds to the historical period (1980–2009), while red represents the SSP585 scenario (2014–2099).

3.2.4. Changes in Cut-Off Low Structure

We now examine the spatial structure of COLs in western North America by compositing COLs for each season and model, generating a multi-model mean centered at the time of maximum intensity, based on T42 250-hPa vorticity. Figure 7 shows the comparison of 250-hPa horizontal winds and total precipitation for ERA5 and CMIP6 historical simulations. While the wind patterns are consistent between ERA5 and CMIP6, the models slightly underestimate the wind magnitude. However, the CMIP6 models accurately capture the asymmetric horizontal wind distribution around the COLs, with the strongest winds concentrated in the southern quadrant, and successfully reproduce the maximum at 5° to 8° radius from the COL center.

In ERA5, COL-associated precipitation is concentrated primarily ahead of the COL center within a 5° radius, where strong ascent and precipitation typically occur in COLs [36]. The largest precipitation occurs in winter and spring (Figures 7a and 7b), while a secondary precipitation zone is observed further south in summer and autumn (Figures 7e and 7f), which may be associated with convective storms linked to the North American monsoon [45]. The composites for CMIP6 shows that the total precipitation biases differ spatially (Figures 7c, 7d, 7g and 7h). There is an underestimation near and east of the COL center, while overestimations are found in the surrounding regions, particularly to the east and south. In summer and autumn, the CMIP6 models struggle to reproduce the observed precipitation patterns near the region of maximum precipitation in ERA5. These discrepancies may be due to differences in vertical motion and the release of latent heat, especially in the eastern region of the COL where substantial precipitation occurs. None of the CMIP6 models accurately capture the summer and autumn precipitation in this critical area (not shown).

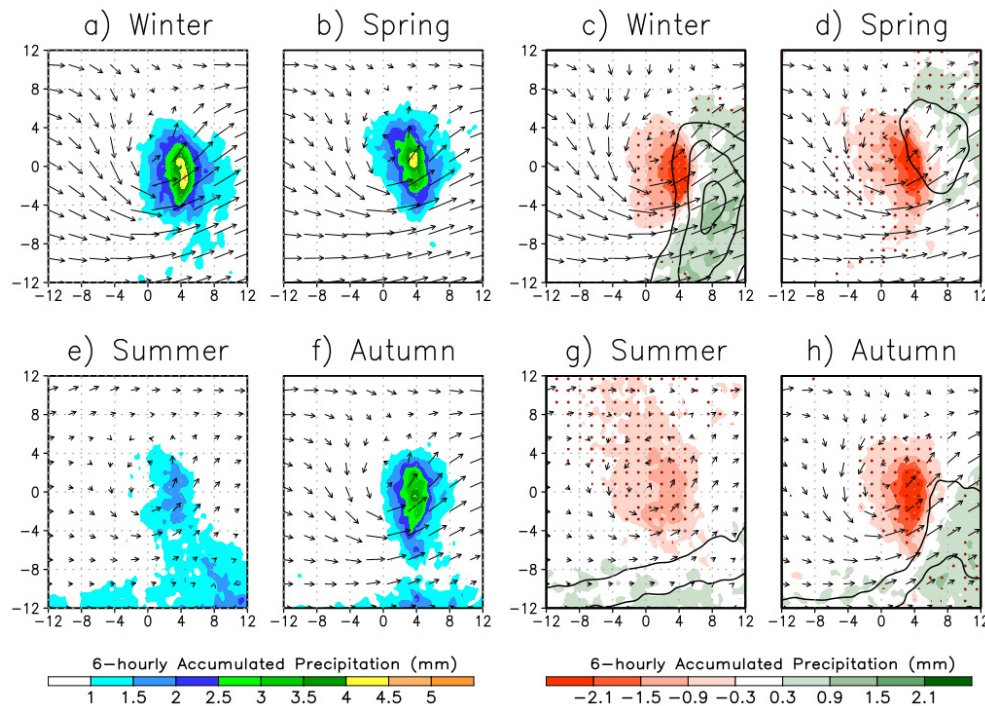


Figure 7. Composites of 250 hPa horizontal wind vectors and total precipitation (mm/6h) during COL events over western North America for Winter (a, c), Spring (b, d), Summer (e, g), and Autumn (f, h). Panels (a, b, e, f) show ERA5 reanalysis, while panels (c, d, g, h) display the CMIP6 multi-model mean for the historical period (1980–2009). Biases in CMIP6 multi-model mean relative to ERA5 are shaded, with precipitation contours at intervals of 0.5 mm starting at 1.0 mm. Stippling in (c), (d), (g), and (h) indicates areas where 80% of models agree on the bias sign compared to ERA5.

We now examine the future response of COLs in terms of 250-hPa relative vorticity and precipitation, as shown in Figure 8. Projected changes exhibit clear seasonal and spatial variations. Precipitation is primarily expected to increase in the eastern sectors of COLs, where ascent is strongest, particularly during winter and spring. In winter, precipitation is projected to rise by over 0.7 mm/6h (~50–70%) in the eastern areas, while decreases are projected in the southern and western regions. Robust model agreement in these areas of increased precipitation supports these projections, despite the overall decline in COL frequency, as shown in Table 1 and Figures 3. Although stronger intensities near the COL center are expected in winter and spring, model agreement on this intensification is lower for winter.

Autumn is the only season where all models consistently project an overall increase in precipitation across the composite domain, with the largest increases occurring in ACCESS-CM2 (+42.9%), NorESM2-LM (+35.5%), and MRI-ESM2-0 (+21.6%). This widespread increase, particularly in the southern sector of COLs, along with a decrease in COL intensity, suggests a potential link to changes in large-scale atmospheric circulation. The projected equatorward shift of COLs in autumn (Figure 3) could strengthen the tropical moisture source, enhancing latent heat release, which in turn induces an anticyclonic potential vorticity anomaly aloft, weakening the upper-level cyclonic circulation [46,47]. Thermodynamic factors likely affect these precipitation changes, reinforcing the role of diabatic processes in weakening COLs [23,48].

The precipitation changes in summer are weaker and more spatially localized. There is a slight weakening of the COL center, with minor increases in precipitation in the east and south quadrants. The changes in vorticity and precipitation during summer are not as pronounced as those observed

in other seasons, potentially due to weaker intensities typical of summer. This may contribute to the relatively low model agreement.

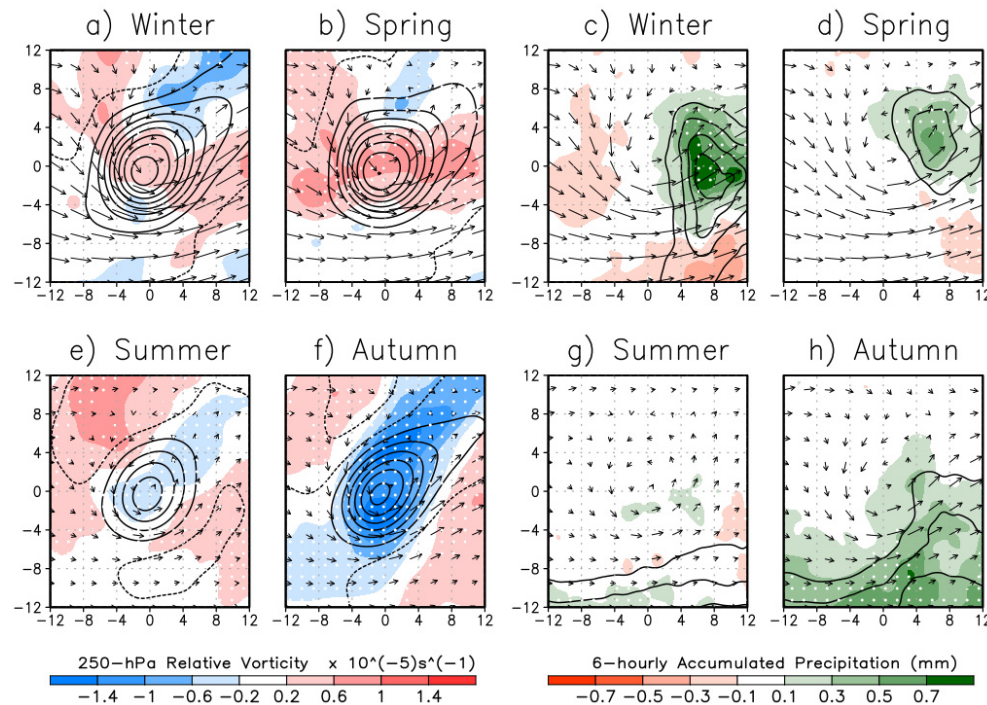


Figure 8. CMIP6 multi-model mean seasonal composites of western North America COLs for the response to future warming (SSP585 minus historical simulations). Panels (a, b, e, f) are 250-hPa horizontal wind vectors and relative vorticity ($\times 10^{-5} \text{ s}^{-1}$), with contours showing the mean and shading the change (contour interval: $1 \times 10^{-5} \text{ s}^{-1}$). Panels (c, d, g, h) are 250-hPa horizontal wind vectors and total precipitation (mm/6h), with multi-model mean values as contours (interval: 0.5 mm, starting at 1.0 mm) and changes as shaded areas. Stippling in (c, d, g, h) indicates regions where 80% of models agree on the sign of change.

3.2.5. Effect of Climate Warming on the Extreme Precipitation in Cut-Off Lows

As shown in Figure 7, the CMIP6 models tend to underestimate precipitation within the core of high precipitation areas while overestimating it along the southern flanks of the COL. The ERA5 reanalysis for extreme COL events (95th percentile) (Figure 9a) indicates that the most intense precipitation occurs to the east and northeast of the COL center, a pattern similar to those shown in the seasonal analysis but with higher values. Precipitation biases in CMIP6 historical simulations relative to ERA5 (Figure 9b) are concentrated in these regions, indicating greater uncertainties in the representation of extreme COL precipitation. The pronounced underrepresentation of extreme precipitation could arise from limitations in physical parameterizations, coarse model resolution, or insufficient ensemble size, which is useful to reduce uncertainties and provide an accurate ensemble mean [49,50].

In future projections (Figure 9c), extreme precipitation shifts farther from the COL center, particularly to the east and south, which might be related to enhanced moisture inflow from lower latitudes. Model agreement on the projected increase in extreme precipitation provides reasonable confidence, though uncertainties remain regarding the magnitude due to persistent biases in extreme event representation.

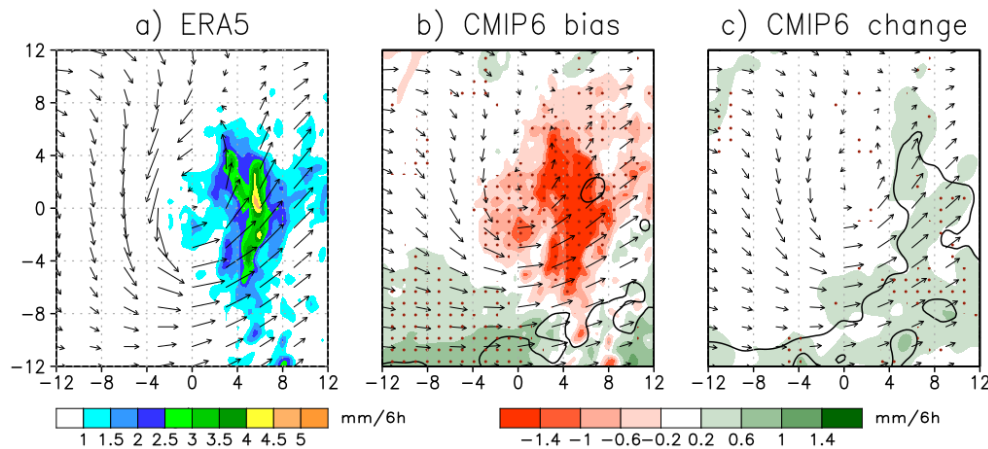


Figure 9. Composites of 250-hPa horizontal wind vectors and total precipitation (shaded, mm/6h) in extreme COLs in western North America (95th percentile) for (a) ERA5, (b) CMIP6 bias, and (c) CMIP6 change (SSP585 minus historical simulations). Contours in (b, c) indicate the total precipitation (mm/6h, interval: 0.5 mm, starting at 1.0 mm) for historical simulations and future projections, respectively. Biases of CMIP6 multi-model mean are relative to ERA5. Stippling indicates areas where 80% of models agree with the sign of bias compared to ERA5.

We have demonstrated that CMIP6 models effectively capture precipitation patterns in COLs in the western North America, despite some systematic biases. We now examine potential changes in lifecycle precipitation under a warmer climate, focusing on extreme COL events (95th percentile) to understand the behavior of extreme precipitation. Each COL is centered on the time it reaches peak intensity, based on the T42 250-hPa vorticity, using historical simulations and future projections. Figure 10a shows minimal variation in vorticity for extreme COLs, with only a slight increase in peak intensity in future projections. However, when considering all COL systems (not shown), a more pronounced increase in intensity is observed under future climate conditions.

Figure 10b indicates that, consistent with previous studies on Southern Hemisphere COLs [7,36,51], maximum precipitation occurs after peak intensity. In future projections, the rate of precipitation increase is more pronounced than in the historical period, with an average rise of about 44% across all models during the decay stage. Notably, more precipitation is projected for systems with shorter lifetimes in a warmer climate, suggesting a shift toward more intense, short-duration precipitation events, potentially heightening weather-related hazards. However, the mechanisms driving this rapid increase in precipitation remains unclear.

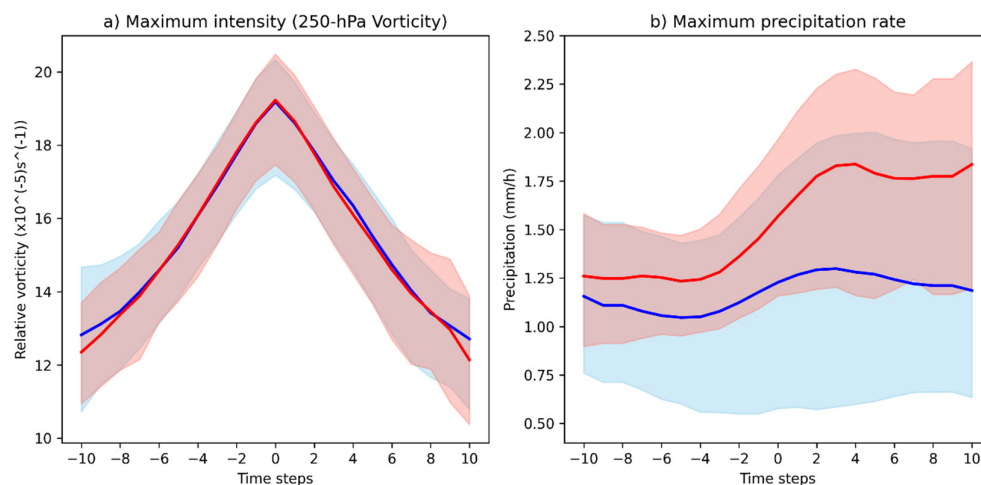


Figure 10. Composite lifecycle of western North America COL-associated precipitation for CMIP6 models. Precipitation values are averaged within a 5° geodesic radius from the COL center, relative to the time of maximum intensity in the T42 250-hPa relative vorticity. Blue and red lines represent the ensemble mean precipitation for historical simulations (1980-2009) and future projections (2070-2099), respectively, with the shaded areas indicating the spread among models in each period.

4. Discussion and Conclusions

This study has evaluated the capabilities of CMIP6 models in simulating the characteristics and impacts of COLs in western North America, a region vulnerable to COLs and changes in precipitation in the future. These results corroborate previous findings that CMIP6 models offer a more realistic representation of COL features compared to CMIP5 models, particularly in terms of track density and intensity. While systematic biases persist—most notably in the equatorward displacement of COL tracks—CMIP6 models exhibit improved performance, especially in capturing the frequency of COLs in regions of high activity.

In terms of COL structure, CMIP6 models accurately capture the asymmetric horizontal wind distribution around COL systems, though they show limitations in representing precipitation patterns. Specifically, precipitation is underestimated near and east of the COL center and overestimated in surrounding regions. CMIP6 models also tend to produce proportionally stronger precipitation systems compared to ERA5. These discrepancies may be a result of poor representation of vertical motion and latent heat release, which could be improved through higher-resolution simulations, especially with convection-permitting models.

Seasonal responses of COLs to climate change reveal distinct patterns. A decline in COL frequency is projected for winter, while an increase is expected in summer. These changes are suggested to correlate with shifts in the upper-tropospheric jet streams, where intensification in winter is linked to reduced COL activity, whereas weakening in summer is associated with increased COL activity. Winter precipitation, particularly in eastern areas, is projected to rise by approximately 50-70%. Autumn also shows consistent increases, especially in the southern sector of COLs, despite a decline in vorticity intensity. This suggests a potential link to changes in large-scale atmospheric circulation, with the equatorward shift of COLs enhancing tropical moisture source. This shift could induce an anticyclonic potential vorticity anomaly aloft, reinforcing the role of diabatic processes in weakening COLs.

Although the overall number of COL tracks is projected to decrease slightly under future climate conditions, associated precipitation is expected to increase. Extreme precipitation events (those exceeding the 95th percentile) are projected to nearly double by the end of the century, with robust model agreement on this trend. The increased extreme precipitation will likely occur farther from the COL center, particularly toward the east and south, potentially due to enhanced moisture inflow from lower latitudes. A significant increase in precipitation is projected during the decay stage, with an average rise of 44%. Additionally, COLs with shorter lifetimes are expected to produce larger amounts of precipitation in a warmer climate, indicating a shift toward more intense, short-duration precipitation events, thereby increasing weather-related hazards. These findings are consistent with previous studies [27–29], which also indicate an increase in COL-related precipitation under warming scenarios. However, uncertainties remain in the representation of extreme precipitation associated with which may arise from limitations in physical parameterizations, coarse model resolution, or the insufficient size of model ensembles.

For future research, employing higher-resolution models, such as HighResMIP (~25–50 km) and kilometer scale models with horizontal resolutions of less than 10 km [52], is crucial for better representing the physical processes and minimizing uncertainties in projections of COL behavior under climate change. Furthermore, examining how variations in jet stream strength and position affect COL dynamics is essential. As we move toward the next climate assessment (CMIP7), addressing these issues is key to refining projections of COL intensity and precipitation patterns. Additional investigations are needed to clarify the implications of these projected changes for regional weather hazards and overall climate impacts.

Supplementary Materials: The following supporting information can be downloaded at the website of this paper posted on Preprints.org, Table S1: Percentage change in the number of identified 250-hPa COL tracks for each CMIP6 model across the seasons (Spring, Summer, Autumn, Winter) within the region of 20°-50°N and 140°-60°W. The percentage change is calculated as the difference between the future period (2070-2099) and the historical period (1980-2009).; Table S2: Same as Table S1 but for the change in the 250-hPa COL track density.

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Data Availability Statement: The data used in this study are sourced from the World Climate Research Programme's Working Group on Coupled Modelling and are freely available through the Earth System Grid Federation (ESGF) at <https://esgf-index1.ceda.ac.uk/search/cmip6-ceda/>. ERA5 data were obtained from the Copernicus Climate Change Service's Climate Data Store, accessible at <https://cds.climate.copernicus.eu/>, accessed on 04 November 2024.

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