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## Article

# Developing a New ANN Model to Estimate Daily Actual Evapotranspiration Using Limited Climatic Data and Remote Sensing Techniques

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**Abstract.** Accurate estimations of actual evapotranspiration (ET<sub>a</sub>) are fundamental for various environmental issues. Artificial Intelligence-based models are considered a promising alternative to the most common direct ET<sub>a</sub> estimation techniques and indirect methods by remote sensing (RS)-based surface energy balance models. Artificial Neural Networks (ANN) are proven to be suitable for predicting reference evapotranspiration and ET<sub>a</sub> based on RS data. The aim of this study is to develop a methodology based on ANNs for estimating daily ET<sub>a</sub> values using NDVI and land surface temperature, coupled with limited site-specific climatic variables at a large irrigation catchment. Two scenarios were implemented by the ANN model. Data from only the 38 days of satellite overpass dates was selected in Scenario-I, while the 769-day data included the satellite overpass dates and other days located between two satellite overpass acquisitions in Scenario-II. An irrigation scheme in the Mediterranean region of Turkiye was selected, and a total of 38 Landsat images, and local climatic data were used in 2021 and 2022. Results showed that R<sup>2</sup> values in Scenario-I and II were acknowledgeably high for training (0.79 and 0.86), testing (0.75 and 0.81), and the entire dataset (0.76 and 0.84), respectively. Results of the new model in two scenarios showed acceptable agreement with ET<sub>a</sub>-METRIC values. The proposed ANN model demonstrated the potential of obtaining daily ET<sub>a</sub> using limited climatic data and RS imagery.

**Keywords:** Evapotranspiration; artificial neural networks (ANNs); remote sensing (RS); METRIC; NDVI; Akarsu Irrigation District

## 1. Introduction

Actual evapotranspiration (ET<sub>a</sub>) measurements in the field or estimation by any respective conventional method (Kumar et al. 2011) are needed in irrigation scheduling at large-scale irrigation schemes as well as in a farmyard, water balance works in hydrologic studies, the accurate design of hydraulic structures related to irrigation schemes, etc. By introducing a two-step procedure (Cetin 2020), ET<sub>a</sub>, in other words, crop evapotranspiration ET<sub>c</sub>, is estimated under standard conditions (Allen et al. 1998), i.e., well-watered and optimal agronomic conditions. On the other hand, ET<sub>a</sub> is an essential component of the water budget, hydrological modeling, and irrigation water management in arid and semi-arid regions of the world. Nevertheless, it is not an easy task to acquire a representative ET<sub>a</sub> value for it can be determined directly using either a lysimeter or water balance work/approach which are rather labor-intensive, time-consuming, in turn, very expensive. As stated clearly by Allen et al. (1998), Rawatet al. (2017), Gharbia et al. (2018), and Alsenjar et al. (2023a), among others, ET<sub>a</sub> describes the physical processes of the amount of water that can occur either evaporation or transpiration according to climatic conditions, crop types, and soil status. Since direct methods of ET<sub>a</sub> calculation focus only on one point or parcel, in turn, they have limitations in showing the variation of ET<sub>a</sub> spatially and temporally at a large scale (Zhang et al. 2011), particularly in areas with large irrigation schemes of >100 000 ha as in Turkiye. Therefore, accurately estimating ET<sub>a</sub> is of utmost importance as well as a critical issue in water balance methods and agricultural water management at the irrigation scheme level. In this regard, the remote sensing (RS)-based surface energy balance models are one of the indirect methods to estimate spatially ET<sub>a</sub> over large-

scale irrigation catchments with a high spatial and temporal resolution (Alsenjar et al. 2023b; Alsenjar and Cetin 2023c; Cetin et al. 2023a, 2023b).

Several researchers have estimated ET<sub>a</sub> based on the RS-based surface energy balance models such as the Surface Energy Balance Algorithm for Land (SEBAL) (Bastiaanssen et al. 1998a, 1998b), Mapping EvapoTranspiration at high Resolution and with Internalized Calibration, i.e., METRIC (Allen et al. 2007a, 2007b), and Surface Energy Balance System (SEBS) (Su 2002). These models normally use Landsat satellite Imagery or Moderate Resolution Imaging Spectroradiometer (MODIS) which include more required RS data alongside weather parameters (Allen et al. 2007a; Alsenjar et al. 2023b; Cetin et al. 2023a, 2023b), among others. However, as pointed out by Bachour et al. (2014), among others, RS technology has some limitations. Therefore, some of these data cannot be provided due to cloud cover and the unavailability of all relevant climatic data. For example, the implemented METRIC model to estimate ET<sub>a</sub> requires more of the input parameters of Landsat satellite data (net radiation flux (R<sub>n</sub>), soil heat flux (G), sensible heat flux (H), latent heat flux (LE), leaf area index (LAI), land surface temperature (LST), surface albedo ( $\alpha$ ), normalized difference vegetation index (NDVI), to name but a few).

In addition, H calculation in the METRIC model is based on the anchoring pixels, i.e., hot and cold, so there are complexities involved in sensitivity selecting anchor pixels for improved estimation of H and LE fluxes (Singh and Irmak 2011). Moreover, Landsat satellite data with 30 m by 30 m spatial resolution are normally available every 16 days for satellite scenes (Alsenjar et al. 2023b; Cetin et al. 2023a, 2023b). As such, the methodology obtained by Allen et al. (2007a) has been used to estimate ET<sub>a</sub> based on the METRIC model for each pixel during the satellite image overpass date. In this regard, to determine ET<sub>a</sub> values based on the METRIC model for a predetermined period, i.e., month or season, Allen et al. (2007a) have applied the procedure by making the interpolation of daily reference evapotranspiration fraction (ET<sub>rf</sub>) or crop coefficient (K<sub>c</sub>) values between two satellite images and multiplying by reference evapotranspiration, i.e., E<sub>To</sub>, for each day and then integrated for a specific period of time. However, the accepted method by Allen et al. (2007a) has some gaps or limitations in showing the effect of precipitation or irrigation precisely on the daily growth stages and NDVI of crops between two satellite image overpasses. As known, both ET<sub>a</sub> and E<sub>To</sub> have a nonlinear character in nature (ASCE 2000a), and in turn, is a complex phenomenon. Therefore, more nonlinearity exists in the evapotranspiration process due to its stochastic behavior (ASCE 2000b). Thus, over the past few decades, artificial neural networks (ANNs) have been successfully utilized in modelling reference evapotranspiration, i.e., E<sub>To</sub> (ASCE 2000a, 2000b; Kumar et al. 2011). As such, the idea of artificial neural networks usage in engineering applications goes back to the 1940s (ASCE 2000a), it has been widely used in hydrological practices, particularly in the estimations of reference evapotranspiration, since the early nineties. Contrary to this, the literature review revealed that there have been hardly ever studies on estimating METRIC-based actual evapotranspiration values by ANN models.

The most significant merit of the ANN models is to solve complex problems using fewer inputs by adjusting the weights to be able to predict the correct output of the input parameters. Recently, the ANN algorithm has been applied to estimate reference evapotranspiration (E<sub>To</sub>) and crop evapotranspiration (ET<sub>c</sub>) for wheat, maize, and potato, in different regions of the world (Bruton et al. 2000; Odhiambo et al. 2001; Kumar et al. 2002; Dehbozorgi and Sepaskhah 2011; Khoshhal and Mokarram 2012; Abrishami et al. 2019; Yamac and Todorovic 2020). As for ANN-based ET estimation, rarely studies used land surface parameters calculated from RS data, including but not limited to land surface temperature, vegetation indices, etc., and limited meteorological parameters as inputs in the ANN model (Virnodkar et al. 2020).

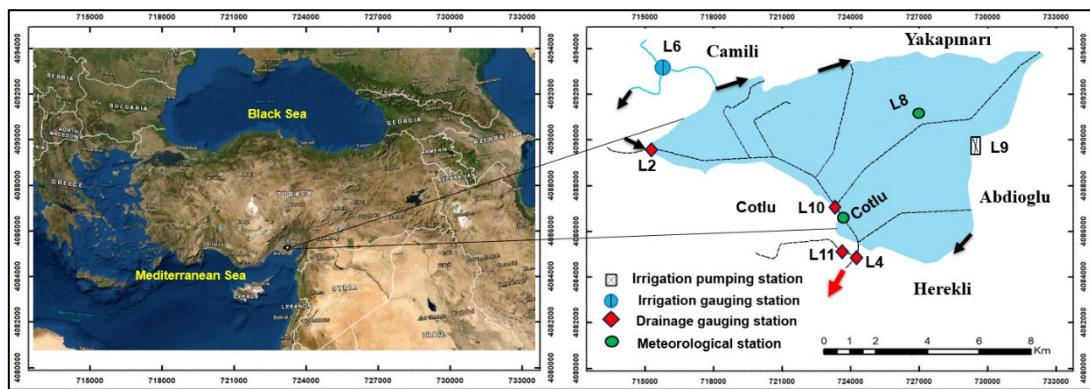
ET<sub>a</sub> estimation by the ANN model can be beneficial and powerful for using it as an input factor in water balance calculations at large-scale irrigation catchments since, as claimed by Kumar et al. (2011), theoretically, ANN is expected to produce better results than a regression model for the same data length. The novelty of this study is to establish a new methodology for generating daily actual evapotranspiration (ET<sub>a</sub>) series based on the ANN model using some of the parameters of MODIS data coupled with available daily moderate spatial and less input of weather variables as compared

to the existing methods of ETa-based surface energy balance models estimation. Therefore, this study aims to use an artificial neural network (ANN) approach to estimate daily actual evapotranspiration (ETa) values in large-scale irrigation catchments using two parameters of MODIS data, i.e., NDVI and LST, coupled with limited site-specific climatic variables at a large-scale irrigation scheme located in the Lower Seyhan Plain irrigation project area with >210 000 ha land (Cetin et al. 2023b). Furthermore, this paper is the first attempt to generate a new model using the ANN algorithm as an alternative to the existing methods of actual evapotranspiration estimation in a large-scale irrigation district in the Eastern Mediterranean Region of Turkiye. Moreover, this methodology can be generalized for estimating daily ETa using the ANN model to different climate regions and zones of the world.

## 2. Materials and Methods

### 2.1. Study area And Its Characteristics

The research area that has gained popularity under the name Akarsu Irrigation District (hereafter, AID,  $A=9495$  ha $\approx 95$  km $^2$ , Figure 1) in the studies carried out so far is located in the *Lower Seyhan Plain (LSP)* in the southeastern part of the Mediterranean region of Turkiye. The LSP shows typical characteristics of a deltaic plain with a rather flat topography (a slope of 1% or less) and a large-scale irrigation and drainage network (Alsenjar et al. 2023a, b; Cetin et al. 2023a, b). The Mediterranean climate, characterized by warm and rainy in the winter season whereas dry and hot in the summer season, prevails utterly in the LSP and, in turn, in the study area. The average annual precipitation of the basin is around 650 mm (Cetin et al. 2020). In the LSP of Turkiye, there is a very remarkable difference in temperature and evaporation values in the irrigation season, in July and August in particular, compared to those in the winter season, specifically in December, January, and February (Alsenjar et al. 2023b; Cetin et al. 2020, 2023a). As reported by Alsenjar et al. (2023b), due to meteorological and geographical factors, the definition of the water year, i.e., hydrological year, varies from region to region. It has been defined as the period, with a length of 365-day, between October 1st of one year and September 30th of the next, as late September to early October is the time for many drainage areas or catchments in Turkiye to have the lowest stream flows and consistent groundwater levels.



**Figure 1.** The study area is located in the southeastern Mediterranean region of Turkiye. Meteorological stations are located at L8 and Cotlu. Irrigation water is diverted from L6 and L9 locations into the AID; L2 and L11 stand for drainage water inputs and L4 is the drainage outlet of the catchment.

### 2.2. Remote Sensing Data Used

#### 2.2.1. Landsat Satellite Imagery

To run the METRIC model and for the actual evapotranspiration (ETa) estimations, a total of 38 clear-sky Landsat satellite images were downloaded from the USGS website

(<http://earthexplorer.usgs.gov>) (path 175, row 34) and used in this research (Table 1). These images are Landsat 7, Landsat 8, and Landsat 9 with 30 m by 30 m spatial resolution. General characteristics of the Landsat satellite images are given in Table 1. The Environment for Visualizing Images (ENVI) software program was applied to perform a cloud mask for one satellite image on May 1, 2022, i.e., DOY 121, of grayscale fill, as shown in Table 1.

**Table 1.** Availability of Landsat 7, Landsat 8, and Landsat 9 scene information in the 2021 and 2022 water years: names of scenes, acquisition dates, and overpass time.

Image	Day of the year (DOY)	Landsat scene-ID	Satellite type	Cloud cover (%)	Acquisition dates	Overpass local time (AM)
1	260	LC81750342020260L GN00	Landsat 8	1	16.09.2020	11:15:56.5028510
2	300	LE71750342020300S G100	Landsat 7	8	26.10.2020	10:38:56.1154274
3	316	LE71750342020316N PA00	Landsat 7	3	11.11.2020	10:37:51.0228172
4	364	LE71750342020364N PA00	Landsat 7	1	29.12.2020	10:34:21.8153233
5	22	LC81750342021022L GN00	Landsat 8	9	22.01.2021	11:15:49.9861710
6	54	LC81750342021054L GN00	Landsat 8	7	23.02.2021	11:15:43.2139690
7	79	LE71750342021078S G100	Landsat 7	5	19.03.2020	10:28:24.8443048
8	118	LC81750342021118L GN00	Landsat 8	8	28.04.2021	11:15:15.8809360
9	134	LC81750342021134L GN00	Landsat 8	1	14.05.2021	11:15:15.9098560
10	158	LE71750342021158S G100	Landsat 7	1	07.06.2021	10:21:43.6865663
11	182	LC81750342021182L GN00	Landsat 8	4	01.07.2021	11:15:35.1871370
12	190	LE71750342021190S G100	Landsat 7	3	09.07.2021	10:19:04.5579664
13	198	LC81750342021198L GN00	Landsat 8	5	17.07.2021	11:15:36.9021800
14	214	LC81750342021214L GN00	Landsat 8	0	02.08.2021	11:15:45.3409259
15	230	LC81750342021230L GN00	Landsat 8	2	18.08.2021	11:15:51.0643500
16	262	LC81750342021262L GN00	Landsat 8	9	19.09.2021	11:15:59.0216650
17	278	LC81750342021278L GN00	Landsat 8	1	05.10.2021	11:16:04.4435930
18	294	LC81750342021294L GN00	Landsat 8	0	21.10.2021	11:16:07.4309270
19	326	LC81750342021326L GN00	Landsat 8	6	22.11.2021	11:16:02.0432969

20	358	LC81750342021358L GN00	Landsat 8	4	24.12.202 1	11:15:59.4307040
21	001	LE71750342022001N PA00	Landsat 7	13	01.01.202 2	10:03:01.7753300
22	017	LE71750342022017N PA00	Landsat 7	3	17.01.202 2	10:01:30.1854341
23	049	LE71750342022049N PA00	Landsat 7	6	18.02.202 2	09:58:18.1083401
24	081	LE71750342022081N PA00	Landsat 7	19	22.03.202 2	09:55:10.7155745
25	089	LC81750342022089L GN00	Landsat 8	7	30.03.202 2	11:15:25.6965150
26	113	LC91750342022113L GN00	Landsat 9	2	23.04.202 2	11:15:26.7327310
27	121	LC81750342022121L GN00	Landsat 8	60	01.05.202 2	11:15:28.7371250
28	140	LE71750342022140S G100	Landsat 7	12	20.05.202 2	09:50:56.6332014
29	157	LE71750342022157S G100	Landsat 7	23	06.06.202 2	09:50:06.2292853
30	169	LC81750342022169L GN00	Landsat 8	1	18.06.202 2	11:15:53.3072380
31	186	LE71750342022186S G100	Landsat 7	0	05.07.202 2	09:42:29.8002719
32	201	LC81750342022201L GN00	Landsat 8	1	20.07.202 2	11:15:58.6410700
33	209	LC91750342022209L GN00	Landsat 9	0	28.07.202 2	11:15:42.4933480
34	220	LE71750342022220S G100	Landsat 7	29	08.08.202 2	09:39:50.3921524
35	237	LE71750342022237S G100	Landsat 7	9	25.08.202 2	09:38:18.9640686
36	249	LC81750342022249L GN00	Landsat 8	8	06.09.202 2	11:16:15.2109079
37	271	LE71750342022271S G100	Landsat 7	3	28.09.202 2	09:34:51.9536731
38	297	LC81750342022297L GN00	Landsat 8	0	24.10.202 2	11:16:18.0041120

### 2.2.2. Moderate Resolution Imaging Spectroradiometer (MODIS) Products

Two parameters of MODIS data, i.e., normalized difference vegetation index (NDVI) and land surface temperature (LST) were downloaded by the Google Earth engine ('MODIS/MOD09GA\_006\_NDVI'; 'MODIS/061/MOD11A1\_LST') to apply the artificial neural network (ANN) for estimating daily ET<sub>a</sub> values for the entire study area. Typical characteristics of MODIS data are illustrated in Table 2 along with the spatial and temporal resolutions. Daily NDVI and LST data sets were used as casual variables in the modelling practice.

**Table 2.** MODIS data used in the research area.

MODIS standard products	Parameter	Spatial resolution	Temporal resolution
MOD09GA-Terra	NDVI	500 m by 500 m	Daily

MOD11A1.061-Terra	LST	1000 m by 1000 m	Daily
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### 2.3. In-Situ Meteorological Observations

In this research, hourly and daily climatic variables (minimum and maximum temperatures ( $T_{\min}$ ,  $T_{\max}$ ), wind speed ( $U$ ), solar radiation ( $R_s$ ), minimum and maximum relative humidity values ( $RH_{\min}$  and  $RH_{\max}$ ), and precipitation ( $P$ )) acquired from two meteorological stations, i.e., L8 and Cotlu meteorological stations established in the AID, were used (Figure 1). Before using any climatic data observed in L8 and Cotlu meteorological stations, quality controls, i.e., QC (gaps in the data, outliers, constant values, jumps, etc.) were checked thoroughly. No consistency was found in the meteorological data sets. Climatic data sets cover from September 16, 2020, to October 24, 2022, i.e., 769 daily data sets.

### 2.4. Reference Evapotranspiration ( $ETo$ ) Estimation

Reference evapotranspiration ( $ETo$  in  $\text{mm day}^{-1}$  unit) is defined and computed using the FAO-Penman-Monteith approach given by Allen et al. (1998). Equation 1 was developed for short grass; the albedo was 0.23, whilst the aerodynamic resistance was  $70 \text{ s m}^{-1}$ .

$$ETo = \frac{0.408 \Delta (R_n - G) + \gamma \frac{900}{T + 273} u_2 (e_s - e_a)}{\Delta + \gamma (1 + 0.34 u_2)} \quad (1)$$

where  $ETo$  is the reference evapotranspiration ( $\text{mm day}^{-1}$ ),  $R_n$  is the net radiation at the crop surface ( $\text{MJ m}^{-2} \text{ day}^{-1}$ ),  $G$  is the soil heat flux density ( $\text{MJ m}^{-2} \text{ day}^{-1}$ ),  $T$  is the mean daily air temperature at 2 m height ( $^{\circ}\text{C}$ ),  $u_2$  is the wind speed at 2 m height ( $\text{m s}^{-1}$ ),  $e_s$  is the saturation vapor pressure ( $\text{kPa}$ ),  $e_a$  is the actual vapor pressure ( $\text{kPa}$ ),  $e_s - e_a$  is the saturation vapor pressure deficit ( $\text{kPa}$ ),  $\Delta$  slope ( $\text{kPa}^{\circ}\text{C}^{-1}$ ) is the vapor pressure curve and  $\gamma$  the psychrometric constant ( $\text{kPa}^{\circ}\text{C}^{-1}$ ).

### 2.5. METRIC Model

The METRIC model was applied to estimate a) surface energy balance components ( $SEB$ ), i.e., latent heat ( $LE$ ), net radiation ( $R_n$ ), the sensible heat ( $H$ ), and soil heat flux ( $G$ ) in Equation 2, b)  $ET_a$  for each pixel, and the whole study area by using Landsat satellite imagery and meteorological stations, i.e., L8 and Cotlu at the time of satellite overpass, primarily based on Allen et al. (2007a, 2007b) through R-METRIC model using a water package in the *R* program (Olmedo et al. 2016) and LandMOD ET mapper-MATLAB (Bhattarai et al. 2017).

$$LE = R_n - G - H \quad (2)$$

All the fluxes are in the unit of watt per meter square (i.e.,  $\text{W m}^{-2}$ ). Further information on the METRIC model and equations, i.e., step-by-step  $ET_a$  calculation, as well as the FAO-Penman-Monteith approach are given by Allen et al. (2007a).

### 2.6. Developing an ANN Model for Actual Evapotranspiration ( $ET_a$ ) Estimation

Artificial Neural Networks (ANNs) are mathematical models that resemble biological neural networks. ANNs can learn from examples and adapt solutions over time by recognizing patterns in data, along with rapidly processing information (Jain et al. 2008). In essence, ANNs are tools to mimic the underlying likely relationship between input and output variables in the hand adequately.

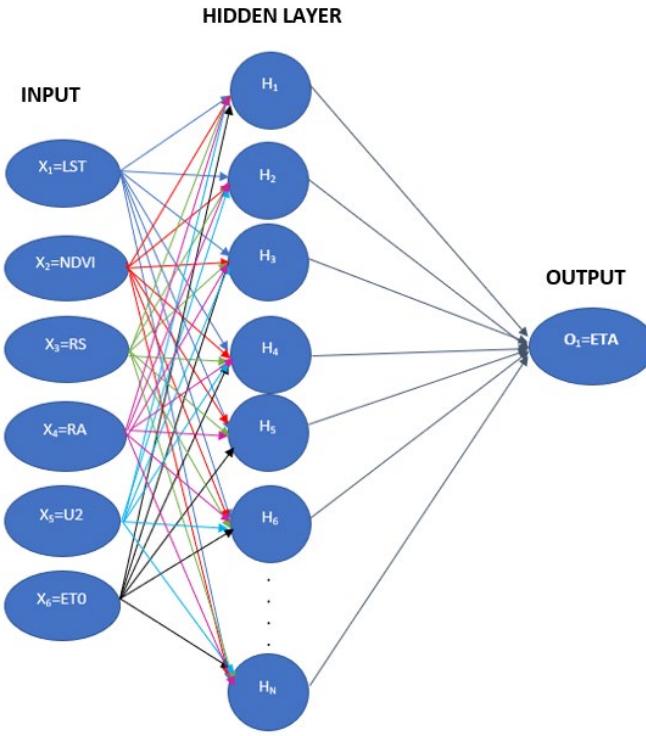
Water resources and hydrological processes are often complex, multivariable, and nonlinear. ANNs exhibit a flexible structure to address these complex relationships, making them capable of learning and integrating complex relationships by using various input data. Therefore, in recent times, ANNs have been increasingly utilized in hydrology and water resource management. ANNs might be considered as flexible modeling tools and can theoretically model any type of relationship with good accuracy. With ANNs, there is no need to make specific assumptions about the models and the underlying relationships; the underlying relationship is determined solely through data mining procedures.

This data-driven approach is one of the most significant advantages of ANNs in solving various complex real-world prediction problems. ANNs have been used in a whole range of hydrological applications, including reference evapotranspiration estimations and predicting groundwater levels (Coppola et al. 2003; Daliakopoulos et al. 2005), flood forecasting simulations (Karahan et al. 2014), rainfall and streamflow modeling (Luk et al. 2000), and aquifer parameter estimations (Garcia and Shigidi 2006; Karahan and Ayvaz 2006).

Models for calculating ETo and plant water requirements involve a myriad of variables such as meteorological data, soil properties, plant type, and climate conditions. ANNs have a significant advantage in handling these complexities due to their ability to use a large amount of data. Particularly, when trained with large datasets, these networks have a better capacity to learn complex relationships and patterns. Predicting future changes in water resources and plant water requirements due to climate change is becoming increasingly important. In ET modeling, machine learning algorithms are being used more and more as an alternative to traditional methods (Granata 2019; Tikhmarine et al. 2019). These algorithms can be used as alternatives to traditional equations for ETc and/or ETo predictions. They also provide insights into how ET behaves over time and space (Madugundu et al. 2017; Tang et al. 2018). ANNs can be trained and adapted to be used in different geographical areas, allowing for customized predictions based on different plant species and climatic conditions.

Despite their mentioned advantages, ANNs have some important disadvantages. They require a significant amount of data to learn complex relationships and to determine the optimal network architecture (Karahan and Ayvaz 2008). To create the network structure, the number of hidden layers in the model and the optimal number of neurons in each layer need to be determined. In most of the ANN studies in the literature, a trial-and-error procedure is used to determine the network architecture, which is a time-consuming process. According to Maier and Dandy (2000), in most water resources problems, using a single hidden layer is sufficient. Therefore, in this study, a single hidden layer is used in ETa estimation.

To determine the optimal number of neurons in the hidden layer, 80% of the data was used for training the network, and 20% was used for testing. This process was repeated for 100 different randomly selected training and testing datasets, and the Mean Squared Error (MSE) value was calculated. This process was repeated in a loop from 1 to the maximum number of neurons, which is 30 in this study, and the number of neurons that yielded the minimum MSE value was selected as the optimal number of neurons, and the analyses were conducted accordingly. Figure 2 illustrates the typical structure of multi-layer ANNs. The input layer consists of 6 nodes, 4 parameters acquired by the two meteorological stations (Figure 1) installed in the study area and two variables downloaded from MODIS satellite data. In situ-climatic observations are solar radiation (Rs), extraterrestrial radiation (Ra), wind speed (u2) at 2 m height, and the reference evapotranspiration values, i.e., ETo, were calculated by the standard FAO-Penman-Monteith approach through following Allen et al. (1998). The connections between the input layers and 30 hidden layers take different weights and are trained depending on the required output of daily ETa.



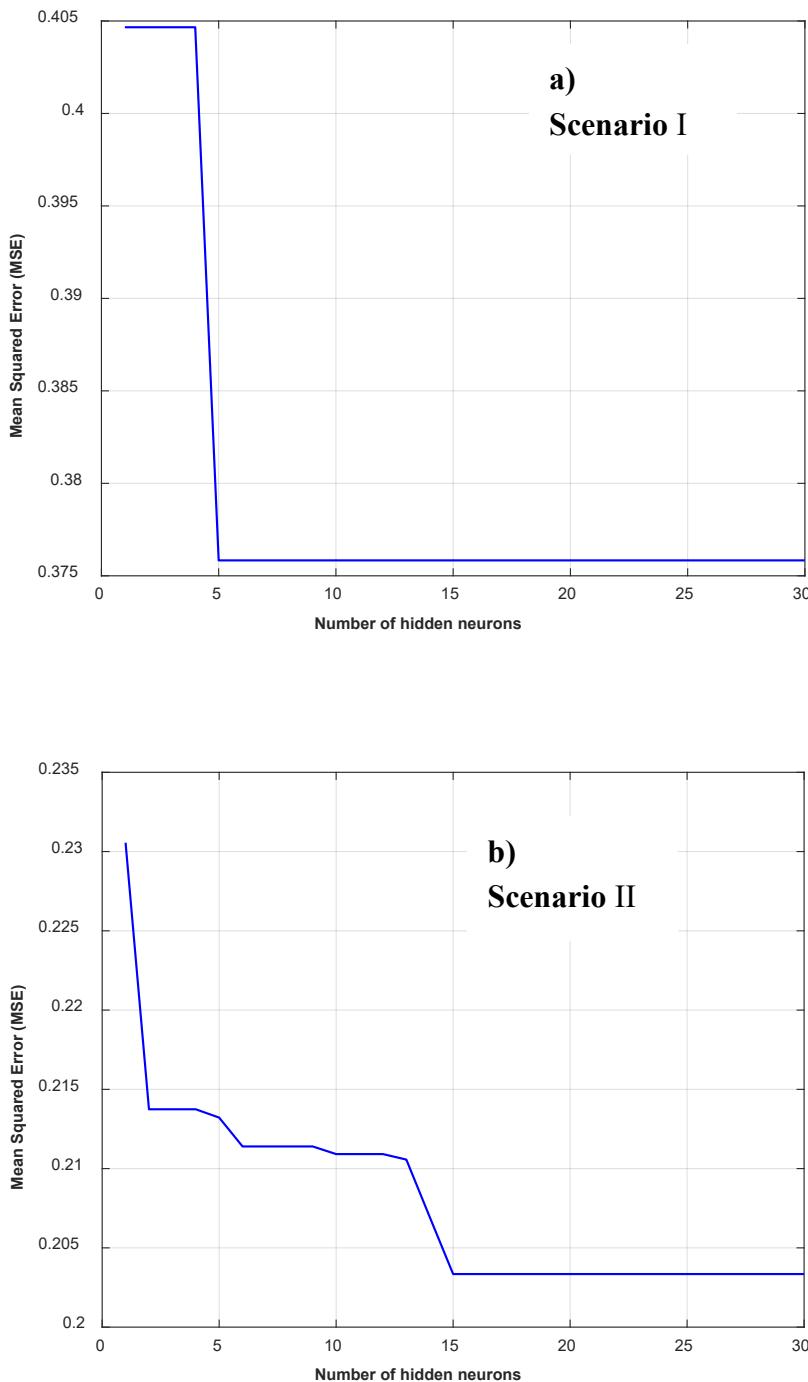
**Figure 2.** The typical structure of multi-layer ANNs used in this study.

### 3. Results

#### 3.1. Implementation of the ANN Model

The ANN model has been applied to two different scenarios. In the training of Scenario I, data from only the 38 days of satellite observations listed in Table 1 were used, while the data from the other days were used as test data. In Scenario II, 80% of the total data set was used for training, and the remaining 20% was used for testing.

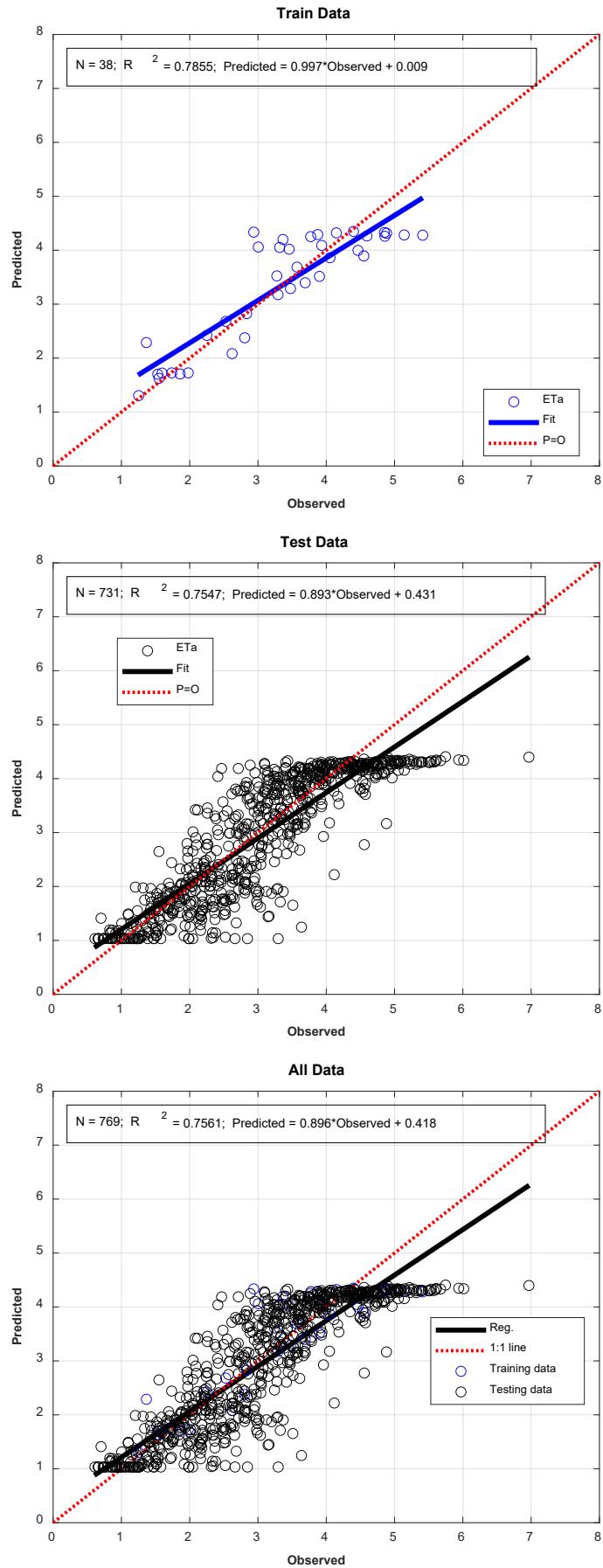
For both scenarios, firstly, the network architecture was created, and the change in the MSE value concerning the number of neurons is presented in Figures 3a and 3b. As seen from Figure 3, the optimum number of neurons obtained for Scenario I and II is 5 and 15, respectively.



**Figure 3.** A three-layer feed-forward ANN.

### 3.1.1. Scenario I

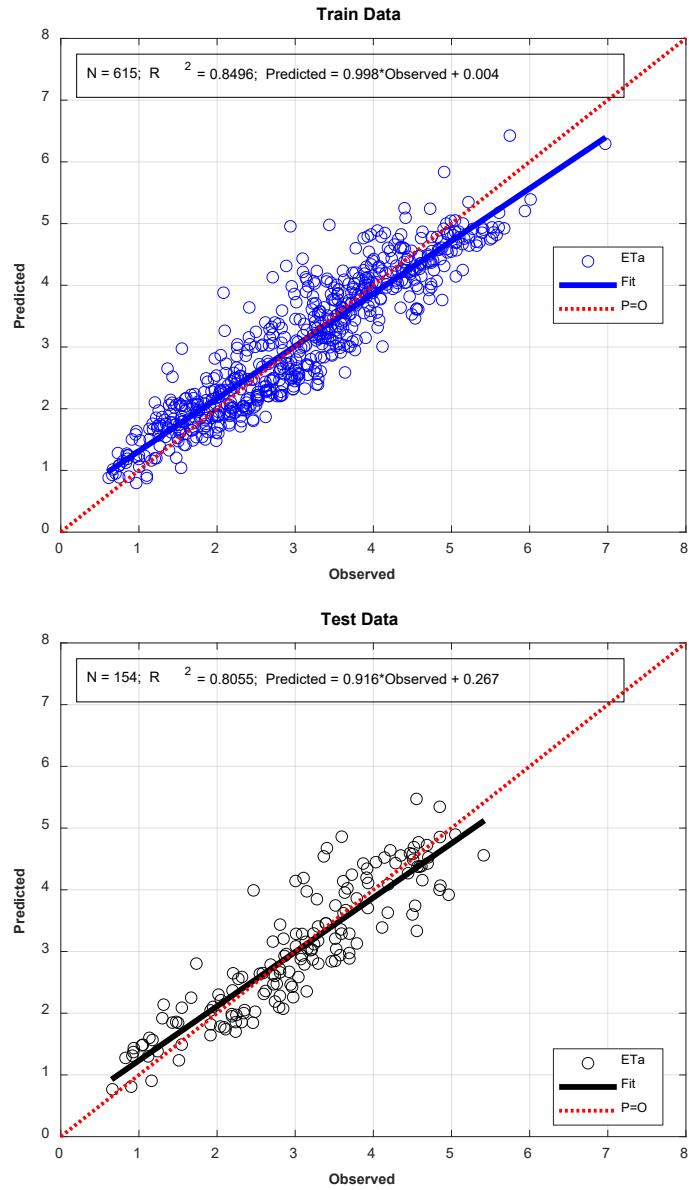
The developed model was applied to Scenario I, and the model results are summarized in Figure 4. As seen from Figure 4, even though the data used in the model training accounts for approximately 5% of the total data, the R-squared value is 0.7547 for the test data and 0.7561 for the total data. This indicates that the developed model has quite good with respect to the learning stage and prediction ability; however, it also shows an increased error rate in predicting high ET<sub>a</sub> values. This result is a natural consequence of the need for a substantial amount of data for ANN models to learn complex relationships in complex problems, as mentioned above. In the following section, this situation will be evaluated in more detail.

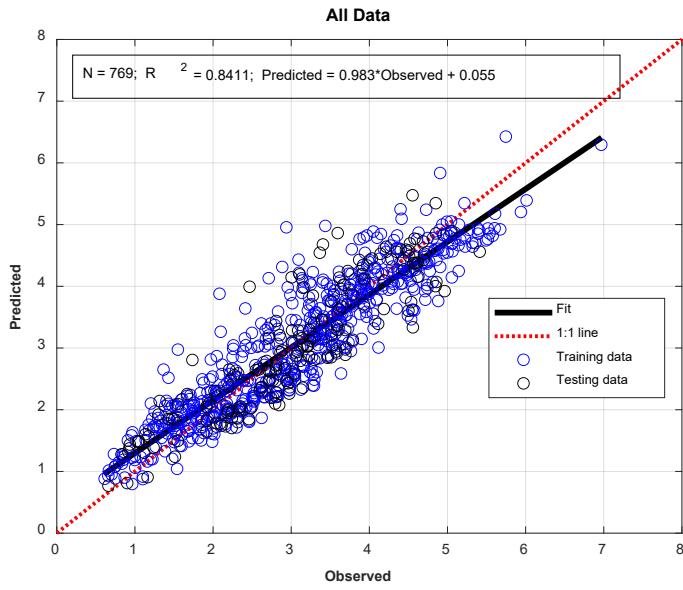


**Figure 4.** Model results for Scenario I.

### 3.1.2. Scenario II

Figure 5 shows the model results of the developed model for Scenario II. As seen in Figure 5, the number of data points used in the model training is 615, which constitutes 80% of the total data. During the training phase, the R-squared value is 0.8496, while for the test data, it is 0.8055, and for the total data, it is 0.8411. These results indicate that the developed model has a very good learning and prediction ability.





**Figure 5.** Model results for Scenario II.

#### 4. Discussion

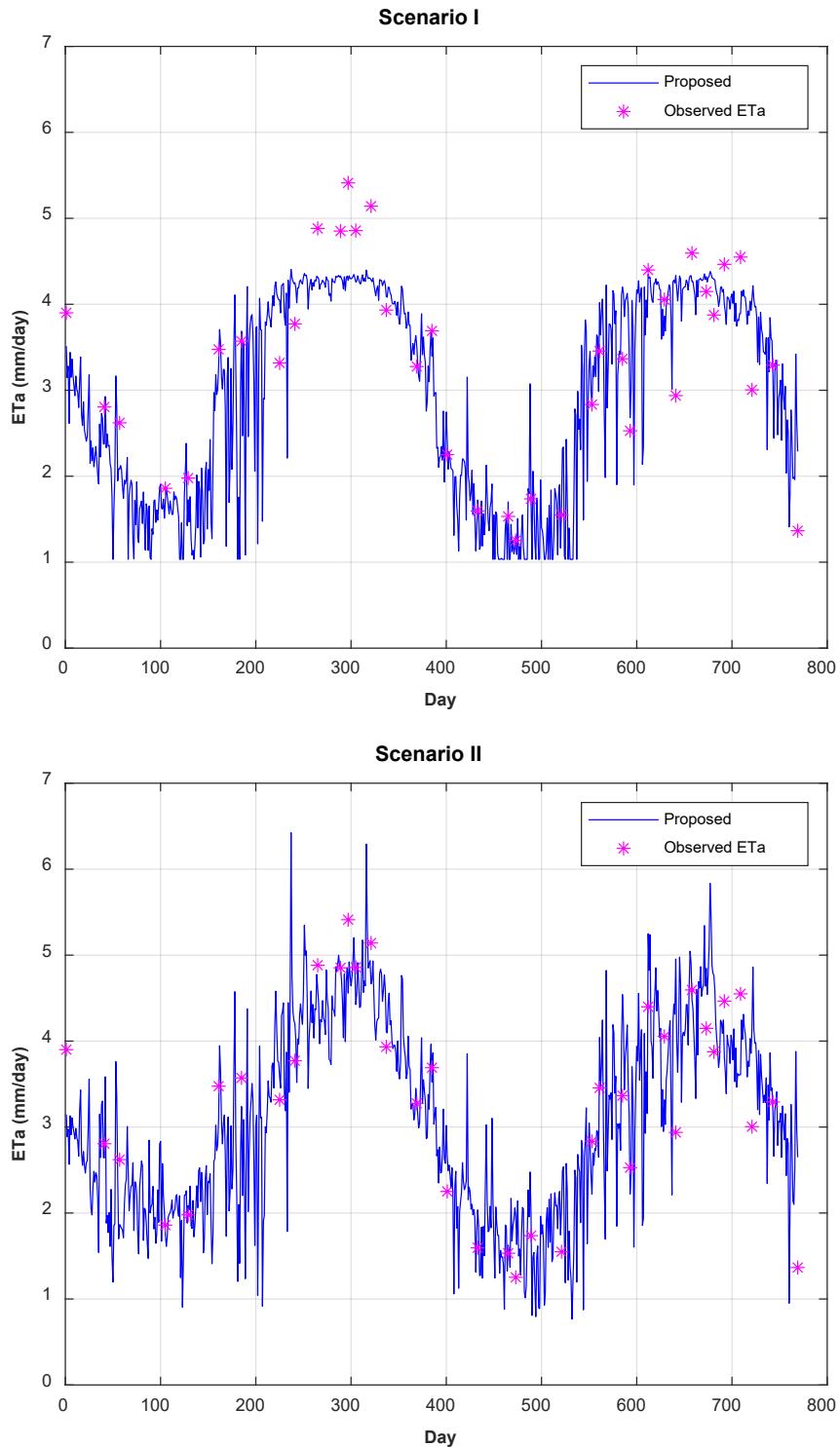
The primary objective of this study was to develop a methodology for estimating daily ET<sub>a</sub> values for large irrigation areas based on artificial neural networks using some parameters of MODIS data and climatic variables with ET<sub>0</sub> values. It is important to highlight that ET<sub>0</sub> estimations will be an easy task provided that climatic data required by the standard FAO-Penman-Monteith approach are available at the study site. In our study, there are two meteorological stations for collecting data. However, the problem is that Landsat satellite images are available with a 16-day repeat cycle. Therefore, ET<sub>a</sub> values for the days of no satellite overpass should be estimated by using an appropriate methodology such as ANNs. To this end, the developed methodology consists of two main parts. In the first stage of the methodology, daily ET<sub>0</sub> values calculated using the FAO-Penman-Monteith approach detailed in Allen et al. (1998) are combined with MODIS data to create the input data for the model. These values are based on observations from a meteorological station established in the study area. Additionally, observed ET<sub>a</sub> values are calculated based on Allen et al. (2007a, 2007b) using L8 and Cotlu meteorological station data for each pixel at satellite overpass times and other times. This completes the data set required for the training and testing stages of the developed ANN model.

After completing the data set, calculations are performed in two different scenarios. In Scenario I, only data acquired for the days with satellite imagery are used for model training, while in Scenario II, a randomly selected 80% portion of the total data set is used. It should be noted that the input and output layers of both scenarios are the same, as shown in Figure 2. The optimal number of neurons in the hidden layer is determined to be 5 for Scenario I and 15 for Scenario II, as detailed in the previous section. After determining the appropriate network architecture, detailed analyses were conducted for both scenarios.

In the model training, the goal was to minimize the MSE values as the objective function, and the results obtained for training, testing, and the entire dataset were provided graphically in the previous section. In the relevant graphs, the number of data points used, the R<sup>2</sup> value, and the regression relationships between predicted values and observed values are presented.

As seen in Figure 4, in Scenario I, even though approximately 5% of the total dataset was used for model training, high R<sup>2</sup> values of 0.7855, 0.7547, and 0.7561 were achieved for training, testing, and the entire dataset, respectively. However, it can be observed from the relevant graphs that despite the overall high model performance, the model's prediction ability in Scenario I decreases unexpectedly if the ET<sub>a</sub> values are around 4 mm day<sup>-1</sup> or greater, indicating conspicuous underestimates at high for high actual evapotranspiration rates. In Scenario I, the model fails to mimic

the high actual evapotranspiration rates. This can be seen more clearly in Figure 6, where the model results are presented graphically for both scenarios in comparison with ET<sub>a</sub> values calculated based on Allen et al. (2007a, 2007b). As clearly visible in the respective Figure, although the general trend is captured in both scenarios, it is noteworthy that the error rate in predicting values, especially those greater than 4, increases in Scenario I. It is believed that this is due to the low number of satellite observations (only 38 daily data) available for training the ANN model. Scatter plots provided in Figure 5 for Scenario II, as well as the temporal ET<sub>a</sub> values presented in Figure 6, support this hypothesis. Most studies in the literature also support this argument by allocating 80% of the total dataset for training and the remaining 20% for testing when training ANN models.



**Figure 6.** Model results for Scenario I and II.

## 5. Conclusions

This study presents a novel approach that allows daily ET<sub>a</sub> estimation using a new ANN model in a large-scale irrigation scheme using limited climatic data and MODIS satellite data. The study demonstrated that the daily ET<sub>a</sub> predicted values are comparable with the ET<sub>a</sub> which is estimated by the RS-based surface energy balance models. Furthermore, this methodology is the first attempt to generate a new model using the ANN algorithm as an alternative to the existing methods of ET<sub>a</sub> estimation in a large-scale irrigation district in the Eastern Mediterranean Region of Turkiye. In this context, the results of the new model in two implemented scenarios in this study showed acceptable agreement with ET<sub>a</sub> values estimated by the METRIC model at the large irrigation scale. Thus, this work is considered a significant contribution to obtaining reliable results without the need for lengthy and labor-intensive processes for complex equations as in RS-based surface energy balance models. This study revealed that the proposed model is a powerful tool for estimating daily actual ET using the limited meteorological observations and some of the parameters of remote sensing in arid and semi-arid regions as well as in different climate regions and zones of the world.

**Research Impact Statement:** The proposed methodology is promising to generate a new ANN model for estimating daily ET<sub>a</sub> as an alternative to the conventional ET<sub>a</sub> estimation procedures in large-scale irrigation schemes.

**Data Availability:** Data is available on request from the authors.

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