

Review

Not peer-reviewed version

A Review of User Perceptions of Drought Indices and Indicators Used in the Diverse Climates of North America

[Richard R. Heim Jr.](#)*, [Deborah J. Bathke](#), [Barrie Bonsal](#), Ernest W. T. Cooper, Trevor Hadwen, Kevin Kodama, [Dan McEvoy](#), Meredith Muth, John W. Nielsen-Gammon, Reynaldo Pascual Ramirez, [Holly R. Prendeville](#), Brad Rippey, [David B. Simeral](#), Richard L. Thoman Jr., Michael S. Timlin

Posted Date: 28 November 2023

doi: 10.20944/preprints202311.1770.v1

Keywords: user-inspired science; drought impacts; Canada; United States; Mexico; Pacific Islands; Caribbean Islands; user engagement; Köppen climate zones; climate change; drought in tropical climates; drought in polar climates



Preprints.org is a free multidiscipline platform providing preprint service that is dedicated to making early versions of research outputs permanently available and citable. Preprints posted at Preprints.org appear in Web of Science, Crossref, Google Scholar, Scilit, Europe PMC.

Copyright: This is an open access article distributed under the Creative Commons Attribution License which permits unrestricted use, distribution, and reproduction in any medium, provided the original work is properly cited.

Disclaimer/Publisher's Note: The statements, opinions, and data contained in all publications are solely those of the individual author(s) and contributor(s) and not of MDPI and/or the editor(s). MDPI and/or the editor(s) disclaim responsibility for any injury to people or property resulting from any ideas, methods, instructions, or products referred to in the content.

Review

A Review of User Perceptions of Drought Indices and Indicators Used in the Diverse Climates of North America

Richard R. Heim Jr. ^{1,*}, Deborah J. Bathke ², Barrie Bonsal ³, Ernest W. T. Cooper ⁴, Trevor Hadwen ⁵, Kevin Kodama ⁶, Dan McEvoy ⁷, Meredith Muth ⁸, John W. Nielsen-Gammon ⁹, Reynaldo Pascual Ramirez ¹⁰, Holly R. Prendeville ¹, Brad Rippey ¹², David B. Simeral ¹³, Richard L. Thoman Jr. ¹⁴ and Michael S. Timlin ¹⁵

¹ National Oceanic and Atmospheric Administration/National Centers for Environmental Information, 151 Patton Avenue, Asheville, NC 28801-5001 U.S.A.; Richard.Heim@noaa.gov

² National Drought Mitigation Center, University of Nebraska-Lincoln, 3310 Holdrege St., Lincoln, NE 68583-0988 U.S.A.; dbathke2@unl.edu

³ Watershed Hydrology and Ecology Research Division, Environment and Climate Change Canada, 11 Innovation Boulevard, Saskatoon, SK, Canada, S7N 3H5; barrie.bonsal@canada.ca

⁴ E. Cooper Environmental Consulting, 5612 47a Avenue, Delta, British Columbia, Canada, V4K 3Y2; cooper_ernie@hotmail.com

⁵ National Agroclimate Information Service, Agriculture and Agri-Food Canada, 2010 12th Avenue, Regina, Saskatchewan, Canada S4P 0M3; trevor.hadwen@canada.ca

⁶ National Oceanic and Atmospheric Administration/National Weather Service, Honolulu Forecast Office, 2525 Correa Road, Suite 250, Honolulu, HI 96822 U.S.A.; kevin.kodama@noaa.gov

⁷ Desert Research Institute/Western Regional Climate Center, 2215 Raggio Parkway, Reno, Nevada 89512 U.S.A.; Daniel.McEvoy@dri.edu

⁸ National Oceanic and Atmospheric Administration/National Integrated Drought Information System, 325 Broadway, Boulder, Colorado 80305 U.S.A.

⁹ Southern Regional Climate Center, Dept. of Atmospheric Sciences, Texas A&M University, College Station, Texas 77843-3150 U.S.A.; n-g@tamu.edu

¹⁰ National Meteorological Service of Mexico/National Water Commission, Av. Observatorio 192, Miguel Hidalgo, Mexico City, Mexico; reynaldo.pascual@conagua.gob.mx

¹¹ United State Department of Agriculture Northwest Climate Hub, 1220 SW 3rd Avenue, Suite 1400, Portland, Oregon 97204 U.S.A.; holly.prendeville@usda.gov

¹² U.S. Department of Agriculture, Office of the Chief Economist, 1400 Independence Avenue SW, Washington, D.C. 20250 U.S.A.; brad.rippy@usda.gov

¹³ Desert Research Institute/Western Regional Climate Center, 2215 Raggio Parkway, Reno, Nevada 89512 U.S.A.; David.Simeral@dri.edu

¹⁴ Alaska Center for Climate Assessment and Policy, International Arctic Research Center, University of Alaska Fairbanks, Fairbanks, Alaska 99775 U.S.A.; rthoman@alaska.edu

¹⁵ Midwestern Regional Climate Center, Illinois State Water Survey, Prairie Research Institute, University of Illinois, Champaign, Illinois 61820 U.S.A.; mtimlin@illinois.edu

¹⁶ Cooperative Institute for Research in Environmental Sciences, University of Colorado-Boulder, Boulder, Colorado, 80309 U.S.A.; elizabeth.weight@noaa.gov

* Correspondence: Richard.Heim@noaa.gov

Abstract: Drought monitoring and early detection have improved greatly in recent decades through the development and refinement of numerous indices and indicators. However, a lack of guidance, based on user experience, exists as to which drought monitoring tools are most appropriate in a given location. This review paper summarizes the results of targeted user engagement and the published literature to improve the understanding of drought across North America, and to enhance the utility of drought monitoring tools. Workshops and surveys were used to assess and make general conclusions about the perceived performance of drought indicators, indices and impacts information used for monitoring drought in the five main Köppen climate types (Tropical, Temperate, Continental, Polar Tundra, Dry) found across Canada, Mexico, and the United States. In Tropical, humid Temperate, and southerly Continental climates, droughts are perceived to be

more short-term (less than 6 months) in duration rather than long-term (more than 6 months). In Polar Tundra climates, Dry climates, Temperate climates with dry warm seasons, and northerly Continental climates, droughts are perceived to be more long-term than short-term. In general, agricultural and hydrological droughts were considered to be the most important drought types. Drought impacts related to agriculture, water supply, ecosystem, and human health were rated to be of greatest importance. Users identified the most effective indices and indicators for monitoring drought across North America to be the U.S. Drought Monitor (USDM) and Standardized Precipitation Index (SPI) (or another measure of precipitation anomaly), followed by the Normalized Difference Vegetation Index (NDVI) (or another satellite-observed vegetation index), temperature anomalies, crop status, soil moisture, streamflow, reservoir storage, water use (demand), and reported drought impacts. Users also noted the importance of indices that measure evapotranspiration, evaporative demand, and snow water content. Drought indices and indicators were generally thought to perform equally well across seasons in Tropical and the colder Continental climates, but their performance was perceived to vary seasonally in Dry, Temperate, Polar Tundra, and the warmer Continental climates, with improved performance during warm and wet times of the year. The drought indices and indicators, in general, were not perceived to perform equally well across geographies. This review paper provides guidance on when (time of year) and where (climate zone) the more popular drought indices and indicators should be used. The paper concludes by noting the importance of understanding how drought, its impacts, and indicators are changing over time as the climate warms, and by recommending ways to strengthen the use of indices and indicators in drought decision-making.

Keywords: user-inspired science; drought impacts; Canada; United States; Mexico; Pacific Islands; Caribbean Islands; user engagement; Köppen climate zones; climate change; drought in tropical climates; drought in polar climates

1. Introduction

Drought is a hydrometeorological phenomenon that has significant economic, social, and environmental impacts, and that occurs in all climates of the world regardless of international boundaries [1,2]. A 2021 World Meteorological Organization report noted that drought caused more than \$US 250 billion in economic losses and the death of more than 650,000 people globally during the last 50 years [3]. According to statistics from the Center for Research on the Epidemiology of Disasters (CRED), in 2022 there were 2601 deaths worldwide attributed to drought, 106.9 million people were affected, and drought caused \$US 34.2 billion in economic losses [4]. Drought is essentially an imbalance between water supply (precipitation) and water demand (in nature, evapotranspiration), but it is a very difficult phenomenon to define because it often develops slowly, operates on multiple time scales, is difficult to recognize until impacts are apparent, and produces effects which often accumulate slowly over time and may linger for years after the event ends [5–7]. Drought occurs on multiple timescales. It has been described as a “creeping” phenomenon [5,8,9] and experiencing its effects, in the words of a New Mexico rancher, is “like being slowly strangled” [10]. Some droughts can develop rapidly when lower-than-normal rates of precipitation are accompanied by abnormally high temperatures, high winds, high insolation, and low humidity; such droughts are referred to as “flash droughts” [11,12]. The climatological community has consequently defined drought in general terms such as a “prolonged absence or marked deficiency of precipitation,” a “deficiency of precipitation that results in water shortage for some activity or for some group,” or a “period of abnormally dry weather sufficiently prolonged for the lack of precipitation to cause a serious hydrological imbalance” [13,14], and identified five general types of drought: meteorological, agricultural, hydrological, socioeconomic, and ecological drought. Meteorological drought is associated with the atmospheric conditions resulting in the absence or reduction of precipitation and can develop quickly and end abruptly. Agricultural drought is usually associated with short-term

dryness in surface soil layers (root zone) which occurs at a critical time during the growing season; flash drought can be an important factor in the development of agricultural drought. Precipitation deficits over a prolonged period that affect surface or subsurface water supply result in hydrological drought and are associated with reduced streamflow, groundwater, reservoir, and lake levels; a hydrological drought will persist long after a meteorological drought has ended [5]. Agricultural impacts are more often associated with drought that occurs at short timescales, while hydrological impacts are more often associated with drought that occurs at longer timescales [15–17]. Socioeconomic drought associates the supply and demand of some economic goods with elements of meteorological, hydrological, and agricultural drought. Ecological drought is a recently-identified type of drought defined as an episodic deficit in water availability that drives ecosystems beyond thresholds of vulnerability, impacts ecosystem services, and triggers feedbacks in natural and/or human systems [18].

Numerous drought indices and indicators¹ have been developed over the decades to quantify drought. Early drought indices were limited by the data that were readily available at the time, primarily precipitation, and were tailored to specific applications for specific regions or measured specific types of drought. In 1965, Wayne Palmer developed an index that was the first to address the total moisture status by integrating water supply (precipitation), water demand (evapotranspiration), and water storage (soil moisture) [5,20]. Drought indices and models became more sophisticated as technology improved, more data became available (including remotely sensed data), and computer processing capabilities advanced. Some were developed with the intent to be more universally applicable, but it has been determined that no single drought index can adequately detect all types of drought for all locations [5,15,19,21–24]. A “convergence of evidence” approach for drought monitoring, in which all relevant drought indices and indicators are examined to determine the level of drought, was pioneered in the U.S. with the Drought Monitor to address this difficulty [25]. In order for the convergence of evidence (i.e., Drought Monitor) approach to work, the relevancy of the drought indices and indicators to accurately depict drought in specific regions needs to be known. The World Meteorological Organization (WMO) has compiled a list of many drought indices and indicators used globally in the *Handbook of Drought Indicators and Indices* [19], but the Handbook only lists the indices and indicators -- it provides no guidance on when and where it is appropriate to use them -- and it is not an exhaustive list.

An index or indicator is only as good as its correlation with actual or potential drought impacts and its utility in supporting drought management decisions. Two ways of testing this are objective comparisons with quantitative impact metrics and qualitative assessment through extended use by diverse practitioners. Quantitative scientific analyses have been conducted [15,22,23], and others are ongoing, to assess the applicability of drought indices. Meanwhile, the climatological community has been employing a “use-inspired science” philosophy in recent years that links climate tools and services to sector-specific users and requirements by engaging with users and stakeholders through various venues [26–33]. Users have been engaged at workshops and other meetings, such as the annual Drought Monitor Forums and workshops sponsored by the U.S. National Oceanic and Atmospheric Administration’s (NOAA) National Integrated Drought Information System (NIDIS), U.S. Department of Agriculture (USDA), and U.S. Department of the Interior’s (DOI) Geological Service (USGS) to address user needs, improve drought communication, and improve drought monitoring tools and methods. These user-engagement activities have resulted in an improved regional understanding of drought in those climate zones which are densely populated and have a long history of economic activity. But for regions that have limited climate data and limited economic activity, such as polar regions and tropical islands, the understanding of drought is limited. In the last decade, work has been done to address these gaps. This work includes several workshops to improve the understanding of drought and identify drought monitoring tools appropriate for use in polar and tropical climates, and a comprehensive 2-year study run by the Commission for

¹ Like the World Meteorological Organization [19], we define indicators as variables or parameters used to describe drought conditions, while indices are typically computed numerical representations of drought severity, assessed using climatic or hydrometeorological inputs that may include indicators.

Environmental Cooperation² (CEC) to assess which drought indices and indicators are appropriate to use across all the diverse climates of North America.

The Drought Monitor methodology has been adopted for operational drought monitoring in the United States (experimentally in 1999 and operationally beginning in 2000), and in Canada and Mexico a few years later. The corresponding U.S. Drought Monitor³ (USDM), Canadian Drought Monitor⁴ (CDM), and Mexican Drought Monitor⁵ (MDM) products are merged to create a monthly North American Drought Monitor⁶ (NADM) product through a collaborative process involving agencies of the three countries [34,35]. The Drought Monitor methodology is being adopted in other parts of the world as well. The drought indices and indicators utilized in the Drought Monitors vary with country but include many discussed here. These Drought Monitor products will benefit from the work summarized in this paper as well as from ongoing work to assess the applicability of drought indices.

This review paper integrates and summarizes the results of these recent workshops and the CEC study, and supplements it with published research that objectively analyzes the effectiveness of drought indices. Based on this integrated information, the paper serves as a toolbox by providing guidance on when and where in North America it is appropriate to use the drought indices and indicators listed in the WMO Handbook as well as other indicators that were not listed in the Handbook and/or were developed since the release of the Handbook.

The paper is organized into the following sections. Section 1 (this Introduction) discusses the background leading to this review. Section 2 discusses user engagement activities (Appendix S-A describes the meetings and workshops in more detail). Section 3 discusses drought definitions. Section 4 summarizes the user-engagement data on the usefulness of drought indices and indicators discussed by Köppen climate type (Appendices S-C and S-D discuss results from these meetings and workshops in more detail). Section 5 consolidates the discussion of the indicator and index assessment results from user-engagement activities and published literature, discusses important implications of the user-engagement process, identifies some priorities for future work on indicators, and presents some general recommendations. Details of the meetings, workshops, published research, and the Köppen climate classification system, and tables of acronyms, are provided as appendices (S-A to S-F) in Supplementary Material.

2. User Engagement Data

North American users were engaged to assess their drought index and indicator needs through a series of meetings and activities. Tropical Pacific and Caribbean islands are included, as their drought conditions are operationally assessed in the USDM. Appendix S-A describes the meetings and workshops that have been held over the past decade that engaged users in the Pacific and Caribbean protectorates of the U.S., Hawaii, and Alaska. The results of these meetings and workshops [36–42] include drought indices and indicators, impacts, and drought monitoring methodologies that are relevant to these regions.

The CEC funded and managed a project during 2019–2020 that engaged users across North America to assess the performance of drought indices and indicators used to monitor drought in the

² The CEC was established by Canada, Mexico, and the United States to implement the North American Agreement on Environmental Cooperation, the environmental side-accord to the North American Free Trade Agreement. Its mission is to facilitate effective cooperation and public participation to conserve, protect and enhance the North American environment in support of sustainable development for the benefit of present and future generations.

³ The U.S. Drought Monitor can be accessed here: <https://droughtmonitor.unl.edu/>

⁴ The Canadian Drought Monitor can be accessed here: <https://www.agr.gc.ca/eng/agriculture-and-the-environment/drought-watch-and-agroclimate/canadian-drought-monitor/?id=1463575104513>

⁵ The Mexican Drought Monitor can be accessed here: <https://smn.conagua.gob.mx/es/climatologia/monitor-de-sequia/monitor-de-sequia-en-mexico>

⁶ The North American Drought Monitor can be accessed here: <https://www.ncdc.noaa.gov/temp-and-precip/drought/nadm/>

diverse climates of the continent [43]. The project was guided by a steering committee composed of representatives of agencies and institutions from the U.S., Mexico, and Canada. User engagement included an online survey, plus consultation during two virtual online webinars, one in English on 8 October 2020 and the other in Spanish on 13 October 2020. Participants included federal, state, provincial, and non-governmental users, including academic and tribal users, in the three countries who were involved in drought monitoring, communications, hazard mitigation or disaster resilience, environmental and natural resources planning, government research, or comprehensive/long-range planning. There were 164 survey responses, of which 145 respondents identified which country they worked in—84 worked in the U.S., 33 in Canada, and 28 in Mexico [43]. A small number of respondents had responsibility for drought monitoring across the entire continent. The indices and indicators that were evaluated included those listed in the WMO *Handbook of Drought Indicators and Indices* as well as 22 indicators that were not included in the WMO handbook. The results of the study were published by the CEC in a *Guide to Drought Indices and Indicators Used in North America* [43]. Tables summarizing the survey results are available online from the NOAA National Centers for Environmental Information (NCEI)⁷. The results of the meetings and workshops discussed in Appendix S-A are data that supplement or add to the data collected through the CEC survey and consultations.

3. The Definition of Drought

The climatological community has grappled with the subject of the definition of drought for decades [44]. Very specific definitions have been proposed for specific applications in specific regions, where drought is quantitatively defined using specific criteria [5,45]. Efforts to develop a universal definition of drought result in broad definitions like those referenced above in the Introduction.

The definition of drought was a discussion topic at the CEC consultations held on 8 October and 13 October 2020. Several participants defined drought using specific criteria of specific indices, with the criteria and indices used varying by region and application. A broad definition was offered (“drought is insufficient water to meet the needs of some activity or group”), but the importance of drought impacts was also noted as well as regulatory actions triggered by drought.

The subject of drought impacts needs to be treated carefully. As noted during the CEC consultations, some might argue that if there are no drought impacts, then there is no drought. If this proposition is to be accepted, there are at least four issues that need to be addressed. The first is that comprehensive, long-term drought impact data are lacking. We cannot state with confidence that there are no impacts if a drought is occurring without these data and thus in these instances there is an unclear relationship with impacts. Groups such as the National Drought Mitigation Center (NDMC) and Community Collaborative Rain, Hail & Snow (CoCoRaHS) Network have citizen science activities⁸ to collect drought impact data in efforts to address this data gap. The second is the fact that many drought impacts are notoriously hard to quantify. For example, impacts of drought on crops in a given year are difficult to isolate from impacts of temperature, insects, and other phenomena. The third centers around types of drought. For agricultural, socioeconomic, and ecological drought, respectively, crops can be damaged, economic activities can be affected, and elements of the ecosystem (natural vegetation, wildlife, etc.) can be harmed. Hydrologic droughts can reduce natural streamflow, lake levels, and groundwater; these are special drought impacts that can be observed with quantitative measurements and therefore can be used as indicators (or even indices if there is a sufficient record), so they can both be impacts and indicators. A more likely hydrological impact would be reduced water supply to a user group. Meteorological drought is measured more by the lack of precipitation than any impacts. The fourth issue concerns mitigation strategies that have been put in place to reduce or postpone the impacts of drought. An example of this is the reservoir system in the western U.S. The reservoir system in California, for example,

⁷ The CEC survey results are summarized at: <https://www.ncei.noaa.gov/access/monitoring/nadiia/>

⁸ Drought Impact Reporter web page: <https://droughtreporter.unl.edu/map/>. Condition Monitoring Observations web page: <https://droughtimpacts.unl.edu/ConditionMonitoringObservations.aspx>

enables the state to endure several years of drought with reduced impacts to urban and irrigated agricultural users. The fact that these users are experiencing adequate water supply during a time of severe meteorological and ecological drought does not mean there is no drought, it just means that the impacts of drought have been mitigated for some—not all—water users. But there are situations in which prolonged drought can overwhelm even these mitigation strategies -- water supplies become inadequate over time due to very long-term drought (and, in the case of the Colorado River, overallocation of water supplies), the mitigation system no longer functions, and impacts are significantly amplified. Clearly, as will be seen in subsequent sections of this paper, drought can cause significant impacts. Therefore, impacts may be a sufficient condition for drought in some cases, but they are not a necessary condition.

In the absence of a single definition of drought suitable for all situations and locations, the assessments of indices and indicators in this paper should be interpreted as relative to whatever individual participants considered drought to be.

4. North American Users' Assessment of Drought Indices and Indicators

North America has climates spanning the range from tropical to polar, arid to humid, and maritime to continental. Very few drought indices and indicators are likely to be effective across all climate zones. Section 4.2 provides an overview of the CEC survey results, while Appendices S-C and S-D in the Supplementary Material contain a detailed discussion of drought indices and indicators used in the various climate zones summarized from the CEC survey and various user engagement meetings and workshops. Section 5.3.1 provides an overview of published objective research that assesses the effectiveness of drought indices by climate zone, with a more detailed discussion of the published research provided in Appendix S-E in the Supplementary Material. An exploration of why the indices and indicators are or are not effective and how they are used is in Section 5 (5.2 and 5.3.2).

4.1. Köppen Climate Types

In the CEC survey, the Köppen climate classification was used to assess the drought indices and indicators because it is the most widely accepted global climate classification [43]. Much of the published research referenced in this paper analyzed drought indices based on Köppen climate zones. Therefore, much of the discussion in this paper is broken down by Köppen climate type. Köppen was a Russian-born biologist who sought to relate vegetation types to climate. His classification scheme is based on monthly temperature, precipitation, and their seasonal characteristics [46]. The major Köppen climate types are: A (Tropical climates that are hot year-round), B (Dry climates), C (Temperate climates—warm temperature wet climates with mild winters), D (Continental climates with cold winters), and E (Polar climates).

The major climate types are indicated by the first letter of the climate zone codes shown in Figure 1. The Köppen analysis and map were provided to the authors by Ricardo Llamas Barba⁹ through the Commission for Environmental Cooperation (CEC). In the CEC study and in this paper, Tropical climates include southern and coastal areas of Mexico, southern Florida, Puerto Rico and the U.S. Virgin Islands (USVI), the U.S. Affiliated Pacific Islands (USAPI), and coastal areas of Hawaii. Dry climates include parts of northern Mexico, much of the western U.S., and western portions of the Great Plains in the U.S. extending into parts of the southern Canadian Prairies. Temperate climates are located in parts of central Mexico, the west coasts of the U.S. and Canada, higher elevations in Hawaii, parts of the panhandle and southern coast of Alaska, and much of the southern Plains to Southeast in the U.S. Continental climates are located in the U.S. from the central and northern Plains to Northeast, across parts of the Northwest, and in much of Alaska; across most of Canada south of

⁹ The Köppen climate map was primarily derived from a pre-produced dataset published by Beck et al. [47]. Ricardo Llamas Barba and his team made some adjustments, including resampling the pixel size and reprojecting the data to match the North American coordinate reference frame used by CEC. The North American climate zones map, along with the relevant metadata, are accessible through the CEC Environmental Atlas (<http://www.cec.org/north-american-environmental-atlas/>).

the Arctic Circle; and areas of higher elevation. In North America, Polar (E) climates are located in northern parts of Alaska, the Canadian far north, and higher elevations in the western U.S. and western Canada

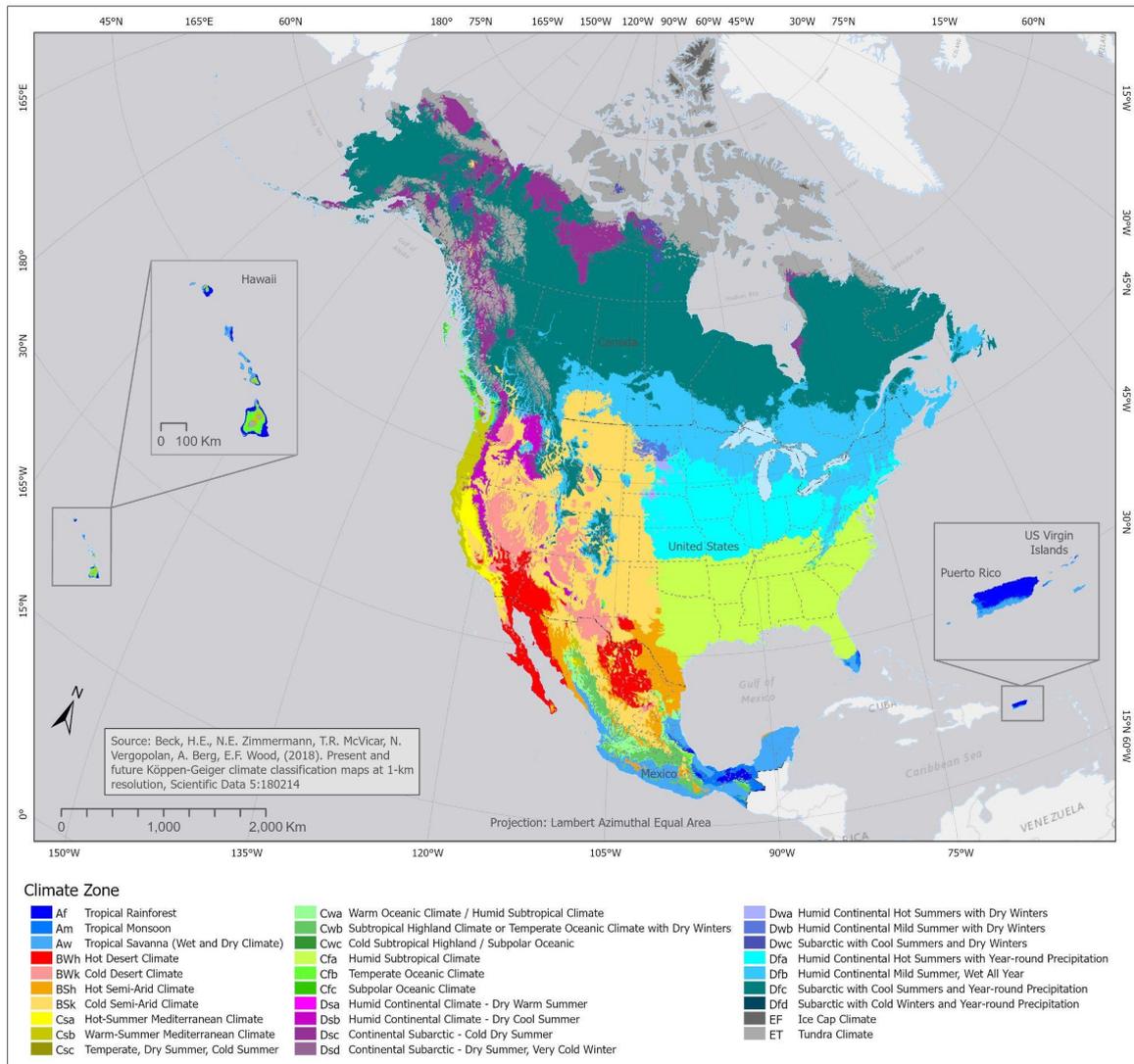


Figure 1. Map showing Köppen climate zones for North America, Hawaii, Puerto Rico, and the U.S. Virgin Islands. Map prepared by Ricardo Llamas Barba for the Commission for Environmental Cooperation (CEC), 2021. Printed with permission from the CEC. The U.S.-Affiliated Pacific Islands (not shown) are classified as an Af climate except for the northernmost islands in the Republic of the Marshall Islands and Commonwealth of the Northern Mariana Islands, which are classified as Aw.

The five Köppen climate types are divided into subzones designated by secondary and tertiary letters (descriptive names associated with each subzone are listed in Table S1 in Appendix S-B). A more detailed discussion of the major climate types and their subzones, including the mathematical parameters defining each, can be found in Appendix S-F.

4.2. CEC Survey Results

The CEC survey included questions about the overall geographical and seasonal performance of the drought indices and indicators, factors affecting the choice of indices and indicators used, and characteristics of drought, in addition to questions rating the effectiveness of the drought indicators and indices (see the CEC's *Guide to Drought Indices and Indicators Used in North America* for a list of the survey questions). It should be noted that, in the CEC survey, respondents were asked about drought,

drought indices, and drought indicators as they pertained to their area of responsibility. The respondents also indicated which climate zones were included in their area of responsibility. In the discussion that follows, the responses were applied to all of the climate zones within the respondent's area of responsibility. For some respondents, the geographical area of responsibility was small and included just one or a few climate zones; for others, the geographical area of responsibility was larger and may have included several climate zones. The presentation of the results in Section 4 was done this way to make the results easier to digest, but this should be kept in mind when interpreting the percentages.

The CEC survey listed all possible Köppen climates and subzones. Some of the subzones are not present on the North American continent, yet survey respondents identified indices and indicators that they felt were (or would be) appropriate for monitoring drought in those subzones. Examples of such subzones include EF (glacial climates found only in Greenland and Antarctica) and Dwd (found only in Siberia) and others. This may mean that those respondents have an interest in these subzones, so the survey results for these subzones were included in the discussion.

The sample size (number of respondents) varied among the climate types and especially among the climate subzones¹⁰. The number of respondents for each of the climate types is: A (49), B (133), C (116), D (138), and E (7) [43]. The number of respondents for the A-D climate subzones ranged from 4 in subzones Cwc, Dsc, Dsd, and Dfd, to 37 in subzone BSk. For Polar climates, ET had five respondents and EF had two¹¹. With such a small respondent sample size, the survey answers may not be representative for EF climates. The discussion in this section will summarize responses by climate type and, in some cases, by climate subzone. For the climate type summaries, the responses were not simple averages of the subzone responses, they were weighted by the number of respondents in each of the subzones. The CEC study [43] presented the results in data tables; the results are summarized in graphical form in this section for clarity of presentation and easier identification of geographical or climatological relationships.

Drought type was rated on a scale of 1 (not important) to 5 (very important). Agricultural, hydrological, meteorological, and ecological droughts were rated as important to very important (ratings 4 or 5) by half or more of the respondents in all five climate types (Figure 2). Socioeconomic drought was so rated by half or more of the respondents in D and E climates. The ratings were similar for the subzones, except for the following: socioeconomic drought was rated as important to very important by half or more of the respondents in the Af, Am, BSh, BWk, Cfb, Cfc, Csa, Csb, and Cwc subzones; and socioeconomic drought was so rated by *less than half* of the respondents in the Dfc subzone.

The duration of a typical drought was perceived by half or more of the respondents to be predominantly short-term (< 6 months) in A and D climates, and long-term (> 6 months) in B, C, and E climates (Figures 3 and 4), with nearly two-thirds of the respondents in A climates identifying 3-6 months as the typical length of droughts. The ratings varied by a wide margin in A climates and by narrower margins in B, C, D, and E climates. There were deviations from the climate type assessment at the subzone level, including: the majority of respondents perceived drought duration to be typically short-term instead of long-term in Cfa and Cwb subzones, and the majority of respondents perceived drought duration to be typically long-term instead of short-term in the Dfc, Dsb, Dwa, Dwb, and Dwd subzones.

¹⁰ The number of respondents for each of the subzones of each climate type can be found at the NCEI webpage: <https://www.ncei.noaa.gov/access/monitoring/nadiia/>

¹¹ The reader is referred to the NCEI webpage (<https://www.ncei.noaa.gov/access/monitoring/nadiia/>) for the full summary of survey responses for all climate subzones, indices and indicators, drought lengths, drought types, impacts, and factors affecting choice of index. Only those identified as most important by the survey respondents are discussed in the manuscript.

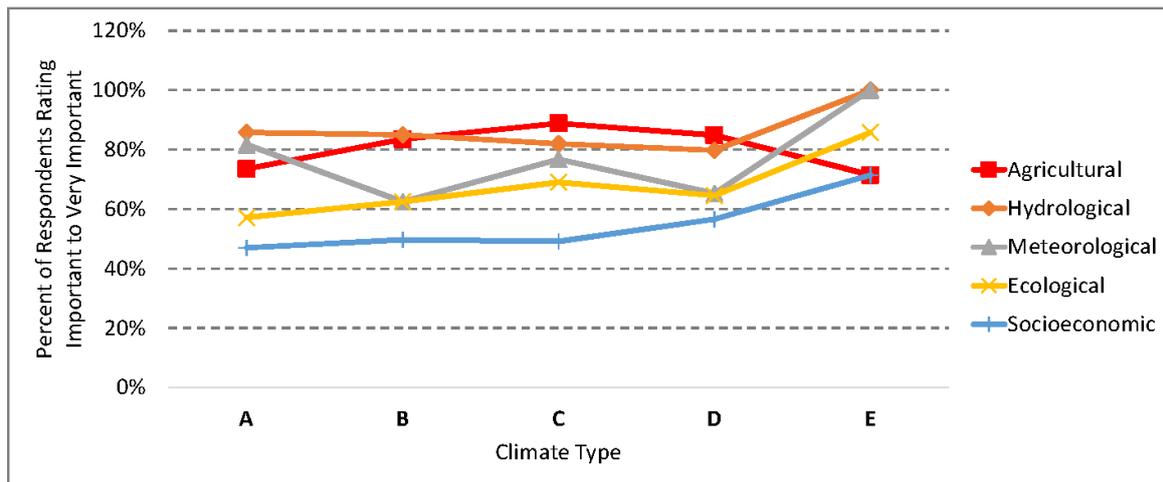


Figure 2. Percent of respondents in each climate type rating the five types of drought as important to very important. Climate Type: A = Tropical, B = Dry, C = Temperate, D = Continental, E = Polar Tundra.

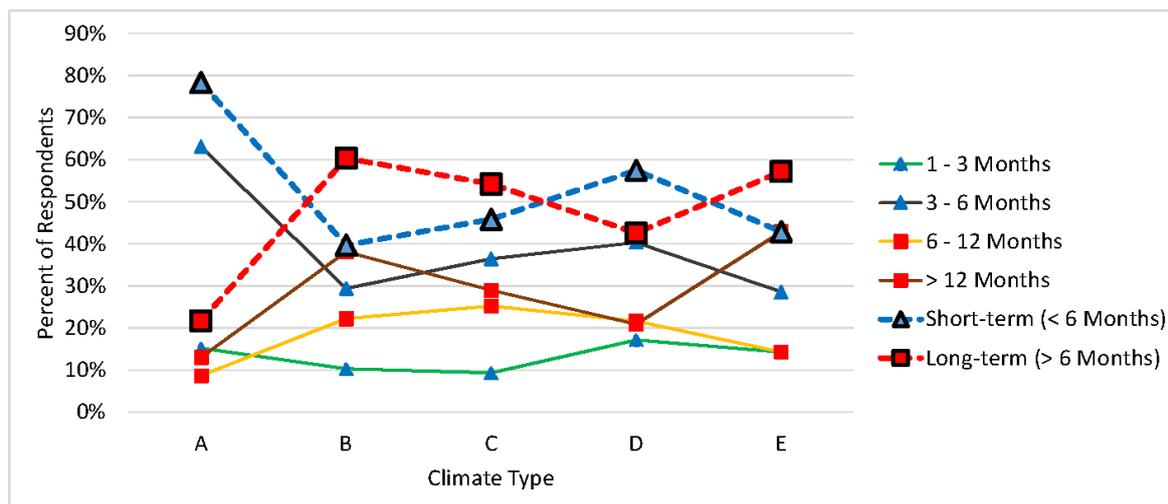


Figure 3. Percent of respondents in each climate type perceiving the typical drought duration to be short-term (triangle with dashed blue line) or long-term (square with dashed red line), and 1-3 months (triangle with solid green line), 3-6 months (triangle with solid gray line), 6-12 months (square with solid amber line), or more than 12 months (square with solid brown line). Climate Type as in Figure 2.

The factors affecting the choice of indicators were rated on a scale of 1 (not important) to 5 (very important). Two factors were rated as important to very important by wide margins across all five climate types: relevance of the index or indicator to my area/region, and availability of relevant and required data to calculate the index or indicator (Figure 5). Familiarity with the specific index or indicator was also rated as important to very important across all five climate types, but by narrower margins. The history of indices and/or indicators used previously in my area/region was important to very important for the majority of respondents in C, D, and E climates, but with narrower margins; this factor was so rated by 50% of the respondents in A climates and 49% in B climates. Complexity or difficulty of the required calculation was rated as important to very important by less than half of the respondents in all five climate types.

The ratings at the subzone level were consistent across most subzones for availability of relevant and required data to calculate the index or indicator—only the Cwc and Dsc subzones departed from the climate type statistics with less than 50% of the respondents rating this factor as important to very important, but it should be noted that each of these two subzones had only three respondents. Relevance of the index or indicator to my area/region was rated important or very important by more

than half of the respondents in each subzone except Dfd (which had only 4 respondents) and Dsc. Familiarity with the specific index or indicator was rated important or very important by more than half of the respondents across all A, B, and E climate subzones, but not consistently so rated across the C and D subzones with a wide variation in responses (those subzones so rating this factor included Cfa, Cfb, Csa, Dfa, Dfb, Dfc, Dfd, Csa, Dsc, Dwa, Dw b, and Dwd). Half or more of the respondents in about half of the subzones (Af, Am, BSk, Cfa, Cfb, Cfc, Csa, Dfa, Dfb, Dfc, Dsa, Dwa, Dw b, Dwd, EF, and ET) rated history of indices and/or indicators used previously in my area/region as important or very important, while about half did not. Complexity or difficulty of the required calculation was so rated only in Am, BSh, BWh, Cfb, and Cwb subzones.

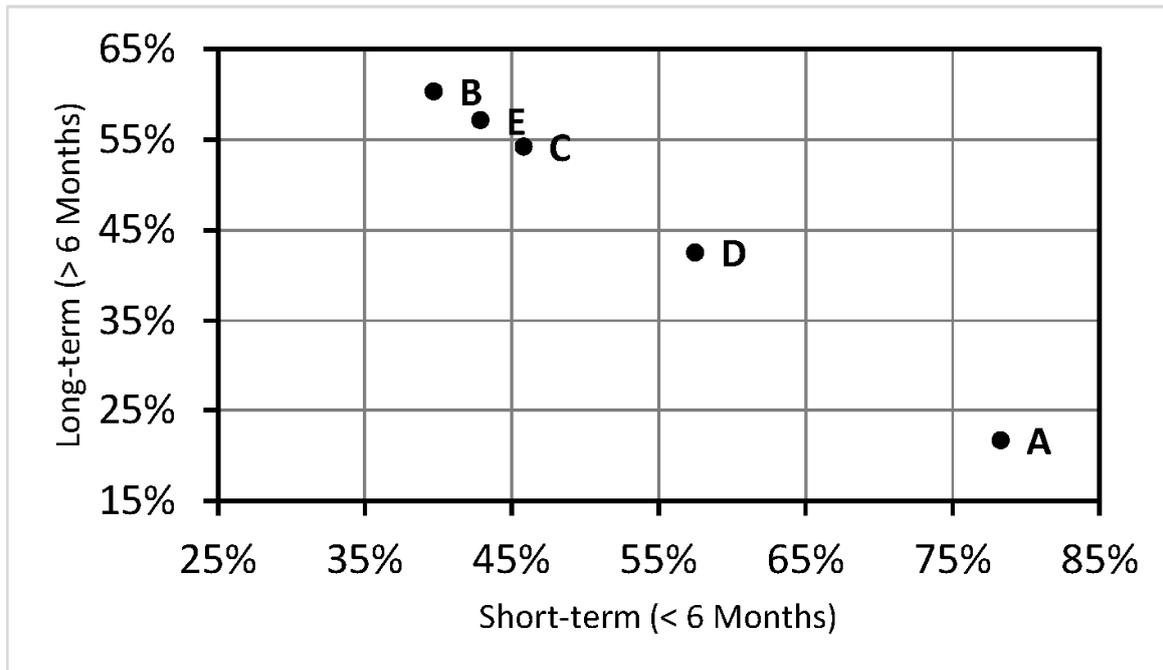


Figure 4. Percent of respondents in each climate type perceiving the typical drought duration to be short-term (less than 6 months, abscissa) or long-term (more than 6 months, ordinate). Climate Type as in Figure 2.

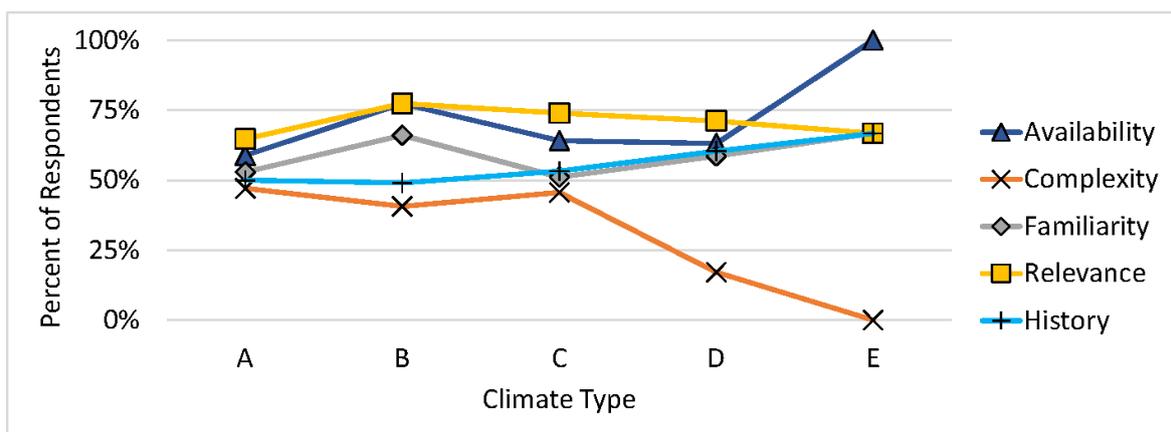


Figure 5. Percent of respondents in each climate type rating five factors as important or very important to their choice of drought indices and indicators. The factors are: availability of relevant and required data to calculate the index or indicator (triangle with dark blue line), complexity or difficulty of the required calculation (X with brown line), familiarity with the specific index or indicator (diamond with gray line), history of indices and/or indicators used previously in my

area/region (+ with light blue line), and relevance of the index or indicator to my area/region (square with orange line). Climate Type as in Figure 2.

The perceived performance of drought indices and indicators, in the aggregate, was evaluated spatially and across seasons. Figure 6 plots the percent of the respondents who indicated the indices performed equally well across seasons versus the percent of the respondents who indicated the indices performed equally well geographically, for the five climate types. Indices in A climates had the highest combined performance. In C climates, indices had the second-best performance geographically. Indices in D climates had the poorest overall performance across seasons, although the B, C, and D evaluations clustered near each other on the seasonal performance scale. Indices in B climates had the poorest overall performance geographically. Indices in E climates had the best overall seasonal performance, although it should be pointed out that the sample size for E climates was small (7 total respondents) and this may skew the results.

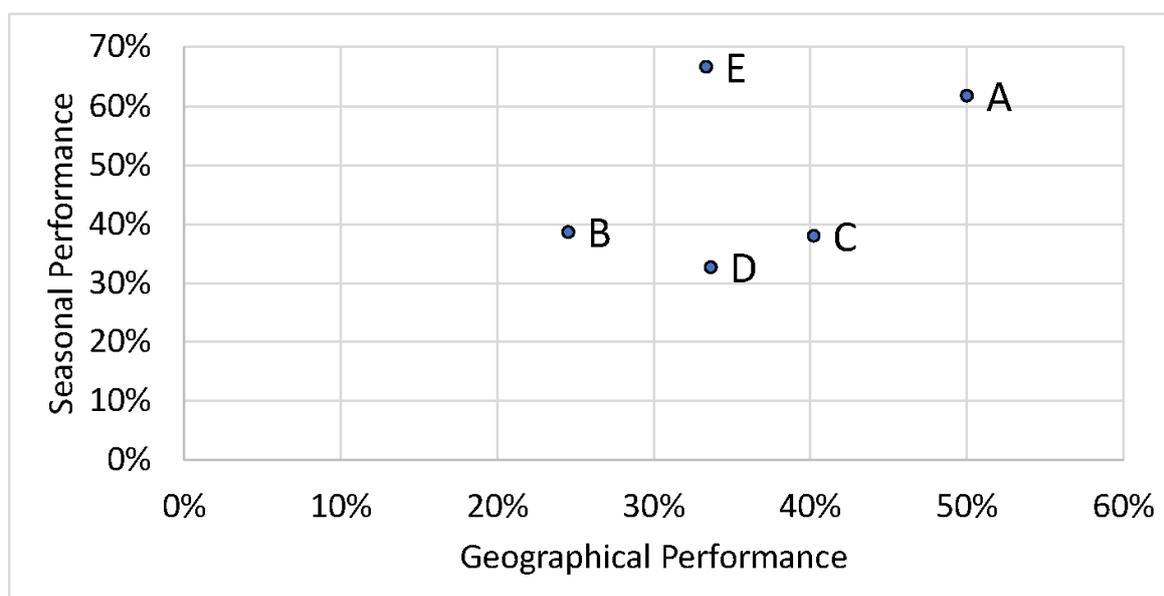


Figure 6. Percent of respondents who indicated the drought indices and indicators, in general, performed equally well across seasons (ordinate) versus percent of the respondents who indicated the drought indices and indicators, in general, performed equally well geographically (abscissa), by climate type. Climate Type as in Figure 2.

The seasonal vs. geographical perceived performance within the climate subzones is similar to the performance of the climate types for the A and B climates, but varies widely for the C, D, and E climates (Figure 7). Sample size of the subzones in the A and B climate types ranges from 9 respondents (Am) to 30 (BSk). Sample size for the C climate subzones ranges from 3 (Cwc) to 18 (Cfa), for the D climate subzones ranges from 3 (Dsc) to 24 (Dfb), and for the E climate subzones includes 2 (EF) and 4 (ET). Some of the subzones with a low sample size had high overall performances (e.g., Cwa, Cwc, Dfd) while some had low overall performances (e.g., Dwa, Dw b, Dsd), so if sample size was a factor in subzone performance, its influence was not consistent.

The survey results for the individual drought indices, indicators and drought impacts, by climate subzone, are presented in the CEC paper [43] and in the Supplementary Materials in Appendix S-C; they are summarized in Section 5 of this paper.

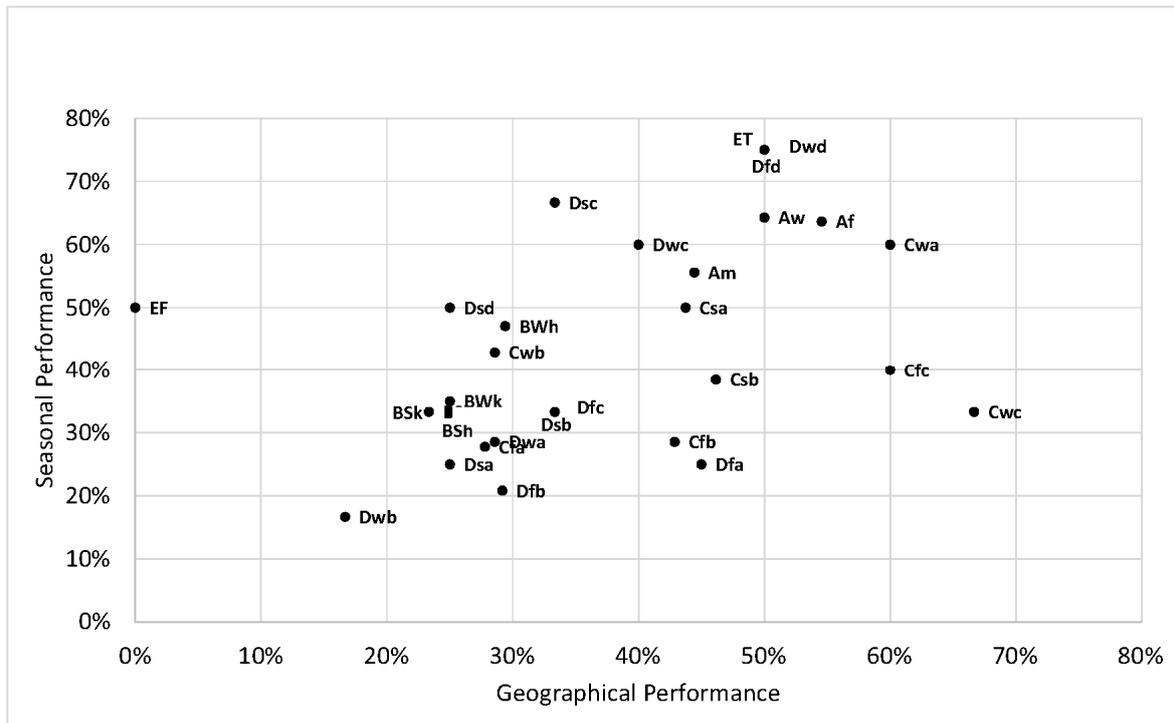


Figure 7. Same as Figure 6, except by fully-specified climate subzone.

4.3. User-Engagement Workshop Data

The data provided by user engagement at the workshops held over the past decade are grouped and discussed by climate type. The discussion can be found in the Supplementary Materials in Appendix S-D. The workshop data are supplemented by published research [36,46,48–78]. These workshop and published data supplement the CEC survey results and include discussion of drought indices, drought monitoring methodologies, and drought impacts. They are summarized in Section 5 of this paper.

5. Discussion

This section summarizes the survey and workshop results and published literature to conclude with recommendations on the use of drought indices and indicators in the climate zones of North America. It includes a discussion of drought impacts and vulnerabilities, evolving drought indicators in a warming climate, and embedding indicators within robust decision systems.

5.1. Consolidated Summary of the Survey and Workshop Results

While recognizing that those who participate in surveys and workshops represent only a portion of all drought indicator users, some conclusions about drought and the performance of the drought indices and indicators can be drawn from the CEC survey and workshop results. In Tropical (A) climates, the more humid Temperate (C) climates (those with precipitation year-round or wet season in the summer), and more southerly Continental (D) climates, respondents generally perceived droughts to last less than 6 months. This may be related to the greater chance of drought-ending rains occurring during the warm season when evapotranspiration is high. In Polar Tundra (ET) climates, Dry (B) climates, Temperate (C) climates with dry warm seasons, and the more northerly Continental (D) climates, respondents generally perceived droughts to last longer than 6 months. This may be related to the smaller chance of drought-ending rains occurring during the warm season when evapotranspiration is high (dry summer climates), and the climatological deficit of water supply (precipitation) related to water demand (evapotranspiration) in Dry climates. In Temperate climates

with no dry or wet season (f qualifier), there is potentially ample opportunity for drought-ending precipitation any time of the year, thus resulting in shorter-duration droughts.

Agricultural drought and hydrological drought were considered to be the most important drought types across all five of the major climate types (A, B, C, D, E) and all of their subzones. This is not surprising considering the historical importance of an adequate food supply and adequate water for sustaining civilization. Since our society places an emphasis on drought impacts to agriculture and water supplies, staff in charge of drought monitoring will reflect those priorities by stating in a survey that those are most important. Meteorological and ecological drought were also important in all five of the major climate types and most of the subzones, but by smaller percentages. Socioeconomic drought was important in E climates, most D climate subzones, and some A, B, and C subzones.

Considering the importance of agricultural and hydrological drought, impacts related to crop and ranching losses and impacts related to water supply (reservoirs, lakes, and ponds) were rated high in all of the major climate types. Also of concern in all of the major climate types were various ecosystem impacts (wetlands, forests, wildlife habitat, health of wildlife) and specifically the threat of wildfires. In four of the climate types (A, C, D, E), drought impacts on human health were an elevated concern.

The CEC respondents across all five of the major climate types identified the most important factors affecting their choice of drought index as availability of relevant and required data to calculate the index or indicator and relevance of the index or indicator to their area/region.

The performance of the drought indices and indicators across seasons and geographically was rated in the aggregate. The majority of respondents indicated that indices and indicators performed equally well across seasons in Tropical (A), Polar Tundra (ET), and the colder Continental (D) climates, but did not perform equally well in Dry (B), Temperate (C), and the warmer Continental (D) climates. Respondents indicated that indices and indicators did not perform equally well geographically in four of the main climate types (B, C, D, E), with responses almost evenly divided for Tropical climates.

The WMO *Handbook of Drought Indicators and Indices* listed 50 of the most commonly used drought indices and indicators that are being applied across drought-prone regions [19] and the Handbook's authors noted that it was not an exhaustive list. As seen in Appendix S-C of the Supplementary Material, only a subset of those indicators and indices are widely used in North America -- basically those the survey respondents and workshop participants were familiar with and that they considered were relevant to their area or region. It should be noted that the WMO *Handbook* listed indicators and indices used worldwide, and some may not be used in North America, so the respondents were more likely to rate the indices that they are familiar with more highly. This is a form of respondent bias. Preceding paragraphs have already noted the importance of agriculture to human society, and this is reflected by the fact that approximately half of the CEC survey respondents worked in the field of agriculture [43]. Thus, in the following discussion evaluating drought indices and indicators, it should be kept in mind that the results of this study reflect human perceptions more than an objective comparison of the drought indices.

Of the indices and indicators not included in the WMO *Handbook*, in general, the following were rated as effective or very effective for at least one subzone in all five major climate types: soil moisture; reported drought impacts; crop status; reservoir storage; vegetation greenness; groundwater depth; streamflow; precipitation departures from normal; precipitation percentiles; precipitation ranks; temperature departures from normal; temperature ranks; and water use (demand). In addition, several newer indices that are not in the WMO *Handbook* are becoming more widely used. These include the Evaporative Demand Drought Index (EDDI) and snow water equivalent (SWE), which are used often in the western U.S. As noted in Supplementary Material Appendix S-D, the February 2020 Alaska workshop also identified SWE and EDDI as appropriate for Southcentral and southern Interior regions of the state.

The most effective drought indices and indicators for either short-term or long-term drought, or both, across most of the subzones in all of the five major climate types, according to the CEC survey results and workshop participants, were the USDM, Standardized Precipitation Index (SPI),

precipitation percentiles, another measure of precipitation (such as precipitation ranks, departure from normal), Normalized Difference Vegetation Index (NDVI) (or another satellite-observed vegetation index), a measure of temperature anomaly (departure from normal, ranks), crop status, soil moisture, streamflow, reservoir storage, water use (demand), and reported drought impacts. Percent of Normal Precipitation (PNP) was rated in the most effective category in four of the major climate types (all except E). The Standardized Precipitation Evapotranspiration Index (SPEI) was rated in the most effective category in subzones of four of the major climate types (all except A). Groundwater depth was rated in the most effective category in all of the major climate types except D and E. Wildfire locations/reports were rated in the most effective category in subzones of all of the major climate types except E. Other indices and indicators that rated highly in some subzones of one or two major climate types include the Evaporative Stress Index (ESI) (A and C climates); Vegetation Drought Response Index (VegDRI) (C, B, and E climates); Soil Moisture Anomaly (SMA) (C climates); Soil Moisture Deficit Index (SMDI) (A and C climates); Crop Moisture Index (CMI) (C climates); SWE (B and D climates); EDDI and ephemeral ponds (D climates); and Surface Water Supply Index (SWSI) and water quality (E climates).

Evapotranspiration-related indices and indicators (such as SPEI or EDDI) did not appear to be as important in climates where evapotranspiration is high year-round and exhibits little variation that could be picked up by these indices and indicators.

The Palmer Drought Severity Index (PDSI) was rated as effective in Tropical climates for long-term drought. It was also rated effective in Polar climates, but this may be a spurious result considering the small sample size for Polar climates, and the fact that a few of the CEC respondents had responsibility for drought-monitoring continent-wide and the survey was not structured to individualize responses from such individuals (the indicators they selected were applied to all of the climate subzones in their area of responsibility).

5.2. General Insights Regarding the Use of the Drought Indices and Indicators

As noted by survey respondents and workshop participants, seasonal variations in precipitation and temperature impact the performance of indices and indicators that are based on these variables. The indices and indicators include SPI, various measures of precipitation anomaly (percentiles, ranks, departure from normal, percent of normal), SPEI, EDDI, ESI, and various measures of temperature anomaly (departure from normal, ranks). Precipitation during the wet season will have a greater impact on drought evolution than precipitation during the dry season, so precipitation-based indices are more effective during the wet season. Likewise, evapotranspiration during the warm season will have a greater impact on drought evolution than during the cold season (for those climates with pronounced seasonal swings in temperature), so temperature and evapotranspiration-based indices will be more effective during the warm season.

Continental and Polar climates typically have cold winters during which the ground becomes frozen for a period of time. The state of the ground affects water flow, so soil moisture-based indices and indicators are most useful during the warm season when the ground is unfrozen. These indices and indicators include SMA, SMDI, and soil moisture observations and model output (from the CEC survey results), as well as satellite-based indices not mentioned in the survey results, such as SMOS (Soil Moisture and Ocean Salinity satellite) soil moisture anomaly and SMAP (Soil Moisture Active Passive) volumetric measurements. Hydrological indices such as streamflow and ephemeral ponds are most useful during the warm season when the streams and ponds are unfrozen.

In cold climates (E, D, and some C and B climates), vegetation becomes dormant during the winter, which renders indicators and indices that incorporate vegetation health less useful during winter. These include NDVI, Enhanced Vegetation Index (EVI), CMI, and VegDRI. Crop status for annual crops (in B, C, D, and E climates) is most effective during the growing season, while crop status for perennial crops (in A climates) is effective year-round.

The SPI was rated highly in many climate regions. This index is an effective measure of the water supply component in the drought equation and provides a statistically precise historical context. The SPEI is a better measure than the SPI, especially for flash droughts, because it measures the total

moisture status -- i.e., both water supply and water demand. But SPEI was not rated as highly as SPI in the CEC survey. This may be due to precipitation data being more widely available than temperature data (only precipitation data is needed to compute SPI while both precipitation and temperature data are needed to compute SPEI). At a December 2009 workshop¹² in Lincoln, Nebraska, the organizers recommended that the SPI be used to characterize meteorological droughts around the world [79], so this may also be a factor in the SPI's popularity.

Percent of normal precipitation (PNP) was also rated highly in many climate regions. Many reasons were given by participants at the 8 October CEC webinar for why they used the PNP over the SPI. These include:

- greater familiarity with PNP than SPI,
- PNP may be in a regional drought response plan while the SPI is not,
- period of record may not be long enough to compute the SPI but is long enough for PNP.

It was noted at the webinar that, unlike the SPI or precipitation percentiles, PNP by itself only relates precipitation to some base period mean, it does not provide a historical context for the precipitation anomaly. SPI is therefore a better index than PNP for assessing the severity of a drought. PNP can be supplemented with historical information to overcome this shortcoming.

In situ-based drought indicators and indices provide information for point locations and are based on observed hydrometeorological data. Satellite-based indices and indicators provide better spatial coverage than *in situ* data in areas where station networks are sparse, but they are derived from algorithms based on radiometric data and the satellite record is generally shorter. The shortcomings for both *in situ* and remotely-sensed data can be overcome by merging the datasets. For example, statistical or machine learning techniques could be applied to place satellite-based indices into a firmer historical context. One example of this sort of middle ground is McRoberts and Nielsen-Gammon [80], which combines the spatial coverage of radar with the historic record of *in situ* stations.

The USDM was scored in the CEC study as consistently effective to very effective across most climate types and subzones for both short-term drought and long-term drought. This is likely due to several factors:

Its composite nature -- The USDM, like the CDM, MSM, and NADM, integrates the individual drought indices and indicators that are appropriate for the region and season being analyzed. It also integrates drought impact information that is provided by local field observers.

The USDM has some longevity and is widely accepted as an official product. With the majority of the survey respondents from the U.S., and the USDM a recognized official product with extensive media exposure, the respondents are likely more familiar with the USDM than they are with some of the other indicators and indices.

The U.S. borders on both Canada and Mexico with climatic conditions overlapping the borders, making the USDM useful to all three countries, especially in areas near the border. For example, the vast majority of the Canadian population lives in southern Canada, near the U.S. border.

The importance of ecological drought—particularly in Temperate and Continental climates—was highlighted in numerous engagements described earlier in this paper. To more effectively monitor ecological drought and drought impacts to ecosystem goods and services, further research is needed to understand ecological drought sensitivity and ecosystem responses to drought and to identify drought indices and indicators that are most appropriate for ecological drought. These and other key ecological drought research gaps and needs were identified through a series of four webinars co-organized by NIDIS and then USGS National Climate Adaptation Science Center in 2021.

¹² The Inter-Regional Workshop on Indices and Early Warning Systems for Drought was sponsored and organized by the WMO, NDMC, University of Nebraska-Lincoln School of Natural Resources, NOAA/NIDIS, USDA, and the United Nations Convention to Combat Desertification Secretariat.

5.3. Recommendations on the Use of Drought Indices and Indicators

Vicente-Serrano et al. [81] noted “the necessity of testing and comparing the local performance of different drought indices to select the most appropriate one according to the variable of interest” (emphasis added) and “studies comparing the performance of several drought indices, like those evaluated here, would be preferable to determine the best drought index for identifying a certain drought type and its impacts on different systems.” It was in this spirit that the CEC study and the user engagement workshops described herein have been conducted—to identify those drought indices that are most appropriate for local climatic conditions, here based on user preference. But it is acknowledged here that further research employing empirical studies based on objective data is still needed. In this section we combine the survey results with a literature review of some of the existing objective research. It is hoped that this effort provides a guide for selecting the appropriate drought index for specific climate zones.

5.3.1. Overview of Published Objective Research

This section provides an overview of published literature that has objectively evaluated the effectiveness of drought indices and indicators. A more detailed discussion of the literature can be found in Appendix S-E in the Supplementary Material.

AghaKouchak et al. [82], Anderson et al. [83], and Otkin et al. [84] evaluated vegetation cover condition, as sampled by remotely sensed shortwave vegetation indices, and concluded that they may be effective in detecting rapidly changing conditions related to flash drought. Houborg et al. [85], McDonough et al. [86], Tavakol et al. [87], Ma et al. [88], and Beck et al. [89] evaluated new indices that have been developed to monitor soil moisture using orbital sensors. The GRACE (Gravity Recovery and Climate Experiment), SMOS (Soil Moisture and Ocean Salinity), and SPoRT (Short-term Prediction Research and Transition) products show promise in measuring soil moisture during the warm season.

Several researchers [15–17,51,66,81,90–103] have evaluated the effectiveness of the SPI and SPEI. Dai [51], Ellis et al. [94], Quiring [97], and White and Walcott [95] noted that the SPI is based only on precipitation and does not consider evapotranspiration, which makes it more suited to monitoring meteorological and hydrological droughts rather than agricultural drought, although Quiring [97] and Vicente-Serrano et al. [81] concluded that it is more effective than PDSI, Palmer Z Index, Effective Drought Index (EDI), and percent of normal precipitation. Quiring [97] also noted that, since percent of normal precipitation only relates precipitation to a base period and not to the historical variation of precipitation, it cannot be used to compare drought conditions over space or time. Naumann et al. [90], Vicente-Serrano et al. [92], and Homdee et al. [98] determined that the SPEI is more effective when compared with the SPI, but Vicente-Serrano et al. [91] concluded that the SPEI shows different sensitivity to precipitation and reference evapotranspiration as a function of the climatology and Faiz et al. [99] noted that the SPEI is not suitable for colder regions where winter temperatures are mostly below zero and potential evapotranspiration is essentially zero. Seasonality is also important in the application of the SPI [22,66,101], with the SPI more effective during the wet season than during the dry season. Yihdego et al. [15] noted that the SPI's short-term time scales (e.g., 1-month SPI) are more useful for evaluating soil moisture and crop stresses, while long time scales (e.g., 12-month SPI) have been tied to streamflow, reservoir, and groundwater levels. Dai [51] noted that the PDSI and Palmer Hydrological Drought Index (PHDI) do not work well over mountainous and snow-covered areas, but PDSI can be used as a drought index over the low and middle latitudes.

Multiple index comparisons were done by several researchers [21,22,91,104,105]. Keyantash and Dracup [21] found that Rainfall Deciles and the SPI were the best drought indices for monitoring meteorological drought, Total Water Deficit was best for hydrological drought, and Computed Soil Moisture was best for agricultural drought, and that the precipitation anomaly was not especially informative because of its lack of historical context. Wanders et al. [22] determined that all indicators, except those based only on precipitation, have difficulty in EF climates, where temperatures are below freezing for much of the year; this is especially true for streamflow and soil moisture, but also problematical for precipitation which accumulates as snow that will melt only when temperatures

rise above freezing. Vicente-Serrano et al. [91] showed that the highest correlations of SPEI and Reclamation Drought Index (RDI) with precipitation occurred in Mexico, the CONUS, and western and southern Canada. The highest correlations with reference evapotranspiration occurred from northern Mexico to the western CONUS and across western Canada. The lowest correlations with both precipitation and reference evapotranspiration occurred over parts of Alaska and northern and eastern Canada (boreal regions of North America). Low correlations with reference evapotranspiration also occurred in equatorial regions. In a study to quantify the time taken for drought to evolve from precipitation deficits to deficits in soil moisture or streamflow, Gevaert et al. [105] determined that drought propagation is strongly related to climate type. They concluded that 1) droughts propagate slower in dry and continental climates and quicker in tropical climates, and 2) winter season drought propagation tends to be slower than in the summer, especially in tropical savanna and continental climates.

5.3.2. Recommendations Summary

The indices and indicators recommended for use in the five climate zones, based on the references cited above and in Appendix S-E, plus Jain et al. [106], Li et al. [107], Lweendo et al. [108], and Morid et al. [109], are summarized in Table 1. The discussion following the table is based on these studies and the CEC study and workshops discussed previously.

Table 1. Recommended drought indices and indicators, based on published literature. Climate Zone: A = Tropical, B = Dry, C = Temperate, D = Continental, E = Polar Tundra.

Climate Zone	Drought Index or Indicator
A	SPI useful
	SPEI useful in monsoon climates, but is more dependent on variation of precipitation than evapotranspiration
	PDSI can be useful
	EDI useful in monsoon climates
B	SPEI useful, but is more dependent on evapotranspiration than precipitation
	SPI and deciles/percentiles have difficulty for seasons/years with no precipitation
	SPI more useful during the wet season
	EDI useful
	CMI & NDVI from mid-June to mid-July for BS climates
In BW climates, soil moisture and hydrological drought indicators should not be used	
C	PDSI can be useful; SPI more useful than PDSI; and SPEI more useful than SPI, especially in summer
	SPI and deciles/percentiles have difficulty for seasons/years with no precipitation
	SPI and SPEI time scales of 1 to 3 months are relevant for agricultural applications and longer time scales (e.g., 6 to 12 months) for hydrological applications
	ESI effective in detecting rapidly-evolving agricultural drought situations
	Soil moisture indices and Palmer Z Index are effective for monitoring agricultural drought
	EDI useful
D	SPI and deciles/percentiles have difficulty for seasons/years with no precipitation
	SPI more useful during the wet season
	SPEI more useful than SPI, especially in summer
	SPEI less useful during winter in cold climates

	PDSI can be useful in mid-latitude D climates
	Vegetative Health Index (VHI) should be used with caution in cold climates and in winter
	ESI effective in detecting rapidly-evolving agricultural drought situations
	Soil moisture indices and Palmer Z Index are effective for monitoring agricultural drought
E	VHI should be used with caution
	SPEI less useful in E climates than other climates
	Streamflow and soil moisture drought indicators face difficulties in EF climates

Tropical (A) climates located near the equator (such as the USAPI) have high evapotranspiration year-round, which eliminates any advantage potential evapotranspiration-based indices have over precipitation-based indices. Weekly or monthly minimum precipitation thresholds are used as a trigger for drought in the USAPI; once a drought has been established, SPI values, precipitation percentiles or ranks, and impacts are used to determine the drought intensity. For tropical climates with precipitation year-round, the indices generally perform well throughout the year. Tropical climates located farther from the equator typically experience strong seasonal variations in precipitation, with the dry season occurring in winter. For such Aw climates, precipitation-based indices are most useful during the wet season.

Arid and semiarid (B) climates experience a chronic climatological lack of precipitation. In southern Alberta and southwestern Saskatchewan, and much of the western CONUS, winter is the wet season when large-scale synoptic systems bring widespread heavy precipitation, and temperatures are cold enough for much of the precipitation to fall as snow in the mountains. The winter mountain snowpack serves as a “water tower”, and melting of the mountain snowpack during spring and summer feeds streams which provide a water source for the Southwest, Intermountain West, and much of the Pacific Northwest and West Coast during the summer dry season. Reservoirs collect streamflow and also serve as an important water bank for the West, where most of the agriculture is irrigated. Given the distinct dry season in the U.S. West and the extensive system of man-made reservoirs created to store and manage water supply, it is of utmost importance to use multi-year timescales for indices like SPI and SPEI when considering hydrologic drought. Extreme water shortages often take two years or longer to develop; a single year of drought can be buffered with little impact if large reservoirs are already near full. This has led to the development of drought monitoring tools that offer access to drought index time scales up to 72 months (e.g., [110,111]).

In addition, a recent study led by the University of Arizona highlights the complexities of monitoring drought in irrigated agriculture in the southwestern U.S. The study found that, where water supplies for irrigated crop production are in a different location than the land on which crops are produced, effective monitoring of drought impacts to agriculture may require the use of drought indicators, e.g., SPI and SPEI, for the watershed of the water source location, rather than the crop production area. This study emphasized the importance of considering the geographic and temporal scope of drought indices, the physical infrastructure and institutions that manage irrigation water, and the dynamic between water sources and the agricultural operations that use them in irrigated agricultural production.

Summer monsoon precipitation is critical for the U.S. Southwest and summer is the wettest time of year for many locations. As with Aw climates, in B climates, precipitation anomalies during the dry season can result in substantial SPI values but corresponding precipitation departures are small, so (as noted earlier) precipitation anomalies during the wet season are more likely to result in drought development, intensification, or amelioration. This makes drought indices based on precipitation amount, such as the SPI, less useful during the dry season and underscores the importance of using indices that incorporate atmospheric evaporative demand (like SPEI) or rely solely on evaporative demand (like the EDDI) at monthly-to-seasonal timescales. The dry season in the U.S. West is also the fire season and precipitation-based indices give little information with regards to in-season short-

term fire danger. Drought indices that incorporate evaporative demand are therefore more appropriate for fire danger monitoring and fire management applications [112].

The chronic lack of available water in arid climates hinders the growth and development of vegetation, and limits actual evapotranspiration. When precipitation does occur, as from summer monsoon cells in the Southwest, vegetation can rapidly green up. This makes satellite-based greenness indicators important, not only to assess productivity of grasses growing, but also as an indicator of where isolated monsoon cells have dumped rainfall and the spatial extent of that rainfall, which is often missed due to low spatial density of station networks and radar gaps.

The temperature and precipitation regimes of the Temperate (C) climate zone make it ideally suited for growing annual crops. Sufficient precipitation is needed to maintain adequate soil moisture at certain times in the crop's growth cycle, and high evapotranspiration can severely damage the crop, so precipitation-based and soil moisture-based drought indices are important. Satellite-based indicators can monitor the health, not only of crops, but of vegetation in the broader ecosystem. Temperate climates experience seasonal extremes in temperature (which contributes to evapotranspiration; it is noted that insolation, humidity, and wind speed also contribute to evapotranspiration [113,114]) and some have pronounced wet and dry seasons. Precipitation anomalies during the dry season are typically smaller than during the wet season; as a result, precipitation-based indices (PNP, SPI, and SPEI) are more effective during the wet season. Soil moisture measurements and indices are most important for agriculture during the growing season. Drought indices that measure some component of evapotranspiration (such as SPEI and ESI) are most effective when evapotranspiration is most significant, i.e., during the warm season. The combination of low precipitation and high evapotranspiration can contribute to the rapid development of "flash droughts", which makes a combination of precipitation-based and evapotranspiration-based indices especially important. Streamflow provides the water source for filling reservoirs, which makes streamflow and reservoir storage important drought indicators for much irrigated agriculture in Cs climates.

The temperature and precipitation regimes of the warmer (a and b) and wetter (during the growing season) (f and w) subzones of the Continental (D) climate zone make it, like the Temperate climate zone, well-suited for growing annual crops. These D subzones are primarily limited to the northern U.S. and southern Canada. Unlike C climates, cold winter temperatures in D climates typically freeze the soils, with winter precipitation falling in frozen form (snow) and natural vegetation largely going dormant. The frozen soil and dormant vegetation limit the utility of soil moisture-based and vegetation-based indicators to the warmer times of the year. Below-freezing temperatures in winter reduce the amount of moisture the air can hold relative to the summer, so winter precipitation amounts are generally less than summer precipitation amounts (this is especially the case for the w subzones). The drier winters, and the fact that the snow accumulates during the winter and does not enter into the hydrological system until spring snowmelt, limit the usefulness of precipitation-based indices (such as SPI, PNP, and precipitation percentiles) during the cold season. The water content of the snowpack (SWE) is a better indicator of drought conditions during the winter than SPI, PNP, or total precipitation percentiles, but even then SWE percentiles at the peak of the snow season (typically April) are a better measure than SWE percentiles earlier in the snow accumulation season. The PDSI integrates water supply (precipitation) and water demand (potential evapotranspiration) with a soil moisture component. The CEC survey results indicate it is more effective in the warmer D subzones than the colder ones. The factors above have been behind many of the criticisms of the PDSI [5]. Modifications of the PDSI that address frozen soils and precipitation type (such as the Alberta model [115,116]) have improved its performance in D climates, and the self-calibrating PDSI (scPDSI) [117] addresses issues with how the weighting factor K is computed in the original version.

Unique astronomical, geographical, hydrological, and meteorological factors complicate drought monitoring in the more northern D climate subzones of northern Canada and much of Alaska. The high latitude location minimizes sunlight during the winter (zero solar insolation at the winter solstice north of the Arctic Circle), and long summer days maximize solar insolation and, consequently, potential evapotranspiration, during the warm season. This results in a short but intense growing season for

those limited agricultural areas. There is no significant irrigation for agriculture in Alaska, so precipitation during June-August is crucial, and the short growing season magnifies impacts on crops. Snow cover is typically continuous from 2 to 8 months, reducing the importance of short term (i.e., monthly or less) precipitation deficits. During the warm season, streamflow is fed by snowmelt, rainfall, and groundwater, as well as meltwater from mountain glaciers. Streamflow needs to be used with caution as a drought indicator since a summer warm/dry spell could result in above-normal glacier-fed streamflow as opposed to below-normal streamflow which would be the case in the warmer A-C climate zones. In the winter, rivers are frozen which renders streamflow ineffective as a drought indicator during this time of the year. The permafrost layer is frozen year-round, but soil moisture in the active layer (the region where the ground thaws out above the permafrost layer) can be a factor in the warm season. As the soil thaws, even if it doesn't rain, the active layer will be wet; but it can dry out later in the warm season if precipitation is deficient. Frozen soils and continuous, long duration snow cover render remotely-sensed and modeled soil moisture indicators ineffective during the cold season, so they are most useful during the warm season.

5.4. Correlation of Indices with Drought Impacts and Vulnerabilities

As noted earlier in this paper, an index or indicator is only as good as its correlation with actual or potential drought impacts. The importance of correlating an indicator with actual conditions in the field was noted by many CEC survey respondents. Identification of drought indicators and indices that objectively correlate with drought impacts requires robust, sufficiently comprehensive databases of long-term drought impact data. While selected drought impact information is gathered for some locations and research projects have quantified and/or modeled the impacts of drought on specific sectors, significant improvements are needed in documenting drought impacts. For example, the 2021 workshop's regional listening sessions in Alaska highlighted the impacts of a changing climate on local communities whose livelihoods are dependent on local resources. The loss of sea ice, changing animal migration patterns, and increased mortality of traditional foods, including seals, birds, caribou, fish, berries, and various plants, are a direct threat to the livelihoods and cultures of these communities.

In addition, drought impacts are modulated or exacerbated by ecological, socioeconomic, cultural, historical, and other assets and vulnerabilities. Important research questions center on improving our understanding of drought vulnerabilities and impacts and linking this information to drought indices to enable more reliable estimations of current and future impacts. This research is critical for effective mitigation of impacts and for targeted support for the most vulnerable, e.g., those subsistence-based communities who do not currently qualify for relief programs in the U.S. Some key gaps in understanding drought impacts include vulnerabilities of and impacts to socio-cultural and ecological systems; water security; community livelihoods and culture; economic sectors, including recreation; tourism; energy; and health.

5.5. Evolving Drought Indicators for a Warming Climate

Future drought may be substantially different from current and historical drought in frequency, severity, and extent [118,119], so it is important to understand how drought impacts and the nature of drought itself are changing over time [120] in order to evolve drought indicators. For example, in the southwestern U.S., average annual precipitation has not changed significantly over the last 50 years, but aridity (lack of water availability) is increasing across the region due to temperature increases which increase evaporation and contribute to early snow melt. Recent research [121] characterizing the intensity, duration, and distribution changes of snow drought (SWE deficits) from 1980 to 2018 found that snow droughts intensified, lengthened, and became more common across the western U.S. and are projected to continue to increase across western North America (e.g., [122]). Even when precipitation is at or above normal at higher elevations, warming temperatures can result in the precipitation falling as rain rather than snow, thereby weakening the SWE-precipitation relationship. Given the vital role of snow in the western U.S. and Canada for water supply, hydropower, agricultural production, ecosystems, etc., the study concludes that "characterizing

snow deficits (snow droughts) in a changing climate has emerged as a critical knowledge gap". That characterization is essential to adapt drought indicators and indices for the region and to incorporate accurate SWE measures into drought indices.

Similarly, at the 2021 drought workshops in Alaska, where temperatures are rising, resulting in rain rather than snow, and permafrost is rapidly degrading, participants noted that increased temperatures and the increasing length of the warm season (from 4 to 6 months) will increase evapotranspiration, thereby altering the signatures of drought. However, drought in high latitude environments is complicated and not well understood. For example, measuring evapotranspiration in Alaska is complex given the prolonged day-length in summer and short day-length in winter, and the melting of permafrost will affect the amount and depth of available soil moisture. Finally, there is limited data on key drought parameters, so improved instrumentation, continued maintenance of existing stations, and additional research are needed to understand how these and other variables will change in a warming climate and to adapt drought indices accordingly.

Studies cited by the IPCC [123] indicate that there is an increase in the frequency and intensity of agricultural and ecological droughts and, in some parts of the world, meteorological and hydrological droughts. Also, there is a trend toward more short-term droughts in the U.S. than long-term droughts [124] and flash droughts are becoming more frequent globally [125]. Climate change studies indicate that these trends will continue [123]. Hasegawa et al. [126] and Gusyev et al. [127] evaluated how drought will change under climate change scenarios using a comparative approach (calculating comparative standardized indices of future climates using present climate as a reference). But perhaps more significant than the fact that the nature of drought will change is the impact of climate change on stationarity of data, for this affects the foundation of all statistically-based drought indices [128]. In order to put current conditions into a historical context, raw values are standardized into anomalies using probability statistics. A reference period of record of at least 50-60 years is needed in order to empirically compute 50-year return periods (or a probability of recurrence of 0.02 which is the threshold for USDM exceptional [D4] droughts). Hoyleman et al. [128] argue that drought metric error and bias may be introduced where climate has shifted substantially from the time-integrated period-of-record distribution. They conclude that non-stationarity in datasets caused by anthropogenic climate change needs to be accounted for in drought assessments based on statistically-derived drought indices.

As temperatures warm and precipitation regimes change with a changing climate, the location and shape of the Köppen climate zones will change. This is not expected to affect the applicability of specific drought indices to specific climate zones, as only the geographic location of the climate zones will change, not the definition of the zones.

5.6. Embedding Indicators Within Robust Decision Systems

The goal of improving drought indicators is to inform and improve decisions that increase resilience to drought and reduce the impacts of drought. Recommendations for strengthening the use of indicators in drought decision-making (drought management and adaptation strategies) include:

Drought managers and decision-makers, who use drought indices and indicators in their operations, are encouraged to follow the guidance in Section 5.3.2 on the use of drought indices and indicators. This review paper, and especially Section 5.3.2, is intended to serve as a "toolbox" to enable users to better utilize drought indices and indicators that are appropriate for their climate zone and their applications.

In several user engagements described in Supplementary Materials Appendices S-A and S-D, it was noted that indicators need to be embedded in decision processes where a wide range of decision makers -- including government, water providers, sectors impacted by drought, etc. -- regularly communicate about drought in the region before drought hits. This is important for all regions, including those with limited drought experience. For example, participants at the 2021 Alaska drought webinar series noted that communities and decision-makers associate drought with other regions of the country, such as the southwestern U.S., not Alaska. This highlights the need for targeted drought communication that is specific to Alaska. Improving awareness of the nature of

drought and communicating about drought, its impacts, and the changing nature of drought with climate change across all climate zones is fundamental to catalyze proactive drought management.

While some indices have been designed to be universally applicable, no single drought index adequately describes all types of drought for all locations. The importance of location-specific indices was reflected in CEC survey responses, in which users commented that a major factor in their choice of indicator or indices was how “location-specific” the indicator is. Location-specific indices may capture the nuances of drought in relation to the local geography and soil types, water use and management, cultural considerations, sectors of importance locally, etc. However, given that droughts often extend over large geographic areas and multiple jurisdictional boundaries and impact shared surface and groundwater resources, interconnected economic sectors, social structures, etc., it is important for location-specific drought monitoring to be embedded within or linked to larger-scale drought monitoring and decision systems. Collaboration and communication among entities responsible for drought monitoring, mitigation, and response from local to national scales is important to ensure that local monitoring data, impact data, and information is compiled across regions to inform effective mitigation, response and relief strategies at all scales.

Establishing a central, online location for drought indicators and indices could support improved access to and use of these tools for monitoring drought. An online resource could also support continued refinement and improvement of drought indicators and indices, especially to address issues caused by the changing climate such as the non-stationarity issue. Finally, it could allow users to share insights regarding the relevance and applicability of indicators and indices in specific locations, for specific sectors, etc. as a form of peer-to-peer learning that strengthens drought management decisions.

References

1. Abbas, A.; M. Waseem, W. Ullah, C. Zhao, and J. Zhu, 2021: Spatiotemporal Analysis of Meteorological and Hydrological Droughts and Their Propagations. *Water* 13, 2237. <https://doi.org/10.3390/w13162237>
2. Elahi, E., Z. Khalid, M.Z. Tauni, H. Zhang, and X. Lirong, 2022: Extreme weather events risk to crop-production and the adaptation of innovative management strategies to mitigate the risk: A retrospective survey of rural Punjab, Pakistan. *Technovation*, 117, 102255. <https://doi.org/10.1016/j.technovation.2021.102255>.
3. World Meteorological Organization, 2021: WMO Atlas of Mortality and Economic Losses From Weather, Climate and Water Extremes (1970-2019). WMO No. 1267, WMO, 90 pp.
4. CRED (Center for Research on the Epidemiology of Disasters), 2023: 2022 Disasters in numbers. Brussels: CRED. This document is available at: https://cred.be/sites/default/files/2022_EMDAT_report.pdf
5. Heim, R.R., Jr., 2002: A Review of Twentieth-Century Drought Indices Used in the United States. *Bull. Amer. Meteor. Soc.*, 83:1149-1165. <https://doi.org/10.1175/1520-0477-83.8.1149>
6. Sheffield, J. and E. F. Wood, 2011: Drought: Past Problems and Future Scenarios. Earthscan, London. 210 pp.
7. Wilhite, D. A., 2000: 1. Drought as a Natural Hazard. In Wilhite, D. A. (ed.), 2000: Drought: A Global Assessment, Volume I. Routledge, London, pp. 3-18.
8. Gillette, H. P., 1950: A creeping drought under way. *Water and Sewage Works*, March, 104-105.
9. Wilhite, D. A., 2011: National Drought Policies: Addressing Impacts and Societal Vulnerability. In Sivakumar et al. (eds.), 2011: Towards a Compendium on National Drought Policy: Proceedings of an Expert Meeting, July 14-15, 2011, Washington DC, USA. AGM-12/WAOB-2011, World Meteorological Organization, Geneva, pp.13-22.
10. Albuquerque Journal, 11-14-20: New Mexico ranchers face historic drought. <https://www.abqjournal.com/1518091/new-mexico-ranchers-face-historic-drought.html>
11. Otkin, J. A., M. Svoboda, E.D. Hunt, T.W. Ford, M.C. Anderson, C. Hain, and J.B. Basara, 2018: Flash droughts: A review and assessment of the challenges imposed by rapid-onset droughts in the United States. *Bulletin of the American Meteorological Society*, 99(5), 911-919.
12. Lisonbee, J., M. Woloszyn, and M. Skumanich, 2021: Making sense of flash drought: definitions, indicators, and where we go from here. *Journal of Applied and Service Climatology*, vol. 2021, issue 001, 19 pp. DOI: doi.org/10.46275/JOASC.2021.02.001
13. World Meteorological Organization, 1992: International Meteorological Vocabulary. 2d ed. WMO No. 182, WMO, 784 pp.
14. American Meteorological Society, 1997: Meteorological drought—Policy statement. *Bull. Amer. Meteor. Soc.*, 78:847–849.

15. Yihdego, Y., B. Vaheddoost, and R.A. Al-Weshah, 2019: Drought indices and indicators revisited. *Arab J Geosci* 12, 69 (2019). <https://doi.org/10.1007/s12517-019-4237-z>
16. Stagge, J.H., I. Kohn, L.M. Tallaksen, and K. Stahl, 2015: Modeling drought impact occurrence based on meteorological drought indices in Europe. *Journal of Hydrology*, 530, 37-50. <https://doi.org/10.1016/j.jhydrol.2015.09.039>.
17. Stefanidis S., D. Rossiou, and N. Proutsos, 2023: Drought Severity and Trends in a Mediterranean Oak Forest. *Hydrology*. 10(8):167. <https://doi.org/10.3390/hydrology10080167>
18. Crausbay, S.D., A.R. Ramirez, S.L. Carter, M.S. Cross, K.R. Hall, D.J. Bathke, J.L. Betancourt, S. Colt, A.E. Cravens, M.S. Dalton, J.B. Dunham, L.E. Hay, M.J. Hayes, J. McEvoy, C.A. McNutt, M.A. Moritz, K.H. Nislow, N. Raheem, and T. Sanford, 2017: Defining Ecological Drought for the Twenty-First Century. *Bull. Amer. Meteor. Soc.*, 98(12), 2543-2550. Doi: <https://doi.org/10.1175/BAMS-D-16-0292.1>
19. World Meteorological Organization and Global Water Partnership, 2016: Handbook of Drought Indicators and Indices (M. Svoboda and B.A. Fuchs). Integrated Drought Management Programme (IDMP), Integrated Drought Management Tools and Guidelines Series 2. Geneva.
20. Palmer, W.C., 1965: Meteorological drought. U.S. Weather Bureau Research Paper 45, 58 pp. [Available from NOAA Library and Information Services Division, Washington, DC 20852.]
21. Keyantash, J. and J.A. Dracup, 2002: The Quantification of Drought: An Evaluation of Drought Indices. *Bull. Amer. Meteor. Soc.*, 83:1167-1180. <https://doi.org/10.1175/1520-0477-83.8.1167>
22. Wanders, N., H.A.J. Van Lanen, and A.F. Van Loon, 2010: Indicators for Drought Characterization on a Global Scale. WATCH Technical Report 24. <https://library.wur.nl/WebQuery/wurpubs/fulltext/160049> (accessed 6-27-2022).
23. Zargar, A., R. Sadiq, B. Naser, and F.I. Khan, 2011: A review of drought indices. *Environmental Reviews*. 19(NA): 333-349. <https://doi.org/10.1139/a11-013>
24. Peters-Lidard, C.D., D.M. Mocko, L. Su, D.P. Lettenmaier, P. Gentine, and M. Barlage, 2021: Advances in Land Surface Models and Indicators for Drought Monitoring and Prediction. *Bull. Amer. Meteor. Soc.*, 102(5), E1099-E1122. Published online: 01 Jun 2021. <https://doi.org/10.1175/BAMS-D-20-0087.1>
25. Svoboda, M., D. LeCompte, M. Hayes, R. Heim, K. Gleason, J. Angel, B. Rippey, R. Tinker, M. Palecki, D. Stooksbury, D. Miskus, and S. Stephens, 2002: The Drought Monitor. *Bull. Amer. Meteor. Soc.*, 83:1181-1190.
26. NOAA, 2016. Strategic Plan: NOAA's National Environmental Satellite, Data, and Information Service. NOAA/NESDIS Rep. Available online: www.star.nesdis.noaa.gov/star/documents/matrix/NESDIS_Strategic_Plan_2016.pdf
27. Norman, D, 1988. *The Design of Everyday Things*; Doubleday: New York, NY, USA.
28. Schuler, D.; Namioka, A. (Eds.), 1993. *Participatory Design: Principles and Practices*; Lawrence Erlbaum Associates: Hillsdale, NJ, USA.
29. Abras, C.; Maloney-Krichmar, D.; Preece, J., 2004: User-centered design. In *Berkshire Encyclopedia of Human-Computer Interaction*; Bainbridge, W., Ed.; Sage Publications: Thousand Oaks, CA, USA; Volume 2, pp. 763-768.
30. Bødker, K.; Kensing, F.; Simonsen, J., 2004: *Participatory IT Design: Designing for Business and Workplace Realities*; MIT Press: Cambridge, MA, USA.
31. Spinuzzi, C., 2005: The methodology of participatory design. *Tech. Commun.*, 52, 163-174.
32. Oakley, N.S.; Daudert, B., 2016: Establishing best practices to improve usefulness and usability of web interfaces providing atmospheric data. *Bull. Am. Meteor. Soc.*, 97, 263-274.
33. Kruk, M.C.; Vose, R.; Heim, R.; Arguez, A.; Enloe, J.; Yin, X.; Wallis, T., 2018: Drought amelioration: An engagement-to-implementation success story. *Bull. Amer. Meteor. Soc.*, 99, 2457-2462.
34. Lawrimore, J., R.R. Heim Jr., M. Svoboda, V. Swail, and P.J. Englehart, 2002: Beginning a New Era of Drought Monitoring Across North America. *Bull. Amer. Meteor. Soc.*, 83:1191-1192.
35. Heim, R.R., Jr. and M.J. Brewer, 2012: The Global Drought Monitor Portal: The Foundation for a Global Drought Information System. *Earth Interact.*, 16:1-28. doi: <http://dx.doi.org/10.1175/2012EI000446.1>
36. Heim, R.R., Jr., C. Guard, M.A. Lander, and B. Bukunt, 2020: USAPI USDM: Operational Drought Monitoring in the U.S.-Affiliated Pacific Islands. *Atmosphere* 2020, 11, 495 (9 pages), <https://doi.org/10.3390/atmos11050495>
37. University of Maryland, 2018: Drought in the U.S. Affiliated Pacific Islands—Climate Change Impacts and Community Resiliency across a Diverse Landscape. University of Maryland Center for Environmental Science, Integration, and Application Network Newsletter 617. Available online: https://ian.umces.edu/pdfs/ian_newsletter_617.pdf (accessed on 10 June 2021).
38. Borja, J., J. Deenik, A. Frazier, and C. Giardina, 2019: Drought in the U.S Affiliated Pacific Islands: Impacts to Agriculture. Workshop Report. Available online: <https://www.sciencebase.gov/catalog/item/5cf81feee4b07f02a70465e2> (accessed on 10 June 2021).

39. Cordell, S., A. Frazier, C. Trauernicht, and Y.-P. Tsang, 2019: Drought in the U.S Affiliated Pacific Islands: Impacts to Ecosystems. Workshop Report. Available online: <https://www.sciencebase.gov/catalog/item/5cf820c0e4b07f02a70465eb> (accessed on 10 June 2021).
40. Anthony, S.S., 2019: Drought in the U.S Affiliated Pacific Islands: Impacts to Water Resources. Workshop Report. Available online: <https://www.sciencebase.gov/catalog/item/5cf82192e4b07f02a70465f6> (accessed on 10 June 2021).
41. Fuchs, B., R. Heim, and D. Simeral, 2019: Drought in the U.S. Affiliated Pacific Islands: Drought Monitoring. Available online: <https://www.sciencebase.gov/catalog/item/5cf82308e4b07f02a7046602> (accessed on 10 June 2021).
42. Bathke, D.J., H.R. Prendeville, A. Jacobs, R. Heim, R. Thoman, and B. Fuchs, 2019: Defining Drought in a Temperate Rainforest. *Bull. Amer. Meteor. Soc.*, 100:2665–2668. <https://doi.org/10.1175/BAMS-D-19-0223.1>
43. CEC, 2021: Guide to Drought Indices and Indicators Used in North America. Montreal, Canada: Commission for Environmental Cooperation. 62 pp.
44. Redmond, K. T. (2002). The depiction of drought: A commentary. *Bulletin of the American Meteorological Society*, 83(8), 1143-1147.
45. Dracup, J.A., K.S. Lee, and E.G. Paulson, Jr., 1980: On the Definition of Droughts. *Water Resources Research*, 16:2, 297-302.
46. Critchfield, H.J., 1974: *General Climatology*, Third Edition. Prentice-Hall, Inc., Englewood Cliffs, New Jersey. 446 pp.
47. Beck, H.E., N.E. Zimmermann, T.R. McVicar, N. Vergopolan, A. Berg, and E.F. Wood, 2018: Present and future Köppen-Geiger climate classification maps at 1-km resolution. *Sci Data* 5, 180214. <https://doi.org/10.1038/sdata.2018.214>
48. Ravelo, A.C., R. Sanz-Ramos, and J.C. Douriet-Cárdenas, 2014: Detecting, assessing and forecasting droughts in the region of the North Pacific Watershed Organism, Mexico. *Agriscientia* 31 (1): 11-24.
49. Romero, D., E. Alfaro, R. Orellana, M.-E. Hernandez Cerda, 2020: Standardized Drought Indices for Pre-Summer Drought Assessment in Tropical Areas. *Atmosphere*, 11, 1209. <https://doi.org/10.3390/atmos11111209>
50. Van Loon, A.F. and H.A.J. Van Lanen, 2013: Making the distinction between water scarcity and drought using an observation-modeling framework, *Water Resources Research*, 49:1483-1502.
51. Dai, A., 2011: Drought under global warming: A review. *Wiley Interdiscip. Rev.: Climate Change*, 2, 45–65, <https://doi.org/10.1002/wcc.81>; Corrigendum, 3, 617, <https://doi.org/10.1002/wcc.190>.
52. Paulo, A.A., and L.S. Pereira, 2006: Drought Concepts and Characterization: Comparing Drought Indices Applied at Local and Regional Scales. *Water Internat* 31:37-49. https://www.researchgate.net/publication/233154227_Drought_Concepts_and_Characterization
53. Vaughn, D.M., 2005: Arid Climates. In: Oliver J.E. (ed), *Encyclopedia of World Climatology*. Encyclopedia of Earth Sciences Series. Springer, Dordrecht. https://doi.org/10.1007/1-4020-3266-8_16
54. Glickman, T.S. (ed), 2000: *Glossary of Meteorology*, Second Edition. American Meteorological Society, Boston.
55. Durrenberger, R.W., 1987: Arid Climates. In: Oliver J.E. and R.W. Fairbridge (eds), *The Encyclopedia of Climatology*. Van Nostrand Reinhold Company, New York, ISBN 0-87933-009-0.
56. Thornthwaite, C.W., 1948: An Approach toward a Rational Classification of Climate, *Geographical Review*, 38:55-94.
57. Bonsal, B.R., X. Zhang, and W.D. Hogg. 1999. Canadian Prairie growing season precipitation variability and associated atmospheric circulation. *Climate Res.*, 11, 191-208.
58. Creswell, R. and F.W. Martin, 1993: *Dryland Farming: Crops & Techniques for Arid Regions*. ECHO Technical Note. Published 1993, revised 1998 by the Education Concerns for Hunger Organization (ECHO).
59. Golla, B., 2021: Agricultural production system in arid and semi-arid regions. *J Agric Sc Food Technol* 7(2): 234-244. DOI: 10.17352/2455-815X.000113
60. National Drought Mitigation Center (NDMC), 2021: Drought Plans/State Plans. <https://drought.unl.edu/Planning/DroughtPlans/StatePlans.aspx> (accessed on 28 December 2021).
61. Nguyen, T. and L. Moeller, 2010: California Drought Contingency Plan. California Department of Water Resources, Sacramento, California. <https://drought.unl.edu/Planning/DroughtPlans/StatePlanning.aspx?st=ca> (accessed 28 December 2021).
62. Oregon Water Resources Department, 2016: Drought Annex: State of Oregon Emergency Operations Plan. Oregon Office of Emergency Management, Oregon Water Resources Department, Salem, Oregon. https://drought.unl.edu/archive/plans/drought/state/OR_2016.pdf (accessed on 28 December 2021).
63. Washington Department of Ecology, 2018: Washington State Drought Contingency Plan. Publication No. 18-11-005, Department of Ecology, Water Resources Program, Olympia, Washington. https://drought.unl.edu/archive/plans/drought/state/WA_2018.pdf (accessed on 28 December 2021).
64. Utah Department of Natural Resources and Division of Emergency Management, 1993: Utah Drought Response Plan, revised 2003 and 2013. Utah Department of Natural Resources and Division of Emergency

- Management. <https://water.utah.gov/wp-content/uploads/2020/04/Drought-Response-Plan.pdf> (accessed on 28 December 2021).
65. Arizona Governor's Drought Task Force, 2004: Arizona Drought Preparedness Plan: Operational Drought Plan. Governor's Drought Task Force. https://drought.unl.edu/archive/plans/drought/state/AZ_2004.pdf (accessed on 28 December 2021).
 66. Colorado Department of Natural Resources, 2018: Colorado Drought Mitigation and Response Plan. Colorado Department of Natural Resources, Colorado Water Conservation Board. https://drought.unl.edu/archive/plans/drought/state/CO_2018.pdf (accessed on 28 December 2021).
 67. Montana Drought Advisory Committee, 1995: The Montana Drought Response Plan. Montana Drought Advisory Committee, Office of the Governor. https://drought.unl.edu/archive/plans/drought/state/MT_1995.pdf (accessed on 28 December 2021).
 68. Quiring, S.M., 2009a: Developing Objective Operational Definitions for Monitoring Drought. *Journal of Applied Meteorology and Climatology*, 48(6), 1217-1229. Retrieved Jun 27, 2022, from <https://journals.ametsoc.org/view/journals/apme/48/6/2009jamc2088.1.xml>
 69. Hanesiak, J.M., R.E. Stewart, B.R. Bonsal, P. Harder, R. Lawford, R. Aider, B.D. Amiro, E. Atallah, A.G. Barr, T.A. Black, P. Bullock, J.C. Brimelow, R. Brown, H. Carmichael, C. Derksen, L.B. Flanagan, P. Gachon, H. Greene, J. Gyakum, W. Henson, E.H. Hogg, B. Kochtubajda, H. Leighton, C. Lin, Y. Luo, J.H. McCaughey, A. Meinert, A. Shabbar, K. Snelgrove, K. Szeto, A. Trishchenko, G. van der Kamp, S. Wang, L. Wen, E. Wheaton, C. Wielki, Y. Yang, S. Yirdaw, and T. Zha, 2011: Characterization and Summary of the 1999–2005 Canadian Prairie Drought. *Atmosphere-Ocean*, 49(4), 421–452. <https://doi.org/10.1080/07055900.2011.626757>
 70. Mather, J.R., 1974: *Climatology: Fundamentals and Applications*. McGraw-Hill Book Company, New York. 412 pp.
 71. Woetzel, J., D. Pinner, H. Samandari, H. Engel, M. Krishnan, N. Denis, and T. Melzer, 2020: Will the world's breadbaskets become less reliable? McKinsey Global Institute, Climate Risk and Response/Case Study: Breadbasket failure, 18 May 2020. <https://www.mckinsey.com/business-functions/sustainability/our-insights/will-the-worlds-breadbaskets-become-less-reliable> (accessed on 28 December 2021).
 72. Shulski, M. and G. Wendler, 2007: *The Climate of Alaska*. University of Alaska Press, Fairbanks, Alaska. 216 pp.
 73. Quiring, S.M., and T.N. Papakryiakou, 2003: An evaluation of agricultural drought indices for the Canadian prairies. *Agricultural and Forest Meteorology*, 118, 49-62. [https://doi.org/10.1016/S0168-1923\(03\)00072-8](https://doi.org/10.1016/S0168-1923(03)00072-8).
 74. Peña-Gallardo, M., S.M. Vicente-Serrano, F. Domínguez-Castro, S. Quiring, M. Svoboda, S. Beguería, and J. Hannaford, 2018: Effectiveness of drought indices in identifying impacts on major crops across the USA. *Climate Research*, 75 (3). 221-240. <https://doi.org/10.3354/cr01519>
 75. ACIA, 2004: *Impacts of a Warming Arctic: Arctic Climate Impact Assessment*. ACIA Overview report. Cambridge University Press. 140 pp. <https://www.amap.no/documents/doc/impacts-of-a-warming-arctic-2004/786>
 76. Beck, P.S.A., G.P. Juday, C. Alix, V.A. Barber, S.E. Winslow, E.E. Sousa, P. Heiser, J.D. Herriges, and S.J. Goetz, 2011: Changes in forest productivity across Alaska consistent with biome shift. *Ecology Letters*, 14(4), 373-379. <https://doi.org/10.1111/j.1461-0248.2011.01598.x>
 77. Churakova Sidorova, O.V., C. Corona, M.V. Fonti, S. Guillet, M. Saurer, R.T.W. Sigwolf, M. Stoffel, and E.A. Vaganov, 2020: Recent atmospheric drying in Siberia is not unprecedented over the last 1,500 years. *Sci Rep* 10, 15024. <https://doi.org/10.1038/s41598-020-71656-w>
 78. Van Loon, A.F., S.W. Ploum, J. Parajka, A.K. Fleig, E. Garnier, G. Laaha, and H.A.J. Van Lanen, 2015: Hydrological drought types in cold climates: quantitative analysis of causing factors and qualitative survey of impacts. *Hydrol. Earth Syst. Sci.*, 19, 1993–2016. <https://doi.org/10.5194/hess-19-1993-2015> (accessed 6/28/2022).
 79. Hayes, M., M. Svoboda, N. Wall, and M. Widhalm, 2011: The Lincoln Declaration on Drought Indices: Universal Meteorological Drought Index Recommended, *Bull. Amer. Meteor. Soc.*, 92(4), 485-488. https://journals.ametsoc.org/view/journals/bams/92/4/2010bams3103_1.xml
 80. McRoberts, D.B., and J. W. Nielsen-Gammon, 2012: The use of a high-resolution SPI for drought monitoring and assessment. *J. Appl. Meteor. Clim.*, 51, 68-83, doi:10.1175/JAMC-D-10-05015.1
 81. Vicente-Serrano, S.M., S. Beguería, J. Lorenzo-Lacruz, J.J. Camarero, J.I. Lopez-Moreno, C. Azorin-Molina, J. Revuelto, E. Moran-Tejeda, & A. Sanchez-Lorenzo, 2012: Performance of Drought Indices for Ecological, Agricultural, and Hydrological Applications. *Earth Interact.*, 16, <https://doi.org/10.1175/2012EI000434.1>.
 82. AghaKouchak, A., A. Farahmand, F.S. Melton, J. Teixeira, M.C. Anderson, B.D. Wardlow, and C.R. Hain, 2015: Remote sensing of drought: Progress, challenges and opportunities, *Rev. Geophys.*, 53, 452–480, doi:10.1002/2014RG000456.
 83. Anderson, M.C., C. Hain, J. Otkin, X.Zhan, K. Mo, M. Svoboda, B. Wardlow, and A. Pimstein, 2013: An Intercomparison of Drought Indicators Based on Thermal Remote Sensing and NLDAS-2 Simulations with U.S. Drought Monitor Classifications. *Journal of Hydrometeorology*, 14:1035-1056.

84. Otkin, J.A., M.C. Anderson, C. Hain, I.E. Mladenova, J.B. Basara, and M. Svoboda, 2013: Examining Rapid Onset Drought Development Using the Thermal Infrared-Based Evaporative Stress Index. *Journal of Hydrometeorology*, 14, 1057-1074. [10.1175/JHM-D-12-0144.1](https://doi.org/10.1175/JHM-D-12-0144.1).
85. Houborg, R., M. Rodell, B. Li, R. Reichle, and B. Zaitchik, 2012: Drought indicators based on model assimilated GRACE terrestrial water storage observations. *Wat. Resour. Res.*, 48, W07525, [10.1029/2011WR011291](https://doi.org/10.1029/2011WR011291).
86. McDonough, K.R., S.L. Hutchinson, J.M.S. Hutchinson, J.L. Case, and V. Rahmani, 2018: Validation and assessment of SPoRT-LIS surface soil moisture estimates for water resources management applications. *J. Hydrology*, 566, 43-54. <https://doi.org/10.1016/j.jhydrol.2018.09.007>.
87. Tavakol, A., V. Rahmani, S.M. Quiring, and S.V. Kumar, 2019: Evaluation analysis of NASA SMAP L3 and L4 and SPoRT-LIS soil moisture data in the United States. *Remote Sensing of Environment*, 229, 234-246.
88. Ma, H., J. Zeng, N. Chen, X. Zhang, M.H. Cosh, and W. Wang, 2019: Satellite surface soil moisture from SMAP, SMOS, AMSR2 and ESA CCI: A comprehensive assessment using global ground-based observations. *Remote Sensing of Environment*, 231, 111215. <https://doi.org/10.1016/j.rse.2019.111215>.
89. Beck, H.E., M. Pan, D.G. Miralles, R.H. Reichle, W.A. Dorigo, S. Hahn, J. Sheffield, L. Karthikeyan, G. Balsamo, R.M. Parinussa, A.I.J.M. van Dijk, J. Du, J.S. Kimball, N. Vergopolan, and E.F. Wood, 2021: Evaluation of 18 satellite- and model-based soil moisture products using in situ measurements from 826 sensors. *Hydrol. Earth Syst. Sci.*, 25, 17-40. <https://doi.org/10.5194/hess-25-17-2021>
90. Naumann, G., E. Dutra, P. Barbosa, F. Pappenberger, F. Wetterhall, and J.V. Vogt, 2014: Comparison of drought indicators derived from multiple data sets over Africa. *Hydrol. Earth Syst. Sci.*, 18:1625-1640. www.hydrol-earth-syst-sci.net/18/1625/2014/[doi:10.5194/hess-18-1625-2014](https://doi.org/10.5194/hess-18-1625-2014)
91. Vicente-Serrano, S.M., G. Van der Schrier, S. Beguería, C. Azorin-Molina, and J.-I. Lopez-Moreno, 2015: Contribution of precipitation and reference evapotranspiration to drought indices under different climates. *J. Hydrol.*, 526, 42–54, <https://doi.org/10.1016/j.jhydrol.2014.11.025>.
92. Vicente-Serrano, S.M., S. Beguería, and J.I. López-Moreno, 2010: A Multiscalar Drought Index Sensitive to Global Warming: The Standardized Precipitation Evapotranspiration Index. *J. Climate*, 23, 1696–1718, <https://doi.org/10.1175/2009JCLI2909.1>.
93. Ma, M., L. Ren, F. Yuan, S. Jiang, Y. Liu, H. Kong, & L. Gong, 2014: A new standardized Palmer drought index for hydro-meteorological use. *Hydrol. Process.* <http://dx.doi.org/10.1002/hyp.10063>.
94. Ellis, A.W., G.B. Goodrich, and G.M. Garfin, 2010: A hydroclimatic index for examining patterns of drought in the Colorado River Basin. *Int. J. Climatol.* 30, 236-255. <https://doi.org/10.1002/JOC.1882>
95. White, D.H. and J. Walcott, 2009: The role of seasonal indices in monitoring and assessing agricultural and other droughts: a review. *Crop & Pasture Science*, 60, 599-616. <https://www.publish.csiro.au/cp/CP08378>
96. Wu, H., M.J. Hayes, D.A. Wilhite, and M.D. Svoboda, 2005: The effect of the length of record on the standardized precipitation index calculation. *Int. J. Climatol.*, 25: 505-520. <https://doi.org/10.1002/joc.1142>
97. Quiring, S.M., 2009b: Monitoring Drought: An Evaluation of Meteorological Drought Indices. *Geography Compass*, 3(1), 64-88. <https://doi.org/10.1111/j.1749-8198.2008.00207.x>
98. Homdee, T., K. Pongput, & S. Kanae, 2016: A comparative performance analysis of three standardized climatic drought indices in the Chi River basin, Thailand. *Agriculture and Natural Resources*. 50. [10.1016/j.anres.2016.02.002](https://doi.org/10.1016/j.anres.2016.02.002).
99. Faiz, M.A., Y. Zhang, N. Ma, F. Baig, F. Naz, and Y. Niaz, 2021: Drought indices: aggregation is necessary or is it only the researcher's choice? *Water Supply*, 1 December 2021; 21 (8): 3987–4002. [doi: https://doi.org/10.2166/ws.2021.163](https://doi.org/10.2166/ws.2021.163)
100. Yang, T., J. Ding, D. Liu, X. Wang, & T. Wang, 2019: Combined Use of Multiple Drought Indices for Global Assessment of Dry Gets Drier and Wet Gets Wetter Paradigm, *Journal of Climate*, 32(3), 737-748. Retrieved Feb 23, 2022, from <https://journals.ametsoc.org/view/journals/clim/32/3/jcli-d-18-0261.1.xml>
101. Hoell, A., T.W. Ford, M. Woloszyn, J.A. Otkin, and J. Eischeid, 2021: Characteristics and Predictability of Midwestern United States Drought. *Journal of Hydrometeorology*, 22(11), 3087-3105. https://www.ssec.wisc.edu/~jasono/papers/hoell_jhm_nov2021.pdf
102. Chen, J., B. Zhang, J. Zhou, and F. Guo, 2023: Temporal and Spatial Changes of Drought Characteristics in Temperate Steppes in China from 1960 to 2020. *Sustainability*. 2023; 15(17):12909. <https://doi.org/10.3390/su151712909>
103. Chong, K.L., Y.F. Huang, C.H. Koo, Ali Najah Ahmed, and Ahmed El-Shafie, 2022: Spatiotemporal variability analysis of standardized precipitation indexed droughts using wavelet transform. *J. Hydrology*, 605, 127299. <https://doi.org/10.1016/j.jhydrol.2021.127299>.
104. Wang Y., C. Zhang, F.R. Meng, C.P. Bourque, and C. Zhang, 2020: Evaluation of the suitability of six drought indices in naturally growing, transitional vegetation zones in Inner Mongolia (China). *PLoS One*. 2020 May 29;15(5):e0233525. doi: 10.1371/journal.pone.0233525. PMID: 32470003; PMCID: PMC7259598.
105. Gevaert, A.I., T.I.E. Veldkamp, and P.J. Ward, 2018: The effect of climate type on timescales of drought propagation in an ensemble of global hydrological models. *Hydrol. Earth Syst. Sci.*, 22, 4649–4665, <https://doi.org/10.5194/hess-22-4649-2018>.

106. Jain, V.K., R.P. Pandey, M.K. Jain, and H.-R. Byun, 2015: Comparison of drought indices for appraisal of drought characteristics in the Ken River Basin. *Weather and Climate Extremes*, 8:1-11. ISSN 2212-0947, <https://doi.org/10.1016/j.wace.2015.05.002>.
107. Li, R., A. Tsunekawa, and M. Tsubo, 2014: Index-based assessment of agricultural drought in a semi-arid region of Inner Mongolia, China. *J Arid Land*, 6(1), 3-15. <https://doi.org/10.1007/s40333-013-0193-8>
108. Lweendo, M.K., B. Lu, M. Wang, H. Zhang, and W. Xu, 2017: Characterization of Droughts in Humid Subtropical Region, Upper Kafue River Basin (Southern Africa). *Water*, 9(4), 242. <https://doi.org/10.3390/w9040242>
109. Morid, S., V. Smakhtin, and M. Moghaddasi, 2006: Comparison of seven meteorological indices for drought monitoring in Iran. *Int. J. Climatol.*, 26 (7), 971-985. <https://rmets.onlinelibrary.wiley.com/doi/pdf/10.1002/joc.1264>
110. Abatzoglou, J. T., McEvoy, D. J., & Redmond, K. T., 2017: The West Wide Drought Tracker: drought monitoring at fine spatial scales. *Bulletin of the American Meteorological Society*, 98(9), 1815-1820.
111. Huntington, J. L., Hegewisch, K. C., Daudert, B., Morton, C. G., Abatzoglou, J. T., McEvoy, D. J., & Erickson, T. (2017). Climate engine: Cloud computing and visualization of climate and remote sensing data for advanced natural resource monitoring and process understanding. *Bulletin of the American Meteorological Society*, 98(11), 2397-2410.
112. McEvoy, D.J., M. Hobbins, T.J. Brown, K. VanderMolen, T. Wall, J.L. Huntington, and M. Svoboda, 2019: Establishing relationships between drought indices and wildfire danger outputs: A test case for the California-Nevada drought early warning system. *Climate*, 7(4), 52.
113. Oliver, J.E., 2005: Evapotranspiration. In *Encyclopedia of World Climatology*, pp. 370-373, Editor: John E. Oliver. Springer: Dordrecht, The Netherlands.
114. Hobbins, M., and J. Huntington, 2016: Evapotranspiration and Evaporative Demand. In *Handbook of Applied Hydrology*, 2nd Edition, Chapter 42, Editors: Vijay P. Singh. McGraw-Hill Publishing: New York.
115. Akinremi, O.O., S.M. McGinn, and A.G. Barr, 1996: Evaluation of the Palmer Drought Index on the Canadian Prairies. *Journal of Climate*, 9(5), 897-905. https://journals.ametsoc.org/view/journals/clim/9/5/1520-0442_1996_009_0897_eotpdi_2_0_co_2.xml
116. Gobena, A.K. and T.Y. Gan, 2013: Assessment of Trends and Possible Climate Change Impacts on Summer Moisture Availability in Western Canada based on Metrics of the Palmer Drought Severity Index. *Journal of Climate* 26(13):4583-4595. Doi: 10.1175/JCLI-D-12-00421.1 <https://journals.ametsoc.org/view/journals/clim/26/13/jcli-d-12-00421.1.xml>
117. Wells, N., S. Goddard, & M.J. Hayes, 2004: A Self-Calibrating Palmer Drought Severity Index. *Journal of Climate*, 17(12), 2335-2351. Retrieved Feb 22, 2022, from https://journals.ametsoc.org/view/journals/clim/17/12/1520-0442_2004_017_2335_aspdsi_2.0.co_2.xml
118. Balting, D.F., A. AghaKouchak, G. Lohmann, and M. Ionita, 2021: Northern Hemisphere drought risk in a warming climate. *npj Climate and Atmospheric Science*, 4:61 ; <https://doi.org/10.1038/s41612-021-00218-2>
119. Stevenson, S., S. Coats, D. Touma, J. Cole, F. Lehner, J. Fasullo, and B. Otto-Bliesner, 2022: Twenty-first century hydroclimate: A continually changing baseline, with more frequent extremes. *Proceedings of the National Academy of Sciences* 2022-03-22 119(12): e2108124119. <https://www.pnas.org/doi/abs/10.1073/pnas.2108124119>
120. Van Loon, A.F., K. Stahl, G. Di Baldassarre, J. Clark, S. Rangescroft, N. Wanders, T. Gleeson, A.I.J.M. Van Dijk, L.M. Tallaksen, J. Hannaford, R. Uijlenhoet, A.J. Teuling, D.M. Hannah, J. Sheffield, M. Svoboda, B. Verbeiren, T. Wagener, and H.A.J. Van Lanen, 2016: Drought in a human-modified world: reframing drought definitions, understanding, and analysis approaches, *Hydrol. Earth Syst. Sci.*, 20, 3631-3650, <https://doi.org/10.5194/hess-20-3631-2016>.
121. Huning, L.S. and A. AghaKouchak, 2020: Global snow drought hot spots and characteristics. *Proceedings of the National Academy of Sciences*, Aug 2020, 117 (33), 19753-19759. DOI: 10.1073/pnas.1915921117
122. Shrestha, R., B. Bonsal, J.M. Bonnyman, A.J. Cannon, and M.R. Najafi. 2021. Heterogeneous snowpack response and snow drought occurrence over northwestern North America under 1.0 °C to 4.0 °C warmings. *Clim. Change*, 164, doi:10.1007/s10584-021-02968-7.
123. IPCC (Intergovernmental Panel on Climate Change), 2021: *Climate Change 2021: The Physical Science Basis. Contribution of Working Group I to the Sixth Assessment Report of the Intergovernmental Panel on Climate Change* [Masson-Delmotte, V., P. Zhai, A. Pirani, S.L. Connors, C. Péan, S. Berger, N. Caud, Y. Chen, L. Goldfarb, M.I. Gomis, M. Huang, K. Leitzell, E. Lonnoy, J.B.R. Matthews, T.K. Maycock, T. Waterfield, O. Yelekçi, R. Yu, and B. Zhou (eds.)]. Cambridge University Press, Cambridge, United Kingdom and New York, NY, USA, 2391 pp. doi:10.1017/9781009157896
124. Heim, R.R., Jr., 2017: A Comparison of the Early 21st Century Drought in the USA to the 1930s and 1950s Drought Episodes. *Bull. Amer. Meteor. Soc.*, 98, 2579-2592, <https://doi.org/10.1175/BAMS-D-16-0080.1>
125. Yuan, X., Y. Wang, P. Ji, P. Wu, J. Sheffield, and J.A. Otkin, 2023: A global transition to flash droughts under climate change. *Science*. 380, 187-191. <https://www.science.org/doi/10.1126/science.abn6301>

126. Hasegawa, A., M. Guryev, and Y. Iwami, 2016: Meteorological Drought and Flood Assessment Using the Comparative SPI Approach in Asia Under Climate Change. *J. Disaster Res.*, 11(6): 1082-1090. <https://doi.org/10.20965/jdr.2016.p1082>
127. Guryev, M., A. Hasegawa, J. Magome, P. Sanchez, A. Sugiura, H. Umino, H. Sawano, and Y. Tokunaga, 2016: Evaluation of Water Cycle Components with Standardized Indices Under Climate Change in the Pampanga, Solo and Chao Phraya Basins. *J. Disaster Res.*, 11(6): 1091-1102. <https://doi.org/10.20965/jdr.2016.p1091>
128. Hoylman, Z.H., R.K. Bocinsky, and K.G. Jencso, 2022: Drought assessment has been outpaced by climate change: empirical arguments for a paradigm shift. *Nature Communications*, 13:2715. <https://doi.org/10.1038/s41467-022-30316-5>

Disclaimer/Publisher's Note: The statements, opinions and data contained in all publications are solely those of the individual author(s) and contributor(s) and not of MDPI and/or the editor(s). MDPI and/or the editor(s) disclaim responsibility for any injury to people or property resulting from any ideas, methods, instructions or products referred to in the content.