

Article

Satellite Retrieval of Downwelling Shortwave Surface Flux and Diffuse Fraction under All Sky Conditions in the Framework of the LSA SAF Program (Part 2: Evaluation)

Dominique Carrer¹, Suman Moparthy¹, Chloé Vincent¹, Xavier Ceamanos¹, Sandra C. Freitas² and Isabel F. Trigo²

¹ Météo-France/CNRM, CNRS/GAME, 42 avenue Gaspard Coriolis, 31057 Toulouse Cedex, France

² Instituto Português do Mar e da Atmosfera (IPMA), Rua C-Aeroporto, 1749-077 Lisboa, Portugal

Abstract: High frequency knowledge of the spatio-temporal distribution of the Downwelling Surface Shortwave Flux (DSSF) and its diffuse fraction (fd) at the surface is nowadays essential for understanding climate processes at the surface-atmosphere interface, plant photosynthesis and carbon cycle, and for the solar energy sector. The EUMETSAT Satellite Application Facility for Land Surface Analysis operationally delivers estimation of the MDSSFTD (Downwelling Surface Short-wave radiation Fluxes – Total and Diffuse fraction) product with an operational status since the year 2019. The method for the retrieval was presented in the companion paper [40]. The part 2 now focuses on the evaluation of the MDSSFTD algorithm and presents the comparison of the corresponding outputs, i.e. total DSSF and diffuse fraction (fd) components, against *in-situ* measurements acquired at four BSRN stations over a seven-month period. The validation is performed on an instantaneous basis. We show that the satellite estimates of DSSF and fd meet the target requirements defined by the user community for all-sky (clear and cloudy) conditions. For DSSF, the requirements are 20Wm^{-2} for $\text{DSSF} < 200\text{Wm}^{-2}$, and 10% for $\text{DSSF} \geq 200\text{Wm}^{-2}$. The MBE and rMBE compared to the ground measurements are 3.618Wm^{-2} and 0.252%, respectively. For fd, the requirements are 0.1 for $\text{fd} < 0.5$, and 20% for $\text{fd} \geq 0.5$. The MBE and rMBE compared to the ground measurements are -0.044 and -17.699%, respectively. The study also provides a separate analysis of the product performances for clear sky and cloudy sky conditions. The importance of representing the cloud-aerosol radiative coupling in the MDSSFTD method is discussed. Finally, it is concluded that the quality of the Aerosol Optical Depth (AOD) forecasts currently available is enough accurate to obtain reliable diffuse solar flux estimates. This quality of AOD forecasts was still a limitation a few years ago.

Keywords: Solar Radiation; Meteosat Second Generation; Validation; Land Surface Modelling

1. Introduction

The downwelling surface short-wave radiation flux (DSSF) refers to the radiative energy in the wavelength interval $[0.3\text{ }\mu\text{m}, 4.0\text{ }\mu\text{m}]$ reaching the Earth's surface per time and area unit. An accurate knowledge of the spatio-temporal distribution of the downwelling solar radiation at the surface is essential not only for understanding climate processes at the surface-atmosphere interface [1, 2], but also for plant photosynthesis and carbon cycle, e.g., [3-5] and for the solar energy sector [6]. Concerning the current status of DSSF modelling in atmospheric models, [7] and [8] found that the National Centers for Atmospheric Prediction (NCEP) and the National Center for Atmospheric Research (NCAR) data consistently overestimated DSSF by 17%–27%. Comparisons with satellite data have also revealed large positive biases in NCEP–NCAR DSSF ranging from 25 to 50Wm^{-2} over the United States [9-10] and from 40 to 80Wm^{-2} over Europe [11]. However, in a recent study, [12] examined the progress made by two new reanalyses in the estimation of surface irradiance (ERA5

and COSMO-REA6) and negative biases of around -5 Wm^{-2} . They showed the largest deviations under clear-sky conditions, which is most likely caused by the aerosol data used.

DSSF essentially depends on the solar zenith angle, cloud coverage, aerosols, and to a lesser extent on atmospheric absorption and surface albedo. Over the past few decades the scientific community has developed computation methods to estimate both downward and net surface solar irradiance from satellite observations [13–29]. In addition to those estimates, two incoming solar radiation products derived from MSG/SEVIRI were also developed, being operated since 2005 by EUMETSAT Satellite Application Facility (SAF) on Land Surface Analysis (LSA; [30]): the MDSSF product (referenced LSA-201) corresponding to instantaneous values, and the DIDSSF product (LSA-203) corresponding to daily accumulated values. Both products consider clear and cloudy skies to provide total shortwave fluxes at the surface. However, all these estimates lack of the repartition of the total flux into its direct and diffuse components (through the diffuse fraction, for example). Moreover, even though these products have proven to be of high quality, [31–33] showed that they still have some limitations under clear sky conditions, especially as they are determined taking as hypothesis a temporally and spatially constant load and type of continental aerosols [34]. The importance of aerosols on the DSSF has been established in numerous studies on some highly polluted regions [35–39]. Thus, an initiative has been conducted by EUMETSAT to upgrade the physics in the scientific algorithms used for the satellite-derived DSSF retrievals and to provide first estimations of the diffuse fraction of the radiation.

The physics of this upgraded algorithm is described in the companion paper [40]. The new product version has been referenced as LSA-207 by EUMETSAT, corresponding to the MSG Downwelling Surface Short-wave radiation Fluxes – Total and Diffuse fraction (MDSSFTD). Two different modules are used to calculate the set of MDSSFTD outputs, one for clear conditions, and the other for cloudy conditions. The two methods are designed to ensure the spatial and temporal continuity of DSSF and diffuse fraction in the LSA-207 product. Details on the methodology as well as the major limitations are given in [40]. The input cloud mask is used to distinguish between the two methods. The summary of the two methods is as follows and described in detailed in [40].

In clear sky conditions, the formulation based on the algorithm SIRAMix [41, 42] is used to estimate the total flux and the diffuse fraction. The atmospheric pressure, water content, ozone content, aerosols vary in time and space and are provided by atmospheric model forecasts. Both direct and diffuse flux terms are estimated by combining pre-computed aerosol transmittances and albedo (computed using radiative transfer models for varying aerosol load, solar zenith angles and water vapour content) from look-up tables and semi-empirical radiative transfer equations [42]. The total flux is the sum of both direct and diffuse flux estimates. The diffuse fraction is obtained as the ratio of the diffuse flux to the total flux estimate. In cloudy sky conditions, the total flux is estimated using simplified radiative transfer equations as described in [40]. The cloud transmittances are estimated from SEVIRI radiances at the top-of-atmosphere (TOA) level. The atmospheric transmittance term used for the estimation of the effective transmittance remains the same as in clear sky conditions. However, an extra cloud transmittance term is added as well as two multiple scattering terms. In the case of cloudy-sky conditions, the diffuse fraction is estimated using an empirical formulation. The clear-sky and cloudy sky methods are designed to provide smooth transitions in the frontiers between clear and cloudy pixels (see [40]). Finally, LSA-207 then includes an estimation of the total incoming solar radiation with an improved modelling of the aerosol impact on the atmospheric transmittance compared to the previous MDSSF product (LSA-201). The diffuse fraction of the radiation for all sky conditions is now also available. Moreover, estimations of auxiliary quantities are also provided: the equivalent Aerosol Optical Depth (AOD) at 550nm, the Opacity Index (OI) characterizing the opacity of the atmosphere (defined in [40]), and a quality flag information (QF).

This study makes an evaluation of the satellite-derived MDSSFTD product. The article is organized as follows. Section 0 presents the data and the metrics used for the validation. Section 0 presents validation results and Section 0 concludes about the performance of the product regarding the users requirements.

2. Data and Metrics

2.1. Requirements

Over India, a satellite-based surface solar radiation dataset called Surface Solar Radiation Data Set-Heliosat (SARAH-E) was developed and evaluated against *in-situ* measurements over a variety of sites. The results indicate an overestimation of the satellite DSSF, with a mean bias of 21.9Wm^{-2} [43]. Study Over Finland and Sweden [44] also discussed the retrieval accuracies of two different satellite-derived DSSF dataset (the polar-orbiting satellite-based dataset – CLARA-A1 – and the geostationary satellite-based dataset – SARAH). They showed comparable accuracies in comparison with ground measurements, in particular 10Wm^{-2} for the monthly means metrics and 15Wm^{-2} for daily means metrics. Over Europe, [31] completed the inter-comparison of satellite-derived incoming solar products from the different Satellite Application Facilities of EUMETSAT and concluded that the products have comparable mean biases ($+4\text{Wm}^{-2}$) and root mean square differences ($80\text{--}100\text{Wm}^{-2}$) for instantaneous metrics. Performances of the historical LSA-SAF DSSF (LSA-201) satellite-derived incoming solar radiation product were also discussed more in detail in [45,46]. On an instantaneous basis, the bias between the satellite product and the ground data was shown to be small with absolute values of less than 10Wm^{-2} over Europe [45], and even lower over France (3.75Wm^{-2} representing 2.5%). The standard deviation of the difference between instantaneous satellite estimates and ground measurements were of the order of 40Wm^{-2} for clear sky data and 110Wm^{-2} for cloudy sky data. Finally, the satellite estimates of DSSF are today ranging in average from 10 to 30Wm^{-2} in absolute bias scale. However, these past works also pointed out that the absolute metrics usually used to evaluate the product performances are not equivalent if the domain (or period) of interest is located in high latitude or in low latitudes (or winter and summer periods).

The characteristics of the LSA-207 MDSSFTD product and the targeted accuracies agreed with EUMETSAT are described in Table 1. These are a compromise between what is currently achievable, given existing observations and algorithm input data, and what would suit most users and applications. In this respect, the ‘threshold’ requirement is then defined as the minimum accuracy which is acceptable for DSSF user needs. This paper assesses the performance of the product by referring to the ‘target’ requirement. However, it may be relevant to note that because the topic of retrieving diffuse fraction from satellite is very recent, there is today no performance requirements defined by the scientific community for this parameter. We therefore have fixed the ‘target’ accuracy to 20%, although a larger uncertainty (e.g. $>30\%$) could have been also considered. The target accuracy metrics used are the Mean Bias Error (MBE) for low values, and the relative MBE (rMBE) for high values of DSSF or fd.

Table 1 Product Requirements for MDSSFTD, in terms of area coverage, resolution and accuracy.
The targeted requirements are indicated in bold.

Product	Coverage	Resolution		Threshold	Accuracy	
		Temporal	Spatial		Target	Optimal
MDSSFTD (LSA-207) DSSF_TOT	MSG disk	15 min	MSG pixel resolution	20%	DSSF<200 W/m²: 20W/m² (MBE) DSSF>=200 W/m²: 10% (rMBE)	5%
MDSSFTD (LSA-207) Diffuse Fraction FRACTION_DIFFUSE (fd)	MSG disk	15 min	MSG pixel resolution	30%	fd <0.5: 0.1 (MBE) fd >=0.5: 20% (rMBE)	10%

2.2. Ground measurements and preprocessing

Four ground stations are used for the validation analyses presented in this document. The stations considered are Carpentras, De Aar, Tamanrasset and Toravere from the BSRN (Baseline Surface Radiation Network, <http://www.bsrn.awi.de>) of the World Climate Research Programme. Their location is presented in Figure 1.

The *in-situ* measurements of instantaneous total and diffuse DSSF as observed at these stations are used as reference. The stations are located in climatically different regions of Europe and Africa. For example, the station Tamanrasset in North Africa is influenced by coarser dust particles and more clear conditions than the other Europe-based stations. Toravere is located at the highest latitude and therefore is related to frequent periods of overcast in the winter and fall. This will help evaluation the MDSSFTD method under cloudy situation and high solar zenith angles.



Figure 1 Location of the ground stations providing *in-situ* measurements.

BSRN derived surface flux values are retrieved at a high temporal frequency going from 1 minute to a few minutes. BSRN data already account for missing or bad measurements for which missing flag values are assigned. The missing values are discarded in the comparison with satellite flux measurements. The direct flux measurements account for the varying solar zenith angle dependence. All measurements (total, diffuse and direct surface solar flux measurements) are discarded when the solar zenith angle is greater than 80 degrees. MSG/SEVIRI satellite-derived products differ on a pixel basis from 0 to 12 minutes with the product time. SEVIRI scan takes 12 minutes for the data acquisition over the MSG disk starting from the South pole and finishing its acquisition in the North pole. For the sake of a fair comparison, the ground measurements are averaged over 15 minutes, centered around the exact acquisition time of the satellite for every SEVIRI pixel. The relationship between the estimated time difference as a function of row number (or latitude) is detailed in other LSA SAF reports (product user manual MDSSF; 2.6V2 at <https://landsaf.ipma.pt/en/products/longwave-shortwave-radiation/>; last time consulted on 2nd of September, 2019). Note that the diffuse fraction is not a direct measurement. This variable is obtained by dividing the diffuse component over the total component that are measured by the ground instruments. The BSRN total fluxes and diffuse fluxes are measured respectively with pyranometers and shaded pyranometers (Kipp & Zonen/CM21 for stations Carpentras and Toravere, Kipp & Zonen/CMP21 for station De Aar, and Eppley/PSP for station Tamanrasset). The accuracy of the BSRN total fluxes measurements, provided the measure is made according to the BSRN protocol, is estimated to 0.5% or 1.5 W m⁻², while the accuracy of the diffuse measurements is estimated to 2% or 3 W m⁻² [47]. In practice, specific analyses on BSRN sites accuracy [48-49] confirmed that the uncertainty of the measures are in agreement with such levels of uncertainties, with some limitations for low sun elevation angles and low radiation fluxes. As the validity of the MDSSFTD products are limited to sun zenith angles below 80°, we consider that the *in situ* measurement of total fluxes and the BSRN-derived diffuse fractions can reasonably be taken as references for our validation analysis.

2.3 CAMS all-sky radiation data

The Copernicus Atmosphere Monitoring Service (CAMS) all-sky radiation data was also used to compare with the MDSSFTD product. CAMS all-sky radiation service distributes global, direct and diffuse irradiances as well as a direct at normal incidence irradiance, for all-sky (clear+cloud), clear-sky only, and cloudy-sky only. These data are provided as timeseries with a temporal resolution of 1-minute, 15-minutes, 1h, 1 day or 1 month, and are made available since February 2004 with a 2-days delay. The spatial coverage corresponds to the Meteosat Second Generation (MSG) disk. To produce the timeseries, radiation data is spatially interpolated to the point of interest from a product available at the native spatial resolution of MSG/SEVIRI. These data are publicly available from the CAMS portal (<https://atmosphere.copernicus.eu/solar-radiation>, last consulted on 18/10/2019).

The CAMS Radiation service relies on the Heliosat-4 method [50], which is composed of two modules: McClear v3 for clear-skies [51, 52] and McCloud for cloudy-skies [50]. The McClear approach, version 3, used by CAMS radiation service, has been upgraded from McClear v2 and now consist in a physical modelling using the radiative transfer model libRadtran [53]. As for the MDSSFTD method, the McClear v3 now also relies on aerosols content and load, as well as gases contents forecasted by ECMWF and distributed by CAMS. Compared to MDSSFTD method, McClear v3 uses a monthly-climatology [54] of the MODIS surface albedos [55] and a similar approach than MDSSFTD for the aggregation of the optical depths of each aerosol species to derive the properties of the aerosol mixture [52]. The McCloud method estimates the cloud properties from MSG measurements using a model adapted from APOLLO (AVHRR Processing scheme Over cLOUDs, Land and Ocean, [56, 57].

For our validation analysis, we used the CAMS global and diffuse radiation data, at a temporal resolution of 15-minutes. For sake of consistency with the evaluation against ground measurements, the product was extracted for the whole validation period (February to October 2017) at the location of the four BSRN stations already considered for the ground measurements analysis (Carpentras, De Aar, Tamanrasset and Toravere).

2.4. Metrics

The target accuracy metrics used are the Mean Bias Error (MBE) for low values of DSSF (<200Wm⁻²) or of fd (<0.5). The target accuracy metrics used are the relative MBE (rMBE) for high values of DSSF (>=200Wm⁻²) or of fd (>=0.5). MBE is computed as

$$MBE = \frac{1}{N} \sum_{i=1}^N (satelliteproduct_i - reference_i)$$

and the relative MBE, noted “rMBE” is a dimensionless metric, expressed in percent units, and defined as:

$$rMBE = \frac{1}{N} \sum_{i=1}^N \frac{satelliteproduct_i - reference_i}{reference_i}$$

where N is the number of points and ‘reference’ corresponds to the ground measurements in our study.

The choice of the metrics was made to ensure consistency with the two other existing LSA-SAF products MDSSF and DIDSSF (LSA-201 and 203; see Section 0) for which the same evaluation strategy was used. The ground measurements are separated into clear and cloudy samples based on the information contained in the cloud mask. For example, if the SEVIRI pixel is defined as cloudy (respectively clear) according to the information contained in the quality flag, the corresponding time slot is then defined as cloudy (respectively clear) for the ground measurements. In the case of clear sky retrievals, the clear sky pixel is excluded when the adjacent time slots (up to 1h, that is, 30 minutes before and 30 minutes after) are defined as cloudy. This is deemed to suppress any residual cloud contamination (or cloud shadow effects) in the clear sky retrievals. Same strategy is applied reversely to identify cloudy-sky pixels with adjacent time slots which are clear-sky.

The aimed requirements are the target accuracies (values in bold Table 1). These metrics will be used in the following to evaluate the performances of the MDSSFTD product for clear-sky, cloudy-sky, and all-sky conditions. However, the user needs expressed to EUMETSAT is to have a MDSSFTD product which meets the target requirements for the all-sky conditions (without distinction according to cloudiness). The performances are evaluated based on the metrics that are obtained from all the available ground measurements (i.e., for all the stations and over the entire period of evaluation). The MDSSFTD product has the spatial resolution of the native SEVIRI grid (3km at the sub-satellite point over Africa and around 5km over Europe). [58] show that there is no major representativeness issue between the local ground-based solar radiation measurements and the satellite estimates (which have kilometer scales).

The evaluation is performed over a period of seven months: from February to October 2017. The stability of the metrics is also examined by splitting on a daily basis the metrics and analysing stability of the metrics from day to day.

3. Results

3.1. Sensitivity study: inter-comparison of models to estimate diffuse flux in cloudy-sky conditions

A critical module in the method used for MDSSFTD is the choice of the empirical formulation used to estimate the diffuse fraction in the cloudy-sky case. In this context, a specific sensitivity study is first made to compare a set of existing solutions. One shall note that this sensitivity study only reflects the impact of the parametrisation choice on the diffuse fraction retrieval under cloudy conditions. Retrieval of the diffuse fraction under clear sky remains unchanged (see [40]).

We detail here this sensitivity study that compared several empirical formulations from the literature. The different formulations are based on [59-61]. All three methods estimate the diffuse part of the total solar irradiance from the clearness index ('Kt'). This allows the calculation of the diffuse fraction by simply dividing the diffuse flux by the total counterpart. Another formulation based on [62] was also considered in this sensitivity study (Section 0). This fourth method, however, estimates the direct component of the solar irradiance based on 'Kt', which is used to infer the direct fraction of the solar irradiance to finally retrieve the diffuse fraction.

Figure 2 displays the density scatter plots for the diffuse fraction retrieved following each of the four formulations mentioned above, all compared to the *in-situ* diffuse fraction in all-sky conditions (clear and cloudy). The statistics shown in Figure 2 are obtained considering the four stations over the entire period of the study. As it can be seen, the statistics from the four formulations are highly similar, with slightly lower performances for the method from [62]. Because formulation based on [60] was validated against several stations over Europe and USA, we decided to use this formulation for our application. Indeed, the other models were derived from flux measurements over more limited areas, which make them less representative at the continental scale made possible by MSG.

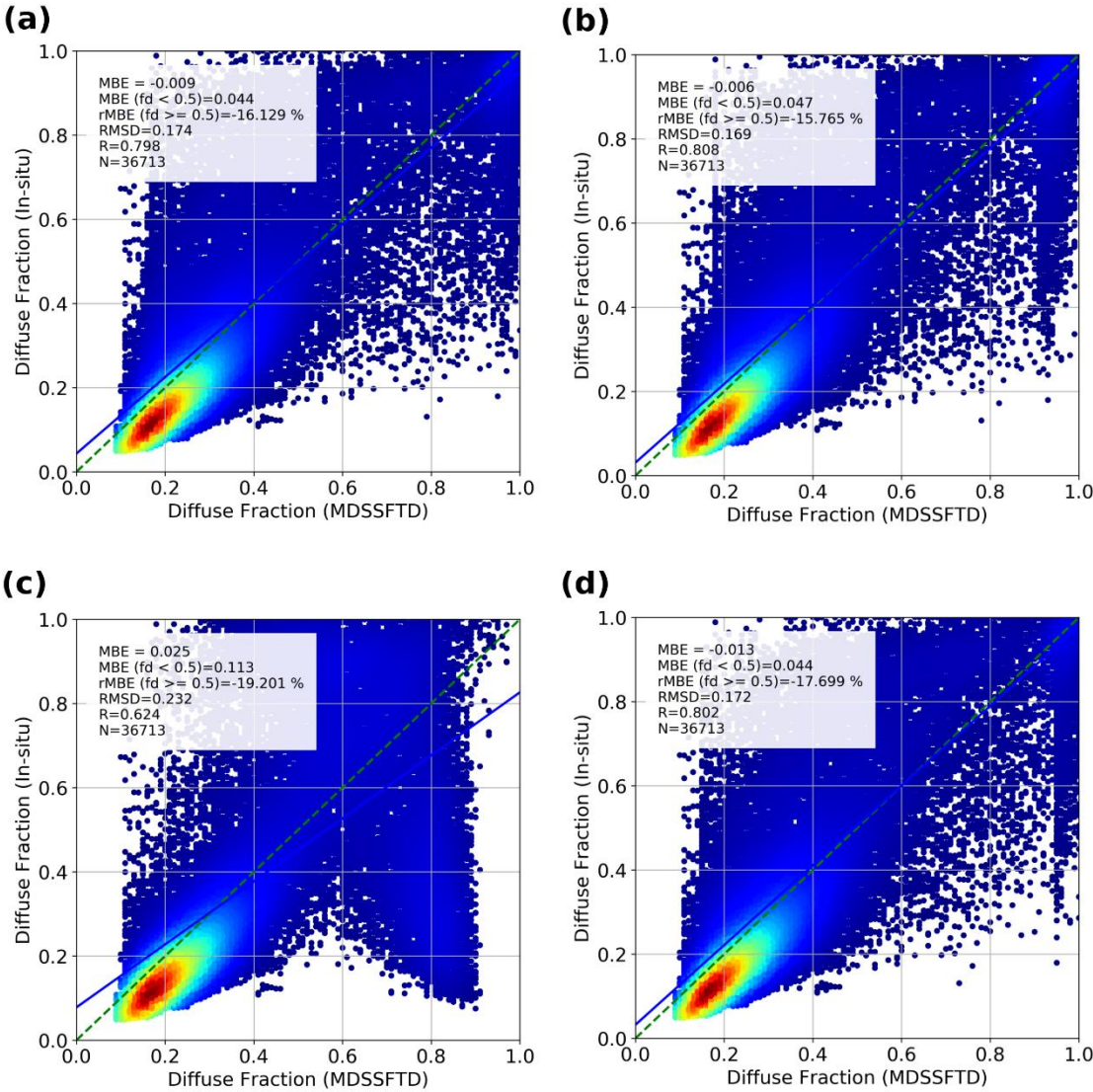


Figure 2 Diffuse fraction components retrieved following four empirical formulations, a) Erbs et al., b) Orgill and Hollands c) Louche et al., d) Reindl et al., and compared to the *in-situ* diffuse fraction component. Blue color corresponds to low density of points and red color corresponds to high density of points. Blue line represents the mean fit across the whole evaluation data.

3.2. Diurnal comparisons for clear-sky and all-sky days

The diurnal total and diffuse down-welling surface flux components from the MDSSFTD product are compared against the same flux components derived from the ground BSRN measurements. As already mentioned, the BSRN data (available at a high temporal frequency) are averaged over 15 minutes and centred around the correct MSG acquisition time slot (see Section 0).

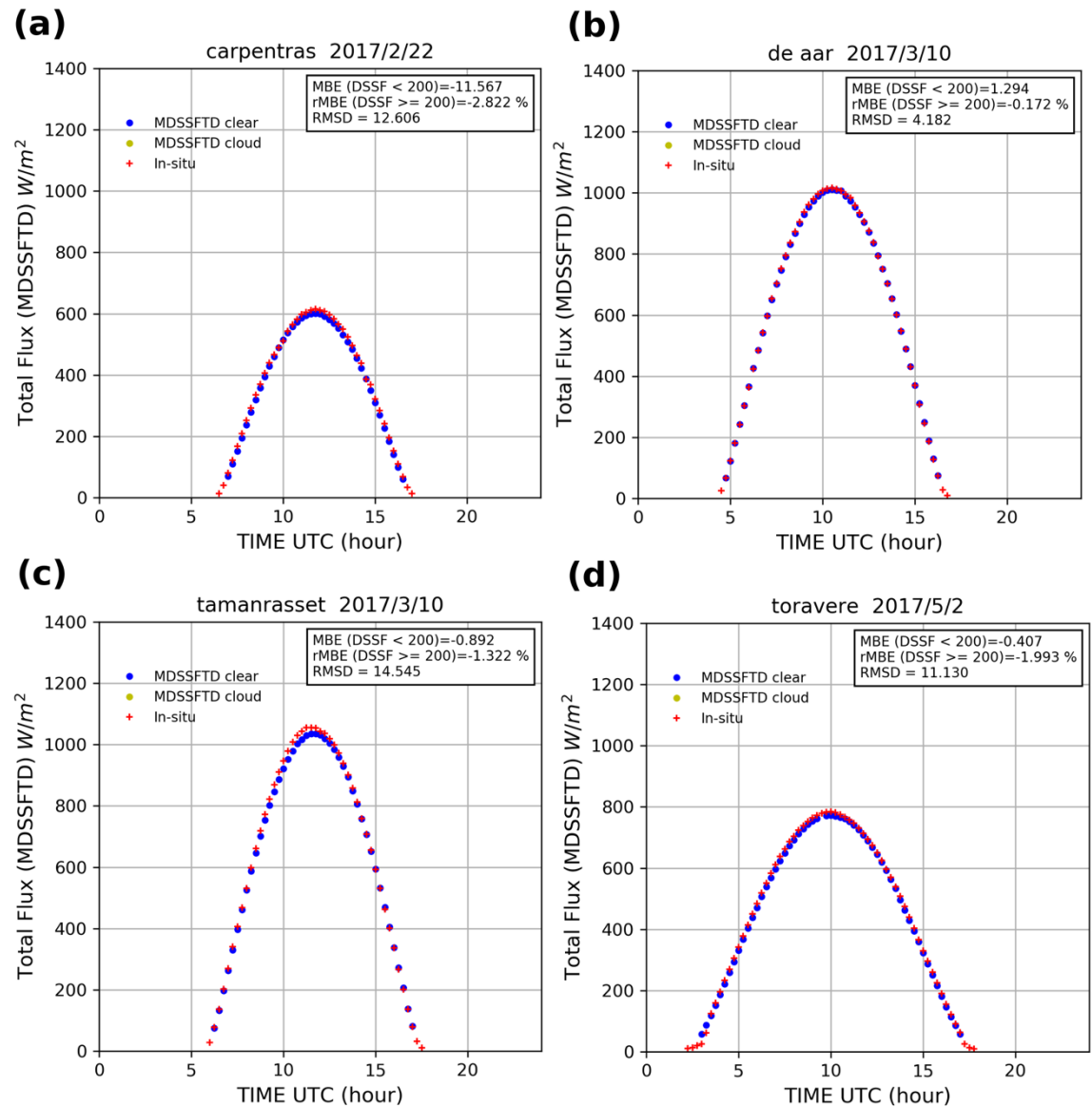


Figure 3 Diurnal variation of the total MDSSFTD component in clear sky conditions compared against *in-situ* measurements for a) Carpentras, b) De Aar, c) Tamanrasset, and d) Toravere for a selected day. Yellow cloudy samples do not appear in this figure as the chosen dates were fully clear.

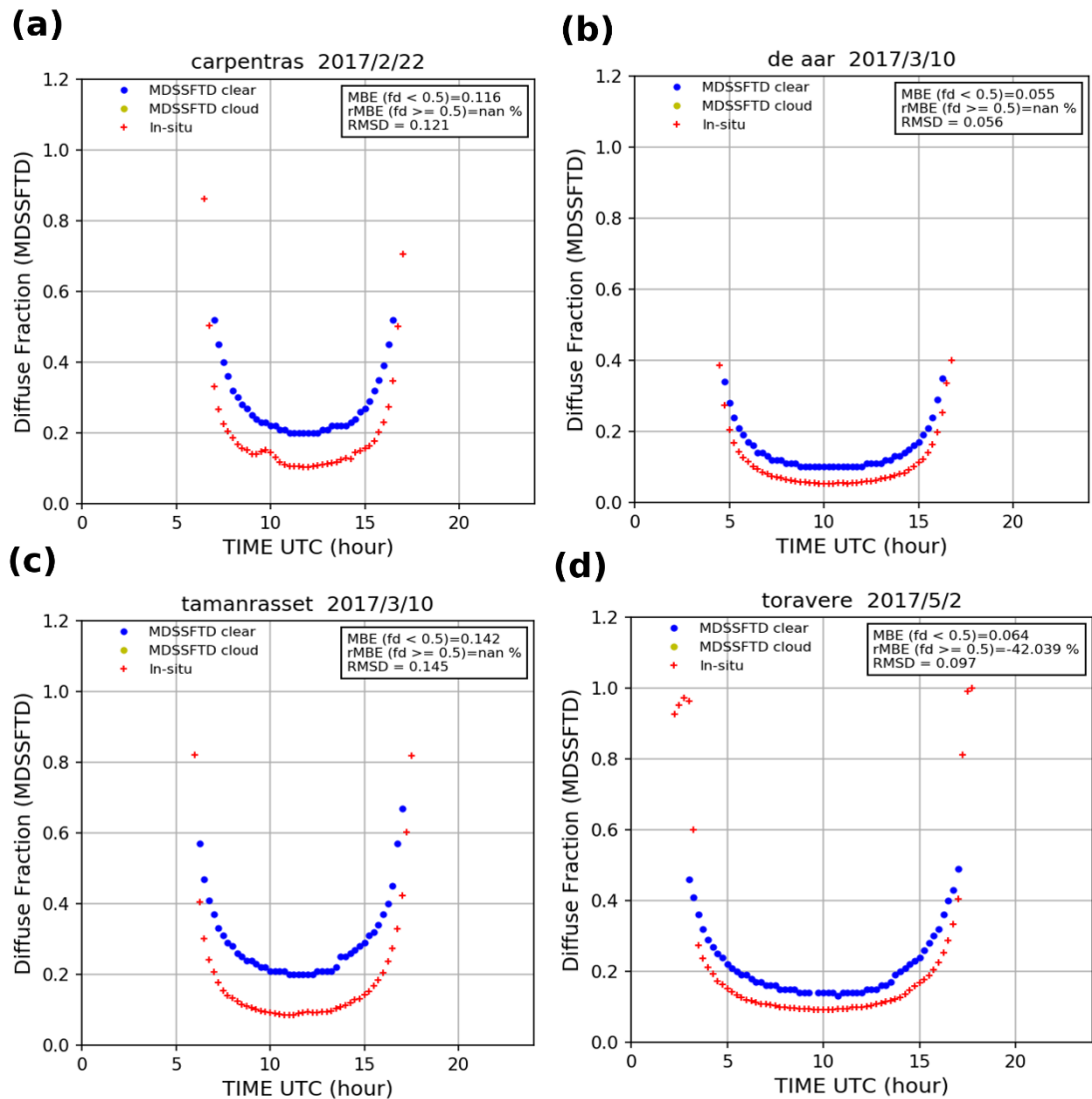


Figure 4 Diurnal variation of the diffuse fraction MDSSFTD component in clear sky conditions compared against *in-situ* measurements for a) Carpentras, b) De Aar, c) Tamanrasset, and d) Toravere for a selected day. Yellow cloudy samples do not appear in this figure as the chosen dates were fully clear.

Figure 3 shows a comparison between satellite-derived estimates and ground measurements of the diurnal cycle of the total flux for clear sky conditions all along the day. It can be observed how the satellite-derived estimates capture well the diurnal variations compared to ground measurements.

Figure 4 shows a comparison between satellite-derived estimates and ground measurements of the diurnal cycle of the diffuse fraction for clear sky conditions. Again, the satellite-derived estimates capture well the diurnal variations compared to ground measurements. In particular, the increase of the diffuse fraction with extreme geometries is well reproduced. For these four days, a slight overestimation between 0.055 and 0.142 is found for $fd < 0.5$ in clear-sky conditions. This overestimation comes from the slight overestimation of the diffuse DSSF by MDSSFTD, which was also found for clear sky situations by [41] when using SIRAMix and the McClear method [51]. These two methods used CAMS aerosol data and GADS aerosol properties, which may point to an overestimation of the highly scattering aerosol components by CAMS or a limited transformation from CAMS to GADS components.

Figures 5 and 6 show the results of similar comparisons that were conducted for dates showing all sky conditions (partially clear and partially cloudy). The diurnal variations of the MDSSFTD

product, including the total DSSF and the diffuse fraction, are compared against ground measurements for selected days in Figures 5 and 6, respectively. A rather satisfactory agreement exists between the satellite derived estimates and the ground measurements. The increase of the diffuse fraction with the cloudiness is generally well represented (e.g., see Figure 6 d).

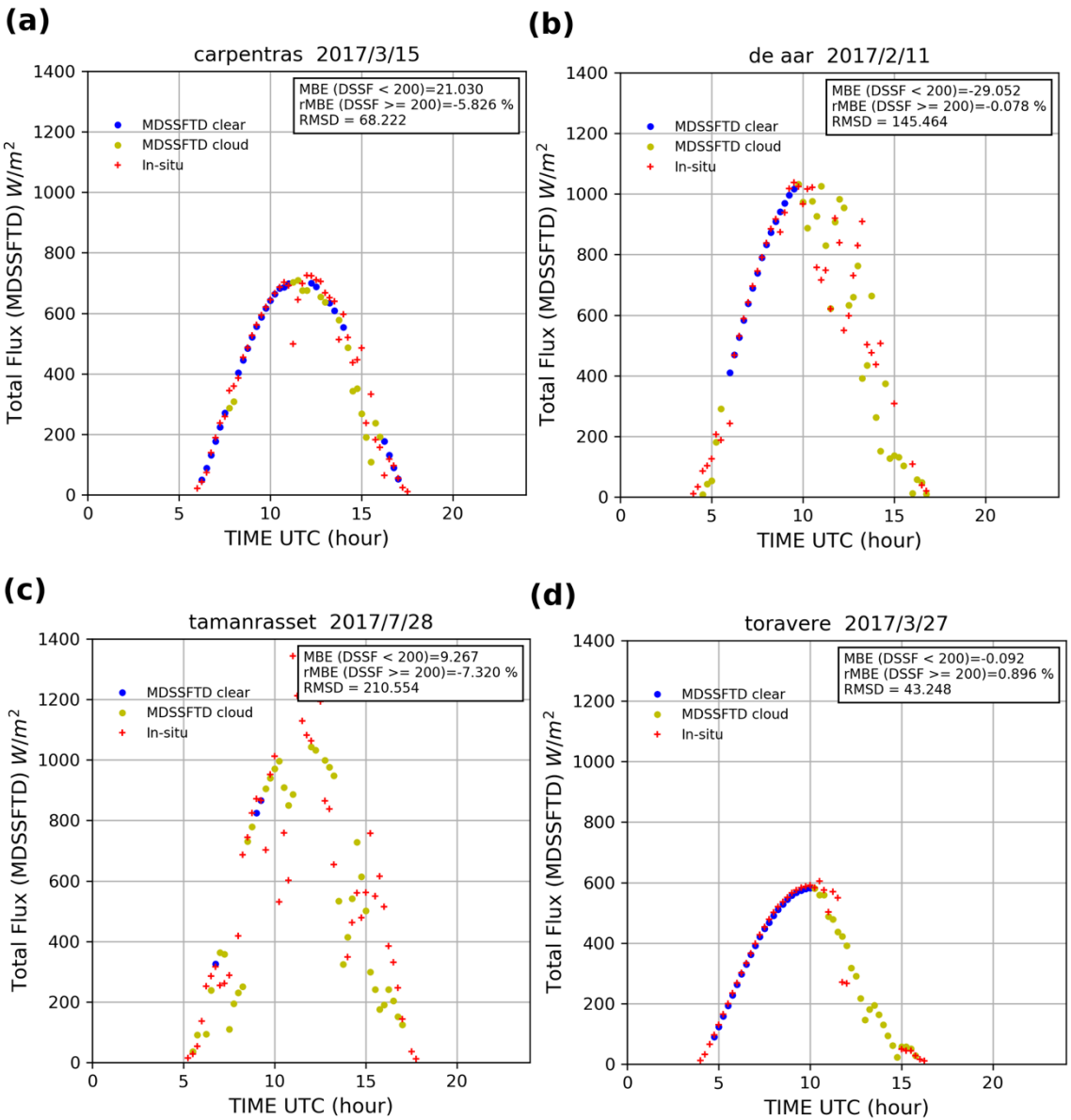


Figure 5 Diurnal variation of the total MDSSFTD component in all sky conditions compared against *in-situ* measurements for a) Carpentras, b) De Aar, c) Tamanrasset, and d) Toravere for a selected day (partially clear and cloudy). The yellow dots represent cloudy retrievals and the blue dots represent clear retrievals.

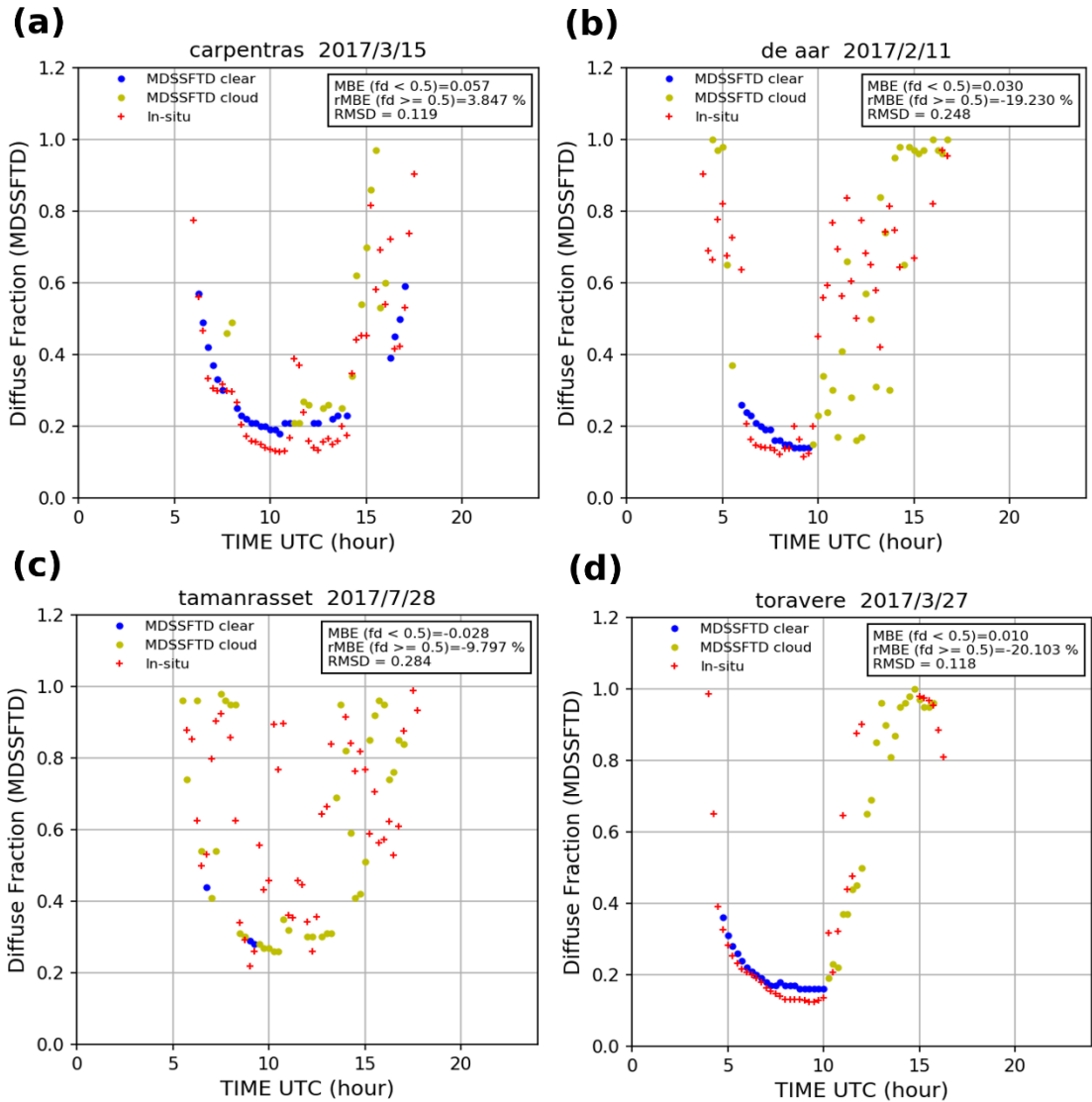


Figure 6 Diurnal variation of the diffuse fraction of MDSSFTD in all sky conditions compared against *in-situ* measurements for a) Carpentras, b) De Aar, c) Tamanrasset, and d) Toravere for a selected days (partially clear and cloudy). The yellow dots represent cloudy retrievals and the blue dots represent clear sky retrievals.

3.3. Global performances

This section details the overall statistics of the MDSSFTD product that are obtained by considering the evaluation over the four ground stations for the entire period of interest, with temporal frequency of every 15min. Statistics are hence discussed successively for clear-sky, cloudy-sky, and all-sky conditions.

3.3.1. Clear-sky conditions

Figure 7 displays the density scatter plot between instantaneous measurements of MSG-derived surface down-welling solar flux measurements for total and diffuse fraction components with their *in-situ* counterparts. Only clear sky retrievals are considered here thanks to the use of the cloud mask used as input in the MDSSFTD algorithm. Figure 7 shows that the satellite estimates of DSSF and fd meet the requirements for total DSSF, which are 20Wm^{-2} for $\text{DSSF} < 200\text{Wm}^{-2}$ and 10% for $\text{DSSF} \geq 200\text{Wm}^{-2}$ (as described in Table 1). The MBE and rMBE compared to the ground measurements are 8.637Wm^{-2} and 0.776%, respectively. On the other hand, the requirements for fd

are 0.1 for $fd < 0.5$ and 20% for $fd \geq 0.5$. The MBE and rMBE compared to the ground measurements in this case are 0.062 and -22.197%, respectively. The statistical scores in terms of MBE and RMSD (root mean square deviation) for the comparison between MDSSFTD total and diffuse fraction components with their *in-situ* counterparts for all four stations are given in Tables 2 and 3. The scores for all stations are in agreement with the DSSF product requirements. The diffuse fraction compares well for all stations except for high values of diffuse fraction ($fd \geq 0.5$). However, only 12 days over the 7-month period of the study have $fd \geq 0.5$ in clear-sky conditions (see Figure 13). Indeed, these values of diffuse fraction correspond to intense aerosol loading, which is relatively infrequent. Therefore, statistics in that case ($fd \geq 0.5$ and clear-sky conditions) cannot be considered as significant from a statistical point of view.

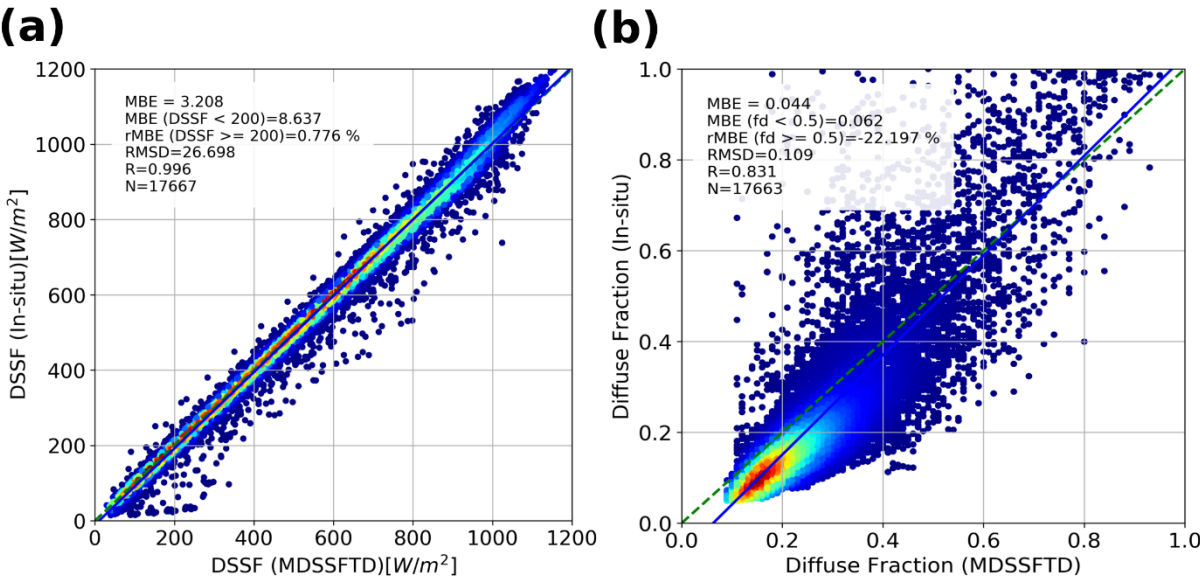


Figure 7 Comparison of instantaneous MSG-derived MDSSFTD measurements for (a) total DSSF, (b) diffuse fraction components with their *in-situ* counterparts for clear-sky retrievals. The retrievals are collected every 15 min. Blue line represents the mean fit across the whole evaluation data. Blue circles corresponds to low density of points and red circles corresponds to high density of points.

Table 2 Statistical scores obtained from the comparison between MDSSFTD derived total flux estimates and ground *in-situ* measurements over the selected BSRN sites for clear sky retrievals. If the value is in bold, the metric does not meet the “target” requirements. If no value appears in bold, all the metrics meet the requirements. R_VAL correspond to the Pearson correlation coefficient. RMSD is the root mean square deviation.

	Lat °N	Lon °E	R_VAL [-]	RMSD [Wm ⁻²]	MBE [Wm ⁻²]	MBE (DSSF<200) [Wm ⁻²]	rMBE (DSSF>=200) [%]
Carpentras	44.08	5.06	0.998	22.809	14.552	6.759	2.623
De Aar	-30.67	23.99	0.996	23.042	-2.946	8.927	-0.797
Tamanrasset	22.79	5.53	0.995	34.096	-0.441	12.336	0.798
Toravere	58.25	26.46	0.996	19.277	1.488	4.297	0.320

Table 3 Statistical scores obtained from the comparison between MDSSFTD diffuse fraction estimates and ground measurements over the selected BSRN sites for clear sky retrievals. If the value is in bold, the metric does not meet the “target” requirements. If no value appears in bold, all the metrics meet

the requirements. R_VAL correspond to the Pearson correlation coefficient. RMSD is the root mean square deviation.

	Lat °N	Lon °E	R_VAL [-]	RMSD [-]	MBE [-]	MBE (fd<0.5) [-]	rMBE (fd>=0.5) [%]
Carpentras	44.08	5.06	0.890	0.069	0.042	0.045	-10.173
De Aar	-30.67	23.99	0.624	0.115	0.065	0.073	-47.698
Tamanrasset	22.79	5.53	0.831	0.134	0.028	0.072	-21.216
Toravere	58.25	26.46	0.768	0.091	0.027	0.039	-31.354

3.3.2. Cloudy-sky conditions

Figure 8 displays the density scatter plot between instantaneous measurements of MSG-derived surface down-welling solar flux measurements for total and diffuse fraction components with their *in-situ* counterparts. Only cloudy sky retrievals are considered for this experiment. Figure 8 shows that the satellite estimates of DSSF and fd meet the target requirements. For DSSF, the requirements are 20Wm⁻² for DSSF<200Wm⁻² and 10% for DSSF>=200Wm⁻². The MBE and rMBE compared to the ground measurements are -6.618Wm⁻² and -2.782%, respectively. For fd, the requirements are 0.1 for fd<0.5 and 20% for fd>=0.5. The MBE and rMBE compared to the ground measurements are 0.027 and -15.796%, respectively. Tables 4 and 5 give the statistical scores in terms of MBE and RMSD for the comparison between MDSSFTD total and diffuse fraction components with their *in-situ* counterparts for all four stations. The vertical patterns that are observed in Figure 8(b) comes from the method that is used for cloudy skies. Indeed, the estimation of the diffuse fraction is estimated using three equations that are selected according to the value of clearness index. More details are given in the companion paper [40].

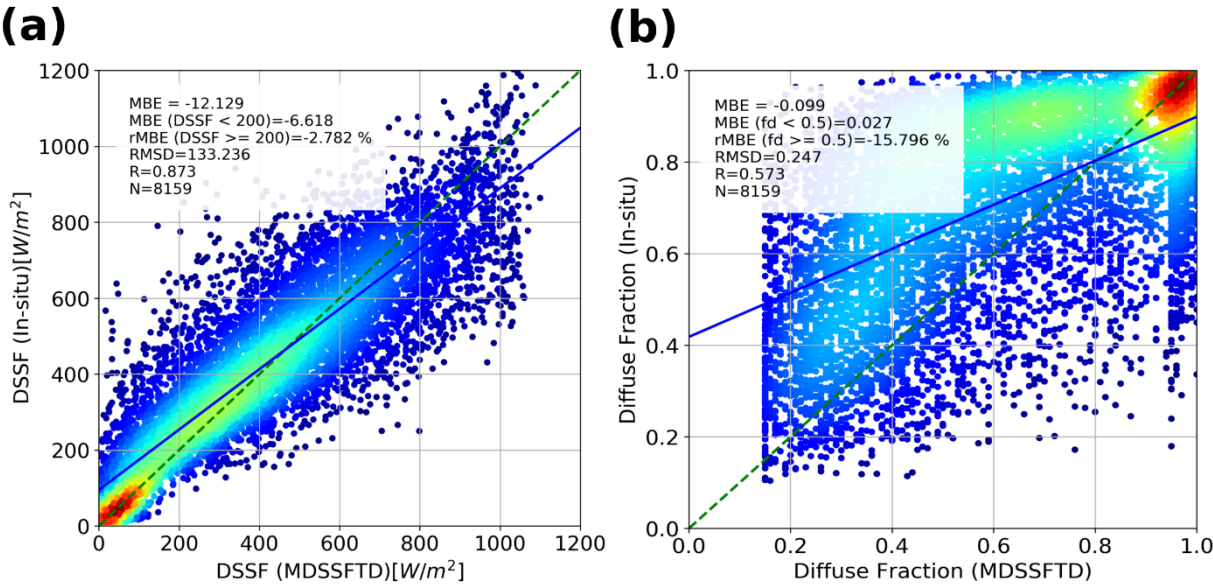


Figure 8 Comparison of instantaneous MSG-derived MDSSFTD measurements for (a) total, (b) diffuse fraction components with their *in-situ* counterparts for cloudy-sky retrievals. The retrievals are collected every 15 min. Blue line represents the mean fit across the whole evaluation data. Blue circles corresponds to low density of points and red circles corresponds to high density of points.

Table 4 Statistical scores obtained from the comparison between MDSSFTD total flux estimates and ground measurements over the selected BSRN sites for cloudy sky retrievals. If the value is in bold,

the metric does not meet the “target” requirements. If no value appears in bold, all the metrics meet the requirements.

	Lat °N	Lon °E	R_VAL [-]	RMSD [Wm ⁻²]	MBE [Wm ⁻²]	MBE (DSSF<200) [Wm ⁻²]	rMBE (DSSF>=200) [%]
Carpentras	44.08	5.06	0.886	124.089	-5.226	-20.663	-1.463
De Aar	-30.67	23.99	0.860	144.928	-22.573	-16.236	-5.585
Tamanrasset	22.79	5.53	0.861	174.924	54.220	-21.255	10.806
Toravere	58.25	26.46	0.880	117.562	-34.572	0.800	-8.025

Table 5 Statistical scores obtained from the comparison between MDSSFTD diffuse fraction estimates and ground measurements over the selected BSRN sites for cloudy sky retrievals. If the value is in bold, the metric does not meet the “target” requirements. If no value appears in bold, all the metrics meet the requirements.

	Lat °N	Lon °E	R_VAL [-]	RMSD [-]	MBE [-]	MBE (fd<0.5) [-]	rMBE (fd>=0.5) [%]
Carpentras	44.08	5.06	0.620	0.246	-0.108	-0.047	-15.082
De Aar	-30.67	23.99	0.489	0.287	-0.109	-0.007	-17.631
Tamanrasset	22.79	5.53	0.418	0.306	-0.198	-0.079	-25.767
Toravere	58.25	26.46	0.647	0.215	-0.060	0.064	-11.693

3.3.3. All-sky (clear and cloudy) conditions

In a similar way, the total DSSF and diffuse fraction from the MDSSFTD product for all sky retrievals are compared against their *in-situ* counterparts in Figure 9. We remind that the metrics obtained for all-sky conditions are those that are used to evaluate the performances of the product in the framework of the LSA SAF program (see Section 0).

Figure 9 displays the density scatter plot between instantaneous measurements of MSG-derived surface downwelling solar flux measurements for total and diffuse fraction components with their *in-situ* counterparts for the all-sky retrievals. Figure 9 shows that the satellite estimates of DSSF and fd meet the requirements. For DSSF, the requirements are 20Wm⁻² for DSSF<200Wm⁻² and 10% for DSSF>=200Wm⁻². The MBE and rMBE compared to the ground measurements are 3.618Wm⁻² and 0.252%, respectively. For fd, the requirements are 0.1 for fd<0.5 and 20% for fd>=0.5Wm⁻². The MBE and rMBE compared to the ground measurements are 0.044 and -17.699%, respectively. The statistical scores in terms of MBE and RMSD for the comparison between MDSSFTD total and diffuse fraction components with their *in-situ* counterparts for all four stations are given in Tables 6 and 7. The scores for all stations are in agreement with the MDSSFTD product requirements. The diffuse fraction compares well for most stations except in De Aar and Tamanrasset if a 20% threshold is considered for fd>0.5.

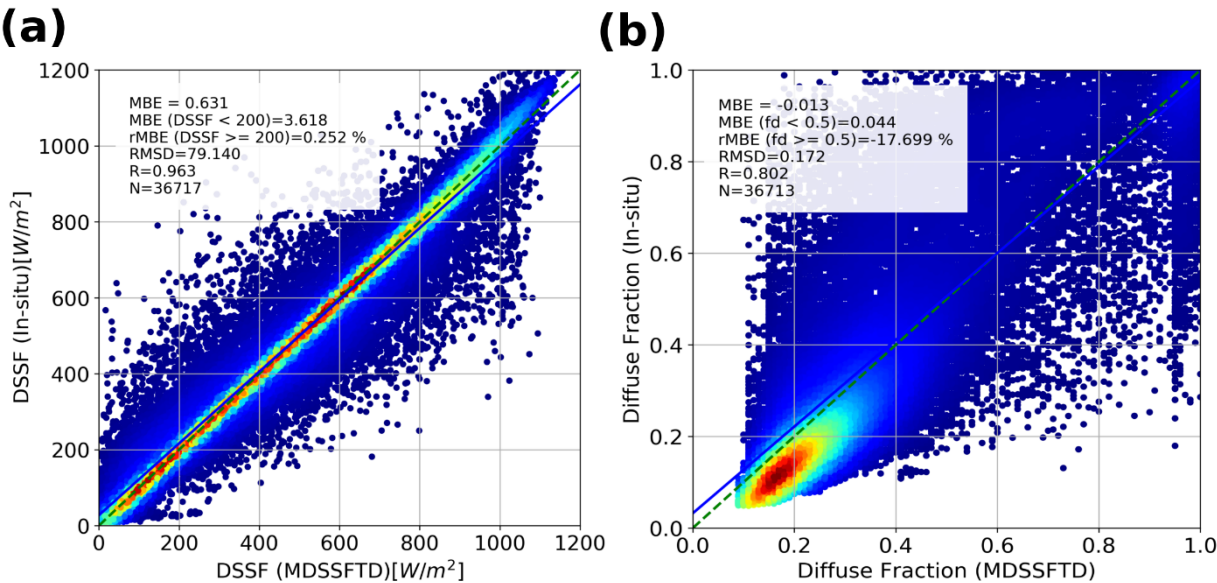


Figure 9 Comparison of instantaneous MSG-derived MDSSFTD measurements for (a) total, (b) diffuse fraction components with their *in-situ* counterparts for all sky (clear and cloudy) retrievals. The retrievals are collected every 15 min. Blue line represents the mean fit across the whole evaluation data. Blue circles corresponds to low density of points and red circles corresponds to high density of points.

Table 6 Statistical scores obtained from the comparison between MDSSFTD total flux estimates and ground measurements over the selected BSRN sites for all (clear and cloudy) sky retrievals. If the value is in bold, the metric does not meet the “target” requirements. If no value appears in bold, all the metrics meet the requirements.

	Lat	Lon	R_VAL [-]	RMSD [Wm ⁻²]	MBE [Wm ⁻²]	MBE (DSSF<200) [Wm ⁻²]	rMBE (DSSF>=200) [%]
Carpentras	44.08	5.06	0.969	69.584	10.790	0.728	2.037
De Aar	-30.67	23.99	0.969	64.833	-4.015	5.891	-0.993
Tamanrasset	22.79	5.53	0.965	86.075	10.722	5.034	2.939
Toravere	58.25	26.46	0.917	94.607	-18.026	3.604	-4.125

Table 7 Statistical scores obtained from the comparison between MDSSFTD diffuse fraction estimates and ground measurements over the selected BSRN sites for all (clear and cloudy) sky retrievals. If the value is in bold, the metric does not meet the “target” requirements. If no value appears in bold, all the metrics meet the requirements.

	Lat	Lon	R_VAL [-]	RMSD [-]	MBE [-]	MBE (fd<0.5) [-]	rMBE (fd>=0.5) [-]
Carpentras	44.08	5.06	0.818	0.151	-0.006	0.029	-14.753
De Aar	-30.67	23.99	0.734	0.161	0.022	0.054	-21.809
Tamanrasset	22.79	5.53	0.755	0.184	-0.034	0.057	-22.949
Toravere	58.25	26.46	0.786	0.191	-0,035	0.038	-13.361

3.4. Stability of the metrics

Here, we present the time series of the mean statistics averaged over all stations for the MDSSFTD total flux products (DSSF and f_d). The goal is to study the temporal evolution of performances with time. The 15-min statistics between the 15-min satellite derived products and the 15-min resampled ground measurements are averaged on a daily basis. Standard deviation of the 15-min statistics are also calculated on a daily basis and reported in the following plots. The statistics are calculated for both DSSF regimes and both outputs: for total flux, DSSF less than 200 Wm^{-2} and greater than 200 Wm^{-2} ; and for the diffuse fraction, f_d lower than 0.5 and greater than 0.5.

First, Figures 10 and 13 show the time series of the metrics for clear-sky conditions. Second, Figures 11 and 14 show the time series of the metrics for cloudy-sky conditions. Finally, Figures 12 and 15 show the time series of the metrics for all-sky (clear and cloudy) conditions. All figures show the daily averages along with the standard deviation, which is related to the variation of values among the different stations. First, it is important to highlight that all these conditions do not have the same level of representativeness due to the varying number of samples in the different cases. The only case that frequently shows values going beyond the requirement limits (i.e. the horizontal green lines) is the cloudy-sky case for high values of f_d ($f_d \geq 0.5$; see Figure 14). In all the other conditions, and especially for all-sky (clear and cloudy) conditions, the average statistics obtained from the product outputs meet the target requirement along the entire period of the analysis.

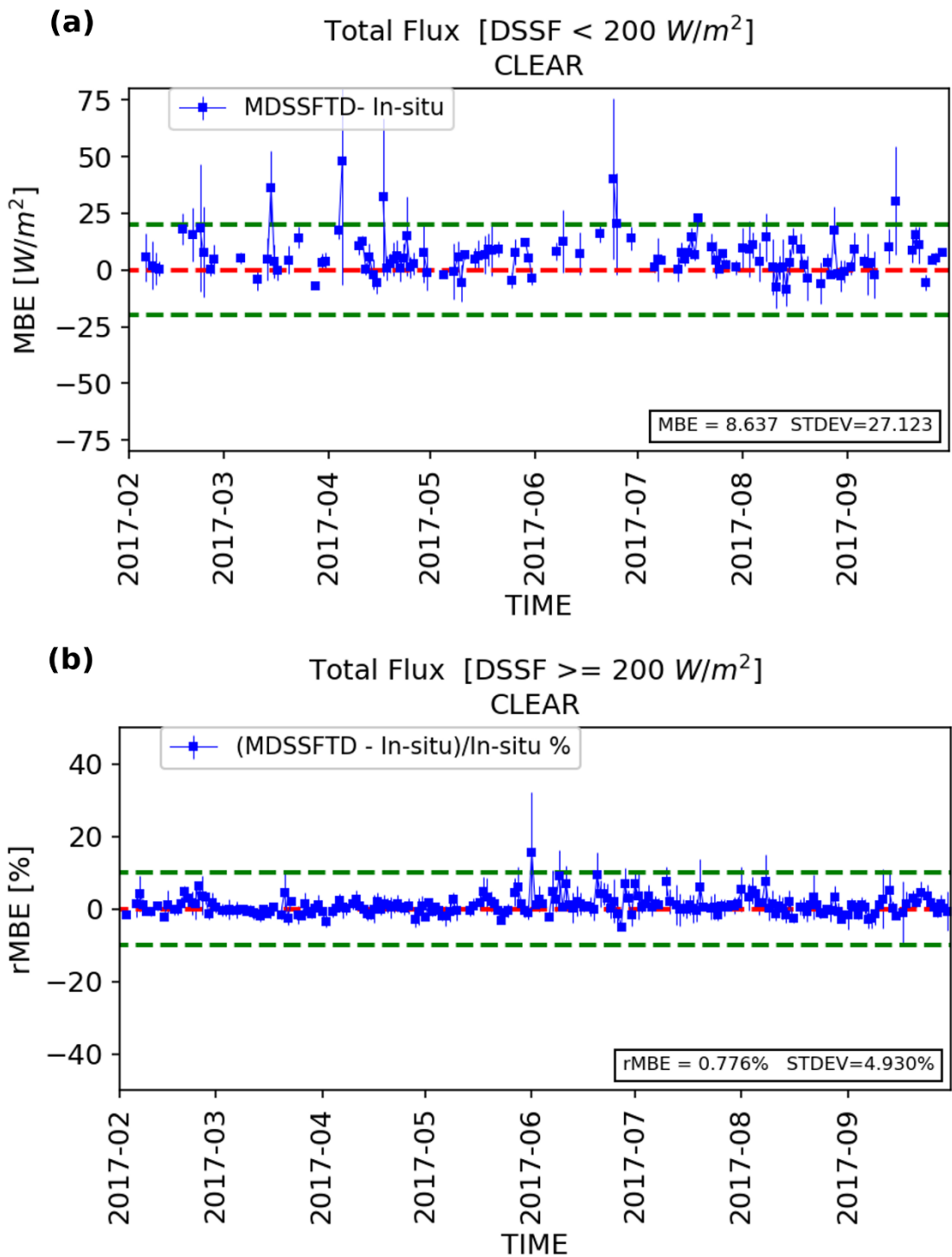


Figure 10 Time series of statistics of difference averaged on a daily basis between in 15 min *in-situ* measurements and 15 min MDSSFTD total flux (blue dots). The daily standard deviation of the absolute and relative statistics are indicated with vertical blue lines. The data points are filtered to keep those a) total flux values less than 200 Wm⁻² (absolute statistics) and b) total flux values greater than 200 Wm⁻² (relative statistics). The comparison is made only for the clear-sky conditions. The green dotted horizontal lines characterize the “target” accuracy requirements.

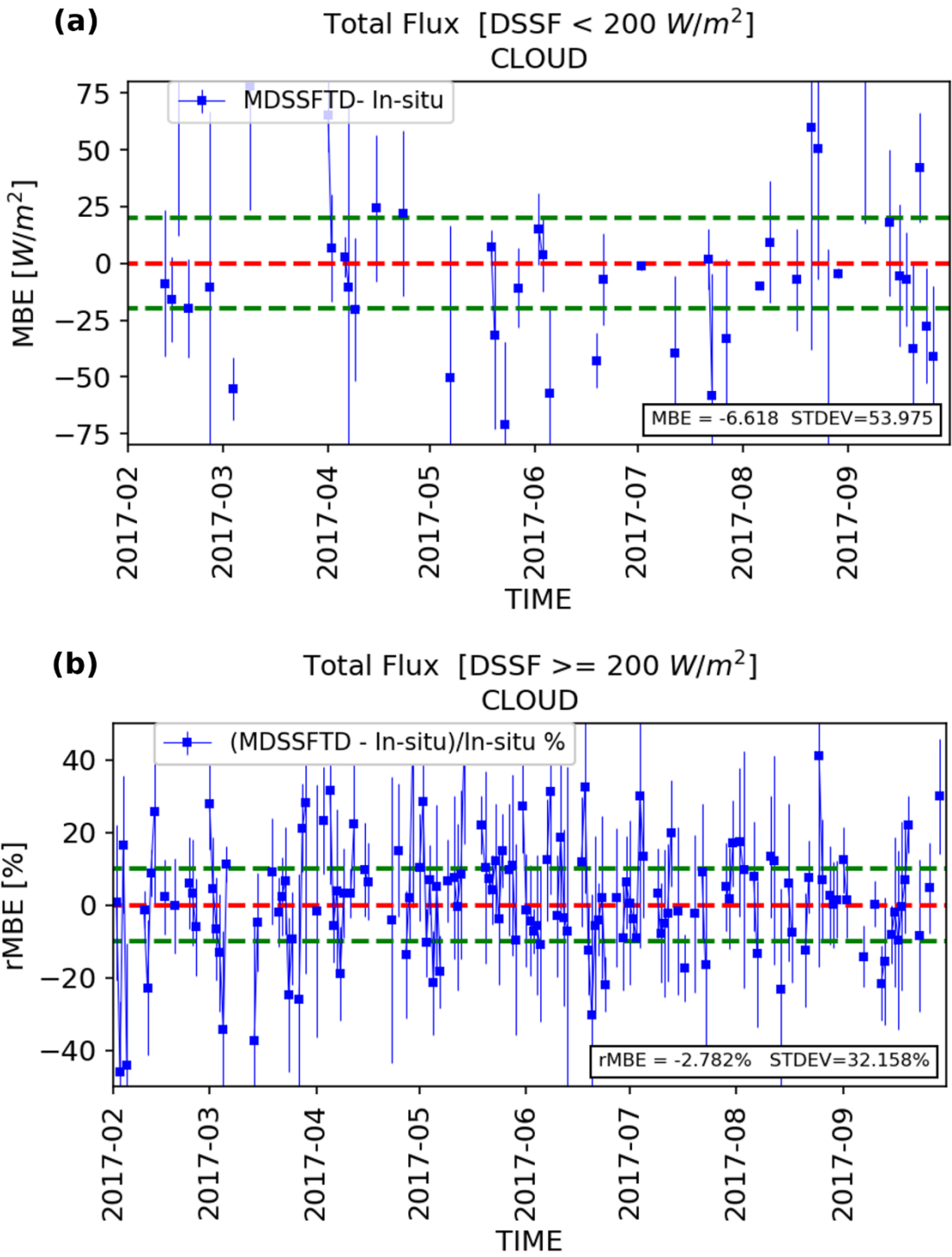


Figure 11 Same than Figure 10 for cloudy-sky conditions.

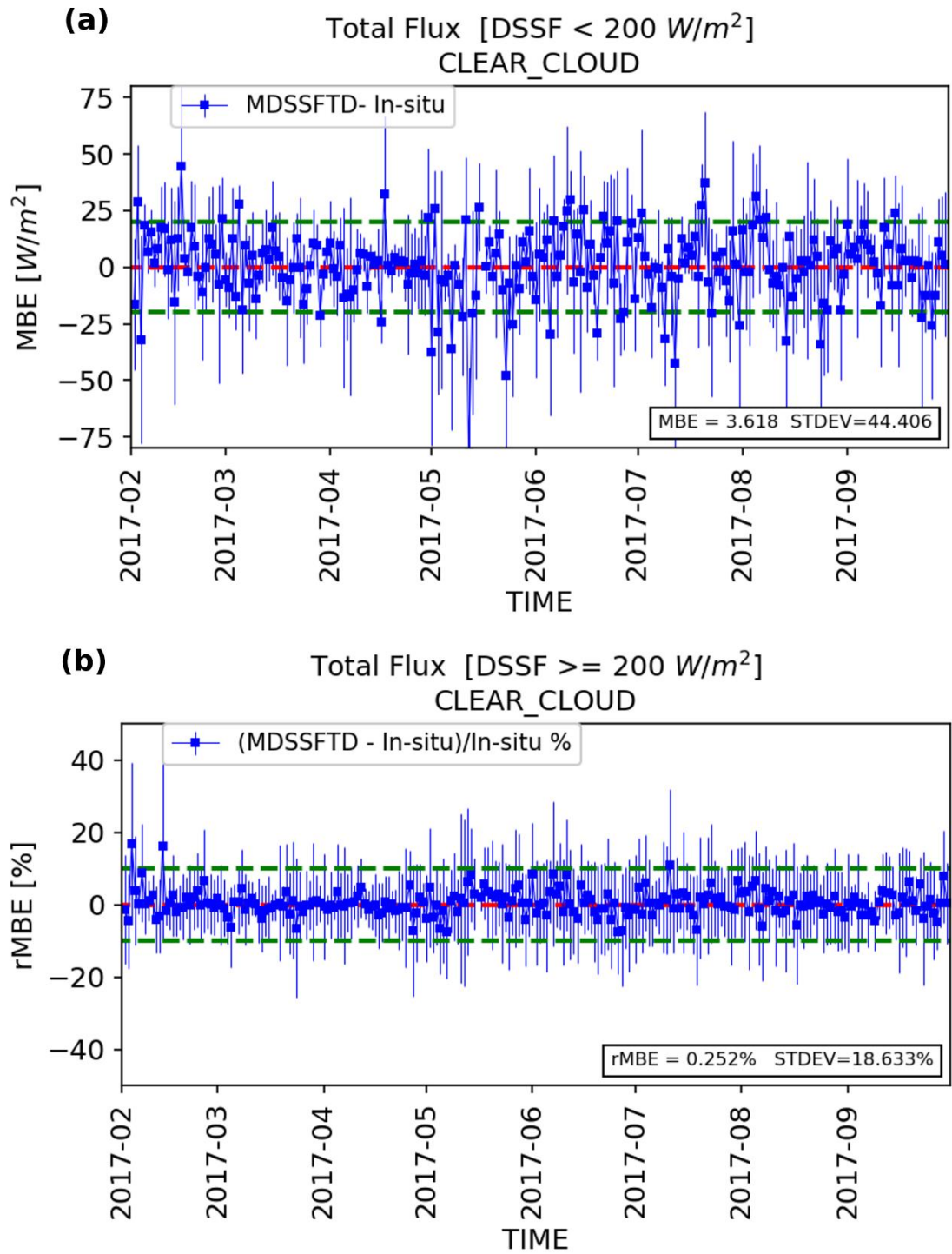


Figure 12 Same than Figure 10 for all-sky (clear and cloudy) conditions.

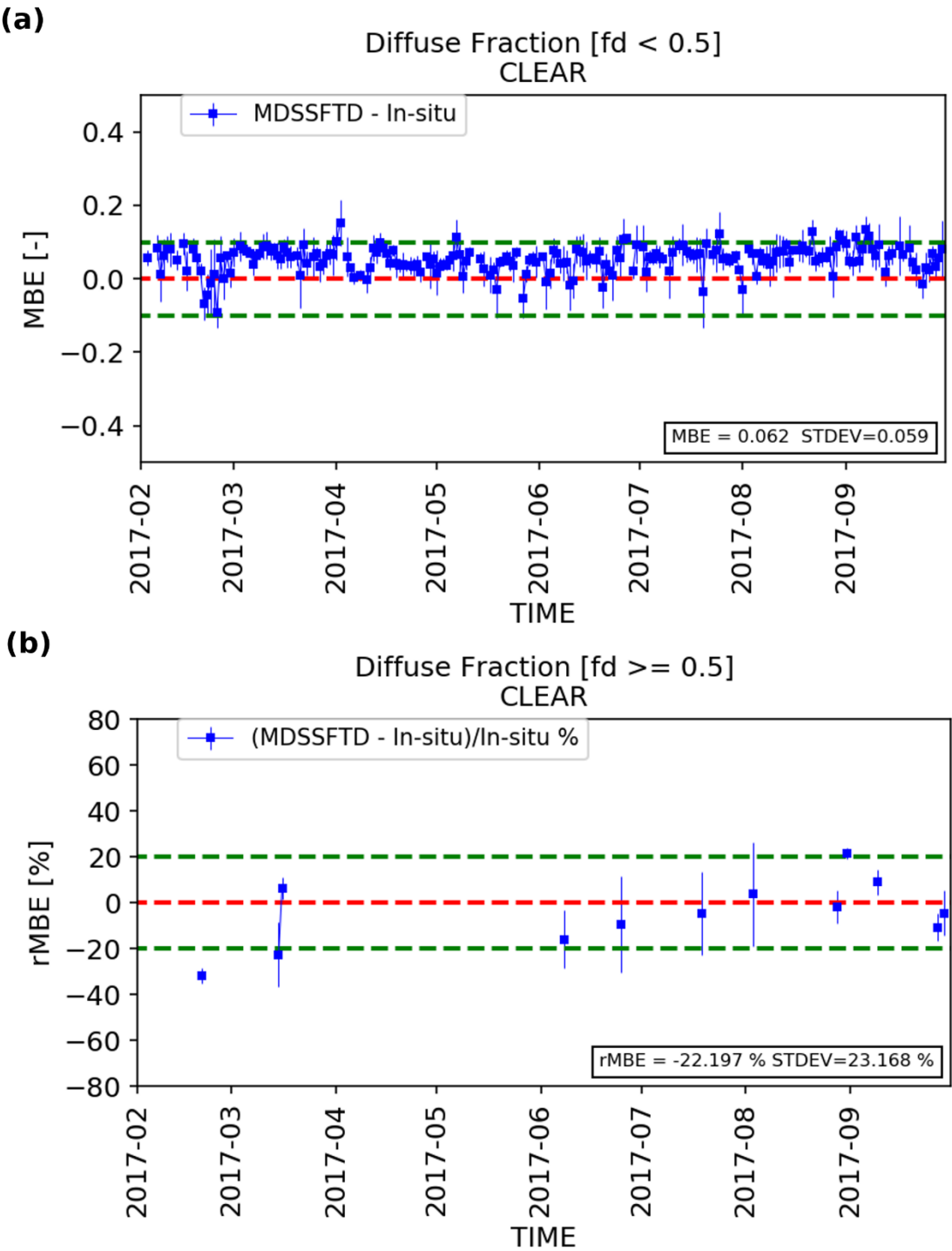


Figure 13 Time series of the relative mean bias for the comparison of the diffuse fraction on the 15 minutes time step basis. Statistics of difference are averaged on a daily basis. The standard deviation of the relative statistics are indicated with vertical blue lines. The comparison is made only for clear-sky conditions. The green dotted horizontal lines characterize the “target” requirements.

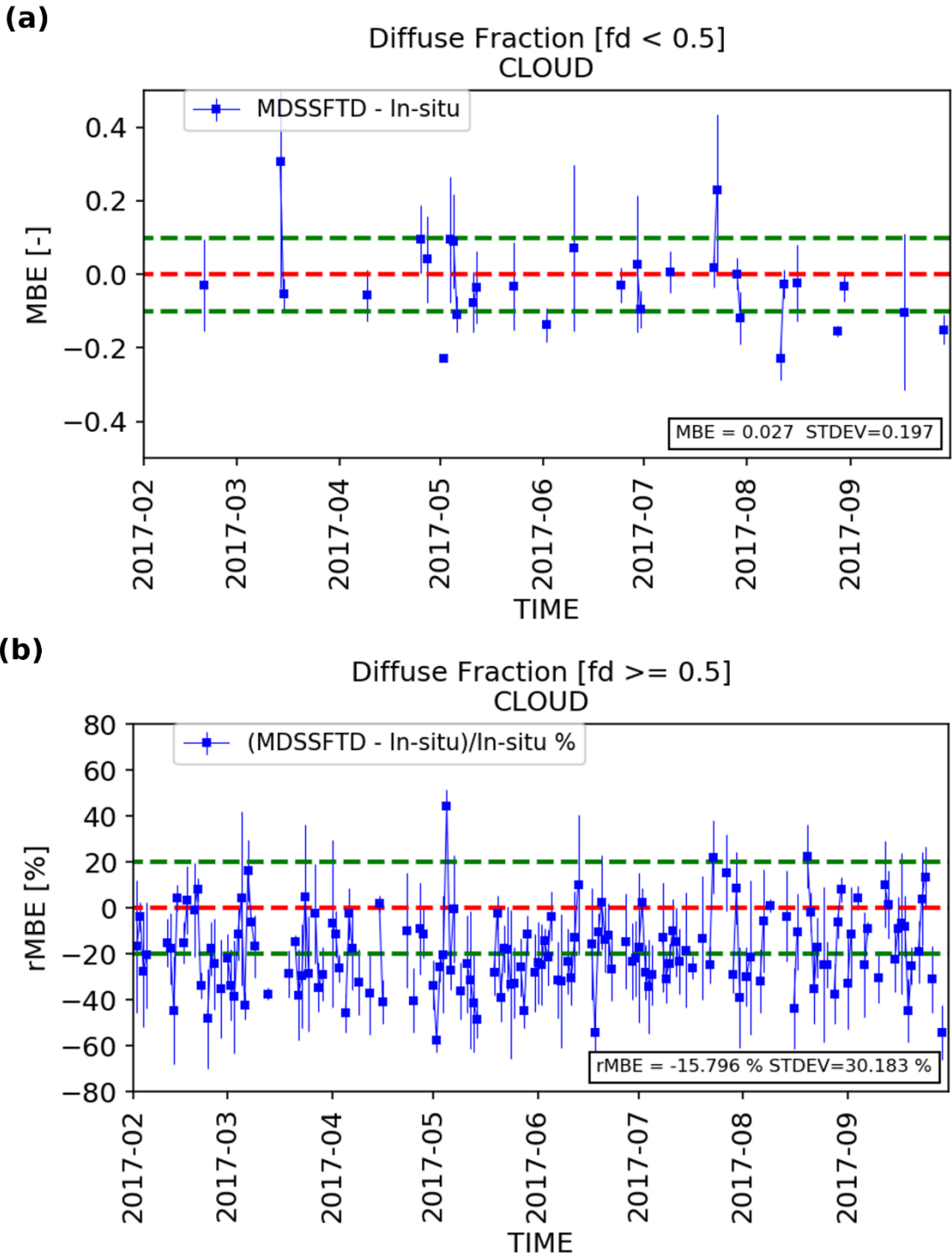


Figure 14 Same than Figure 13 for cloudy-sky conditions.

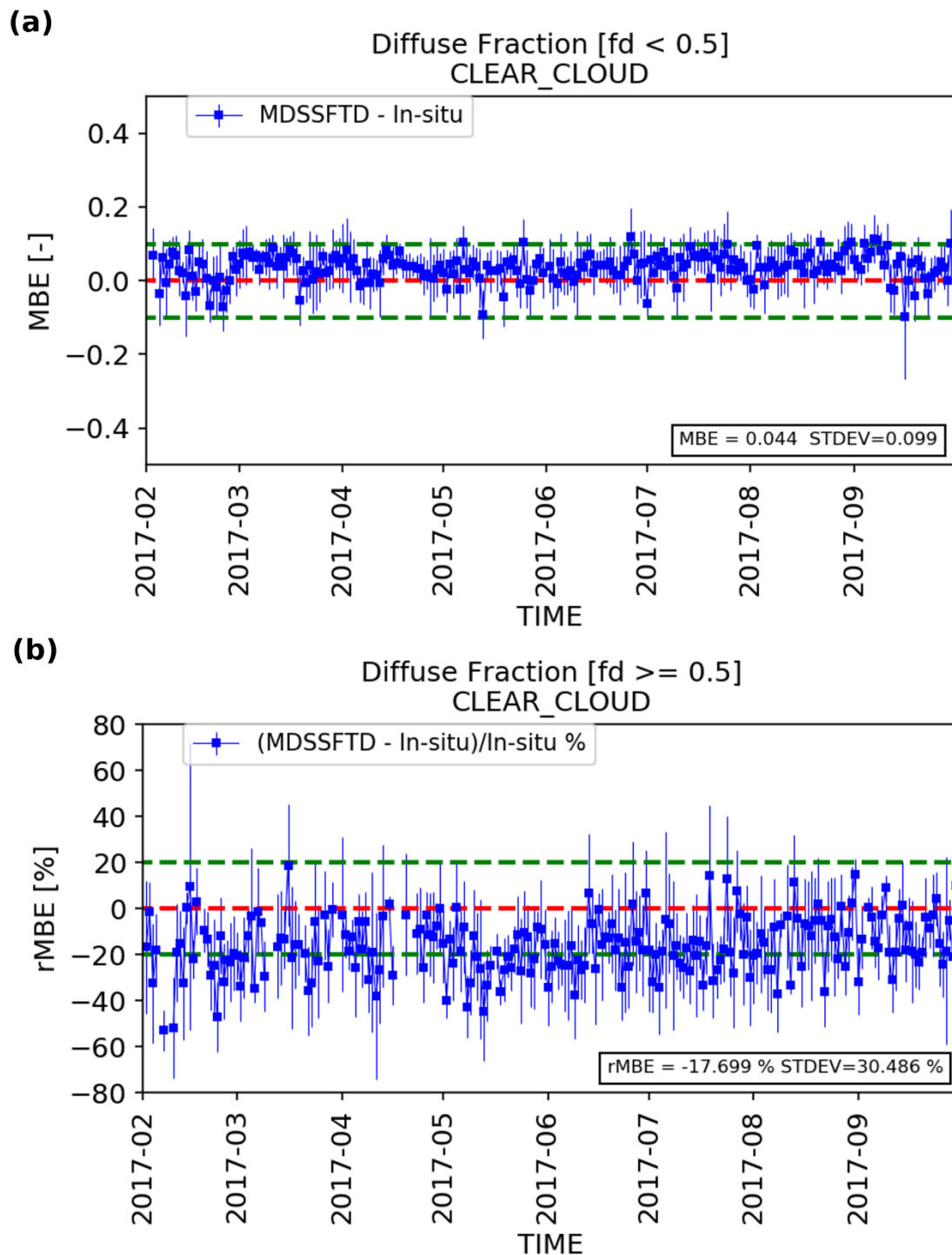


Figure 15 Same than Figure 13 for all-sky (clear and cloudy) conditions.

3.5. Impact of the activation of the cloud-aerosol coupling

The method for DSSF retrieval is using a simple radiative transfer model that takes into account the radiative coupling between aerosols and clouds as described in the companion paper [40] (see fourth term of Eq. 24). Figure 16 gives an example of this cloud-aerosol coupling for a selected day in Carpentras. We clearly observe a better agreement with the *in-situ* measurements around noon in the case of the activation of the cloud-aerosol radiative coupling (yellow dots compared to light blue dots). Even if the AOD is not large (i.e. 0.2), the impact of the cloud-aerosol coupling remains important. We are here clearly in presence of very thin clouds in the high atmosphere.

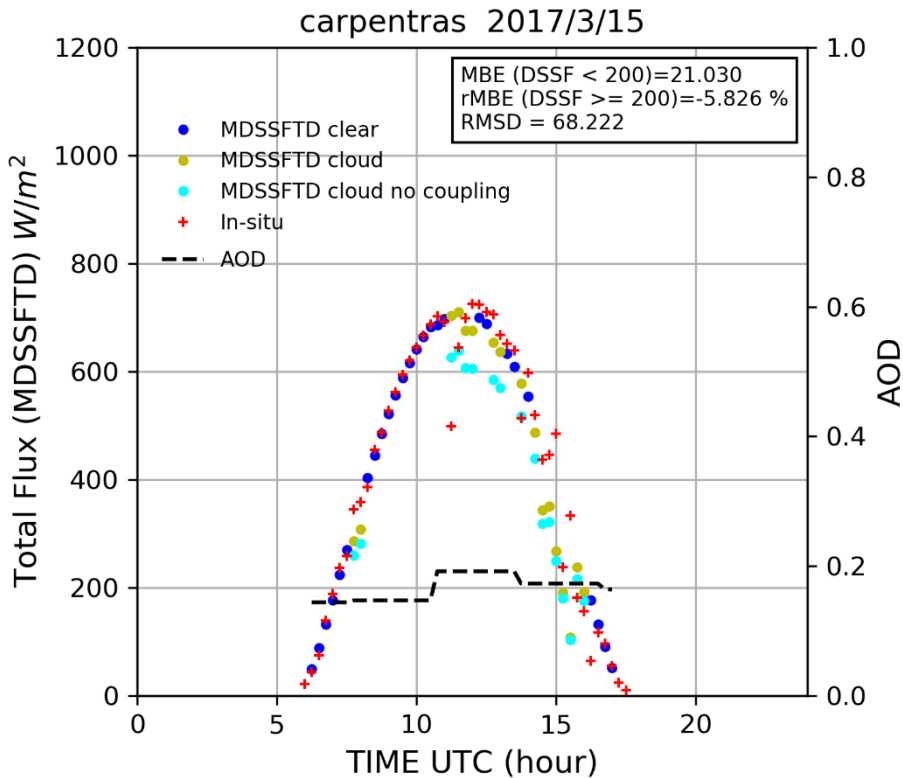


Figure 16 Same diurnal variation than in Figure 5 for Carpentras. Light-blue dots show the estimations of total MDSSFTD with ‘no coupling’ (i.e., no activation of the cloud-aerosol coupling) in cloudy conditions. Yellow dots show same LSA-207 DSSF retrievals in cloudy conditions than in Figure 5. Black dashed line represents the AOD (CAMS). Statistics in the top right corner are those of the MDSSFTD product (with activation of the cloud-aerosol coupling).

Figure 11 showed the performances of the DSSF estimated by the MDSSFTD algorithm, which considers the influence of the cloud-aerosol coupling under cloudy-sky conditions. MDSSFTD satellite estimates are very close to the *in-situ* measurements (MBE=-6.618, rMBE=-2.782%). Figure 17 now shows the same comparison by using the same code after disabling the coupling between cloud and aerosols (by simply removing the fourth term of Eq. 24 in [40]). We clearly observe a large degradation of the performances of the algorithm in this case (MBE=-11.807 Wm⁻², rMBE=-11.272%). The presence of aerosols makes the atmospheric transmittance decrease, and in turn the DSSF becomes lower. However, this atmospheric transmittance decreasing is too large in cloudy conditions. This sensitivity test illustrates the importance of the indirect radiative impacts of clouds on aerosol radiative forcing. In our study, the activation of the coupling improves the performances of about 8%. Clouds induce an increase of the atmospheric transmittance by reflecting, back to the surface, part of the radiation scattered by aerosols. This radiative cloud-aerosol coupling is included in the LSA-207 product.

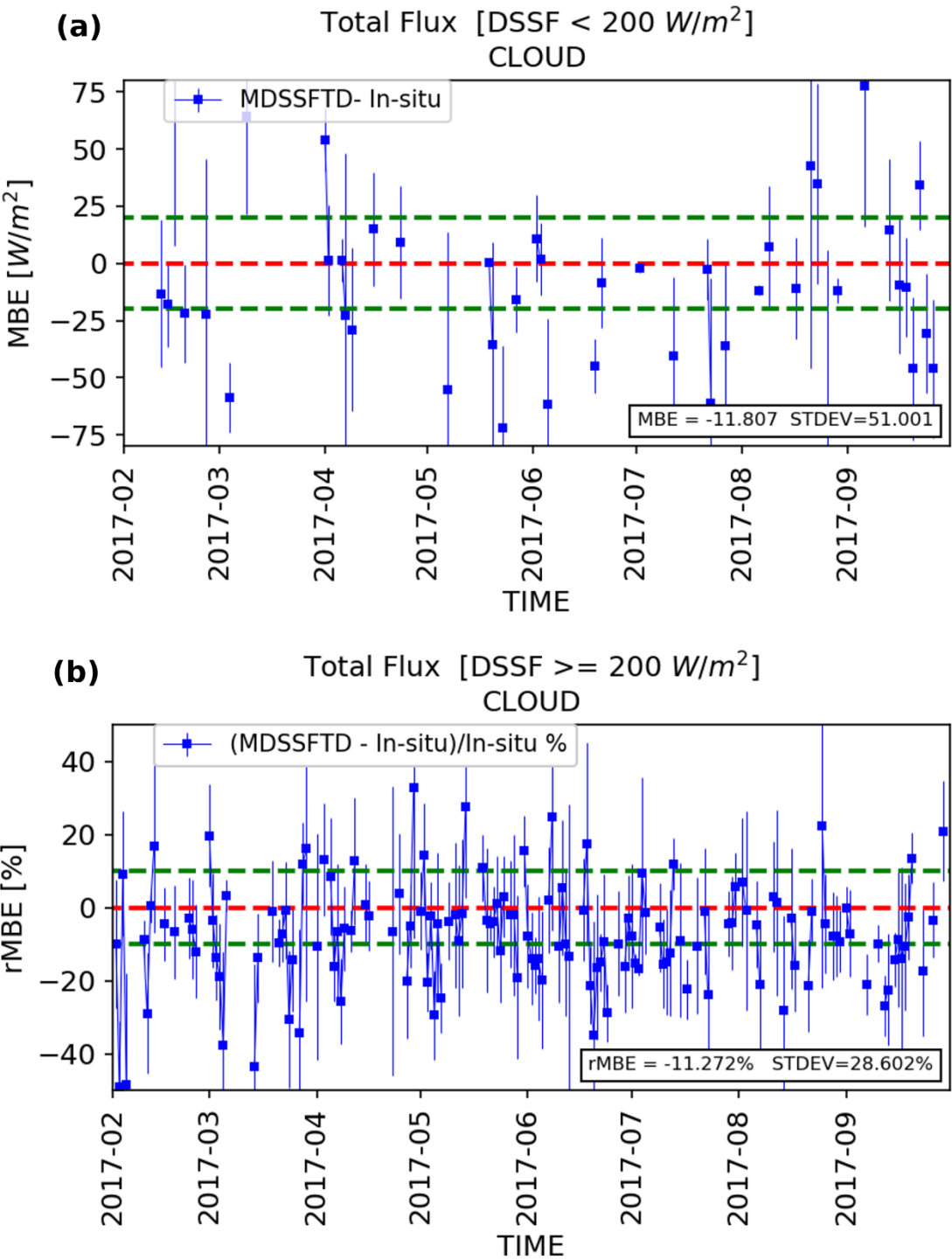


Figure 17 Same than Figure 11 for cloudy-sky conditions but inactivating the coupling between cloud and aerosol.

3.6. Comparison to CAMS radiation product

The retrievals from MDSSFTD are compared against the counterpart estimates from the CAMS radiation product. Figure 18(a) shows the good agreement between the two products for all-sky conditions, with a correlation of 0.966. Figures 18(b) and 18(c) take a further look to the comparison by exploring the clear sky and cloudy sky retrievals separately. The clear sky comparison gives a high agreement between the two products, which are using CAMS aerosol data as input. The comparison for cloudy sky also shows a good agreement (correlation of 0.909)

despite the differences of the retrieval methods for cloudy sky conditions. The higher dispersion is justified by the increased difficulty of the retrieval for cloudy skies.

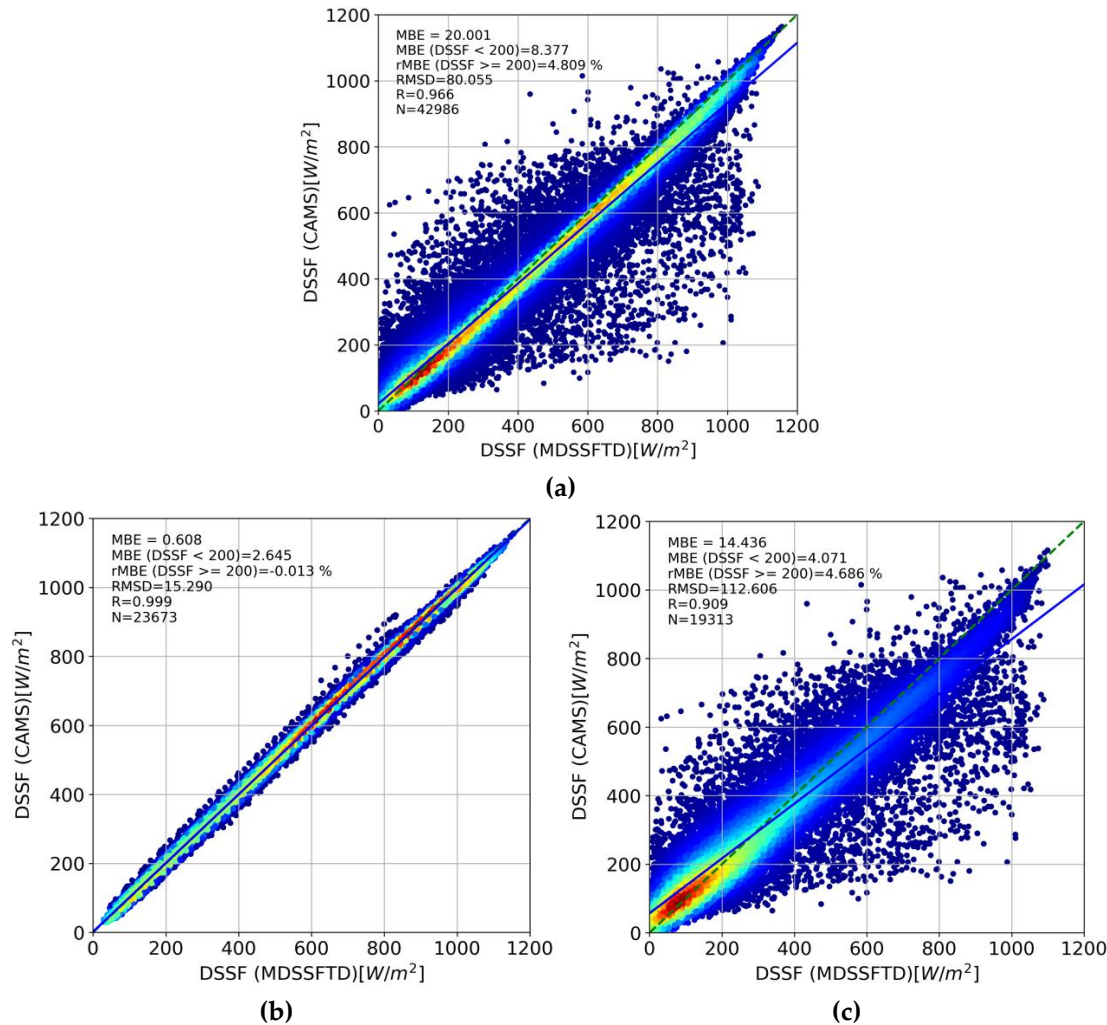


Figure 18 Density scatter plots for the comparison between MDSSFTD and CAMS radiation product for (a) all-sky, (b) clear sky, and (c) cloudy sky conditions.

4. Conclusions

This paper presents the results of the comparison of the LSA-207 MDSSFTD product outputs, namely the total DSSF and diffuse fraction (fd) components, against the *in-situ* measurements acquired at four BSRN stations over a seven-month period. The validation is performed on instantaneous satellite retrievals with MSG/SEVIRI (i.e. acquired every 15 minutes).

The results show that the satellite estimates of DSSF and fd meet the requirements for all-sky (clear and cloudy) conditions. For DSSF, the requirements are 20Wm^{-2} for $\text{DSSF} < 200\text{Wm}^{-2}$ and 10% for $\text{DSSF} \geq 200\text{Wm}^{-2}$. The MBE and rMBE compared to the ground measurements are 3.618Wm^{-2} and 0.252%, respectively. For fd, the requirements are 0.1 for $\text{fd} < 0.5$ and 20% for $\text{fd} \geq 0.5$. The MBE and rMBE compared to the ground measurements are -0.044 and -17.699%, respectively.

A more detailed analysis of the product performances was also performed separately for clear and cloudy sky conditions. For DSSF in clear-sky conditions, the MBE and rMBE compared to the ground measurements are 8.637Wm^{-2} and 0.776%, respectively. For fd, the MBE and rMBE compared to the ground measurements are 0.062 and -22.197%, respectively. Thus, the two products outputs also meet the target requirements if only clear-sky conditions are selected and if we do not consider $\text{fd} \geq 0.5$ case (which is not statistically representative). For DSSF in cloudy-sky conditions, the MBE and rMBE compared to the ground measurements are -6.618Wm^{-2} and 2.782%, respectively. For fd, the MBE and rMBE compared to the ground measurements are 0.027 and -15.796%, respectively. Thus,

the product meets the target requirements for all conditions with only a few exceptions. The major limitations of the retrieval approach described in the companion article [40] are not an obstacle for meeting the required quality. It is noted that the requirements for the product MDSSFTD are defined for the all-sky conditions only.

In an earlier study by [42], it was shown that the use of MACC-II (now CAMS) AOD forecasts as input to the MDSSFTD clear sky method instead of reanalyses significantly decreased the quality of the DSSF products under clear sky conditions. For the last years, the quality of the CAMS AOD forecasts currently available could have improved, which makes the high sensitivity of the MDSSFTD diffuse estimation to the quality of AOD forecasts not to be a limitation anymore. Finally, we show that this AOD information is of primary importance for the estimation of the atmospheric transmittance either in clear or in cloudy conditions. In cloudy-sky conditions, the modelling of the cloud-aerosol radiative coupling allows to reduce the overall bias by around 8%.

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References

[1] Mateos, D., Antón, M., Valenzuela, A., Cazorla, A., Olmo, F. J., & Alados-Arboledas, L. (2013). Short-wave radiative forcing at the surface for cloudy systems at a midlatitude site. *Tellus B: Chemical and Physical Meteorology*, 65(1), 21069.

[2] Van Tricht, K., Lhermitte, S., Lenaerts, J.T.M., Gorodetskaya, I.V., L'Ecuyer, T.S., Noël, B., van den Broeke, M.R., Turner, D.D., van Lipzig, N.P.M., 2016. Clouds enhance Greenland ice sheet meltwater runoff. *Nat. Commun.* 7, 10266.

[3] Mercado, L. M., Bellouin, N., Sitch, S., Boucher, O., Huntingford, C., Wild, M., & Cox, P. M. (2009). Impact of changes in diffuse radiation on the global land carbon sink. *Nature*, 458(7241), 1014.

[4] Carrer, D., Roujean, J. L., Lafont, S., Calvet, J. C., Boone, A., Decharme, B., ... & Gastellu-Etchegorry, J. P. (2013). A canopy radiative transfer scheme with explicit FAPAR for the interactive vegetation model ISBA-A-gs: Impact on carbon fluxes. *Journal of Geophysical Research: Biogeosciences*, 118(2), 888-903.

[5] O'Sullivan, M. et al. Small global effect on terrestrial net primary production due to increased fossil fuel aerosol emissions from East Asia since the turn of the century. *Geophys. Res. Lett.* 43, 8060–8067 (2016).

[6] Yoshida, S., Ueno, S., Kataoka, N., Takakura, H., & Minemoto, T. (2013). Estimation of global tilted irradiance and output energy using meteorological data and performance of photovoltaic modules. *Solar Energy*, 93, 90-99.

[7] Betts, A. K., S.-Y. Hong, and H.-L. Pan, 1996: Comparison of NCEP–NCAR reanalysis with 1987 FIFE data. *Mon. Wea. Rev.*, 124, 1480–1498.

[8] Brotzge, J. A., 2004: A two-year comparison of the surface water and energy budgets between two OASIS sites and NCEP–NCAR reanalysis data. *J. Hydrometeor.*, 5, 311–326.

- [9] Berbery, E. H., K. E. Mitchell, S. Benjamin, T. Smirnova, H. Ritchie, R. Hogue, and E. Radeva, 1999: Assessment of land-surface energy budgets from regional and global models. *J. Geophys. Res.*, 104, 19 329–19 348.
- [10] Schroeder TA, Hamber R, Copps NC and Liang S (2009) Validation of solar radiation surfaces from MODIS and reanalysis data over topographically complex terrain. *J. App. Meteor. Climat.* 48, 2441-2458
- [11] Babst, F., R. W. Mueller, and R. Hollman, 2008: Verification of NCEP reanalysis shortwave radiation with mesoscale remote sensing data. *IEEE Trans. Geosci. Remote Sens.*, 5, 34–37.
- [12] Urraca, R., Huld, T., Gracia-Amillo, A., Martinez-de-Pison, F. J., Kaspar, F., & Sanz-Garcia, A.: Evaluation of global horizontal irradiance estimates from ERA5 and COSMO-REA6 reanalyses using ground and satellite-based data, *Solar Energy*, 164, 339-354, 2018.
- [13] Bishop, J. K. B., and Rossow, W. B., 1991. Spatial and temporal variability of global surface solar irradiance. *Journal of Geophysical Research*, 96, 16389 – 16858.
- [14] Darnell, W., Staylor, W., Gupta, S., and Denn, F., 1988. Estimation of surface insolation using sun-synchronous satellite data. *Journal of Climate*, 1, 820 – 835.
- [15] Dedieu, G. P., Deschamps, P., and Kerr, Y., 1987. Satellite estimation of solar irradiance at the surface of the earth and of surface albedo using a physical model applied to METEOSAT data. *Journal of Climate and Applied Meteorology*, 26, 79 –87.
- [16] Gautier C., Diak G., Masse S., 1980, A simple physical model to estimate incident solar radiation at the surface from GOES satellite data, *J. Climate Appl. Meteor.*, 19, 1005-1012.
- [17] Cano, D., Monget, J. M., Albuissou, M., Guillard, H., Regas, N., Wald, L., 1986: A method for the determination of the global solar radiation from meteorological satellite data. *Solar Energy*, 37, 31±39.
- [18] Gautier, C., & Landsfeld, M. (1997). Surface solar radiation flux and cloud radiative forcing for the Atmospheric Radiation Measurement (ARM) Southern Great Plains (SGP): A satellite, surface observations, and radiative transfer model study. *Journal of the atmospheric sciences*, 54(10), 1289-1307.
- [19] Li, Z., and Leighton, H. G. (1993). Global climatologies of the solar radiation budgets at the surface and in the atmosphere from 5 years of ERBE data. *Journal of Geophysical Research*, 98, 4919 – 4930.
- [20] Masuda, K., Leighton, H. G., & Li, Z. (1995). A new parameterization for the determination of solar flux absorbed at the surface from satellite measurements. *Journal of Climate*, 8(6), 1615-1629.
- [21] Hammer, A., Heinemann, D., Lorenz, E., & Lücke, B. (1999). Short-term forecasting of solar radiation based on image analysis of meteosat data. *Proc. EUMETSAT meteorological satellite data users conference* (pp. 331–337).

- 541 [22] Möser, W., and E. Raschke, 1984: Incident solar radiation over Europe from METEOSAT data. *J. Climate*
542 *Appl. Meteor.*, 23, 166–170.
- 543 [23] Pinker, R., and Ewing, J. (1985). Modeling surface solar radiation: model formulation and validation. *Journal*
544 *of Climate and Applied Meteorology*, 24, 389 – 401.
- 545 [24] Pinker, R., and Laszlo, I. (1992). Modeling surface solar irradiance for satellite applications on a global scale.
546 *Journal of Applied Meteorology*, 31, 194 – 211.
- 547 [25] Tarpley, J. (1979). Estimating incident solar radiation at the surface from geostationary satellite data. *Journal*
548 *of Climate and Applied Meteorology*, 18, 1172 –181.
- 549 [26] Whitlock, C. H., Charlock, T. P., Staylor, W. F., Pinker, R. T., Laszlo, I., Ohmura, A., Gilgen, H., Konzelman,
550 T., DiPasquale, R. C., Moats, C. D., LeCroy, S. R., and Ritchey, N. A. (1995). First global WCRP shortwave surface
551 radiation budget dataset. *Bulletin of the American Meteorological Society*, 76, 905 – 922.
- 552 [27] Romano, F.; Cimini, D.; Cersosimo, A.; Di Paola, F.; Gallucci, D.; Gentile, S.; Gerdali, E.; Larosa, S.; T. Nilo,
553 S.; Ricciardelli, E.; Ripepi, E.; Viggiano, M. Improvement in Surface Solar Irradiance Estimation Using
554 HRV/MSG Data. *Remote Sens.* 2018, 10, 1288.
- 555 [28] Gallucci, D.; Romano, F.; Cersosimo, A.; Cimini, D.; Di Paola, F.; Gentile, S.; Gerdali, E.; Larosa, S.; Nilo, S.T.;
556 Ricciardelli, E.; Viggiano, M. Nowcasting Surface Solar Irradiance with AMESIS via Motion Vector Fields of
557 MSG-SEVIRI Data. *Remote Sens.* 2018, 10, 845.
- 558 [29] Ineichen, P. High Turbidity Solis Clear Sky Model: Development and Validation. *Remote Sens.* 2018, 10,
559 435.
- 560 [30] Trigo, I. F., C. C. DaCamara, P. Viterbo, J.-L. Roujean, F. Olesen, C. Barroso, F. Camacho-de-Coca, D.
561 Carrer, S. C. Freitas, J. García-Haro, B. Geiger, F. Gellens-Meulenberghs, N. Ghilain, J. Meliá, L. Pessanha, N.
562 Siljamo, A. Arboleda (2011), The Satellite Application Facility on Land Surface Analysis, *Int. J. Remote Sens.*, ,
563 32, 2725-2744.
- 564 [31] Ineichen, P., Barroso, C. S., Geiger, B., Hollmann, R., Marsouin, A., and Mueller, R.: Satellite Application
565 Facilities irradiance products: hourly time step comparison and validation over Europe, *Int. J. Remote Sens.*, 30,
566 doi:0.1080/01431160802680560, 2009.
- 567 [32] Roerink, G. J., Bojanowski, J., de Wit, A. J. W., Eerens, H., Supit, I., Leo, O., and Boogaard, H. L. (2012)
568 Evaluation of MSG-derived global radiation estimates for application in a regional crop model. *Agricultural*
569 *and Forest Meteorology*, 160:36–47.
- 570 [33] Moreno, A. & Gilabert, M.A. & Camacho, F. & Martínez, B., 2013. "Validation of daily global solar
571 irradiation images from MSG over Spain," *Renewable Energy*, Elsevier, vol. 60(C), pages 332-342.
- 572

- 573 [34] Bevan, S. L., North, P. R. J., Los, S. O., and Grey, W. M. F.: A global dataset of atmospheric aerosol optical depth and
574 surface reflectance from AATSR, *Remote Sens. Environ.*, 116, 199–210, 2012.
- 575 [35] Jayaraman, A., Lubin, D., Ramachandran, S., Ramanathan, V., Woodbridge, E., Collins, W.D., Zalpuri, K.S.,
576 1998. Direct observations of aerosol radiative forcing over the tropical Indian Ocean during the January-February
577 1996 pre-INDOEX cruise. *J. Geophys. Res. Atmospheres* 103, 13827–13836.
- 578 [36] Satheesh, S.K., Ramanathan, V., 2000. Large differences in tropical aerosol forcing at the top of the
579 atmosphere and Earth's surface. *Nature* 405, 60–63. doi:10.1038/35011039
- 580 [37] Cherian, R., Quaas, J., Salzmann, M., Wild, M., 2014. Pollution trends over Europe constrain global aerosol
581 forcing as simulated by climate models. *Geophys. Res. Lett.* 41, 2176–2181. doi:10.1002/2013GL058715
- 582 [38] Dramé, M. S., Ceamanos, X., Roujean, J. L., Boone, A., Lafore, J. P., Carrer, D., & Geoffroy, O. (2015). On the
583 Importance of Aerosol Composition for Estimating Incoming Solar Radiation: Focus on the Western African
584 Stations of Dakar and Niamey during the Dry Season. *Atmosphere*, 6(11), 1608-1632.
- 585 [39] Kosmopoulos, P.G., Kazadzis, S., Taylor, M., Athanasopoulou, E., Speyer, O., Raptis, P.I., Marinou, E.,
586 Proestakis, E., Solomos, S., Gerasopoulos, E., Amiridis, V., Bais, A. and Kontoes, C. (2017). Dust impact on
587 surface solar irradiance assessed with model simulations, satellite observations and ground-based
588 measurements. *Atmos. Meas. Tech.* 10: 2435–2453.
- 589 [40] Carrer, D., Ceamanos, X., Moparthy, M., Vincent, C., Coehlo, S., Trigo, I., 2019. Satellite retrieval of
590 downwelling shortwave surface flux and diffuse fraction under all sky conditions in the framework of the LSA
591 SAF program (part 1: methodology). Submitted.
- 592 [41] Ceamanos, X., Carrer, D., Roujean, J.-L., An efficient approach to estimate the transmittance and reflectance
593 of a mixture of aerosol components, *Atmospheric Research*, Vol. 137, Feb. 2014a, pp. 125-135
- 594 [42] Ceamanos, X., Carrer, D., Roujean, J.-L., 2014b. Improved retrieval of direct and diffuse downwelling surface
595 shortwave flux in cloudless atmosphere using dynamic estimates of aerosol content and type: application to the
596 LSA-SAF project. *Atmospheric Chem. Phys.* 14, 8209–8232. doi:10.5194/acp-14-8209-2014
- 597 [43] Riihelä, A.; Kallio, V.; Devraj, S.; Sharma, A.; Lindfors, A. Validation of the Sarah-e satellite-based surface
- 598 [44] Riihelä, A.; Carlund, T.; Trentmann, J.; Müller, R.; Lindfors, A.V. Validation of CM SAF Surface Solar
599 Radiation Datasets over Finland and Sweden. *Remote Sens.* 2015, 7, 6663-6682.
- 600 [45] Geiger, B., Meurey, C., Lajas, D., Franchistéguy, L., Carrer, D., & Roujean, J. L. (2008). Near real-time
601 provision of downwelling shortwave radiation estimates derived from satellite observations. *Meteorological*
602 *Applications*, 15(3), 411-420.

- [46] Carrer, D., Lafont, S., Roujean, J.-L., Calvet, J.-C., Meurey, C., Le Moigne, P., and Trigo, I. : Incoming solar and infrared radiation derived from METEOSAT: impact on the modelled land water and energy budget over France, *J. Hydrometeor.*, 13, 504–520, 2012.
- [47] McArthur, L. J. B. "Baseline Surface Radiation Network (BSRN)-Operation Manual Version 2.1." (2005).
- [48] Vuilleumier, L., Hauser, M., Félix, C., Vignola, F., Blanc, P., Kazantzidis, A., & Calpini, B. (2014). Accuracy of ground surface broadband shortwave radiation monitoring. *Journal of geophysical research: atmospheres*, 119(24), 13-838.
- [49] García, R. D., Cuevas, E., Ramos, R., Cachorro, V. E., Redondas, A., & Moreno-Ruiz, J. A. (2019). Description of the Baseline Surface Radiation Network (BSRN) station at the Izaña Observatory (2009–2017): measurements and quality control/assurance procedures. *Geoscientific Instrumentation, Methods and Data Systems*, 8(1), 77-96.
- [50] Qu, Z., Oumbe, A., Blanc, P., Espinar, B., Gesell, G., Gschwind, B., ... & Wald, L. (2017). Fast radiative transfer parameterisation for assessing the surface solar irradiance: The Heliosat-4 method. *Meteorologische Zeitschrift*, 26(1), 33-57.
- [51] Lefèvre, M., Oumbe, A., Blanc, P., Espinar, B., Gschwind, B., Qu, Z., ... & Benedetti, A. (2013). McClear: a new model estimating downwelling solar radiation at ground level in clear-sky conditions. *Atmospheric Measurement Techniques*, 6, 2403-2418.
- [52] Gschwind, B., Wald, L., Blanc, P., Lefèvre, M., Schroedter-Homscheidt, M., & Arola, A. (2019). Improving the McClear model estimating the downwelling solar radiation at ground level in cloud-free conditions—McCclear-v3. *Meteorologische Zeitschrift*.
- [53] Mayer, B., & Kylling, A. (2005). The libRadtran software package for radiative transfer calculations—description and examples of use. *Atmospheric Chemistry and Physics*, 5(7), 1855-1877.
- [54] Blanc, P., B. Gschwind, M. Lefèvre, L. Wald, 2014: Twelve monthly maps of ground albedo parameters derived from MODIS data sets. In *Proceedings of IGARSS 2014*, held 13-18 July 2014, Quebec, Canada, USBKey, pp. 3270-3272.
- [55] Schaaf, C. B., Gao, F., Strahler, A. H., Lucht, W., Li, X., Tsang, T., ... & Lewis, P. (2002). First operational BRDF, albedo nadir reflectance products from MODIS. *Remote sensing of Environment*, 83(1-2), 135-148.
- [56] Kriebel, K. T., Saunders, R. W., & Gesell, G. (1989). Optical properties of clouds derived from fully cloudy AVHRR pixels. *Beiträge zur Physik der Atmosphäre*, 62, 165-171.

- 632 [57] Kriebel, K. T., Gesell, G., Ka" Stner, M., & Mannstein, H. (2003). The cloud analysis tool APOLLO:
633 improvements and validations. *International journal of remote sensing*, 24(12), 2389-2408.
- 634 [58] Hakuba MZ, Folini D, Sanchez-Lorenzo A, Wild M (2013) Spatial representativeness of ground-based solar
635 radiation measurements. *J Geophys Res* 118:8585–8597.
- 636 [59] Erbs, D.G., Klein, S.A., Duffie, J.A., 1982. Estimation of the diffuse radiation fraction for hourly, daily and
637 monthly-average global radiation. *Sol. Energy* 28, 293–302. doi:10.1016/0038-092X(82)90302-4
- 638 [60] Reindl DT, Beckman WA, Duffie JA. Diffuse fraction correlations. *Solar Energy* 1990;45:1–7.
- 639 [61] Orgill JF, Hollands KGT. Correlation equation for hourly diffuse radiation on a horizontal surface. *Solar*
640 *Energy*,1977;19:357–9.
- 641 [62] Louche A, Notton G, Poggi P, Simonnot G. Correlations for direct normal and global horizontal irradiation
642 on French Mediterranean site. *Solar Energy* 1991;46:261–6.