

Article

The DeepWater Horizon Oil Slick: High Resolution Model Simulations of River Front Effects, Initialized and Verified by Satellite Observations

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1 **Abstract:** The effect of river fronts on oil slick transport has been demonstrated using high resolution
2 forcing models and a fully fledged oil drift model, OpenOil. The model system is used to simulate
3 the 2010 DeepWater Horizon oil spill. Metocean forcing data are taken from the GoM-HYCOM
4 1/50° ocean model with realistic river input and ECMWF global forecast products of wind and wave
5 parameters with 1/8° resolution. The simulations are initialized from satellite observations of the
6 surface oil patch. OpenOil includes most of the relevant processes, such as emulsification, evaporation,
7 wave entrainment, stranding and droplet formation. The model takes account of the actual oil type
8 and properties, using the ADIOS oil weathering database of NOAA. The effect of using a newly
9 developed parameterization for oil droplet size distribution is studied and compared to a traditional
10 algorithm. Although the algorithms provide different distributions for a single wave breaking event,
11 it is found that the net difference after long simulations is negligible, indicating that the outcome is
12 robust regarding the choice of parameterization. That indicates that the wave entrainment, vertical
13 mixing and re-surfacing mechanisms that are part of OpenOil are more important for determining the
14 final droplet size spectrum than the spectrum prescribed for individual wave breaking events. In both
15 cases, the size of the droplets controls how much oil is present at the surface and hence are subject
16 to wind and Stokes drift. The effect of removing river outflow in the ocean model is investigated in
17 order to showcase effects of river induced fronts on oil spreading. A consistent effect on the amount
18 and location of stranded oil is found, and considerable impact of river induced fronts is seen on the
19 location of the surface oil patch. During a case with large river outflow (May 20-27, 2010), the total
20 amount of stranded oil is reduced by about 50% in the simulation with no river input. The results
21 compare well with satellite observations of the surface oil patch.

22 **Keywords:** HYCOM; OpenDrift; OpenOil; Oil Spill; Modelling; Simulations; Satellite; Observations;
23 River Fronts; DeepWater Horizon

24 1. Introduction

25 The presence of both shelf and open sea dynamics makes the Northern Gulf of Mexico (NGoM) a
26 topographically and dynamically complex and interesting study area, in the presence of intense oil
27 exploration [1,2]. Interactions of the Mississippi River (MR) plume and the Loop Current (LC) system
28 were found to be important for the transport and fate of oil during the 2010 DeepWater Horizon (DWH)
29 incident [3,4].

30 According to the U.S. Energy Information Administration, more than 45% of total U.S. petroleum
31 refining capacity and 51% of total natural gas processing plant capacity are located along the Gulf coast

[5]. Oil leaks and accidents, such as the explosion on the DWH platform in 2010 (at 28.737°N, 88.366°W), have released significant quantities of hydrocarbons [6,7] in the sensitive marine environment around the MR Delta, and over the Louisiana TEXas (LATEX) and Mississippi Alabama FLoridA (MAFLA) shelves [3] (Fig. 2).

The present study is motivated by recent advances in the understanding of the role that salinity induced fronts and circulation areas dominated by river plume dynamics may play in the transport of hydrocarbons [3,8]. We use a comprehensive high resolution modeling system to demonstrate reliable estimates of hydrocarbon transport under the influence of all circulation driving mechanisms, with an emphasis on correctly accounting for fronts and circulation patterns due to river plume influence. Such fronts are density-driven, but also related to distinct circulation regimes that characterize river plume dynamics and may even control near-surface transport, subject to strong variability under the influence of additional shelf and deep sea flows. The river induced, buoyancy-driven flows include a westward coastal current along the LATEX shelf (“downstream” plume regime, ie. in the direction of Kelvin-wave propagation) and a northeastward flow toward the MAFLA shelf (“upstream”). The former is due to geostrophic balance between Coriolis and cross-shelf pressure gradient (e.g. [9], [10] and [11]) and is enhanced by downwelling-favorable winds; such conditions favor material entrainment and nearshore confinement. The latter is due to the balance between along-shelf pressure gradient and along-shelf acceleration [9]. This coastal flow has a tendency to turn offshore, enhanced by upwelling-favorable winds; such conditions promote material removal away from the coastal zone, while they restrict material from offshore sources to reach the coasts. The third important flow regime, the buoyancy-driven anticyclonic bulge, is often suppressed due to the Mississippi Delta proximity to a steep slope ([12], [13]); however, it does form under high discharge conditions [3]. These flow regimes induce well defined fronts that exhibit strong variability, depending on river discharge, winds and interaction with offshore flows [13]. These are relevant to pathways of any “materials”, such as river-borne (eg. nutrients) or from offshore sources potentially driven within plume influence (eg. hydrocarbons from neighboring exploration sites). River induced fronts and related circulation regimes are hard to monitor and are often not well represented in numerical models. In the Gulf of Mexico (GoM), this is a particularly difficult task, due to complexity in topographic controls and direct influence of large scale oceanic currents (namely the Loop Current, LC, and associated eddy field) on plume induced transport ([13], [3]). Related transport pathways may have great impact on coastal ecosystems in the vicinity of GoM areas with river influence (especially the Northern GoM), but also remote ecosystems that can be reached along the LC (such as the South Florida islands and reefs, eg. [14] and [15]). In the simulations presented herein, we employ a detailed representation of river plume dynamics, so the related circulation features and fronts are well described.

Many studies have dealt with simulations of the DWH spill [3,4,16–20], with focus on both subsurface [20] and surface [4] transport. North *et al.* [16] used a plume model to predict a stratification-dominated near field, in which small oil droplets detrained from the central plume containing faster rising large oil droplets and gas bubbles and became trapped by density stratification. They showed that simulated droplets with diameters between 10 and 50 μm formed a distinct subsurface plume, which was transported horizontally and remained in the subsurface for >1 month. In contrast, droplets with diameters >90 μm rose rapidly to the surface. Le Hénaff *et al.* [4] focused on oil transport on the water surface and found that the wind played a major role in advecting the oil to the northern GOM. Barker [19] conducted Monte Carlo simulations consisting of 500 individual oil trajectory scenarios using historical data of water currents and winds. The results by Barker [19] indicated that, in approximately 75% of the scenarios, oil would be transported out of the GOM by the Loop Current. This means that the actual trajectory of oil from the DWH falls in the 25% of scenarios.

Androulidakis *et al.* [8] carried out a field experiment deploying surface drifters at different times near the Taylor Energy Site which is located in the vicinity of the MR outflow region over the NGOM and near the DWH site (approximately at 28.938°N, 88.978°W, Fig. 2). This multi-platform observational experiment was conducted in April 2017 to investigate the main transport pathways

82 from the Taylor Site and toward the NGoM continental shelves and offshore, toward the Gulf interior.
83 Results indicated that the surface transport was determined by the MR plume extension over the
84 Taylor Energy Site and the river induced fronts in combination with local circulation, prevailing winds
85 and broader regional dynamics (LC system). The drifters deployed during the field experiment in
86 tandem with satellite data, drone imagery, wind measurements, and marine radar derived currents
87 and images described three major transport pathways, in agreement with the three major circulation
88 patterns of the MR plume [10,13,15,21].

89 The drifters deployed by Androulidakis *et al.* [8] followed the two prevailing coastal currents
90 associated with MR plume dynamics (downstream/upstream moving westward/eastward of the
91 Mississippi Delta) and an offshore pathway under the influence of basin-wide circulation. Near the
92 Taylor site, the existence of multiple river fronts influence the fate of oiled waters, preventing the
93 transport of hydrocarbon toward the delta like a natural boom barrier, trapping and directing the oil
94 either westward or eastward in agreement with Kourafalou and Androulidakis [3], who showed a
95 similar interaction during the DWH accident. *In situ* thermohaline measurements around the Taylor
96 Energy Site and across the river front showed that the MR plume near the Taylor Site was 5m to 10m
97 deep, while the clearer ocean water column was characterized by a 40 m upper-ocean homogeneous
98 layer, mainly controlled by temperature.

99 Much progress has been made after the DWH accident in understanding the GoM ecosystem, the
100 physical oceanography and its economic significance [1]. In this study, an open source Lagrangian
101 oil drift model, OpenOil, has been used to simulate the DWH oil spill evolution. OpenOil takes into
102 account major factors that influence the short term drift of surface oil slicks such as metocean forcing
103 (including Stokes drift), emulsification, evaporation, vertical entrainment and mixing. Simulations are
104 initiated from satellite observations and a point source at the sea floor. The initialization of simulations
105 from satellite observations is a relatively new feature in marine oil spill modelling [22]. The effect of
106 two different oil droplet size distribution on the horizontal drift and vertical mixing is discussed. This
107 study showcases how NGoM oil pathways are influenced by river plume dynamics and river induced
108 fronts. It is also investigated whether the use of realistic daily river discharge has a significant effect on
109 the simulated location of the Surface Oil Patch (SOP) and stranding of oil.

110 2. Materials and Methods

111 2.1. Shapefiles of surface oil patch

112 Shape files derived from satellite analysis of the DWH SOP can be accessed through the
113 NOAA-ERMA website [23]. In the present study, oil elements were seeded uniformly within the
114 region enveloping the thick and thin oil slicks with no distinction. The shapefiles were used here for
115 both initialization of the oil drift simulations and for verification of results.

116 2.2. Metocean forcing

117 In the cases presented here, the ocean circulation fields come from a data-assimilative,
118 high-resolution (1/50°, 1.8 km) configuration of the Hybrid Coordinate Ocean Model [24] in the Gulf
119 of Mexico (GoM), developed by the authors. This configuration, which we refer to as GoM-HYCOM
120 1/50°, uses daily river forcing and data assimilation. The HYCOM model has a flexible, hybrid vertical
121 coordinate system, in which the distribution of vertical layers is optimized: they are isopycnal in
122 stratified water columns, terrain-following sigma in regions with sharp topography, and isobaric in
123 the mixed layer and very shallow areas [25]. More information about the HYCOM model is available
124 in the model user's manual [24] and the references therein. The GoM-HYCOM 1/50° covers the entire
125 GoM and uses 32 vertical layers. The model configuration is similar to the one used by Le Hénaff
126 and Kourafalou [26], with the realistic river forcing parameterization developed by Schiller and
127 Kourafalou [10]. The river discharge data were obtained from the Army Corps of Engineers and the
128 U.S. Geological Survey. The model is initialized in October 2009 with fields from the operational

129 Global HYCOM (GLB-HYCOM) simulation run at the Naval Research Laboratory at the Stennis Space
130 Center (GLB-HYCOM expt_90.8, [24], and it is nested at the open boundaries with model fields from
131 the same simulation. The atmospheric forcing is based on the 3-hourly winds, thermal forcing and
132 precipitation forecast fields from the European Centre for Medium-Range Weather Forecasts [27],
133 with spatial resolution of 0.125° (see below). The model assimilates satellite observations of Sea
134 Surface Temperature and Sea Surface Height, and in situ observations of temperature and salinity from
135 buoys, cruises, surface drifters, Argo floats and XBT casts. More details about the model configuration
136 can be found in Le Hénaff and Kourafalou [26] and Androulidakis *et al.* [21]. For the present study,
137 altogether four simulations were performed, two for each period studied: one with the attributes
138 mentioned above, called *Reference* simulation, and one called *No river*, in which the salinity fronts have
139 been removed by shutting off the river discharge, setting precipitation to zero, and turning off the
140 assimilation of salinity profiles. This is a procedure called twin experiments also used by other authors
141 studying the effects of river fronts near MR [28]. All other forcing conditions (e.g. meteorological,
142 boundary) remained the same between the two experiments in order to investigate the impact of an
143 individual forcing mechanism (here the Mississippi buoyant discharge and the related density fronts)
144 on the circulation features and furthermore on the oil spill extensions during the DwH period. The
145 outputs from both simulations are available at the Gulf of Mexico Research Initiative Information and
146 Data Cooperative[29].

147 The ECMWF provides daily global forecasts at 0 and 12 UTC with 0.125° resolution. Recent model
148 upgrades have improved the overall performance of the forecasting system throughout the medium
149 range. Further details on model description and verification can be found e.g. in Ehard *et al.* [30],
150 Haiden *et al.* [31] and at [27]. Here, ECMWF daily forecast products were used as atmospheric upper
151 boundary conditions for the GoM-HYCOM 1/50 as well as for providing air temperature and wind
152 drag for the OpenOil simulations with a 3 hourly time step.

153 Wave properties were downloaded from the ECMWF third generation spectral WAve Model
154 global operational runs [32]. WAM is well known, see, e.g. Group [33] and Haiden *et al.* [31]. The WAM
155 model computes two-dimensional wave distribution, with 25 frequencies and 24 directions. From
156 the two-dimensional spectra, several parameters are computed, including significant wave height,
157 peak wave period, mean wave period, peak wave direction and mean wave direction. The wave
158 parameters are computed for total sea, and for wind sea and swell separately [31]. The operational
159 daily WAM forecasts used here are forced by the ECMWF atmospheric forecasts. WAM model output
160 with 0.125° horizontal resolution are downloaded from ECMWF with 12 hourly time step and used
161 here for estimating horizontal Stokes drift and vertical mixing of the oil with a 3 hourly time step using
162 linear interpolation.

163 Two periods are studied here: 20-27 May 2010 and 2-10 July 2010. According to the ECMWF
164 model, during the first period, wind speed varied between 0.1 and 7.2ms^{-1} and the significant wave
165 height varied between 0.1 and 1.2m . In the second period, the corresponding values were 5 to 12 ms^{-1}
166 and 0.1 to 3.2m . The various forcing data are summerized in Table 3.

167 2.3. The oil drift model OpenOil

168 OpenOil is part of the the OpenDrift trajectory modeling framework [34], developed at the
169 Norwegian Meteorological Institute and available as open source software from [35]. OpenOil has
170 been evaluated against drifter and oil slick observations in the North Sea [36,37]. Details of the element
171 tracking model are given in Dagestad *et al.* [34], and model physics that are specific to oil transport
172 and fate are documented in Röhrs *et al.* [36] and in the following.

173 OpenOil is an integrated oil drift model consisting of sub-models for specific physical processes
174 like wave entrainment of oil [38], vertical mixing due to oceanic turbulence [39], resurfacing of oil
175 due to buoyancy [40], and emulsification taking account for oil properties [41]. The resurfacing is
176 a function of oil density and droplet size following Stokes Law, and thereby the model physics are

177 very sensitive to the specification of the oil's droplet size. Figure 1 shows the sequence of operations
 178 involved in the OpenOil simulations.

179 Oil droplet size distribution

180 Several algorithms are implemented to describe the oil's droplet size distribution, based on
 181 previous published parameterizations. The first option is based on the work of Delvigne and Sweeney
 182 [42] (DS88), manifesting a power-law droplet size number distribution as a function of droplet size,
 183 with an exponent of -2.3, i.e. there are many more small droplets than large droplets. Transferring this
 184 to a volume size distribution, as needed for practical oil spill simulation that follows the mass of the
 185 oil spill, the exponent becomes 0.7, i.e. there is more volume in the few large droplets than in the many
 186 small droplets. The typical droplet sizes range from $1\text{ }\mu\text{m}$ to 1 mm .

187 A second option to describe the droplet size distribution is based on Li *et al.* [43] (Li17), which
 188 takes the oil viscosity and the oil-water interfacial tension into account. This parameterization describes
 189 a log-normal law for the number size distribution, and the resulting volume size distribution exhibits
 190 a peak at an intermediate droplet size of about $100\text{ }\mu\text{m}$, depending on oil type and environmental
 191 conditions. Similar types of droplet size distribution have been developed and observed, confirming
 192 that there is a maximum in oil volume at a particular droplet size [44,45].

193 Following Li *et al.* [43], the volume (V) droplet size spectrum is described by the median droplet
 194 diameter, D_{50}^V , as

$$D_{50}^V = d_o r (1 + 10\text{Oh})^p \cdot \text{We}^q \quad (1)$$

195 with the empirical coefficient $r = 1.791$ and the exponents $p = 0.460$ and $q = -0.518$. The PDF for
 196 the droplet size distribution follows a log-normal distribution around the medium diameter with a
 197 logarithmic base-10 standard deviation of $s = 0.38$ (Eq. 16 in Röhrs *et al.* [46]).

198 The Weber number, We, is a dimensionless number describing the relative importance of inertial
 forces and oil-water interfacial tension. It is a function of the sea water density, ρ_w , the significant
 199 wave height, H_s , and the oil-water interfacial tension, σ_{o-w} , and is given by

$$\text{We} = \frac{\rho_w g H_s d_o}{\sigma_{o-w}}, \quad (2)$$

200 where g is the acceleration of gravity and $d_o = 4\sqrt{\frac{\sigma_{o-w}}{g(\rho_w - \rho_o)}}$ is the Rayleigh-Taylor instability maximum
 201 diameter.

202 The Ohnesorge number, Oh, is a dimensionless number describing the ratio of viscous forces to
 203 inertial and surface tension forces. It is a function of the dynamic oil viscosity, μ_o , oil density, ρ_o , and
 204 oil-water interfacial tension:

$$\text{Oh} = \frac{\mu_o}{\sqrt{(\rho_o \sigma_{o-w} d_o)}}. \quad (3)$$

205 The volume size distribution, following [42], is given by

$$V(d) = d^{-0.7}, d_{min} < d < d_{max} \quad (4)$$

206 where d is the droplet diameter. Minimum and maximum droplet radii are set to $10\text{e-}6$ and $10\text{e-}3$
 207 meters, respectively. The exponent of -0.7 in the volume size distribution corresponds to an exponent
 208 in the number size distribution of 2.3 [40].

209 Droplet sizes are assigned to oil particles each time an element is submerged by breaking waves,
 210 following the wave entrainment algorithm of [38]. The implementation of this algorithm in OpenOil is
 211 described with full detail in [36]. The droplet sizes for individual particles are drawn from a random
 212 distribution according to the chosen size distribution. The size distributions represent conditions for a

211 stochastic wave entrainment event, representing equilibrium conditions during a model time step. It
212 is noted that the overall size distribution of all submerged oil in the simulation is further subject to
213 changes, as weather conditions, the oil's emulsification rate change and oil droplets of various sizes are
214 subject to various resurfacing time scales. Resurfaced elements are considered to be part of a surface
215 slick, and are assigned a new droplet size distribution once they are re-entrained. Oil droplets at the
216 sea surface (slick) are not considered to have a radius.

217 **Droplet size distribution during deep blowouts**

218 For oil elements released at the seafloor (wellhead), a simplistic and pragmatic approach of
219 prescribing random radii in the range 0.5 mm to 5 mm was used, as suggested by Johansen [47].

220 **Horizontal transport**

221 With regard to horizontal drift, three processes are considered: Any element, whether submerged
222 or at the surface, drifts along with the ocean current. Elements are further subject to Stokes drift
223 corresponding to their actual depth. Surface Stokes drift is normally obtained from a wave model, and
224 its decline with depth is calculated as described in Breivik *et al.* [48]. Oil elements at the ocean surface
225 are additionally moved with a factor of 2% of the wind. Together with the Stokes drift (typically 1.5%
226 of the wind at the surface), this sums up to the commonly found empirical value of 3.5% of the wind
227 speed [49]. The magnitude of the wind drift factor is discussed in Jones *et al.* [37] who stated that a 2%
228 wind drift factor was required in OpenOil to reproduce their observations of a SOP in the North Sea.
229 In essence, this is believed to be a compensation factor for the inability of any ocean model to represent
230 the strong shear current in the upper few centimeters/decimeters of the ocean, and not surface oil
231 actually moving relative to the water.

232 The three horizontal drift components may lead to a very strong gradient of drift magnitude and
233 direction in the upper few meters of the ocean. For this reason, it is also of critical importance to have
234 a good description of the vertical oil transport processes.

235 **Vertical transport**

236 Oil elements at the surface, regarded as being in the state of an oil slick, may be entrained into
237 the ocean by breaking waves. The entrainment of oil droplets depends on both the wind and wave
238 (breaking) conditions, but also on the oil properties, such as viscosity, density and oil-water interfacial
239 tension. The buoyancy of droplets is calculated according to empirical relationships and the Stokes law
240 following Tkalich and Chan [40], dependent on ocean stratification based on temperature and salinity
241 from the ocean model, and the viscosities and densities of oil and water.

242 In addition to the wave induced entrainment, the oil elements are also subject to vertical turbulence
243 throughout the water column, described using a random-walk scheme based on the turbulent eddy
244 diffusivity wind speed parameterization from Sundby [50].

245 **Weathering**

246 In order to calculate weathering of the oil, OpenOil interfaces with the open source ADIOS
247 oil library [51], developed by NOAA and [41]. In addition to state-of-the-art parametrization of
248 weathering processes such as evaporation, emulsification and dispersion, this software contains a
249 database of measured properties of almost 1000 oil types from around the world. As oils from different
250 sources or wells have vastly different properties, such a database is of vital importance for accurate
251 results. The ADIOS oil library is also used by the NOAA oil drift model [52].

252 The weathering algorithms describes evaporation and emulsification rate of oil, i.e. the water
253 content. The emulsification and evaporation greatly affect oil density, viscosity and oil-water interfacial
254 tension, and thereby the droplet size distribution through Eqs. 1-3. OpenOil takes into consideration
255 weathering processes that are dominating in the initial oil spill period of 2-3 days. Long-term

256 weathering processes such as sedimentation and microbial degradation are not considered in this
257 study.

258 3. Results

259 A first set of simulations is carried out to investigate the effect of oil droplet size distribution.
260 A second set focuses on the effect of river induced fronts. According to Crone and Tolstoy [6], the
261 average flow rate from the oil well between 22 April and 3 June 2010 was estimated to $0.1 \text{ m}^3 \text{ sec}^{-1}$,
262 assuming a liquid oil fraction of 0.4. Gas and highly volatile compounds are not considered here. After
263 the riser was removed and until the leak was sealed on 15 July, the flow rate increased to $0.12 \text{ m}^3 \text{ sec}^{-1}$,
264 corresponding to $10,368 \text{ m}^3 \text{ day}^{-1}$. The amount of oil present at the sea surface at the initial time of the
265 simulation was estimated based on the simulated removal rate. The residence time of oil at the sea
266 surface depends heavily on oil properties as well as environmental conditions such as temperature of
267 ocean and air, wind and waves, as described above. According to our mass balance calculations for the
268 DWH spill, using the Light Louisiana Sweet oil type from the NOAA oil library and environmental
269 conditions as described above, it seems reasonable to assume that 80% of the oil mass is removed from
270 the surface after 10 days. This is within the range in our simulations that is typically 60 to 95% (see
271 examples of mass balance plots further down). While [53] assumed a constant removal rate of surface
272 oil 20% per day, the removal rate in the present simulation and in reality will vary with wind and
273 wave conditions. The simulations for May 2010 are initialized by seeding 48730 elements in a polygon
274 obtained from NOAA shapefiles. Each element represents initially 1 m^3 oil. A continuous point source
275 at the sea floor seeds an additional 8460 elements (8460 m^3) per day during the simulation. After
276 June 3rd these numbers are increased by 20% to $10368 \text{ m}^3 \text{ day}^{-1}$. The oil elements released at the
277 surface are assigned droplet radii at each entrainment incident, according to the parameterisation of
278 size distributions from respectively DS88 or Li17, see [36] for details. Oil elements at the sea surface
279 (slick) are not considered to have a radius.

280 Around 20-25 May 2010 there was a significant outflow of the Mississippi River (Fig. 3) and part
281 of the SOP was entrained along the LC resulting in a formation popularly referred to as the "tiger
282 tail" (Fig. 4). The OpenOil simulation shown on top in Fig. 4 is carried out using the classical DS88
283 oil droplet size distribution [42]. Fig. 4 lower panel shows the results from repeating this simulation
284 using the new Li17 formulation. In Fig. 5 the mass balance during seven days for the Li17 simulation
285 is shown. There is virtually no difference between DS88 and Li17 and only Li17 is shown here. Both
286 formulations result in about the same fraction of oil at the surface (about 50%) after 7 days, with
287 moderate wind speeds of up to 7.2 ms^{-1} . The light compounds evaporate fast after release hence the
288 more heavy compounds are tracked here ("dead oil"). Patches of thick oil (where the elements retain
289 nearly 100% of their mass) are visible over a larger area (Fig. 4). Larger oil droplets will rise faster to
290 the surface [16,36], and DS88 provides a much higher fraction of large droplet after one hour compared
291 to Li17 (Fig. 6, upper panels). However, it turns out that the DS88 and Li17 provide similar volume
292 distributions after a 24 hr test simulation (Fig. 6, lower panels). Still, the peak in the distribution is
293 at around $100 \mu\text{m}$ for both formulations, and rapid rising to the surface can be expected according to
294 North *et al.* [16]. Due to the small difference between DS88 and Li17, the DS88 simulation results in
295 just marginally more oil stranded after seven days (13.2 vs 12.8 %), particularly west of the Mississippi
296 Delta (Figs. 4). A higher fraction of oil at the surface provides more efficient transport by wind and
297 waves towards the shore and larger likelihood of stranding. For both simulations, the oil droplets
298 quickly loose 40-50% of their mass, mostly through evaporation.

299 Fig. 7 shows the geographical distribution of droplet diameters at the surface. It is apparent that
300 smaller droplets are present outside the MAFLA shelf, probably because the oil droplets have been
301 more subject to wind and hence wave action and natural dispersion in the last 24 hours (Fig. 8).

302 The Li17 formulation is applied for two types of simulations: *Reference* (all forcing data) and
303 No river (all forcing data except for river runoff and precipitation). Two periods were chosen for
304 these experiments: a high river discharge period with variable winds (20-27 May) and a relatively

305 lower discharge period with persistent westward winds (2-10 July). The purpose is to investigate the
 306 effect of the salinity fronts by using the ocean circulation from the *No river* simulation, in which the
 307 precipitation and river discharge are turned off, while atmospheric and wave forcings are kept the
 308 same. In Figs. 9, 10, 11 and 12 it appears how the inclusion of river discharge in the ocean forcing
 309 can have opposite (and sometimes counter-intuitive) effects on oil transport. Table 1 summarizes the
 310 difference of the *Reference* and *No River* simulations for May 20-27, whilst Table 2 shows corresponding
 311 values for the 2-10 July simulations.

312 Kourafalou and Androulidakis [3], based on high-resolution ocean simulations over the NGOM,
 313 showed that the MR discharge peak around 20-30 May led to the formation of downstream (westward)
 314 and upstream (northeastward) plume areas that acted as a conduit for guiding oil toward the LATEX
 315 shelf and away from the MAFLA shelf, respectively (Fig. 10). In the 20-27 May period the river
 316 plume currents created a strong "bulge" that tended to turn waters clockwise around the Delta, with
 317 some waters moving westward. In addition, the offshore GoM circulation (Loop Current and eddies)
 318 removed the riverine waters offshore, toward the GoM interior, forming the so-called "tiger tail"
 319 pathway. The removal of the MR input in the *No River* experiment during this period (Fig. 9) allowed
 320 the spreading of oil toward the western MAFLA shelf due to the absence of this anticyclonic bulge;
 321 the stranded oil along the MAFLA coasts is more apparent in the *No River* case in comparison to
 322 the *Reference* experiment. In contrast, the amount of stranded oil is less along the western coasts
 323 (90W-91W) in the *No River* experiment due to the absence of the downstream MR plume pathway
 324 (Table 1). Moreover, the "tiger tail" signature is weaker in the *No River* simulation. Kourafalou and
 325 Androulidakis [3] showed that the strength of this MR offshore jet could have been an important factor
 326 in forming the "tiger tail" oil distribution pattern as also confirmed from satellite and drifter data [54].

327 The second simulation period (2-10 July) was right after another high discharge period (although
 328 not as high as in May, 3), promoting again a buoyancy-driven downstream current. This tendency
 329 was supported by downwelling-favorable winds [3], resulting in a clear westward transport of both
 330 low-salinity and oil containing waters, along a narrow band (of similar width) close to the LATEX
 331 coast and surrounding the Mississippi Delta; extensive coastal areas of stranded oil are apparent
 332 along the western coasts in the *Reference* experiment (Figs. 11 and 12). The removal of MR input
 333 (*No River* experiment) led to weaker downstream currents both close to the Delta (89.5W) and along
 334 the western coasts (west of 90W) and thus less stranded oil over the same region and more in the
 335 MR delta region (Table 2). The anticyclonic bulge, common in strong discharges and source of the
 336 downstream current [9], is completely absent in the *No River* experiment. As a result, more stranded oil
 337 was present closer to the Delta, inside Louisiana Bight in the *No River* case 2. It seems like the absence
 338 of the anticyclonic bulge that was able to lead surface oiled waters directly west of the Louisiana Bight
 339 allowed the accumulation of oil very close to the Delta. In contrast, smaller differences between the
 340 two experiments are detected over the MAFLA region due to the weaker upstream currents during
 341 early July [3]. During this period of slightly higher wind speed, most elements appear to have lost
 342 70-90% of their mass due to natural dispersion by waves and evaporation shortly after release (Fig. 11).

343 3.1. Figures, Tables and Schemes

Table 1. Percentages of stranded oil elements for the May 20-27 simulations.

	<i>Reference</i> simulation	<i>No River</i> simulation
West of MR Delta	6.9	0.5
MR Delta area	4.0	3.4
East of MR Delta	1.2	2.1
Total	12.1	6.0

Table 2. Percentages of stranded oil elements for the July 2-10 simulations.

	Reference simulation	No River simulation
West of MR Delta	10.6	7.8
MR Delta area	16.7	27.2
East of MR Delta	20.8	20.1
Total	48.1	55.1

Table 3. Summary of Metocean forcing used in the simulation

Model	Parameters	Resolution:	Horizontal	Vertical	Temporal
Model					
NOAA shape files	initial location				
GOM Hycom	horizontal current	1/50°	32 layers	3 hours	
ECMWF atmospheric model	wind velocity, air temperature	1/8°	surface	3 hours	
ECMWF wave model	Stokes drift, wave height and period	1/8°	surface	12 hours	

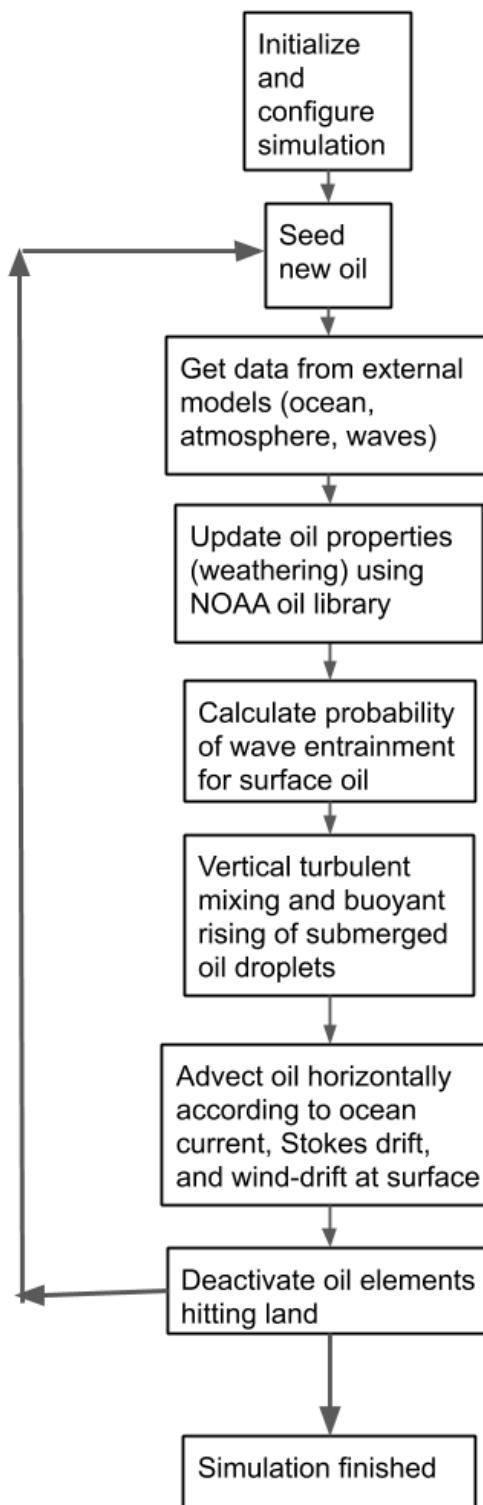


Figure 1. Flow chart showing the sequence of operations involved in the OpenOil simulations.

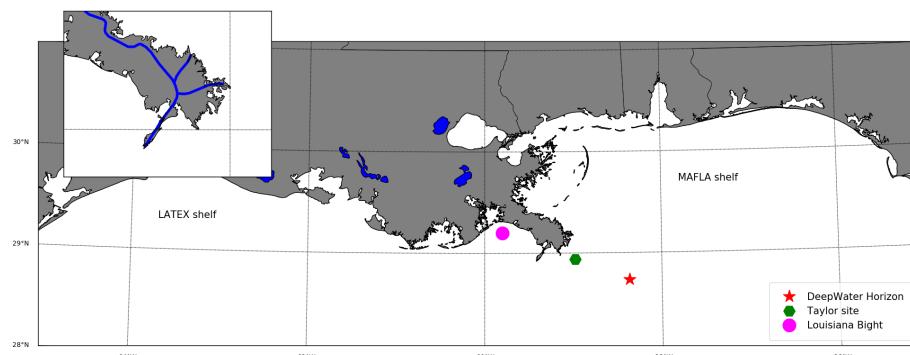


Figure 2. Map of the Northern Gulf of Mexico showing the geographical locations mentioned in the text. The map insert shows the Mississippi River (MR) delta in blue with the three major river passes that release MR water into the Gulf.

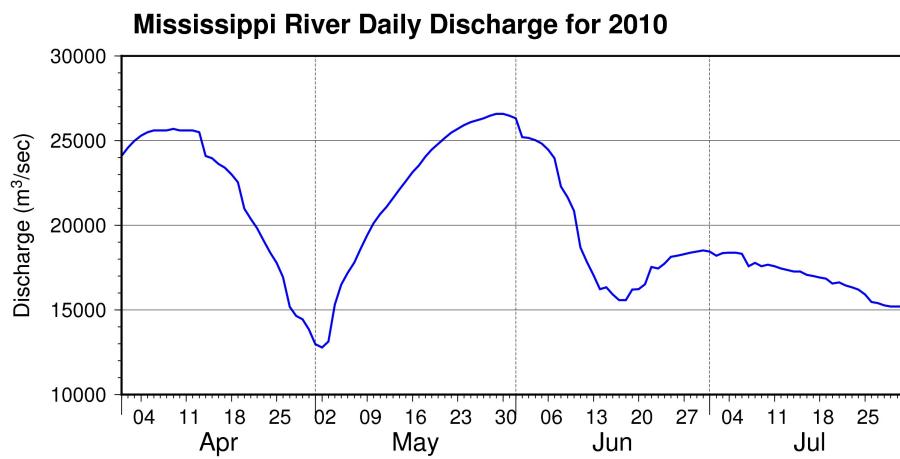


Figure 3. Discharge from Mississippi River in the Northern Gulf of Mexico during late spring and early summer 2010. Data kindly provided by U.S. Army Corps of Engineers.

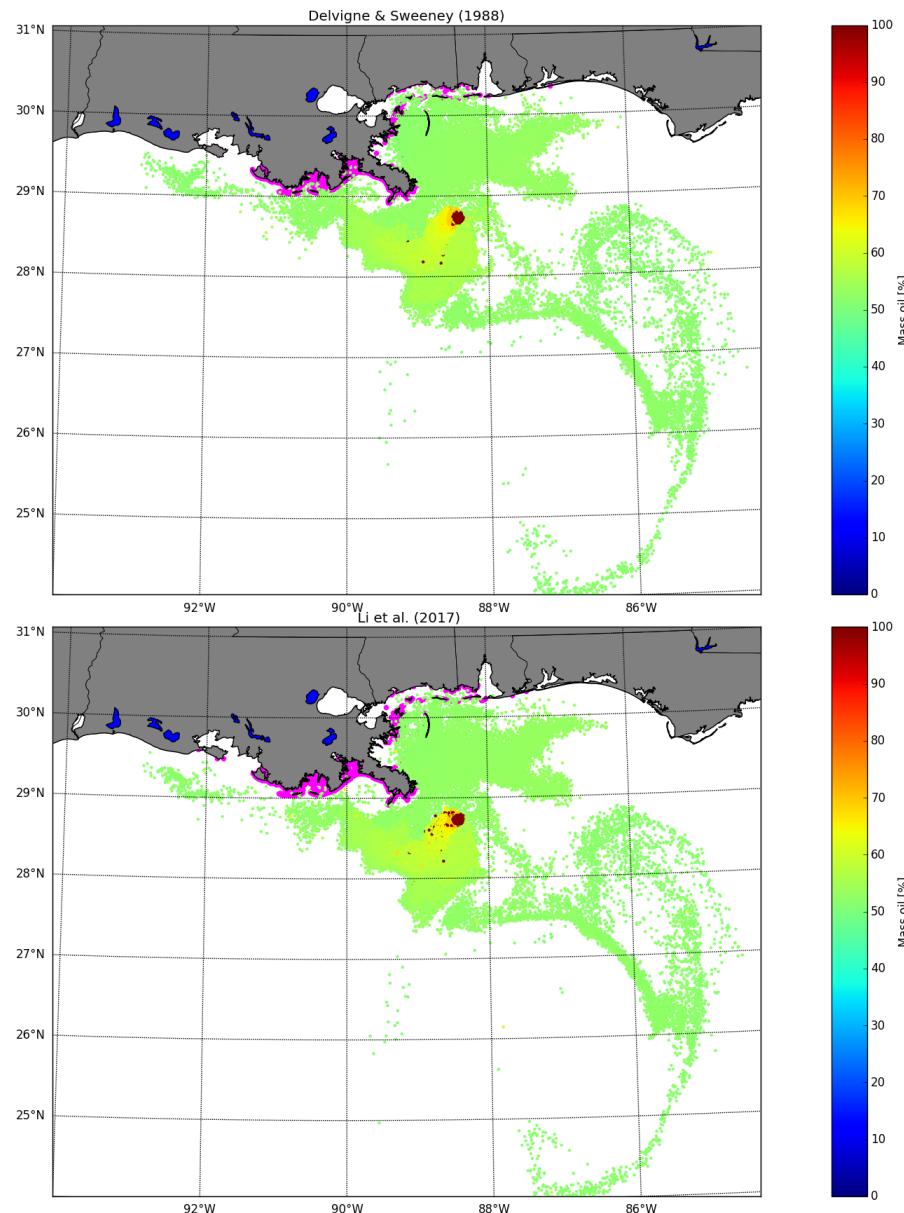


Figure 4. OpenOil simulation for 20-27 May 2010, using the Delvigne and Sweeney [42] oil droplet size distribution (upper - 13.2% stranded oil), and the Li *et al.* [38] distribution (lower - 12.8% stranded oil). Only surface oil elements are shown. Patch colors are the fraction of mass left in the elements. Magenta areas indicate stranded oil.

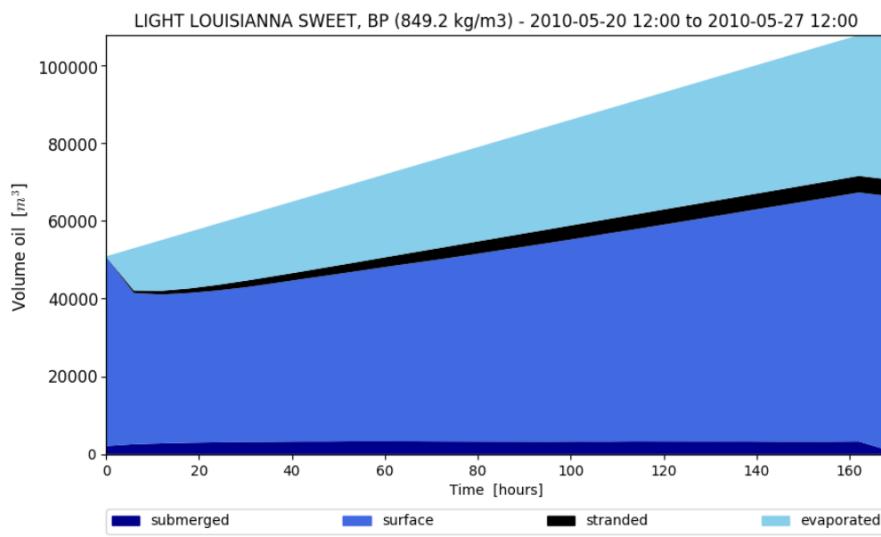


Figure 5. Mass balance of oil in the OpenOil simulation shown in Fig. 2. The Li *et al.* [38] oil droplet size distribution is used.

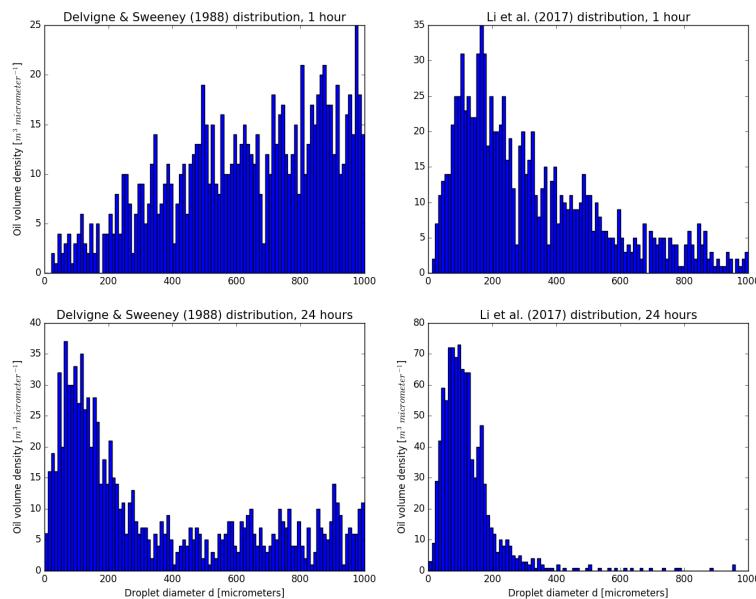


Figure 6. Oil droplet volume histogram for 1000 elements after 1 hours using the *Light Louisiana Sweet*, BP oil type in OpenOil during 8 ms^{-1} wind. The initial condition is a uniform distribution. Delvigne and Sweeney [42] formulation at top left and [38] at top right. Bottom panels show corresponding distributions after 24 hours.

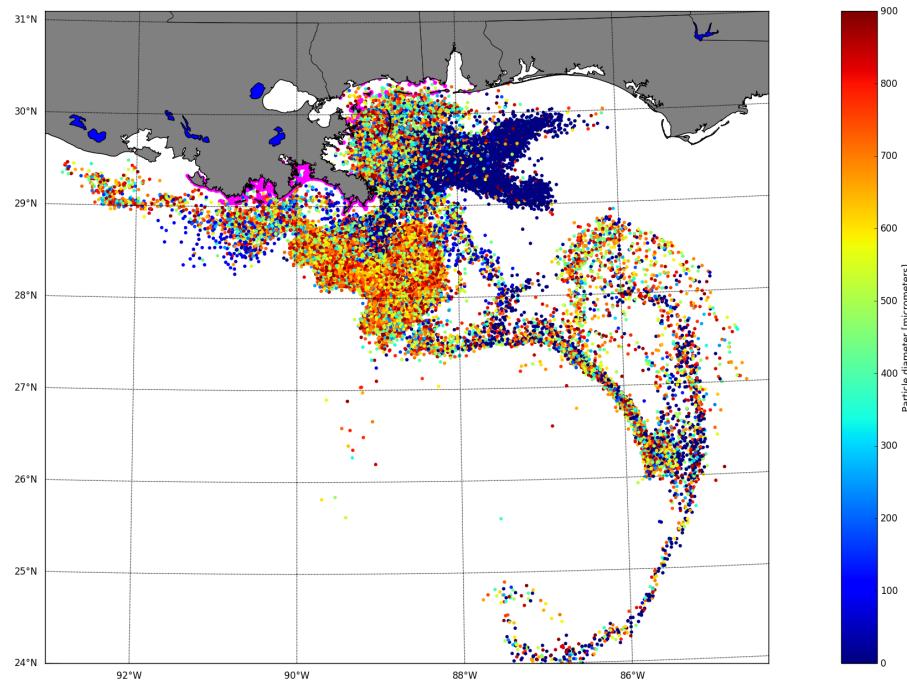


Figure 7. End condition of the *Reference* simulation 20–27 May 2010, showing all active elements (at surface and submerged). Same simulation as in Fig. 4 lower panel. Stranded oil is shown in magenta. The color scale shows diameter of the oil droplets.

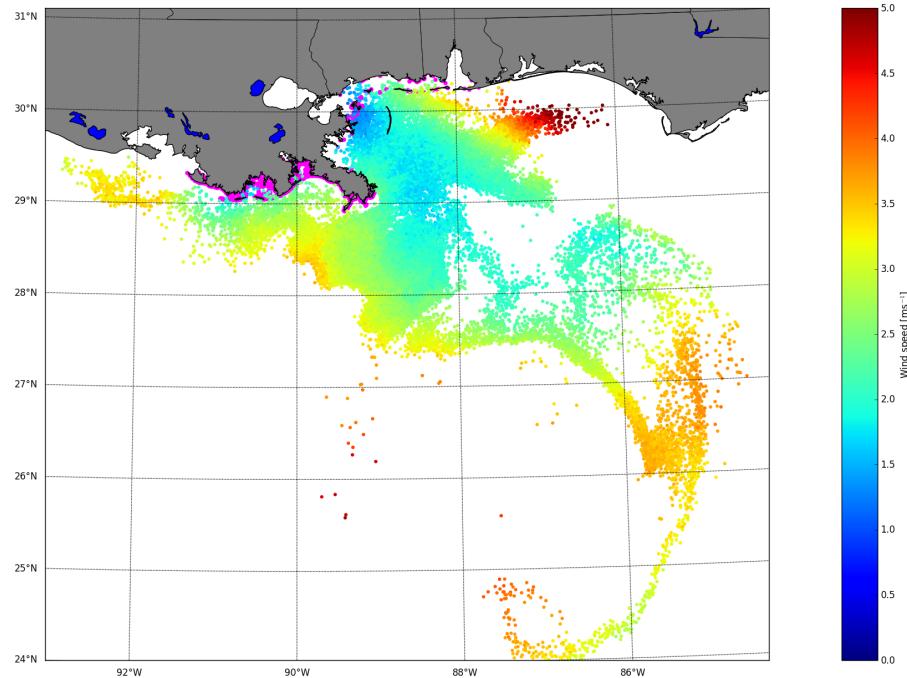


Figure 8. Same as Fig. 7, but the color scale shows the average wind speed experienced by the element during the last 12 hours.

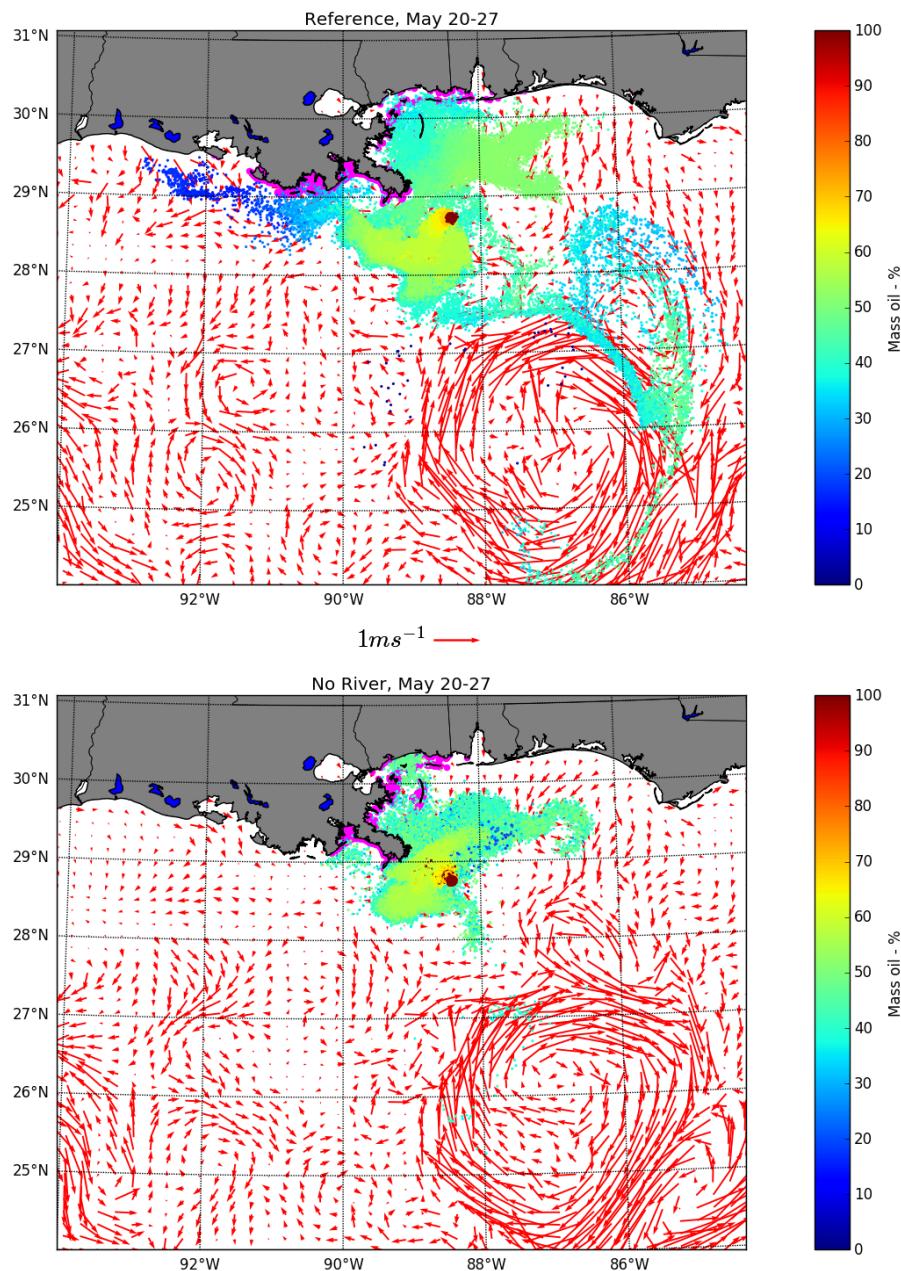


Figure 9. End condition of the simulation 20-27 May 2010, showing active (at surface) and stranded oil elements. Stranded oil is shown in magenta. The color scale indicate how much mass is left in each element. Red arrows are the GoM-HYCOM 1/50 forcing surface currents at the last time step of the simulation (every 5th data point shown). *Reference* simulation at top (12.1% stranded oil), and *no river* simulation below (6 % stranded oil). See also Table 1.

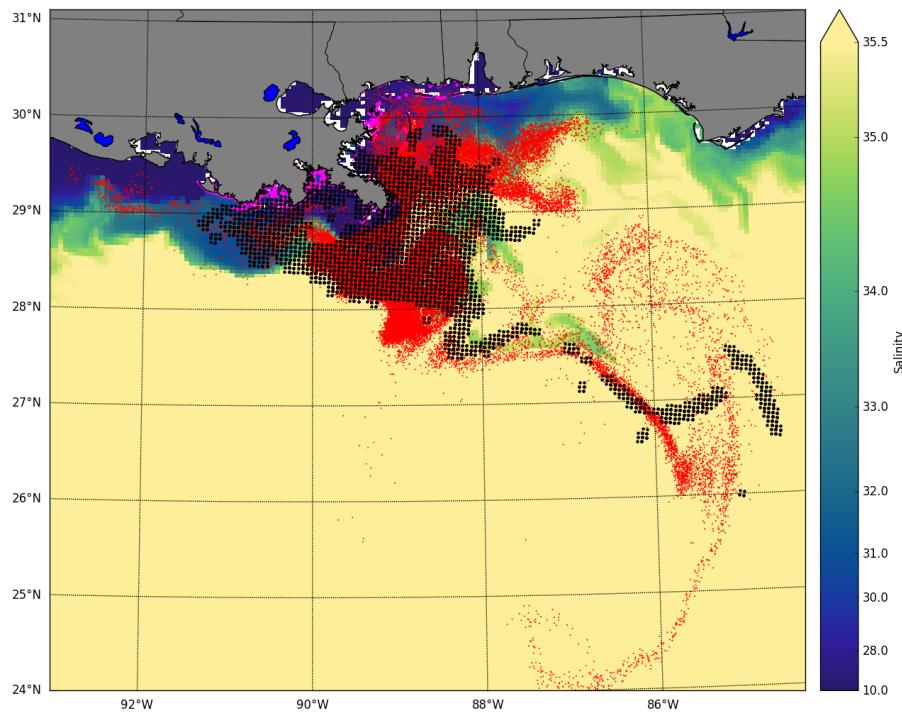


Figure 10. End condition of the *Reference* simulation 20–27 May 2010 (Same simulation as in Fig. 9, upper panel), showing active elements at surface as red dots and the corresponding observed surface oil patch (NOAA shape file) as black dots. Modeled stranded oil is shown in magenta. The color scale shows sea surface salinity in the forcing data.

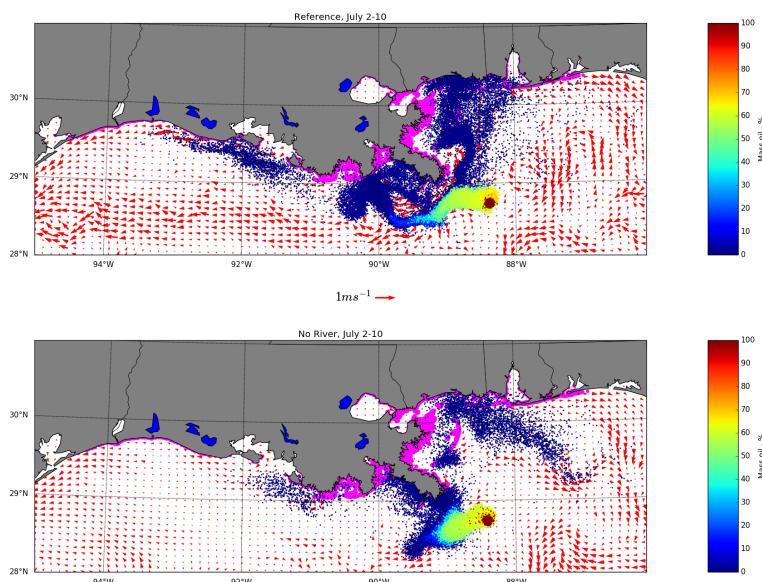


Figure 11. End condition of the simulation 2–10 July 2010, showing active (at surface) and stranded oil elements. Stranded oil is shown in magenta. The color scale indicate how much mass is left in each element. Red arrows are the GoM-HYCOM 1/50 surface currents at the last time step of the simulation (every 2nd data point shown). *Reference* simulation at top (48.1% stranded oil), and *no river* simulation below (55.1% stranded oil). See also Table 2.

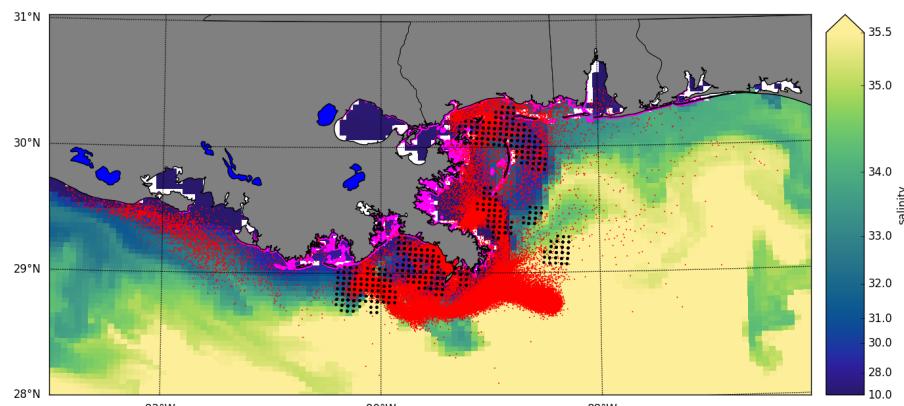


Figure 12. End condition of the *Reference* simulation 2-10 July 2010 (Same simulation as in Fig. 11, upper panel), showing active elements at surface as red dots and the corresponding observed surface oil patch (NOAA shape file) as black dots. Stranded oil is shown in magenta. The color scale shows sea surface salinity in the forcing data.

344 4. Discussion

345 Several simulations of the DWH oil spill have been carried out with high resolution forcing data
346 and a Lagrangian oil spill model. Simulations were initialized from satellite observations of the SOP,
347 and a continuous point source with a realistic spill rate at the sea floor.

348 Our results indicate that the two different formulations for oil droplet size distribution give similar
349 results for both vertical and horizontal distribution of the oil, when wind speeds are typically 5-12
350 ms^{-1} and breaking waves can be expected Fig. 4. Both formulations of the oil droplet size distribution
351 result in the characteristic "tiger tail" shape of the SOP for the period 20-27 May 2010, and significant
352 stranding in the delta west of the Mississippi River mouth in line with the observed SOP [3].

353 Both droplet size formulations that are used here (DS88 and Li17) result in similar size distributions
354 after some time of simulation, as seen in Fig. 6. Li17 prescribes a maximum droplet diameter in the
355 volume distribution as seen in Fig. 6, which is due to two regimes in the size distribution where small
356 droplets are limited by viscous effects, and larger droplets by oil-water inter-facial tension [44]. This
357 causes a peak in the volume distribution as seen in laboratory experiments with repeated mixing
358 and wave breaking from the surface. The DS88 distribution does not prescribe such a maximum,
359 using a power-law that increases towards larger droplets in the volume distribution. However, the
360 time-integrated simulations in OpenOil still produce a maximum in the droplet size distribution.
361 The reason for this is the repeated wave breaking at the surface, which is more pertinent to large
362 droplets that quickly rise to the surface. Hence, the description of buoyancy driven resurfacing and
363 wave breaking in the oil spill model, together with the DS88 droplet size spectrum for individual
364 wave breaking events, produces similar results to a more advanced droplet size distribution that
365 explicitly prescribes a maximum in the volume size distribution. In conclusion, two very different
366 droplet distributions give very similar distributions after some time. That indicates that the wave
367 entrainment, vertical mixing and re-surfacing mechanisms that are part of OpenOil are more important
368 for determining the final droplet size spectrum than the prescribed spectrum for individual wave
369 breaking events.

370 A realistic description of droplet formation is required to describe the effects of an oil spill on
371 the environment [16,36,53]. Fig. 7 shows that the oil spill transport during the DWH spill favors a
372 transport of small droplets towards the northeast, while larger oil droplets follow the paths towards
373 southwest and southeast. As a result of their low buoyancy and turbulent mixing, smaller droplets
374 are mixed into deeper parts of the ocean and subject to ocean currents at depth [36]. Larger droplets
375 experience stronger buoyancy and are subject to surface currents or return to the surface slick. As
376 wind and waves only affect the near-surface drift, the part of the oil slick that forms large droplets is
377 quickly separated from the small droplets which retain at larger depths North *et al.* [16]. This will also
378 impact the effect of the spill on the ecosystem: the parts of the oil spill at the surface is more hazardous
379 to birds and the beach communities, while the small, submerged parts will have a substantially larger
380 surface area to interact with water, fish and plankton [55,56].

381 Next, the effect of realistic river discharge on the simulations is studied. One might expect that
382 removing the river discharge would always bring the oil nearer to the shore, but interactions are
383 complex. The *No River* simulation for May 20-27 showed more stranding oil, in particular close to the
384 Louisiana Bight, but less stranding oil further downstream, along the LATEX shelf. The removal of the
385 MR input reduced the downstream currents that were responsible for the westward transport of oiled
386 waters along the LATEX shelf. The MR plume and the accompanying river fronts were responsible to
387 either entrap oil close to the coasts (e.g. LATEX shelf) or keep oiled waters offshore (e.g. MAFLA shelf)
388 due to the formation of upstream currents (Fig. 10). These results are in line with [3] and the NOAA
389 SOP observations used here and shown in Fig. 9. It is also obvious that, in the *Reference* simulation, the
390 oil elements are guided by the river fronts and they are carried further away from the coast, pushed
391 into the LC south of 28°N and E of 88.5°W Fig. (10).

392 The second simulation period (2-10 July) was right after a second high discharge period, promoting
393 again a buoyancy-driven downstream current. This tendency is supported by downwelling-favorable

394 winds, resulting in a clear westward transport of both low-salinity and oil-containing waters, along a
395 narrow band close to the LATEX coast and surrounding the Mississippi Delta; extensive coastal areas
396 of stranded oil are apparent along the western coasts in the *Reference* experiment. The removal of MR
397 input (*No river* experiment) led to weaker downstream currents both close to the delta (89.5W) and
398 along the western coasts (west of 90W) and thus less stranded oil over the same region.

399 The simulations presented here were initiated by seeding oil elements evenly in a polygon defined
400 by NOAA satellite products [23], in addition to a continuous point source at the sea floor. The next
401 possible step is to initiate the simulations from satellite products which contain information about oil
402 film thickness in addition to area, and hence also quantify the amount of oil at the surface.

403 To the best of our knowledge, this is the first time the importance of the effect of river fronts on
404 oil slick transport in the Gulf of Mexico has been demonstrated using high resolution ocean forcing
405 and a fully fledged oil drift model. It is also a new observation that two very different oil droplet size
406 distributions give similar net results compared to observed surface oil slicks.

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424 Abbreviations

425 The following abbreviations are used in this manuscript:

426 (N)GoM	(Northern) Gulf of Mexico
DWH	DeepWater Horizon
HYCOM	HYbrid Coordinate Ocean Model
ECMWF	European Center for Medium-Range Weather Forecasts
ADIOS	Automated Data Inquiry for Oil Spills
427 MR	Mississippi River
LC	Loop Current
LATEX	LouisianA TEXas shelf
MAFLA	Mississippi Alabama FLoridA shelf
NOAA	National Oceanic and Atmospheric Administration
USGS	US Geological Survey

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