


Article

Evapotranspiration Estimation with UAVs in Agriculture: A Review

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Abstract: Estimating evapotranspiration (ET) has been one of the most important research in agriculture recently because of water scarcity, growing population, and climate change. ET is the sum of evaporation from the soil and transpiration from the crops to the atmosphere. The accurate estimation and mapping of ET are necessary for crop water management. Traditionally, people use weighing lysimeters, Bowen ratio, eddy covariance and many other methods to estimate ET. However, these ET methods are points or location-specific measurements and cannot be extended to a large scale of ET estimation. With the advent of satellites technology, remote sensing images can provide spatially distributed measurements. The satellites multispectral images spatial resolution, however, is in the range of meters, which is often not enough for crops with clumped canopy structure such as trees and vines. And, the timing or frequency of satellites overpass is not always enough to meet the research or water management needs. The Unmanned Aerial Vehicles (UAVs), commonly referred to as drones, can help solve these spatial and temporal challenges. Lightweight cameras and sensors can be mounted on drones and take high-resolution images on a large scale of field. Compared with satellites images, the spatial resolution of UAVs' images can be as high as 1 cm per pixel. And, people can fly a drone at any time if the weather condition is good. Cloud cover is less of a concern than satellite remote sensing. Both temporal and spatial resolution is highly improved by drones. In this paper, a review of different UAVs based approaches of ET estimations are presented. Different modified models used by UAVs, such as Mapping Evapotranspiration at high Resolution with Internalized Calibration (METRIC), Two-source energy balance (TSEB) model, etc, are also discussed.

Keywords: Clumped canopy; evapotranspiration; Unmanned aerial vehicles; METRIC; remote sensing;

1. Introduction

Evapotranspiration (ET) estimation is important in precision agriculture water management. ET is known as the main outgoing water flux from the surface on the earth [1]. Mapping the ET temporally and spatially can be useful for evaluating soil moisture [2], drought monitoring [3], assessing crop water stress [4], etc. It is important to accurately quantify the ET in order to get a better understanding of crop growth. Estimating the ET accurately can also benefit the water resources management and weather forecast [5]. ET is a combination of two separate processes, the evaporation and the transpiration. Evaporation is the process whereby liquid water is converted to water vapour [6]. Then, the water vapour removes from the evaporating surface. Transpiration is the process of the vaporization of liquid water contained in plant tissues and the vapour removal to the atmosphere [6]. The current theory for transpiration is by the following three steps. First, the conversion of liquid phase water to vapor water, the side effect is to cause canopy cooling from latent heat exchange thus canopy temperature can be used as an indicator of ET. Second, diffusion of water molecules from

inside plant stomata on the leaves to the surrounding atmosphere. Third, atmospheric air mixing by convection or diffusion can transport vapor near the plant surfaces to the upper atmosphere or off site away from the plant canopy. Usually, evaporation and transpiration occur simultaneously.

To estimate ET, there are direct and indirect methods. For direct methods, there are lysimeters [7] and water balance methods [8]

$$WB = P + I - D - R - S = ET, \quad (1)$$

where WB (mm day^{-1}) is water balance and P (mm day^{-1}) means precipitation. I (mm day^{-1}) is irrigation, D (mm day^{-1}) means drainage, R (mm day^{-1}) is runoff. S (mm day^{-1}) is the soil moisture storage. ET (mm day^{-1}) is the evapotranspiration of the plant. However, these ET methods are usually point or location-specific measurements and cannot be extended to a large scale of ET estimation because of the heterogeneity of the land surface. And, the hydrologic processes can be very complex, too. Researchers sometimes also need to get variable land surface measurements and parameters [9], which is unlikely to get by these traditional methods. The experiment equipment is also very expensive, such as lysimeter, which is only available for a small group of people.

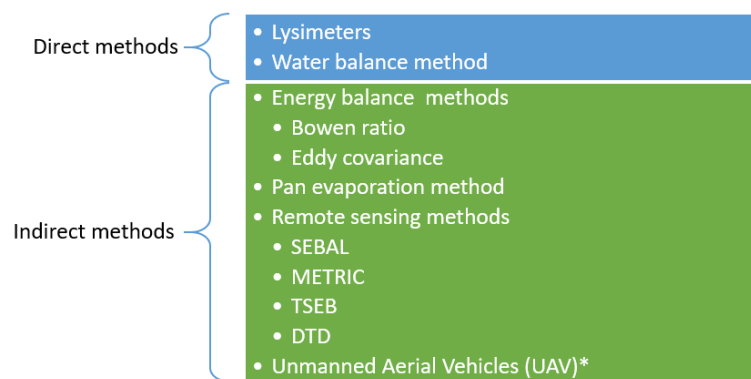


Figure 1. The direct and indirect methods for ET estimation

For indirect methods, there are energy balance methods [1], Pan evaporation methods [8] [10], and remote sensing methods [11]. For energy balance methods, Bowen ratio [12] [13] and eddy covariance [14] have been widely used in ET estimation. Over the past decades, remote sensing techniques have been considered as one of the most powerful methods for mapping and estimating ET [15] [16]. Because they can detect variations in vegetation and soil condition over space and time, remote sensing energy balance models have been useful to account for the spatial variability of ET at regional scales when using satellite platforms such as Landsat and ASTER [17] [18] [19] [20] [21]. Related research has been reviewed a few times [22] [23] [24]. Remote sensing techniques can provide information about Normalized Difference Vegetation Index (NDVI), Leaf Area index (LAI), surface temperature, surface albedo and so on. Several remote sensing models have been developed over the past decades, such as Surface Energy Balance Algorithm for Land (SEBAL) [18] [11], Mapping Evapotranspiration with Internalized Calibration (METRIC) [25], the Dual Temperature Difference (DTD) [26], and the Priestley-Taylor TSEB (TSEB-PT) [27]. There are mainly four energy fluxes in these remote sensing energy balance models, which are net radiation R_n , soil heat flux G , latent heat flux LE and sensible heat flux H . Satellites images can help researchers get spatially distributed measurements, though, the resolution for those images are usually 30 m, which are not good enough for many precision agriculture or ecological applications. Second, the timing of satellites overpass is not always synchronous with research requirements, they usually have a two-week revisit cycle [28]. Third, obtaining high-resolution images can be very expensive [21]. All these disadvantages limit the ET estimation by using traditional satellites images.

As a new potential remote sensing platform, researchers are more and more interested in the potential of UAV data not only in precision agriculture, but also in different scientific and commercial communities [29] [30] [31] [32]. Drones can overcome some of those remote sensing limitations. For example, satellite remote sensing is prone to cloud cover, drones are below the clouds. Compared with satellites, UAVs can be operated at any time if the weather is good. Satellite has fixed flight path, drones are more mobile and more adaptive for site selection. And most drones are cost-effective (thousands of dollars). Mounted on the UAVs, light weight sensors, such as RGB cameras, multispectral cameras, and thermal infrared cameras, can help get higher resolution images. And, higher resolution images are important for recognizing the temporal and spatial variability related to the crops. UAVs have already been widely and successfully used in other agriculture research, such as water stress [33], tree canopy segmentation [34], counting watermelons [35]. Therefore, UAVs have great potential to help estimate ET. Many researchers have already been using UAVs for ET estimation. For example, in [36], the author implemented a remote sensing energy balance (RSEB) algorithms for estimating energy components, such as incoming solar radiation, sensible heat flux, soil heat flux and latent heat flux. Optical sensors were mounted on a UAV to provide high spatial resolution images. By using the UAV platform, experiments results show that the RSEB algorithm can estimate latent heat flux and sensible heat flux with errors of 7% and 5% respectively. It demonstrates that UAV could be used as an excellent platform to evaluate spatial variability in the field.

The rest of the paper is organized as follows. Section II introduces different UAV types being used for ET estimation. Advantages and disadvantages are discussed for them. Several commonly used light weight sensors are also compared in this section. UAV path planning and image processing methods are presented in the second section, too. In section III, the ET estimation methods and models for UAVs agricultural application, such as Two source energy balance (TSEB) model and METRIC model, are discussed. In Section IV, different results of ET estimation methods and models are compared and discussed. Challenges and opportunities for UAVs are mentioned in the conclusion section.

2. Materials and Methods

2.1. Unmanned aerial vehicles (UAV) and light weight sensors

Many kinds of unmanned aerial vehicles are used on different research purposes for ET estimation. Typically, there are two different UAV platforms, fixed-wings, and quadcopter. Fixed-wings can usually fly longer and carry heavy sensors. It can usually fly about 2 hours, which is suitable for a large scale of field. Quadcopters can fly about 30 minutes, which is used for short flight mission in a small scale of field.

In [27], researchers used a fixed-wing UAV to collect thermal data to estimate ET with two source energy balance models. The drone has a 2.2-meter wing span, which can fly 25 minutes when carrying a 2-kilogram payload. The drone was flying at a height of 90 meters with a speed of 60 km/h, which could cover a $400\text{ m} \times 400\text{ m}$ area in one flight. The SkyCircuits ground control station is used to design the flight path and for flight inspection. To collect thermal data, they used an Optris PI 450 LW infrared camera. Based on the specifications, this uncooled thermal camera has an accuracy of $\pm 2\text{ }^{\circ}\text{C}$ or $\pm 2\%$ when the ambient temperature is between $0\text{ }^{\circ}\text{C}$ to $70\text{ }^{\circ}\text{C}$. The thermal images have a spectral range of $7.9\text{ }\mu\text{m}$ and an optical resolution of 382×288 pixels.

In [37], multispectral and thermal images were collected by using an airborne digital system developed by Utah State University. The digital system was installed in a Cessna TU206 aircraft, which also has four ImperX Bobcat B8430 digital cameras. The spectral bands for these cameras are, Red ($0.645\text{ }\mu\text{m} - 0.655\text{ }\mu\text{m}$), Green ($0.545\text{ }\mu\text{m} - 0.555\text{ }\mu\text{m}$), Blue ($0.465\text{ }\mu\text{m} - 0.475\text{ }\mu\text{m}$) and Near-infrared (NIR) ($0.780\text{ }\mu\text{m} - 0.820\text{ }\mu\text{m}$). A ThermalCAM SC640 (FLIR Systems Inc.) is also mounted on the aircraft to collect thermal infrared (TIR) images, the wavelength range is $7.5\text{ }\mu\text{m} - 13\text{ }\mu\text{m}$.

In [21], the author estimated evapotranspiration in a peach orchard by using very-high-resolution imagery from an UAV platform (S1000, DJI, Shenzhen, China). A TIR camera (A65, FLIR Systems Inc.)

and a multispectral camera RedEdge M (MicaSense, Seattle, WA, USA) are also mounted on the drone. The thermal camera A65 has a spatial resolution of 640×512 pixels. It has a Field of View (FOV) of 25° (H) \times 20° (V). The focal length for this TIR camera is 25 mm. For the RedEdge, it has five different bands, Red (668 nm), Green (560 nm), Blue (475 nm), Red edge (717 nm) and Near-infrared (840 nm). RedEdge M has spatial resolution of 1280×960 pixels with a focal length of 5.5 mm.

Quadcopters have been widely used in agricultural research, such as [34] [38], which promises a great potential in ET estimation. In [34], a quadcopter named Hover was used as the UAV platform to collect aerial images, as shown in Figure 2.



Figure 2. The Hover is equipped with a Pixhawk flight controller, GPS, telemetry antennas. Its lithium polymer battery has a capacity of 9500 mAh, which can support a 30-minute flight mission with a camera mounted on it.

The Hover was equipped with a Pixhawk flight controller, GPS, telemetry antennas. And, it can fly over the field by waypoints mode (designed by using Mission Planner software). The Hover's lithium polymer battery has a capacity of 9500 mAh, which can support a 30-minute flight mission with cameras mounted on the Hover drone.

To estimate ET, multispectral images can be collected by Survey 2 (MAPIR, USA) cameras with 4 bands, Blue, Green, Red (RGB) and Near-infrared (NIR). The MAPIR camera has a resolution of 4608×3456 pixels, with a space resolution at 1.01 cm/pixel. The Survey 2 cameras have a faster interval timer, 2 seconds for JPG mode and 3 seconds for RAW + JPG mode. Faster interval timer would benefit the overlap design for UAV flight missions, such as reducing the flight time, enabling higher overlapping. Another multispectral camera being used is Rededge M (MicaSense, USA). The Rededge M has five bands, which are Blue, Green, Red, Near infrared, and Red edge. It has a resolution of 1280×960 pixel, with a 46° field of view. With a Downwelling Light Sensor (DLS), which is a 5-band light sensor that connects to the camera, the Rededge M can measure the ambient light during a flight mission for each of the five bands. Then it can record the light information in the metadata of the images captured by the camera. After the camera calibration, the information detected by the DLS can be used to correct lighting changes during a flight, which usually happens because the clouds cover the sun during a UAV flight.

The thermal camera ICI 9640 P-Series (ICI, USA) was applied for collecting thermal images. The thermal camera has a resolution of 640×480 pixels. The spectral band is from $7 \mu\text{m}$ to $14 \mu\text{m}$. The dimension of the thermal camera is 34 mm \times 30 mm \times 34 mm. The accuracy is supposed to be $\pm 2^\circ\text{C}$. A Raspberry Pi Model B computer was used to trigger the thermal cameras during the flight missions. The SWIR 640 P-Series, which is a shortwave infrared camera, can also be used for ET estimation. The spectral band is from $0.9 \mu\text{m}$ to $1.7 \mu\text{m}$. The accuracy for SWIR camera is $\pm 1^\circ\text{C}$. It has a resolution of 640×512 pixels.

Table 1. Some commonly used sensors on UAV platforms.

Sensors	Function	Resolution	Accuracy
Rededge M	Multispectral	1280 × 960 pixels	8.2 cm/pixel (per band) at 120 m
MAPIR Survey 2	Multispectral	4608 × 3456 pixels	4.05 cm/pixel at 120 m
Mini MCA-6	Multispectral	1280 × 1024 pixels	3.3 cm/pixel at 60 m
Sequoia	Multispectral	4608 × 3456 pixels	17 cm/pixel at 100 m
Cannon S 110	Near infrared	4000 × 3000 pixels	3.5 cm/pixel at 100 m
ICI 9640 P	Thermal infrared	640 × 480 pixel	±1 °C
ICI SWIR 640 P	Short-wave infrared	640 × 512 pixel	±1 °C
Optris PI 450	Thermal infrared	382 × 288 pixels	±2 °C or ±2 %
ThermalCAM SC640	Thermal infrared	640 × 480 pixel	±2 °C or ±2 %
EasIR-9	Thermal infrared	288 × 384 pixel	±2 °C or ±2 %
thermoMAP	Thermal infrared	640 × 512 pixel	14 cm/pixel at 75 m

Some commonly used sensors are listed in the Table 1 for comparison. Compared with traditional remote sensing method, such as satellites, the thermal camera and UAVs make the data collection more flexible and lower cost. Thermal remote sensing images were first used in 1973 to estimate ET [39]. Temperature information is usually converted into land surface characteristics such as albedo, LAI, and surface emissivity. The TIR band is considered as the most important variable because it plays an important role in sensible heat flux and ground heat flux [40]. The cooled thermal cameras are usually more sensitive and accurate than uncooled thermal cameras [41]. But cooled thermal cameras are very big, expensive, and energy consuming [42]. So, they can hardly be used on UAV platform. In contrast, the uncooled thermal camera plays a more and more important role because they are light [43], low power consumption [44] and less expensive than cooled thermal cameras. As mentioned previously, uncooled thermal cameras have been widely used in UAV platform for ET estimation. Not only for ET, it also has been widely used in many other agricultural applications, such as plant disease detection [45], crop water stress estimation [46] [47] and soil moisture detection [48]. And most importantly, they can measure two of the most important parameters of ET estimation, soil temperature and tree canopy temperature, which are widely used in two source energy balance models, SEBAL, and METRIC.

The thermal camera has so many advantages, though, its microlometer is not always sensitive and accurate [42]. Also, most thermal cameras are not always calibrated, which can only measure the relative temperature instead of the accurate value. In precision agriculture, however, most time it's necessary to measure the accurate temperature in many applications [43] such as crop monitoring [49], pest detection [50] and disease detection [45]. Unstable outdoor environmental factors can cause serious measurement drift during flight missions. Post-processing like mosaicking might further lead to measurement errors. We are using thermal camera more and more frequently with its limitations. To answer these two fundamental questions, in [51], the authors finished three experiments to research the best practice of thermal images collection for UAV. To calibrate TIR images, in [21], they used water body and rubber plates as cold and hot features. IR Flash Version 2 is usually used to process thermal UAV images for image format transformation.

2.2. UAV path planning and image processing

In this section, we discussed the potential of path planning and image processing methods for ET estimation. Recently, unmanned aerial vehicles have been widely used in agriculture, such as crop yield estimation [35], soil moisture monitoring, water stress estimation [33] and pest management [52]. Compared with traditional remote sensing tools, such as satellites, drones' flight time can be more flexible and more frequent in the field. And, drones can fly at lower altitude and can take higher spatial and possibly temporal resolution images and thermal images [51] of crops. As a low-cost scientific data collection platform, drones also make data acquisition relatively less expensive. While there are many advantages by using drones for agricultural applications [35], there is still a lot of work for drones when used for estimating evapotranspiration. Many researchers fly the drones in different height,

using specialized equipment, controlling environments and relying on data analysis expertise [53]. Is there any optimal point where the data can be the best representation of crops? In [53], Brandon built a conceptual framework for describing the optimality as a function of spatial, spectral, and temporal factors which represent the best solution. As researchers try to understand the potential of the UAVs, efficient workflow, image processing methods, and better software are still under developing [54] [55] [56] [57].

2.2.1. Pre-flight path planning

Being used as a remote sensing platform, UAVs also create new research problems, such as drone image processing, and flight path planning. Flight missions were programmed by using Mission Planner software. The flight height was setup as 30 m, 60 m, and 120 m in order to compare the image resolution's effect on pomegranate ET estimation. For all the flight missions, the overlap was set up as 75% to make sure the images can be stitched together during images pre-processing. The UAV sensors are designed to take images at nearly 0 nadir angle.

Researchers usually fly drones biweekly to collect data. If there is a drone crash, hardware issues, or unknown reasons, data may not be collected successfully. And, there is only one growing season each year. If data is missed, people have to wait for another year. In order to get enough data, for our research, we flew the drone Hover bi-weekly over the pomegranate field at noon during the growing season in 2017 and 2018. Table 2 is an example of data collection in 2018 by using Survey 2 (MAPIR, USA) and ICI 9460 P (ICI, USA) thermal camera.

Table 2. Multispectral and thermal images collection by UAV during the 2018 growing season at Parlier, CA, USA. -* means there is no data for that day.

Date	Flight time	MAPIR Survey 2	ICI 9460P
4-12-18	12 - 2 pm	30, 60, 120 m	30 m
4-26-18	12 - 2 pm	30, 60, 120 m	30 m
5-9-18	12 - 2 pm	30, 60, 120 m	30 m
5-23-18	12 - 2 pm	30, 60, 120 m	30 m
6-8-18	12 - 2 pm	30, 60, 120 m	-*
6-13-18	12 - 2 pm	30, 60, 120 m	30 m
6-28-18	12 - 2 pm	30, 60, 120 m	-*
7-11-18	12 - 2 pm	30, 60, 120 m	-*
7-25-18	12 - 2 pm	30, 60, 120 m	30 m
8-8-18	12 - 2 pm	30, 60, 120 m	30 m
8-22-18	12 - 2 pm	30, 60, 120 m	30 m
9-14-18	12 - 2 pm	30, 60, 120 m	30 m
10-9-18	12 - 2 pm	30, 60, 120 m	30 m
10-16-18	12 - 2 pm	30, 60, 120 m	30 m
10-30-18	12 - 2 pm	30, 60, 120 m	30 m

2.2.2. Image calibration

In order to minimize the shading effect on the images, the drones are usually flying at noon with clear sky conditions. And, because each pixel in a drone image is a percentage of the reflected light, we need to calibrate the pixel value by using a known reflectance value. Therefore, the images of a color panel were taken right before and after the flight missions, servicing as the reflectance reference.



Figure 3. The images of a color panel were taken right before and after the flight missions, servicing as the reflectance reference.

It is important to take pictures of the reference panel immediately before and after the flight missions, because the solar angle and light intensity can change a lot based on an experiment in [47], which can cause inaccurate experiment results. Then, the digital number of the images were converted to reflectance by an empirical method in [58]

$$\rho_{\lambda} = \frac{DN - DN_d}{DN_w - DN_d}, \tag{2}$$

where ρ_{λ} is the reflectance and DN is the digital number of the raw image. DN_d and DN_w are the dark rellfectance point and white reflectance point in color checker respectively.

2.2.3. Images stitching and orthomosaic image generation

After the flight missions, all of the aerial images were stitched together to generate the orthomosaick images in PhotoScan (Agisoft LLC, Russian). As shown in Table 3, there are usually three steps for image processing, which includes aligning photos, building mesh (optional) and generating orthomosaick. For example, the Agisoft Photoscan settings and workflow for RGB and NIR images (by MAPIR Survey 2) are listed in Table 3.

Table 3. Orthomosaic images generation workflow in Agisoft Photoscan

Step 1 : Align Photos	Step 2 : Build Mesh	Step 3 : Build Orthomosaick
Accuracy: Medium	Surface type: Height field (2.5D)	Type: Planar
Generic preselection: Yes	Source data: Sparse cloud	Projection plane: TOP XY
Key point limit: 40,000	Face count: Medium (30,000)	Rotation angle: 0
Tie point limit: 4,000	Interpolation: Enabled (default)	Surface: Mesh
Adaptive camera model fitting: No	Point classes: All	Blending mode: Mosaic (default)
	Caculate vertex colors: Yes	Enable hole filling: Yes
		Enable back-face culling: No

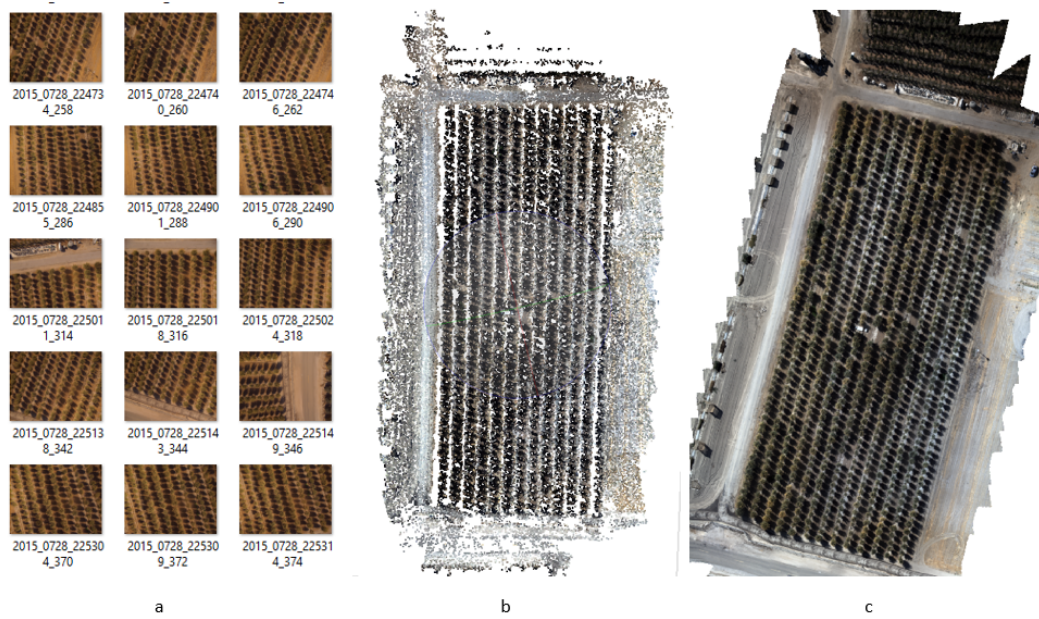


Figure 4. This is a image processing workflow to generate orthomosaick image for a pomegranate field (a) Some drone image examples taken by MAPIR . (b) Tie points images generated by Photoscan. (c) The orthomosaick image for the pomegranate field.

2.2.4. Post-processing calibrated images

In this subsection, a MATLAB code was written in order to calculate the normalized difference vegetation index (NDVI) by using the orthomosaick image. Many studies have used NDVI from remote sensing platform to help estimate crop coefficient values on crops, such as [59] [60] [61]. Then, crop coefficient can be applied to creating local and regional crop evapotranspiration (ET_c) maps. In [62] [63], they used a direct method to estimate ET by using the correlation between the ET, the remote sensing data, and the meteorological data. It is based on the assumption that daily ET has a correlation with the surface and air temperature. According to the relationship between NDVI and surface temperature, [64] [65] proposed a new direct method, which shows that NDVI and surface temperature have effect on the ET quantity. The relationship between the NDVI and surface temperature is strong, though, it changes by seasons [65]. LAI and fractional vegetation cover are also often calculated by using NDVI according to an empirical LAI-NDVI relation [66] [67].

Therefore, it is important to get NDVI values first before proceeding to further ET research. To calculate the NDVI, the following equation is used:

$$NDVI = \frac{\rho_{NIR} - \rho_R}{\rho_{NIR} + \rho_R}, \quad (3)$$

where ρ_{NIR} and ρ_R are the reflectance of near-infrared and red wavebands, respectively. After generating the NDVI image, QGIS was used to generate the pseudocolor map for the NDVI map, as shown in Figure 5.

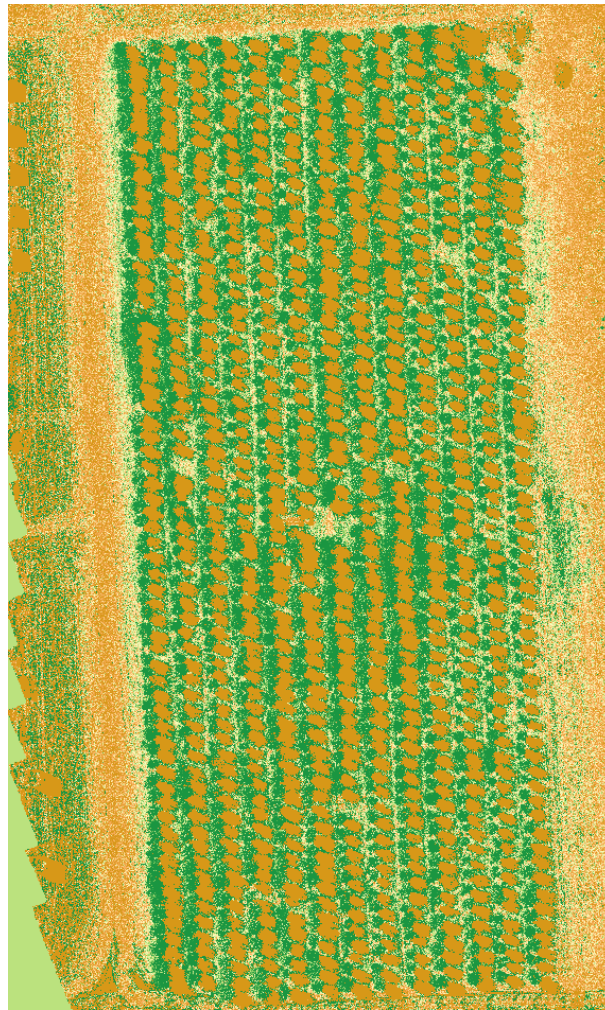


Figure 5. NDVI map for a pomegranate field at USDA

3. ET Estimation Methods for UAV

Most ET estimation methods for UAVs are based on satellites remote sensing methods, the most widely used is energy balance models. TSEB, TSEB-PT, and DTD are introduced in this section. METRIC is also mentioned because of its potential in UAVs application. Researchers need to preprocess the images to input them in the METRIC model, because this model is originally for satellites images. But some research still shows good results [68].

3.1. Two source energy balance (TSEB) models

The TSEB is proposed by Norman [26] and modified by Norman and Kustas to improve the accuracy of LE estimation [24] [69] [70]. Based on [71] [72], the TSEB is effective over homogeneous land and environmental conditions. And, it can reproduce fluxes with similar results to tower-based observations. The TSEB model separates the land surface temperature into soil surface temperature and vegetation surface temperature, and it considers sensible and latent heat fluxes are transferred to the atmosphere from both surface temperature components, as shown in the following equations [37].

$$R_n = R_{ns} + R_{nc} = H + LE + G, \quad (4)$$

$$R_{ns} = H_s + LE_s + G, \quad (5)$$

$$R_{nc} = H_c + LE_c, \quad (6)$$

where R_n is the net radiation (W m^{-2}). The soil net radiation (W m^{-2}) is represented by R_{ns} and R_{nc} is the canopy net radiation (W m^{-2}). H is the sensible heat flux (W m^{-2}). H_s and H_c are soil sensible heat flux (W m^{-2}) and canopy sensible heat flux (W m^{-2}) respectively. LE is the latent heat flux (W m^{-2}). LE_s and LE_c are soil latent heat flux (W m^{-2}) and canopy latent heat flux (W m^{-2}) respectively. G is the soil heat flux (W m^{-2}).

$$H_s = \rho C_p \frac{T_s - T_{ac}}{R_s}, \quad (7)$$

$$H_c = \rho C_p \frac{T_c - T_{ac}}{R_x}, \quad (8)$$

where ρ is the air density (kg m^{-3}), C_p is the specific heat of air ($\text{J kg}^{-1} \text{K}^{-1}$). T_{ac} is the air temperature in the vegetation. R_s is the resistance to heat flux above the soil surface (s m^{-1}). R_x is the boundary layer resistance of the canopy leaves (s m^{-1}).

The net radiation is divided into two parts, the soil net radiation and the canopy net radiation [73] [74].

$$R_{ns} = \tau_l L_d + (1 - \tau_l) \varepsilon_c \sigma T_c^4 - \varepsilon_s \sigma T_s^4 + \tau_s (1 - \alpha_s) S_d, \quad (9)$$

$$R_{nc} = (1 - \tau_l) (L_d + \varepsilon_s \sigma T_s^4 - 2 \varepsilon_c \sigma T_c^4) + (1 - \tau_s) (1 - \alpha_c) S_d, \quad (10)$$

where τ_l and τ_s are the longwave and shortwave radiation transmittances through the canopy. L_d and S_d are the incoming longwave and shortwave radiation (W m^{-2}), which are usually measured from a nearby weather station. σ is the Stefan-Boltzmann constant, which is approximately $5.67 \times 10^{-8} \text{ (W m}^{-2} \text{K}^{-4})$. ε is the surface emissivity. α is the surface albedo. And T is the surface temperature (K). Subscripts "s" and "c" mean soil and canopy.

In the TSEB model, the land surface temperature $T_R(\theta)$ is related to the soil and canopy component temperatures. The fraction of vegetation cover viewing angle θ also has effect on the estimation of $T_R(\theta)$, as shown in the following equation

$$T_R \approx [f_c(\theta) T_c^4 + (1 - f_c(\theta)) T_s^4]^{\frac{1}{4}}, \quad (11)$$

where $f_c(\theta)$ is the fractional vegetation cover observed as angle θ by radiometer. To calculate $f_c(\theta)$, the following equation [75] was used

$$f_c(\theta) = 1 - e^{\frac{-0.5 \Omega(\theta) LAI}{\cos(\theta)}}. \quad (12)$$

The UAV can help get maps of NDVI, LAI, $f_c(\theta)$, and $T_R(\theta)$, which are the most important input data for TSEB model. LE_c is initially estimated by using Priestley-Taylor formulation:

$$LE_c = \alpha_{PT} f_G \frac{\Delta}{\Delta + \gamma} R_{nc}, \quad (13)$$

where α_{PT} is the Priestley-Taylor coefficient. f_G is the LAI fraction. Δ is the slope of the saturation vapor pressure-temperature curve (Pa K^{-1}). γ is the psychrometric constant (Pa K^{-1}).

3.2. TSEB-PT model

In [27], the authors used two source energy balance models, the Priestley-Taylor TSEB (TSEB-PT) and the Dual-Temperature-Difference (DTD) when flying a fixed-wing UAV. The TSEB was first proposed by [76]. In 1995, Norman [26] applied an iterative process in order to derive the canopy and soil temperature. The method is based on the assumption of canopy transpiration, which was shown in Priestley and Taylor potential evapotranspiration [77]. Therefore, this method is called TSEB-PT in order to differentiate it from other TSEB methods. The TSEB-PT method splits the surface temperature into two layers, the canopy T_c and soil T_s temperatures. The calculation of sensible heat flux and latent

heat flux for canopy and soil are separate, which makes the parameterization of resistances easier compared with a single layer model.

3.3. Dual-Temperature-Difference DTD model

The DTD model was proposed in [78], which separates the land surface temperature into vegetation and soil temperatures. Then, it calculates the surface energy balance components with the same procedures as TSEB-PT. And, the DTD model can add one more input dataset, the land surface temperature retrieved one hour after sunrise. Because the energy fluxes are minimal at sunrise, it can help minimize the bias in the temperature estimation. In [79], Guzinski produced surface energy flux successfully by using the DTD model. The author used night observations to substitute for the early morning observation. However, the temporal resolution of the satellite observations is still limited, especially when the weather conditions are bad, such as overcast as satellite thermal infrared observations cannot penetrate clouds. The incapacity to collect data in overcast situations applies to all satellite sensors except for those operating in the microwaves region [79].

The calculation of soil heat flux G is also different because the radiometric temperature is dissimilar between midday and sunrise observations. This difference can be used to estimate the soil surface temperature variations. Soil heat flux is calculated based on the model of [80]. For more details about the TSEB-PT and DTD equations, see [81] [82].

By using the UAV platform, it can get better data in accordance with the DTD requirements and during overcast conditions [83].

3.4. METRIC model

METRIC is originally a satellite images processing model for estimating ET as a residual of the energy balance, which was developed by University of Idaho. METRIC has a self-calibration process which contains ground-based hourly reference ET and the selection of hot, cold pixels [84]. The METRIC is based on SEBAL. SEBAL's innovative component is that the model uses a near-surface temperature gradient, dT . It helps eliminating the need for absolute surface temperature calibration [11]. SEBAL uses T_s , ρ , NDVI and their relationships to calculate the surface fluxes [18]. And SEBAL has been evaluated all over the world [85] [86] [87] [88] [89]. In METRIC, they use the SEBAL to estimate dT . Therefore, there is no need to get accurate aerodynamic surface temperature. In [1], the author summarized three differences between the SEBAL and METRIC. First, for the cold pixel, the METRIC does not consider sensible heat flux as zero. Instead, a surface soil water balance is applied to set ET as zero and 1.05 times reference ET at hot and cold pixels respectively. Reference ET is calculated by using the standardized ASCE Penman-Monteith equation. Second, in METRIC, cold pixels are selected in agricultural settings instead of biophysical characteristics. Third, the extrapolation of instantaneous ET is based on reference ET instead of the actual evaporative fraction.

Allan [17] compared the ET estimation between the METRIC and lysimeter near the Montpellier, Idaho. The difference between METRIC and lysimeter is only 4 %. LE is the rate of latent heat loss from the surface because of the ET. METRIC estimates ET as a residual of the energy balance

$$LE = R_n - G - H, \quad (14)$$

where R_n is the net radiation (W m^{-2}), which is calculated by solving the radiation balance as described in [90]. H is for sensible heat flux (W m^{-2}), which is a function of air density, air specific heat, temperature difference between two canopy heights, the aerodynamic resistance, and temperature gradient. LE is for latent heat flux (W m^{-2}) and G is for soil heat flux (W m^{-2}).

For the R_n , it can be calculated based on the following equation [25]:

$$R_n = (1 - \alpha)R_{s\downarrow} + R_{L\downarrow} - R_{L\uparrow} - (1 - \epsilon_o)R_{L\downarrow}, \quad (15)$$

where $R_{s\downarrow}$ is the incoming short-wave radiation (W m^{-2}). α is the surface albedo, which is dimensionless. $R_{L\downarrow}$ and $R_{L\uparrow}$ are the incoming long-wave radiation (W m^{-2}) and outgoing long-wave radiation (W m^{-2}), respectively. ε_o is the thermal emissivity, which is also dimensionless.

$R_{s\downarrow}$ is usually calculated by the following equation if the sky is clear

$$R_{s\downarrow} = \frac{G_{sc} \cos \theta_{ref} \tau_{sw}}{d^2}, \quad (16)$$

where G_{sc} is the solar constant (1367 W m^{-2}). θ_{ref} is the solar incidence angle. d is the relative distance between the Earth and the Sun. τ_{sw} is the broad band atmospheric transmissivity, which is calculated by using the following equation [91]

$$\tau_{sw} = 0.35 + 0.627 \exp \left[\frac{-0.00146P}{K_t \cos \theta_{hor}} - 0.075 \left(\frac{W}{\cos \theta_{hor}} \right)^{0.4} \right], \quad (17)$$

where P is the atmospheric pressure (kPa). W is the water in the atmosphere (mm). θ_{hor} is the solar zenith angle over a horizontal surface. K_t is the unitless turbidity coefficient. P and W are also calculated based on [91]

$$P = 101.3 \left(\frac{293 - 0.0065z}{293} \right)^{5.26}, \quad (18)$$

$$W = 0.14e_a P_{air} + 2.1, \quad (19)$$

where z is the elevation above the sea level. e_a is the near-surface vapor pressure (kPa).

For the incoming and outgoing long wave radiation $R_{L\downarrow}$ and $R_{L\uparrow}$, they are calculated by using Stefan-Boltzmann equation

$$R_{L\uparrow} = \varepsilon_o \sigma T_s^4, \quad (20)$$

$$R_{L\downarrow} = \varepsilon_a \sigma T_a^4, \quad (21)$$

where ε_a is effective atmospheric emissivity, which is usually calculated by an empirical equation

$$\varepsilon_a = 0.85(-\ln \tau_{sw})^{0.09}, \quad (22)$$

H is computed from surface roughness, wind speed, surface temperature ranges. The sensible heat flux is considered as the most difficult term to calculate in the energy balance equation, in order to calculate the H , an aerodynamic function is used:

$$H = \rho_{air} C_p \frac{dT}{r_{ah}}, \quad (23)$$

where r_{ah} is the aerodynamic resistance (s m^{-1}) between two surface height. In METRIC, r_{ah} is usually calculated by using the wind speed and an iterative stability correction, as shown in the following equation

$$r_{ah} = \frac{\ln(z_2/z_1)}{u_* k}, \quad (24)$$

where z_1 and z_2 are heights above the zero-plane displacement of the vegetation. k is the von Karman constant (0.41). u_* is the friction velocity (m s^{-1}), which is calculated by using

$$u_* = \frac{ku_{200}}{\ln(200/z_{om})}, \quad (25)$$

where u_{200} is the wind speed at a blending height 200m. And z_{om} is the momentum roughness length (m).

dT is the temperature difference between the air and the surface. A strong linear relation between the dT and the surface temperature were found in [25] [18] [92] [93]. The equation for the dT and the surface temperature is shown as

$$dT = a + bT_s, \quad (26)$$

where a and b are derived parameters empirically based on two extreme hot and cold pixels. These hot and cold pixels defined the boundary of the sensible heat flux. The cold pixel represents a well-watered area with no water stress. The H is assumed to be minimum and ET should be maximum. The hot pixel represents a dry and bare field where H is maximum and ET is almost zero. Hot and cold pixels must be selected by experienced users, which makes it difficult for beginner level researchers.

In METRIC, G is the heat storage in soil and vegetation, which can be estimated by R_n surface temperature and vegetation index [85] according to

$$\frac{G}{R_n} = (T_s - 273.15)(0.0038 + 0.0074\alpha)(1 - 0.98NDVI^4). \quad (27)$$

Another equation used for estimating G is proposed by Tasumi [94], who used the data collected by [95].

$$\frac{G}{R_n} = 0.05 + 0.18e^{-0.521LAI} (LAI \geq 0.5), \quad (28)$$

$$\frac{G}{R_n} = 1.80 \frac{(T_s - 273.15)}{R_n} + 0.084 (LAI \leq 0.5). \quad (29)$$

Based on Tasumi's experiment results [94], the two above calculation methods can measure G accurately for irrigated plants in Kimberly, Idaho.

4. Results and Discussion

The comparison between UAV methods and results from other traditional remote sensing studies reveals that the data from the UAV platform and the light weight cameras can estimate the surface energy fluxes with similar accuracy estimated by using satellite data. Therefore, the UAV data can be used for modelling ET estimation with high confidence.

4.1. TSEB models

Mounted on the UAV, multispectral sensors and thermal camera can help obtain high-resolution images. In [37], the authors used a two-source energy balance model for a sub-field and plant canopy scale ET monitoring. It is concluded that the TSEB model can simulate the energy balance components in two vineyards with mean absolute error (MAE) ranging from 15 to 90 W m⁻². They found that the TSEB model is fairly robust and able to calculate LE and ET values under a varying environmental conditions. By using the TSEB, the T_s and T_c have a bias of 0.5 °C and RMSD on the order of 2.5 °C. The accuracy is similar with the following papers [96] [70] [69] [97], in which the RMSD values are between 2.4 to 5.0 °C for T_s and 0.83 to 6.4 °C for T_c .

For a lightweight thermal camera mounted on a UAV, Hoffmann [27] concatenated the LST thermal images into orthomosaick, then applied as the input for TSEB model. Based on the comparison between UAV fluxes and eddy covariance (EC) fluxes, the R_n for TSEB is in good agreement with R_n measured from EC with a RMSE of 44 W m⁻². The sensible heat flux (H) for DTD has RMSE and MAE values of 59 W m⁻² and 49 W m⁻². The soil heat flux (G) are underestimated, which has RMSE and MAE values of 48 W m⁻² and 45 W m⁻². For the latent heat flux, DTD has RMSE and MAE values of 67 W m⁻² and 57 W m⁻². They concluded that the TIR camera placed on a UAV platform can provide high spatial and temporal resolution data for estimating energy balance fluxes of ET. This study show similar results with Guzinski's work [81], who applied the TSEB-PT at the same site but used satellites

images instead of UAV images. In [81], the RMSE is 46 W m^{-2} for R_n , 56 W m^{-2} for H , and 66 W m^{-2} for LE .

In [36], the RSEB algorithm was well implemented and only the climatic parameters, such as T_a , wind speed are measured as the input data. UAV images are used for calculating the NDVI and T_s . The authors evaluate the RSEB algorithm by H and LE . Results show that the algorithm estimates LE and H with errors of 7% and 5%. The RMSE and MAE for LE are 50 and 43 W m^{-2} . And, the RMSE and MAE for H are 56 and 46 W m^{-2} . It also concludes that high spatial resolution maps are capable to detect significant differences between the energy balance fluxes above the tree canopy and the soil surface between rows. UAV could also be used to help the satellite platforms for estimating spatial variability of ET maps.

4.2. SEBAL and METRIC models

In [68], Montibeller used multispectral and thermal camera to collect data for running the SEBAL model. To evaluate the estimated energy fluxes, they used linear regression models, residual plots, RMSE, and MAE methods. The R^2 for the R_n is 0.71, which was underestimated about 17% compared with the flux towers. The RMSE for the R_n is 6.09 W m^{-2} . For the G , the R^2 is 0.17, with a RMSE of 11.23 W m^{-2} . The R^2 for the H is 0.5, with a RMSE of 8.84 W m^{-2} . It overestimate the flux by 5 %. For the LE , R^2 is 0.82m, with a RMSE of 2.67 W m^{-2} . The research also shows that the ET rate is relevant to the crop growth stage. Corn, for example, has higher ET rates up until the tassel appears. However, the relationship between the NDVI and ET is very poor, further study need to be explored. Overall, the research proves that SEBAL model can be used for estimating ET with UAV very well. The algorithms being used by [68] were automated by Python, which facilitates the data processing for ET estimation. This method can also be used for decision making in real time, which can monitor the water consumption of each crop in the field.

5. Conclusions

According to the previous review, it shows that each ET estimation model has its own advantages and disadvantages. For example, METRIC/SEBAL methods, they are more recognized by the remote sensing researchers, but they are based on satellite (Landsat) platforms. It may require more efforts to make it work with UAV images. The TSEB model is less widely known, but it seems that it offers more potential for UAV applications in many crop conditions, for example, tree crops like pomegranate, nectarine, and almonds. When flying a drone, weather condition, field scale, flight time and many other factors should also be considered in order to choose the appropriate algorithms for ET estimation.

ET estimation methods and related agricultural applications have significantly been developing over the past decades. Although remote sensing ET models can help get relatively accurate spatial distribution ET data, important outstanding ET estimation questions are still from local to large scales because of the deficiency of our observation capability. No existing methods can fully satisfy the spatial, temporal, spectral, and accuracy requirements for ET-based science and applications [98]. Therefore, innovative methods or models for ET estimation are required by using UAVs. In [98], the author proposed five requirements to map ET with high fidelity in the future, which are high frequency, high spatial resolution, high temporal resolution, large spatial coverage, and long-term monitoring. High frequency will improve the differentiation of water stress between crops, which enables more efficient water management. High spatial resolution can help detect spatially heterogeneous responses to the water stress. Because ET is highly variable within and among days, high temporal resolution can help detect crops ET in real-time. Large spatial coverage can help detect large scale drought. And long term monitoring will be important to record ET variability.

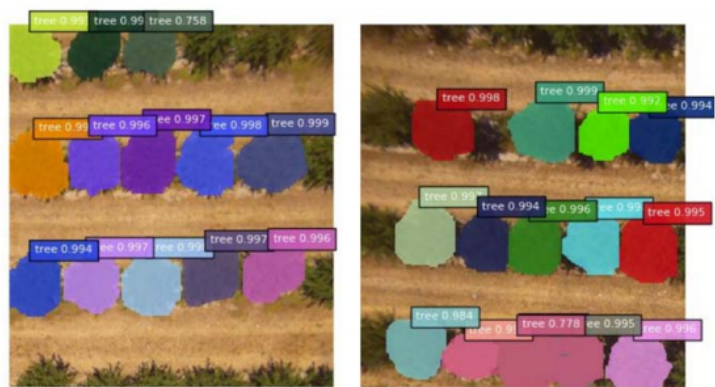


Figure 6. ET estimation in a single tree level

Compared with other satellites based remote sensing methods, UAV platform and light weight sensors can provide better quality, higher spatial and temporal resolution images. UAV can be used to estimate ET in excellent scale and with flexible flight schedules. In the future, tree-by-tree ET estimation would be useful to analyze the crops temporal and spatial variability in the field, as shown in Figure 6. In the [34], they used a deep learning neural network to train UAV images to identify tree canopies at very high accuracy, which have a great chance to be used in the future to estimate ET of a single tree. Also, further research should be focused on remote sensing algorithms and their applications on different crops.

Author Contributions: Haoyu Niu designed the experiments, processed and analyzed the UAV images, and wrote the original manuscript draft. Tiebiao Zhao, Dong Wang, and YangQuan Chen contributed to the results interpretation and discussions and editing of the manuscript.

Funding: This research is supported in part by USDA NIFA grant 2016-38422-25551 (2016-2020) entitled “Innovation and Entrepreneurship Programs to Train Hispanic Students for Future Agricultural and Food Industry”.

Acknowledgments: Thanks go to Dong Sen Yan, Joshua Ahmed, Allan Murillo, Stella Zambrzuski and Christopher Currier for flying drones or collecting field measurements. We gratefully acknowledge the support of NVIDIA Corporation with the donation of the Titan X Pascal GPU used for this research.

Conflicts of Interest: The authors declare no conflict of interest.

Abbreviations

The following abbreviations are used in this manuscript:

MDPI	Multidisciplinary Digital Publishing Institute
ET	Evapotranspiration
UAV	Unmanned Aerial Vehicles
METRIC	Mapping Evapotranspiration at high Resolution with Internalized Calibration
NDVI	Normalized Difference Vegetation Index
SEBAL	Surface Energy Balance Algorithm for Land
RGB	Red Green and Blue
RSEB	Remote sensing energy balance
TSEB	Two source energy balance
TIR	Thermal infrared
FOV	Field of View
NIR	Near infrared
JPEG	Joint Photographic Experts Group
SWIR	Short-wave infrared
MAE	Mean absolute error

Appendix A MATLAB code and Photoscan settings

Appendix A.1 Matlab code for calculating drone image NDVI

The Matlab code is attached here for calculating NDVI of drone images, which can be found at Github: <https://github.com/niuhaoyu16/NDVI-for-drone-images>

Appendix A.2 Agisoft Photoscan image processing settings

Agisoft Photoscan image processing settings is available at Github: <https://github.com/niuhaoyu16/AgisoftPhotoscan-Images-processing-settings>

Appendix B Open source python code for METRIC and TSEB

Appendix B.1 Python code for TSEB

PyTSEB is available at Github: <https://github.com/hectornieto/pyTSEB>

Appendix B.2 Python code for METRIC

PyMETRIC is available at Github: <https://github.com/hectornieto/pyMETRIC>

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