

1 Article

2 Application of SWAT (Soil and Water Assessment 3 Tool) to the Abay River Basin of Ethiopia: The Case 4 of Didessa Sub Basin

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9 **Abstract:** In this study the semi-distributed model SWAT (Soil and Water Assessment Tool), were
10 applied to evaluate stream flow of Didessa sub basin, which is one of the major sub basins in Abay
11 river basin of Ethiopia. The study evaluated the quality of observed meteorological and
12 hydrological data, established SWAT hydrological model, identified the most sensitive parameters,
13 evaluated the best distribution for flow and developed peak flow for major tributary in the sub
14 basin. The result indicated that the SWAT model developed for the sub basin evaluated at multi
15 hydro-gauging stations and its performance certain with the statistical measures, coefficient about
16 determination (R^2) and also Nash coefficient (NS) with values ranging 0.62 to 0.8 and 0.6 to 0.8
17 respectively at daily time scale. The values of R^2 and NS increases at monthly time scale and found
18 ranging 0.75 to 0.92 and 0.71 to 0.91 respectively. Sensitivity analysis is performed to identify
19 parameters those were most sensitive for the sub basin. CN2, GWQMN, CH_K, ALPHA_BNK and
20 LAT_TIME are the most sensitive parameters in the sub basin. Finally, the peak flow for 2-10000
21 returns periods were determined after the best probability distribution is identified in EasyFit
22 computer program.

23 **Keywords:** Didessa Sub basin; SWAT; Simulated Stream flow

25 1. Introduction

26 Hydrological models have been widely applied for comprehending hydrological process
27 within the catchment since the last decades[5, 1]. The models are tools that depict the physical
28 process regulating the conversion of precipitation to stream flow and to represent the catchment
29 process in a simplified way. There are various hydrological models designed to simulate the
30 relationship between rainfall and runoff under different temporal and spatial dimensions. The focus
31 of these models will be to set a relationship between different hydrological components such as
32 precipitation, evapotranspiration, surface runoff, ground water movement. Hence, hydrological
33 models required on a plan must be more robust, and transparent as they would progressively
34 depend on to make informed decisions on sharing and management of limited water resources [2,3].
35 They consider the spatial and temporal changes of different factors [4,6]. Physically based
36 distributed watershed models play a major role in analyzing the impact of land management
37 practices on water, sediment, and agricultural chemical yields in large complex watersheds.

38 The Soil and Water Assessment Tool, a physically-based, semi-distributed, continuous
39 simulation model, is a guaranteeing model which has been broadly used to comprehend water
40 quantity and quality issues over a wide range of watershed scales and environmental conditions
41 [23,27,44-45]. Also, Soil and Water assessment Tool may be a river basin model formed in order to
42 foresee the effect of land management practices on water, sediment, also agricultural substance
43 yields, complex watershed of uneven soils, land use, and management situations for long duration
44 of time. This model has been computationally efficient and easily makes use of available inputs data
45 and enables clients to consider long-term effects. On the other hand, the Soil and water assessment

46 Tool may be capable to simulate a individual watershed or a system of more than one hydrologically
47 joined watersheds, each of which separated into sub basins. The sub basin created should finally
48 partition to hydrologic response units (HRUs) depending on soil classifications and land use
49 distributions. Although Didessa sub basin is less studied sub basin of Upper Blue Nile, there have
50 been several successful SWAT simulations at other sub basins of Blue Nile, Ethiopia. As a further
51 contribution to SWAT simulations in Ethiopia, on this study the Soil and Water Assessment Tool
52 model is established to a sub basin of Upper Blue Nile Basin, namely Didessa sub basin.

53 Likewise anthropogenic activities are dramatically expanding in Ethiopia owed to population
54 growth at alarming rate as well as investment advancement of the country [7 – 9] and model based
55 estimation of watershed outflow is the base in monsoonal climate where, rainfall-runoff relationship
56 of the several landscapes units have complicated hydrology [10]. Generally “watershed variables
57 attributed to the differences in hydrological response of rainfall are soil properties, geology,
58 anthropogenic activities, relief size, local climate and vegetation cover” [10 – 13].

59 Model based approximation of watershed outflow enlarges the prognostic power of watershed
60 hydrology, as this provides a basis for planning of land management issues for developing and
61 securing water resources [14]. Therefore, understanding the watersheds as sources for stream flow
62 is about significant essentialness in the effective usage of water resources, to enhance management
63 activities of water resources, and to mitigate adverse effects of climate change later on. In line with
64 this critical knowledge gap the flood occurrence in 2015 at Arjo Dedessa dam site initiate this study
65 at this particular watershed.

66 The stream of Didessa, the biggest branch of Upper Blue Nile (Abay river of Ethiopia) shares
67 approximately about one fourth of the down stream of Blue Nile [15] has entire catchment area
68 drained toward those waterway is assessed to be 28,000 km². Although the sub basin has
69 comparatively sufficient hydrological and meteorological data series, its hydrological situation have
70 not been well investigated as compared with northern side sub basins of the Blue Nile (Tana sub
71 basin). In relation with its catchment situation, the occurrence of the 2015 flood incidence at the coffer
72 dam site with below normal rainfall season, calling an urgent discussion and evaluation of the
73 hydrological design of the coffer dam and the relief culvert (OWWDSE, 2016). It further invites detail
74 hydrological study to understand and differentiate dominant hydrological processes and parameters,
75 which govern the hydrological condition of the sub basin.

76 Therefore, this study is relevant to fill the knowledge gap of the hydrological situation of the sub
77 basin with finding scientific cause for the flood occurrence and to make sustainable the water
78 resources development activity in the sub basin. The foremost objective of the study is to understand
79 the hydrological situation of Didesa Sub basin that possible elaborated with the following specific
80 objectives: to (1) find the sensitive parameters for the sub basin by using SWAT-CUP algorithm (Sufi-
81 2), (2) evaluate the best probability distributions of flow data using EasyFit, (3) identify the sub
82 catchment which contribute the highest flow, (4) Evaluate the water balance of the main sub
83 catchments of the sub basin using SWAT simulated flow and (5) undertake frequency analysis of to
84 fix extreme Floods with different return periods.

85

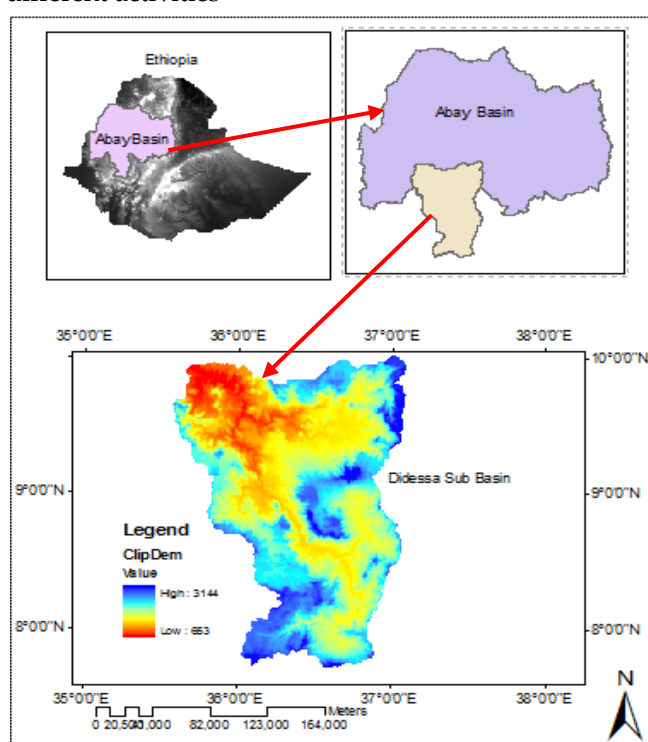
86 2. Materials and Methods

87 2.1. Study Area

88 Didessa sub-basin is located in western part of Ethiopia between latitude 07°40' - 10°0' North
89 and longitude 35°32' - 37°15' East. The overall elevation in the basin varies between 653 and
90 3144 meter above sea level. The absolute catchment coverage drained by the river is projected and
91 delineated to be 28,229 km² initiating from the mount of Gomma in South Western
92 Ethiopia. The SWAT simulated average yearly precipitation of the study area is found to be 1745mm.
93 Most of Didessa sub basin is found in humid tropical climate with heavy rainfall and most of the

94 annual precipitation is received during one season named kiremt. The highest and lowest
 95 temperature ranges amidst 21.3 – 30.90C and 10.9 - 15.10C, respectively.

96 From the assessment of land use/cover, major land use types identified include moderately
 97 cultivated, dense woodland, intensively cultivated land, wooded grassland, open
 98 woodland, natural forest cover, natural forest with coffee, coffee farm with shade trees, riverine
 99 forest, bamboo forest, plantation forest, settlement, shrub land and open grassland. According to
 100 Oromia Water Work Design and supervision (OWWDSE) [47] of Arjo Didessa dam project
 101 feasibility study of 2014 different land use types in different land cover have been identified in the
 102 sub basin. These include mixed cultivation, coffee production, livestock production, subsistence and
 103 commercial forest products utilization, non-timber products utilization, beekeeping, Wildlife
 104 management and utilization, infrastructure development, mining and investment activities on
 105 different activities



106
 107 **Figure 1.** Didessa Sub-basin study site, Abay basin, Ethiopia

108 2.2. Input Dataset and Sources of Data for the model

109 Swat requires daily weather data including precipitation, wind Speed, minimum and maximum
 110 temperature, relative humidity and solar radiation. The main data categories that were utilized in
 111 this study incorporate climate, hydrology, soils, land use/land cover information and more advanced
 112 DEM of 30mx30m spatial resolution. Weather or Climate information was gathered from National
 113 Meteorological Agency of Ethiopia whereas daily flow records were obtained from Ministry of Water,
 114 Irrigation and Energy of Ethiopia. Furthermore, landuse/cover ((MERIS land use land cover, 2009)
 115 were gotten from Oromia Water work Design and Supervision Enterprise (OWWDSE)[47]. Lastly,
 116 Soil shape file was collected from Dr. Belete Berhanu, Soil geo-database of Ethiopia' prepared by
 117 [16]. The meteorological stations are scatter populated and some stations base period are recent with
 118 high missing record. In the case of unavailability of relative humidity, wind speed, and sun shine
 119 hour's data the model might have been run with daily rainfall and temperature. Underneath table 1

120 indicates recorded weather monitoring stations plus accessible information of the time range utilized
121 as input of the study area.

122 Flow records obtained from the Ministry of Water and Energy in Ethiopia at relevant gauging
123 station of Didessa river basin is located near Dembi (Toba),Arjo Didesa near Arjo,Dabana near
124 Abasina,Wama near Nekemte and Angar near Nekemte.

125

126 **Table 1.** List of Selected weather monitoring Stations and Available data sets for rainfall and climatic variables

Station Name	Zone	Station Elevation(m)	Latitude (Deg)	Longitude (Deg)	Data coverage (year)	% of missing Rainfall	% of missing Temp.
Bedele	Illubabor	2011	8.5	36.3	1980-2015	17	24.9
Arjo	Misrak Wellega	2565	8.8	36.5	1989-2015	27	33.1
Shambu	Misrak Wellega	2460	9.6	37.1	1980-2015	14	41.0
Nekemte	Misrak Wellega	2080	9.1	36.5	1980-2014	7	11.5
Gimbi	Mirab Wellega	1970	9.2	35.8	1980-2015	18	41.4
Nedjo	Mirab Wellega	1800	9.5	35.5	1980-2015	20	21.8
Jimma	Jimma	1718	7.7	36.8	1980-2015	5	4.6
Dedessa	Misrak Wellega	1310	9.4	36.1	1980-2015	18	38.1

127

128 **Table 2.** Basic Hydrometric monitoring description for Didesa River Basin

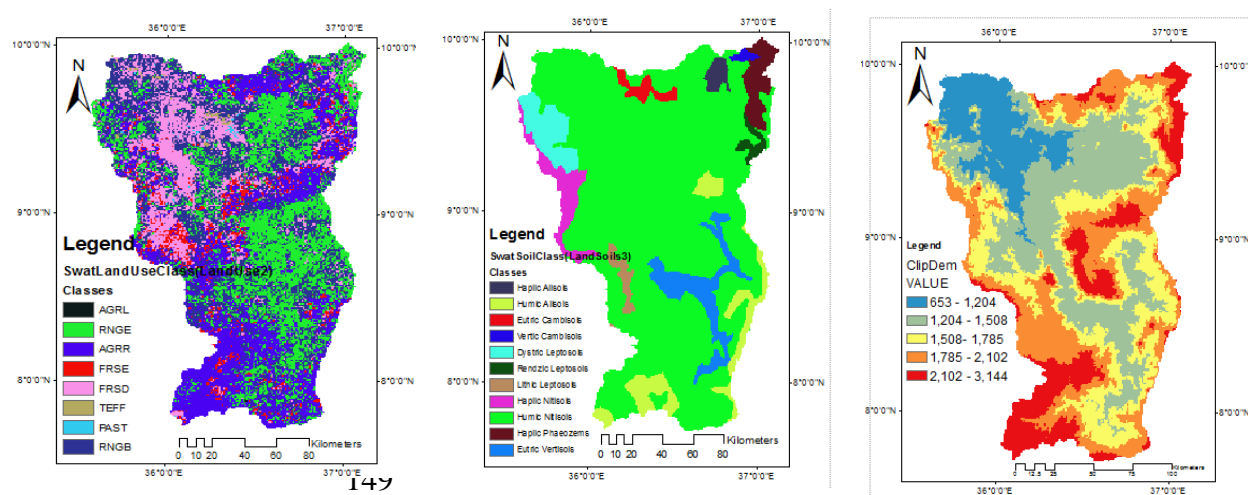
S.N	River	Station	Latitude Deg.	Min	Longitude Deg.	min	Catchment area (km ²)	Data coverage (year)	% Missing
1	Didessa	Arjo	8	41	36	25	9,981	1980-2014	6
2	Anger	Nekemte	9	26	36	31	4,674	1995-2004	7
3	Dabana	Abasina	9	02	36	03	2,881	1980-1985	12
4	Didesa	Nr. Dembi	9	30	36	35	1806	1985-2014	7
5	Wama	Nr. Nekemte	8	47	36	47	844	1980-1985	39

129 Land use is amongst the primary factors influencing loss of soil from surface; likewise
130 evapotranspiration in the catchment. The land use being utilized in this study was that of MERIS
131 (Medium Determination Imaging Spectrometer) based Glob-Cover of 2009. land use map after
132 clipping it to study area and changed to relate for those swat predefined land use grouping. It holds
133 a raster version of the Glob-Cover map with spatial resolution of 30mx30m. Dominate land use or
134 cover for this manuscript was mosaic vegetation or crop lands followed by closed to open shrub land.

135

136

137 Simulation of SWAT necessitates soil composition of different properties like soil textural
 138 property, physical and chemical properties. The soil map utilized in this study was gotten from two
 139 sources. Firstly, the soil data base acquired from Ministry of Water Resource Irrigation and Energy
 140 of Ethiopian has shortage of several soil properties like (available moisture capacity, density,
 141 saturated hydraulic conductivity, percentage of sand, silt and clay) compulsory required in the
 142 model set up were not available in its data base. Secondly, due to the above data base was deficient
 143 in necessary information additional data were substantiated from another source like 'soil geo-
 144 database of Ethiopia' prepared by [11] with spatial resolution of 30mx30m. The soil data base of
 145 Ethiopia which was prepared by Belete Berhanu, contains all the information required for SWAT
 146 simulation. In addition, comparison of this soil database of the Ethiopia with the previously available
 147 FAO soil database (FAO, 1998) indicates that the soil database from the former has more detail
 148 classification than the FAO soil data base.



150 (a) (b) (c)
 151 **Figure 2.** Physiographic data: (a) Didessa sub basin land cover; (b) soil type; (c) DEM

152 2.3. Filling Missing Rainfall and temperature Data

153 Before using climate data for the SWAT software, filling of missed data were handled. XLSTAT
 154 software is used as filling of missing temperature and rainfall data. XLSTAT started in 1995 to make
 155 accessible to anybody a powerful, complete and convenient figures analysis and statistical result. The
 156 XLSTAT is easily available and compatible with all the Microsoft Excel versions that are acclimated
 157 these days, beginning from Excel 97 to Excel 2016, from the interface in numerous languages and
 158 downloaded from the XLSTAT website www.xlstat.com. Some of the application of XLSTAT includes
 159 operations like: completing of missing data using advanced missing value treatment techniques
 160 either take out observations with absent value, mean imputation method, nearest neighbor approach,
 161 NIPALS algorithm, MCMC multiple imputation algorithm. In this paper missing data is treated with
 162 "nearest neighbor approach" because the output of this method gives a hopeful result as the output
 163 treated value are also checked with those were filled with regression method.

164 2.4. Testing of dataset quality

165 Sometimes a significant change may occur in and around a particular rain gauge stations. Such
 166 change occurring in a particular year will start affecting the rain gauge data, being reported from a

167 particular station. In order to detect such inconsistency and to correct and adjust the reported rainfall
168 values, a technique called double mass curve method is generally adopted in this study. In this
169 method a group of 8 adjoining stations are selected in the vicinity of suspicious station. The mean
170 daily rainfall values are serially arranged in reverse chronological order to fix relative consistency.
171 The observations from a certain station were compared with the mean of observations from
172 numerous adjacent stations. In accepted double-mass computations, this testing involves removing
173 from the arrangement the records from an uncertain station and comparing them with the remaining
174 data. Since all the dataset were reliable with the accepted totalities in the area, they are re-combined
175 into the base period station. After the data of each station are arranged in descending order, the
176 accumulative sums, station to be investigated and base station; are plotted against each other and
177 line of best fit were sketched in excel assignment sheet.

178 Seasonal Mann-Kendall Test is adapted to evaluate with a nonparametric test if a trend can be
179 recognized in a series, even when seasonal factor in the sequence. A nonparametric trend test has
180 been primarily suggested via [26] then advanced through [24] finally get enhanced by way of [25]
181 who accustomed to take into account seasonality as well.

182 The null hypothesis H_0 for these tests implies absence of trend in the series. The next three
183 hypotheses indicate presence of non-null, negative, or positive trend. This test depends on values P-
184 value and Kendall's tau. Kendall's tau shows degree of relationship between two samples. P-value
185 measures whether the null hypothesis was accepted or rejected. If the p-value falls below significance
186 level the alternative hypotheses will be accepted and vice versa. If the time series does have a trend,
187 the data cannot be used for frequency analyses or modelling. Those time series with trend cannot
188 incorporate into hydrological or frequency analysis during modeling for hydraulic structures
189 designs.

190 2.5. Selection of parameters for Sensitivity Analysis

191 Before calibration to begin, Parameters those were used in SWAT model to other Upper Blue
192 Nile sub basins were identified from previously published manuscript. Since this not enough to get
193 performance criteria, other parameters were gathered and added from SWAT-CUP manual. About
194 nineteen parameters (CN2, ESCO, SOL_AWC, GW_DELAY, GW_REVAP, REVAPMN, GWQMN,
195 ALPHA_BF, RCHRG_DP, CH_K2, SURLAG, CH_N2, and SOL_K, CH_K2, ALPHA_BNK,
196 SLSUBBSN, OV_N, LAT_TIME, ESCO, EPCO, and HRU_SLP) were incorporated into SWAT-CUP
197 algorithm (Sufi-2) to understand the level their sensitivity. Those parameters along with their
198 sensitivity level were described in (table 6). For the reason that knowing of the more sensitive
199 parameters could make ease the time required for calibration and validation. Furthermore, it is the
200 technique to know the dominant parameters of the watershed those can influence the hydrological
201 balance of the sub basin. The global sensitivity were determined depend on P-value. The smaller the
202 p-value indicates the more sensitive parameter, whereas the larger the p-value point toward the less
203 sensitive for the given watershed[49]. The values close to zero has more significance. According to
204 [30], also sensitivity analysis significantly eases relative sensitivity of parameters identification, rises
205 the accurateness of calibration and lessen uncertainty and the time necessary for it.

206 2.6. Data Processing and Model setup

207 Data processing in this case includes trend test and homogeneity tests for precipitation data of
208 8 stations in Didesa sub basin from 1980 to 2014. Moreover, flow data of the sub basin is also tested

209 for Arjo gauging station from year 1980 to 2014, Dembi station from 1985 to 2014, Angar stations from
 210 1995 to 2004, for Dabana stations from 1982 to 1985 depending availability of flow data. The seasonal
 211 Kendall's tau and p-values for each station is evaluated with XLSTAT software. As p-values are
 212 greater than significance level (alpha), the null hypothesis accepted which shows that the data of all
 213 the stations are free of trend (table 3).

214 **Table 3.** Seasonal Mann-Kendall Test for monthly rainfall of Weather gauging stations

Station name	Kendall's tau	p-value(Two-tailed)	alpha
JIMMA	0.041	0.281	0.05
BEDELE	-0.02	0.93	0.05
NEKEMTE	-0.00138	0.986	0.05
DIDESA	-0.0470	0.228	0.05
GIMBI	0.0451	0.216	0.05
NEJO	0.038	0.268	0.05
ARJO	0.0143	0.694	0.05
SHAMBU	0.057	0.126	0.05

215 Note: p-values are greater than significance level (alpha) i.e. 5%; hence data of all stations are consistent

216 Alexanderson's SNHT (Standard Normal Homogeneity Test) test for Homogeneity is applied
 217 for testing of monthly rainfall. This test i.e. SNHT was established through [48] in order to sense an
 218 alternation in a sequences of precipitation data. The test was recommended to series of ratio of
 219 observations to compare with average of ratio of several stations. After processing of data the output
 220 result shows that the series of precipitation data were remain homogenous. The same procedure was
 221 followed to trend test, in trend test interface of XLSTAT and it is found that data are free of trend (see
 222 table 4 and figure 3).

223 **Table 4.** Alexanderson's SNHT test output value for Homogeneity of rainfall data

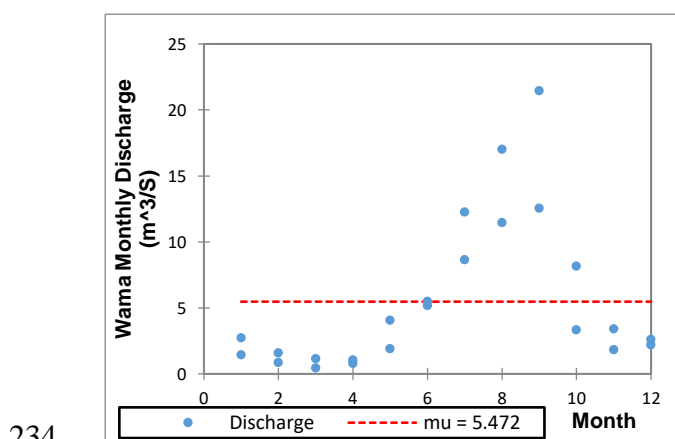
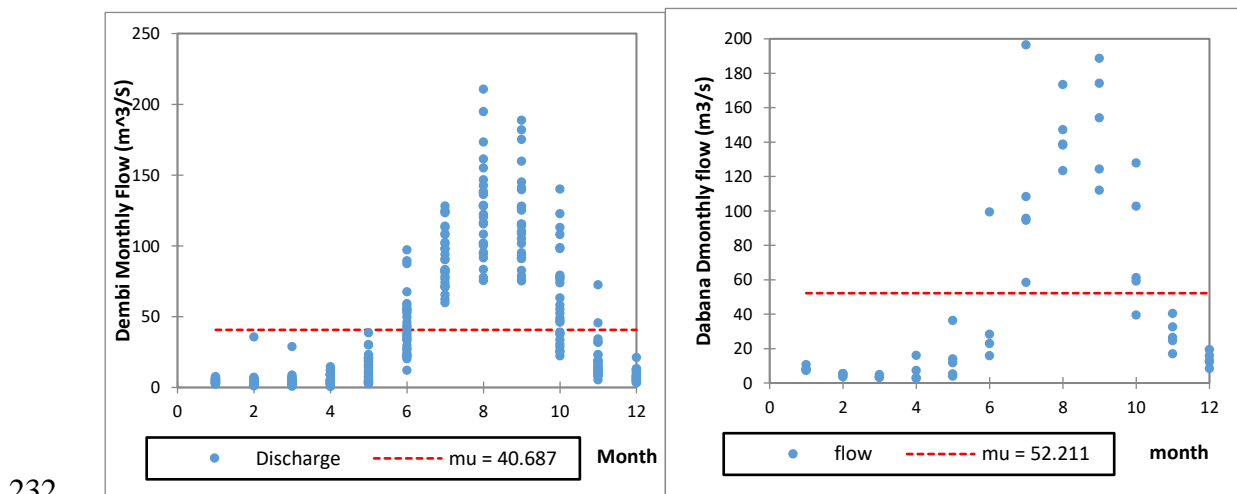
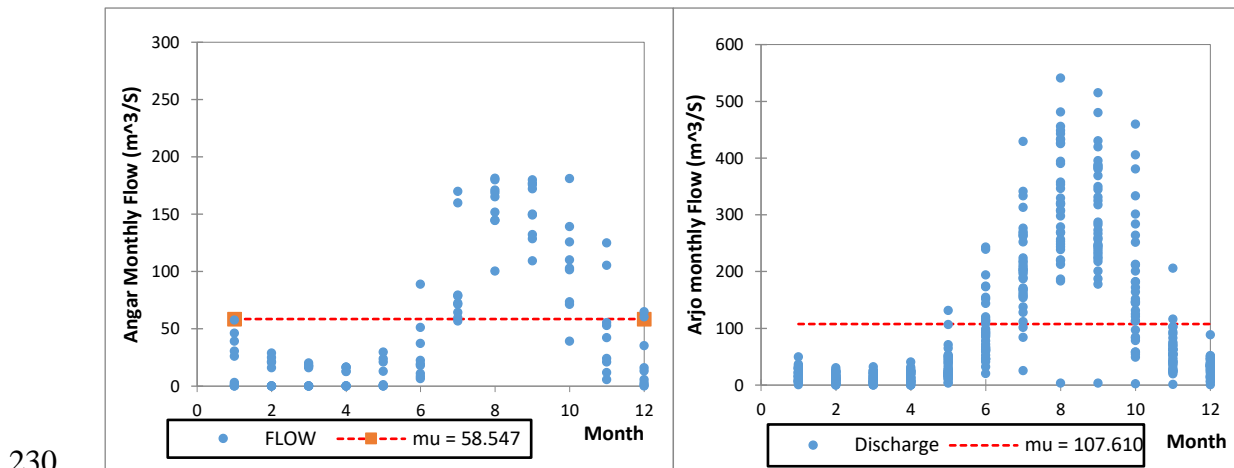
Station Name	Kendall's tau(τ_0)	p-value (Two-tailed)	alpha
JIMMA	7.7	0.146	0.05
BEDELE	2.5	0.93	0.05
NEKEMTE	4.8	0.5	0.05
DIDESA	3.6	0.72	0.05
GIMBI	3.3	0.8	0.05
NEJO	2.81	0.89	0.05
ARJO	2.2	0.96	0.05
SHAMBU	4.1	0.63	0.05

224 As shown in table all stations p-values are greater than alpha (5%), hence data are homogenous.

225

226 Similarly flow series were tested at different gauging stations of Didessa sub basin such as
 227 Didessa near Arjo (1980-2014), Dabana near Abasina (1980-1984), Didessa Near Dembi (1985-2014)

228 and Angar Near Nekemte (1995-2004). And from trend and homogeneity test it is found that the flow
 229 data were homogeneous and no significant trend is found at all discharge gauging stations.



From the graph

235 mu (horizontal red line) indicate the mean of flow data)

236 **Figure3.** XLSTAT final output Graphical representation of homogeneity test of Observed Flow data

237 ArcSWAT version of 2012.10 of ArcGIS 10.1 interface downloaded from website
 238 <http://swat.tamu.edu/software/arcsWat/> is used for watershed delination, HRUs definition and
 239 hydrological simulation. DEM of spatial resolution of 30mx30m was applied in watershed delination.

240 And land use and soil shape file of the same spatial resolution was used in HRUs definitions.
 241 Landuse/soil/slope of thresholds 20/20/20 (%) respectively, produces 604 HRUs and 112 subbains.
 242 Recorded wheather input of daily rainfall and maximum and minimum temperature of 35 years used
 243 as input file to produce the simulation. And three years warm up period is took place to activate the
 244 Swat run step. Finally, the sequencial Uncertainty Fitting i.e SUFI-2 built in SWAT-CUP algorithm
 245 was used to calibrate and validate the model. SWAT-CUP provides algorithms for auto-calibration,
 246 from which Sequential Uncertainty Fitting, Version 2 (SUFI-2) was chosen. SUFI-2 accounts for several
 247 sources of uncertainties such as uncertainty in driving variables e.g. rainfall, conceptual model,
 248 parameters and measured data [49]. It is not a fully automated calibration tool, since it still requires
 249 interaction of the modeller and knowledge about the parameters and their effects on the output [49].
 250 In sufi-2 parameters are given ranges as found in Absolute_swat_values before performing iteration.
 251 Finally, the aligorithm provides the best estimation and optimum parameters range.

252 Model performances were evaluated graphically and statistical procedures with that of quality
 253 criteria. The statistical parameters employed in measure of model quality are Root Mean square
 254 (RMS), coefficient of determination (R^2), Nash Coefficient (NSE)[37] and PBIAS measure the model
 255 quantitatively. In this manuscript model performance was evaluated with values of R^2 , NSE and
 256 PBISAS. The R^2 provides for those extents of the discrepancy between observed and simulated with
 257 linear association using linear regression model. Coefficient of determination values found between
 258 zero and one, which the smaller figure demonstrating more error variation. The figure value more
 259 than 0.5 found in satisfactory range for coefficient of determination [42, 43] and NSE as well. The
 260 Nash Sutcliffe defines the extents of variation between simulated to observed data discrepancy [38].
 261 The Nash value situated between this interval $(-\infty, 1]$ and the value close to zero shows as the model
 262 performance is more suitable [38,39]. PBIAS is taken as a clear quantifier for water balance
 263 errors[40,41] and value close to zero shows the more the value approaches to the acceptable range,
 264 positive number explains under estimation of model and negative illustrates model over estimation.
 265 Model was desired to evaluated based quality criteria as it can give water balance error which can
 266 show that poor model performance[40,41]. For stream flow, the performance rating which ranges
 267 between $0.75 < NSE < 1.00$ and $PBIAS < \pm 10$ is considered as very good for monthly time scale. The
 268 model performance is supposed as good for values ranges between $0.65 < NSE < 0.75$ and $\pm 10 < PBIAS$
 269 $< \pm 15$. Values of $NSE < 0.50$ and $PBIAS > \pm 25$ demonstrates unsatisfactory ranges of performance. The
 270 model performance was considered as satisfactory for an interval value ranges between $0.50 < NSE$
 271 < 0.65 and $\pm 15 < PBIAS < \pm 25$ [40]

272 The R^2 , NSE and PBIAS were evaluated with the equations (1) up to (3) as follows:

273

$$274 \quad NSE = 1 - \frac{\sum_{i=1}^n (Q_i^o - Q_i^s)^2}{\sum_{i=1}^n (Q_i^m - \bar{Q}^m)^2}, \quad (1)$$

$$275 \quad R^2 = \frac{[\sum_{i=1}^n (Q_{o,i} - \bar{Q}_o)(Q_{s,i} - \bar{Q}_s)]^2}{\sum_{i=1}^n (Q_{o,i} - \bar{Q}_{mo})^2 (Q_{s,i} - \bar{Q}_s)^2}, \quad (2)$$

$$276 \quad PBIAS = 100 * \frac{\sum_{i=1}^n (Q_o - Q_s)i}{\sum_{i=1}^n Q_{o,i}}, \quad (3)$$

277 Where, Q_i^o and Q_i^s represent measured and Simulated flow at time step i , correspondingly,
 278 $\overline{Q^o}$ and $\overline{Q^s}$ are average of measured flow, n indicate grand number of paired measured and observed
 279 discharge, o and S are mean measured and simulated discharge, consecutively.

280 3. Results

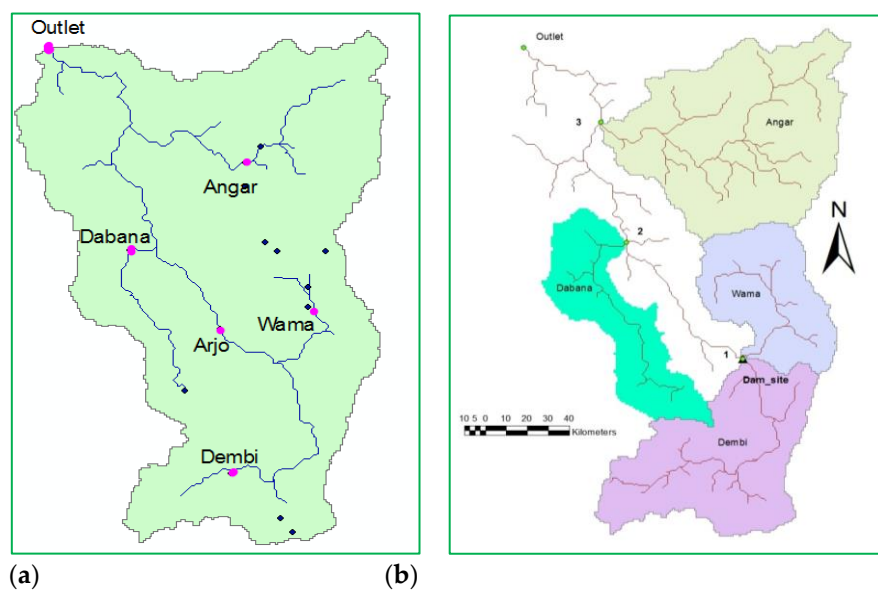
281 3.1. Model Calibration and validation

282 Sensitive parameters identification for the sub basin is alternatively required before starting with
 283 calibration as well as validation of SWAT to simplify the time consumed in the project work. This
 284 procedure involves the determination of sensitive parameters depending on professional or skilled
 285 decision [27]. It was vital to identify crucial parameters those needed for the next task i.e. calibration.
 286 Maximum of the previously published Soil and water Assessment Tool scholarships involves
 287 respective activities of calibration and validation [27] whereas others do calibration without
 288 validation due to scarcity of measured data. For this study, flow gauging stations like: Dabana near
 289 Abasina and Wama near Nekemte left only with calibration as limited observed discharge of less
 290 than 5 years are available after 1980 (starting year of SWAT model run).

291 Initially, calibration was started on monthly base to recognize size of parameters, as well as
 292 seasonal characteristics of flow. After running of SWAT model on monthly time step hydrological
 293 water balance is observed and base flow is overestimated. Through calibration in which the values
 294 parameters range repeatedly changed during this activity the best promising statistical performance
 295 indicator value were attained. Then the daily calibration is succeeded to get more perfect parameter
 296 values about watershed as well as to properly estimate annual flow volumes at necessary point or
 297 junction as shown in figure 4.

298 Parameter values are adjusted to fit observed and simulated hydrograph. To decrease model
 299 overestimated estimated Base flow, the parameters such as (GW_DEEP), (GWQMN), (GW_REVAP),
 300 (REVAPMN) were getting increase to give best correlation between measured and simulated
 301 discharge.

302



303
 304
 305
 306

Figure 4. (a) location of flow-gauging stations, and (b) Junction(1,2,3) at which hydrograph is simulated

307 Model calibration and validation is done depending on location of gauging stations as well as
 308 presence of available gauged discharge data. For Dembi and Arjo daily flow data of 18 consecutive
 309 years with negligible missing is used for calibration and validation. Angar near Nekemte has 10 years
 310 available data within the range of SWAT run period where as Lower Dabana and Wama have 5 years
 311 and 4 years data within the range respectively. Hence due to absence sufficient data calibration is
 312 done without validation at Dabana and Wama discharge gauging stations. Even though model
 313 performance values are not shown in table 5 for these two gauging stations model is calibrated using
 314 observed flow. The calibration statistical value is found in the same table. Didessa near Arjo
 315 gauging station data are classified in to two depending on hydrological change during 2005 in
 316 Didessa sub basin. This change may be due to settlers of Hararghe population, land clearing for sugar
 317 cane production and etc. Flow data from 1997 to 2004 and from 2007 to 2014 are used for this
 318 watershed to understand the basin characteristics within the two range.

319 **Table 5.** Model performance statistics for the Didessa Sub basin at 5 discharge gauging stations.

Time base	Calibration	(1997-2008)	(1997-2001)	(2006-2011)	(1997-2001)
	Criterion	Dembi	Arjo	Arjo	Angar
Daily	R ²	0.66	0.74	0.74	0.8
	NSE	0.6	0.74	0.65	0.8
	PBIAS	-10.5	-6.1	-19.4	0.5
	Validation	(2009-2014)	(2002-2004)	(2012-2014)	(2002-2004)
	R ²	0.70	0.70	0.64	0.62
	NSE	0.66	0.62	0.6	0.61
Monthly	PBIAS	-6	0.0	-4	-16.3
	Calibration	(1997-2008)	(1997-2001)	(2006-2011)	(1997-2001)
	R ²	0.80	0.87	0.75	0.82
	NSE	0.72	0.82	0.71	0.79
	PBIAS	-16.4	-10.5	2.5	-3.5
	Validation	(2009-2014)	(2002-2004)	(2012-2014)	(2002-2004)
R ²	0.86	0.89	0.82	0.92	
NSE	0.83	0.84	0.81	0.91	
PBIAS	-13.3	-20	7.7	-13.4	

320 In both calibration and validation this model shows acceptable statistical values of performance
 321 measurements on the Didessa sub basin at all the gauging stations. It produced acceptable results in
 322 terms of NSE, PBIAS and R² whether on daily or monthly time steps. In both daily and monthly,
 323 model calibration and validation returned a NSE ≥ 0.6 which shows that the figure is acceptable. In
 324 most stations daily performance are greater than that of monthly in terms of PBIAS. Specifically, in
 325 some of the stations very good values of NSE and R² were obtained and were >0.75 on monthly time
 326 scale. Likewise, it might be recognized that the model performance is slightly low on daily time step
 327 than monthly as it was indicated in table 5. Only one station i.e. Arjo showed low performance in the
 328 case of percent of biased (PBIAS) of -20 even though NS and R² are very good at monthly time basis.
 329 So as daily calibration and validation demonstrates lower percent of biased, for water balance and
 330 flow hydrograph simulation and prediction daily was used. The calibrated and validated result
 331 shows that, for all stations; higher percent of biased was found in daily time base. Furthermore,
 332 since average of data is simulated on monthly time step, it is not good as daily time step in water
 333 balance prediction at necessary junction to estimate flow for design purpose. The SWAT parameters
 334 used in in calibration, its optimum value and variation methods are indicated in table 6 below. For
 335 all the watersheds flow is well reproduced for both wet and dry season. The graphical representation

336 of Calibration and validation done in sufi-2 of SWAT-CUP algorithm shows that the peak and dry
 337 season flows were well developed as indicated in figure 5 and figure 6 below on monthly time scale
 338 and daily time scale respectively.

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341

342 **Table 6** Parameters used in ncalibration of SWAT model and its optimum values

Parameters	Fitted Values							Variation Methods
	Range	Dembi	Arjo(1997-2004)	Arjo(2006-2014)	Angar	Dabana	Wama	
V_ALPHA_BF	0-1	0.45	0.357	0.511	0.721	0.741	0.545	Replacement
V_ALPHA_BNK	-0.1-1	-0.01	0.745	0.327	0.457	0.175	0.755	Replacement
V_CH_K2	0-500	125.68	488.78	309.204	285.0	451.31	486.2	Replacement
V_CH_N2	-0.01-0.3	0.273	0.191	0.056	0.389	0.276	0.213	Replacement
R_CN2	-0.25-0.25	0.11	0.124	0.15	0.107	0.153	0.1	Relative
V_EPCO	0-0.9	0.062	0.529	0.213	0.042	0.11	0.103	Replacement
V_ESCO	0-0.1	0.002	0.001	0.003	0.0022	0.0023	0.0021	Replacement
V_GW_DELAY	0-0.3	0.01	0.604	0.234	-0.001	0.01	0.005	Replacement
V_GW_REVAP	0-0.2	0.199	0.183	0.17	0.181	0.177	0.179	Replacement
V_GWQMN	0-5000	3044.4	4850	4616.7	1683.2	3705.5	4596.8	Replacement
R_HRU_SLP	-1-0.4	-0.871	0.784	-0.112	0.287	0.196	-0.665	Relative
V_LAT_TTIME	0-25	5.934	12.643	28.4	4.957	5.217	1.141	Replacement
V_OV_N	0-15	6.237	7.426	5.776	1.502	1.577	6.718	Replacement
V_RCHRG_DP	0-1	0.258	0.222	0.124	0.422	0.265	0.832	Replacement
V_REVAPMN	0-500	147.78	299.45	125.7	149.06	303.0	485.9	Replacement
V_SLSUBBSN	0-180	123.12	87.891	169.2	57.057	123.0	91.9	Replacement
R_SOL_AWC	-02-0.51	-0.049	-0.101	0.025	0.356	0.052	0.077	Relative
R_SOL_K	-0.1-0.25	-0.136	0.009	0.121	0.166	0.038	-0.02	Relative
V_SURLAG	0-30	16.74	20	10.4	4.711	6.258	12.5	Replacement

"V" replaces the existing value with the given value,"R" multiplies the existing value with (1+the given value)

343

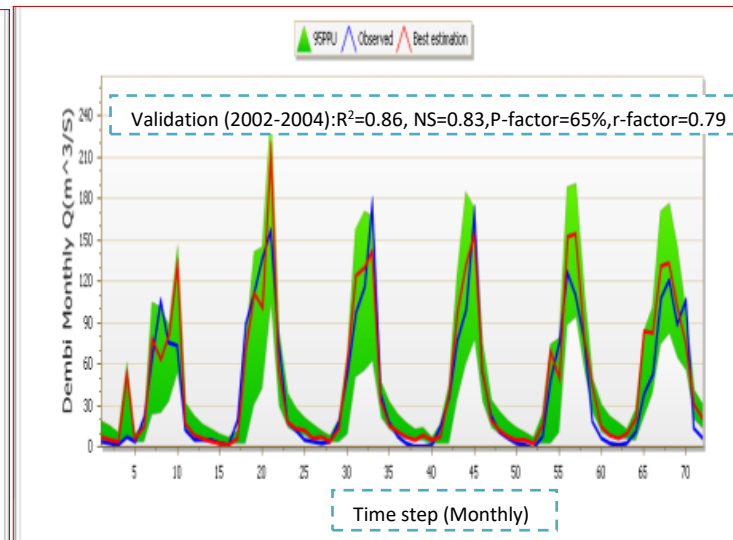
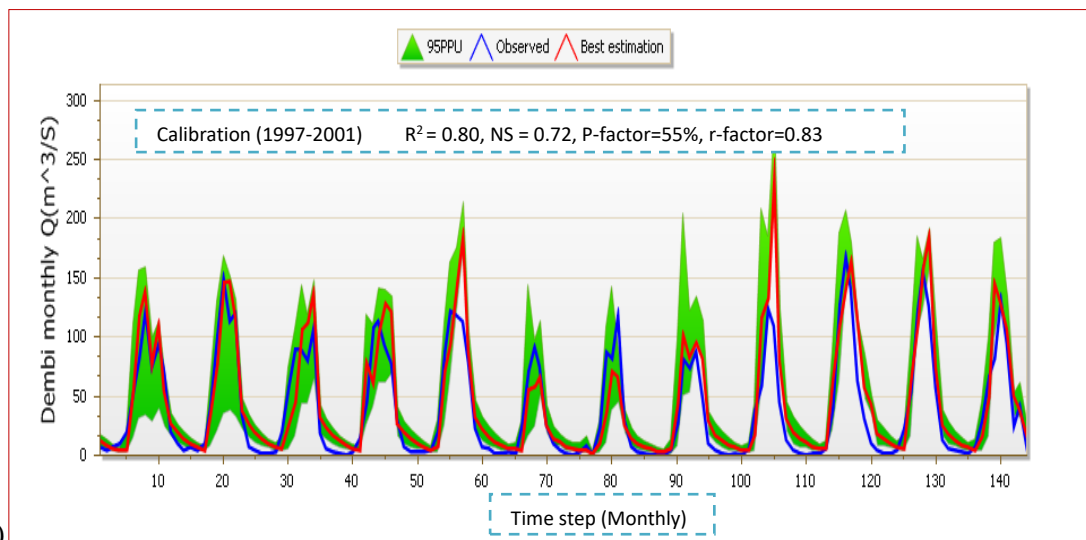
344 3.2. Global sensitivity Analysis

345 The sensitivity of parameters was computed to simplify the time consumed in model calibration
 346 and validation. Identification of sensitivity level of those parameters of the sub basin is also crucial
 347 to understand their effect on watershed to produce surface runoff. The ranks of sensitivity, which
 348 depends on p-value obtained during iteration in SWAT-CUP algorithm of Sufi-2 were tabulated in
 349 (Table A1) for gauging stations involved in this study. The parameter's rank given in the table was
 350 obtained during calibration on daily time step. Some of the sensitive parameters are similar to those
 351 found by [28]sensitive parameters for Abbay basin. Among nineteen parameters selected for model
 352 calibration the CN₂ (curve number), (SOL_AWC(available water capacity), HRU_SLP(average slope
 353 steepness), (SOL_K)saturated hydraulic conductivity, GW_REVAP(ground water reevaporation
 354 coefficient),GWQMN (threshold water depth in the shallow aquifer for flow), and ALPHA_BF (base
 355 flow alpha factor) were found to be those with higher rank at different outlet.

356 New sensitive parameters those can affect flow are Manning's "n" value for the main channel
357 (CH_N2), LAT_TIME Lateral flow travel time, and OV_N (Manning's roughness coefficient) for
358 overland flow were found. The level of parameter's sensitivity was different at different outlets.
359 Nevertheless the CN₂(curve number) was the main sensitivity parameter for all outlets.
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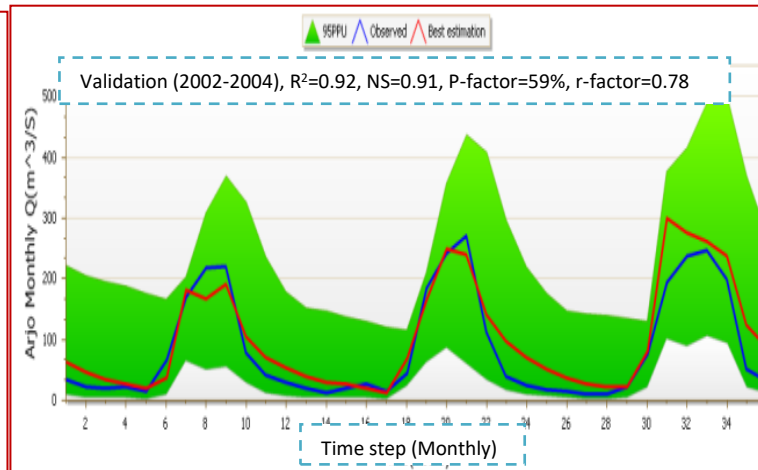
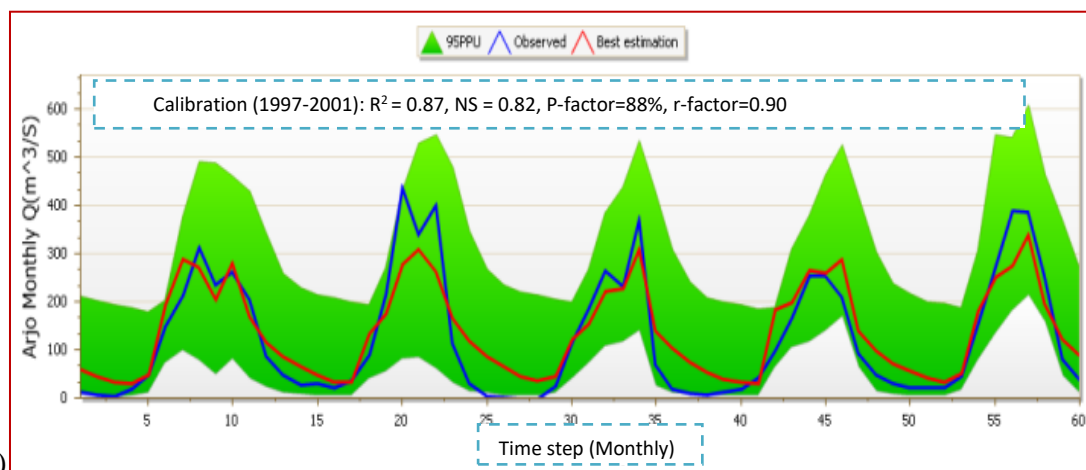
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(a)



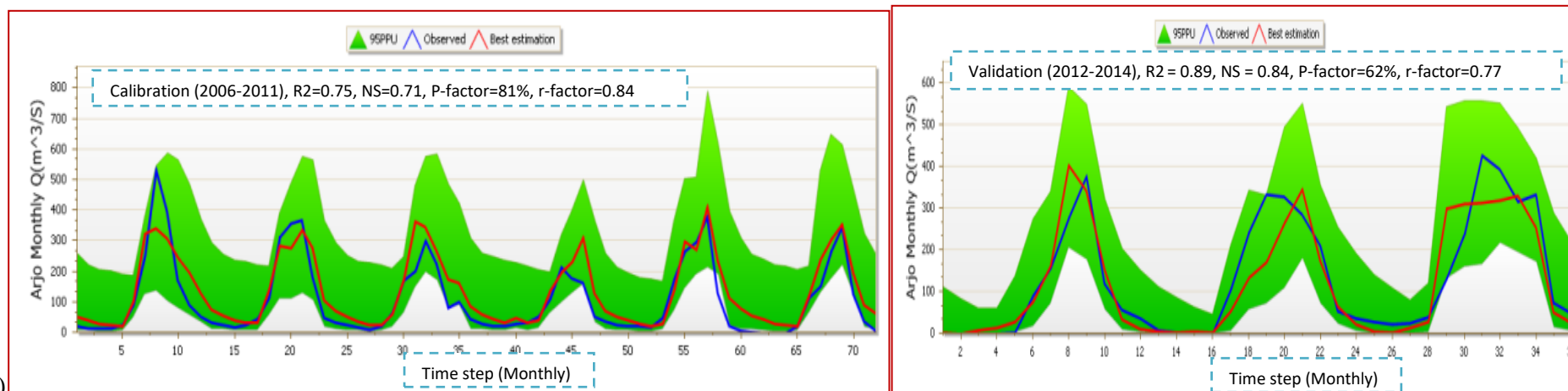
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(b)



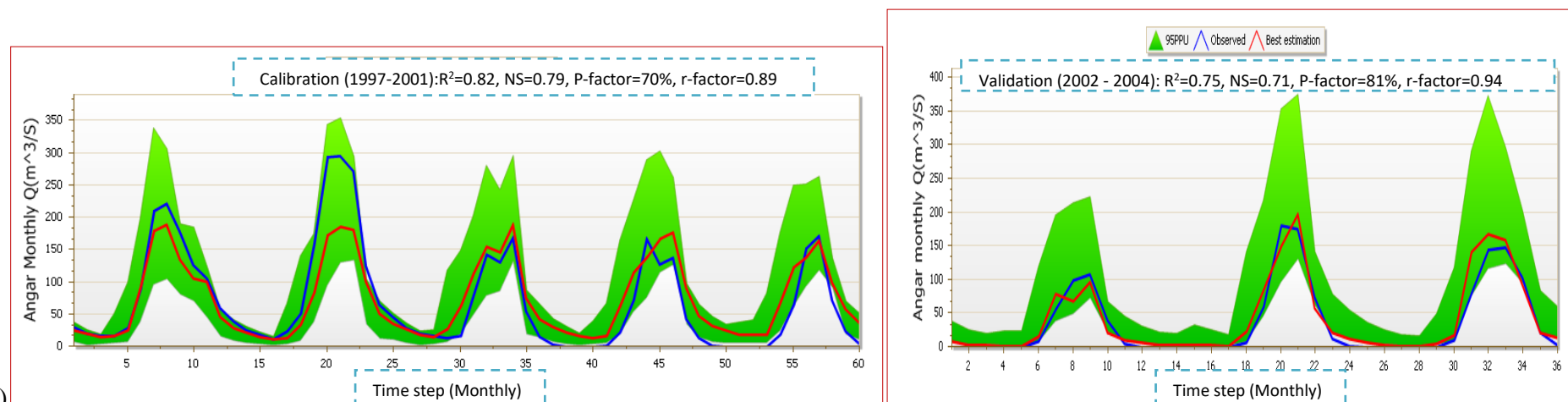
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(c)



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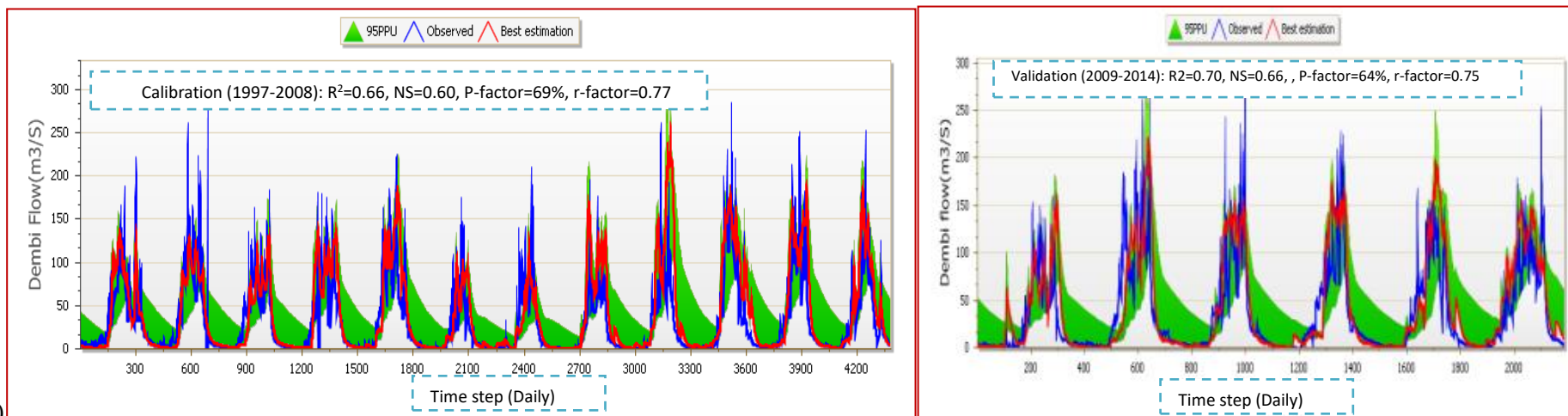


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Figure 5. Simulated and observed hydrographs at the 3 flow gauging station (a) Dembi, (b, c) validation and, (d) Angar at monthly time steps.

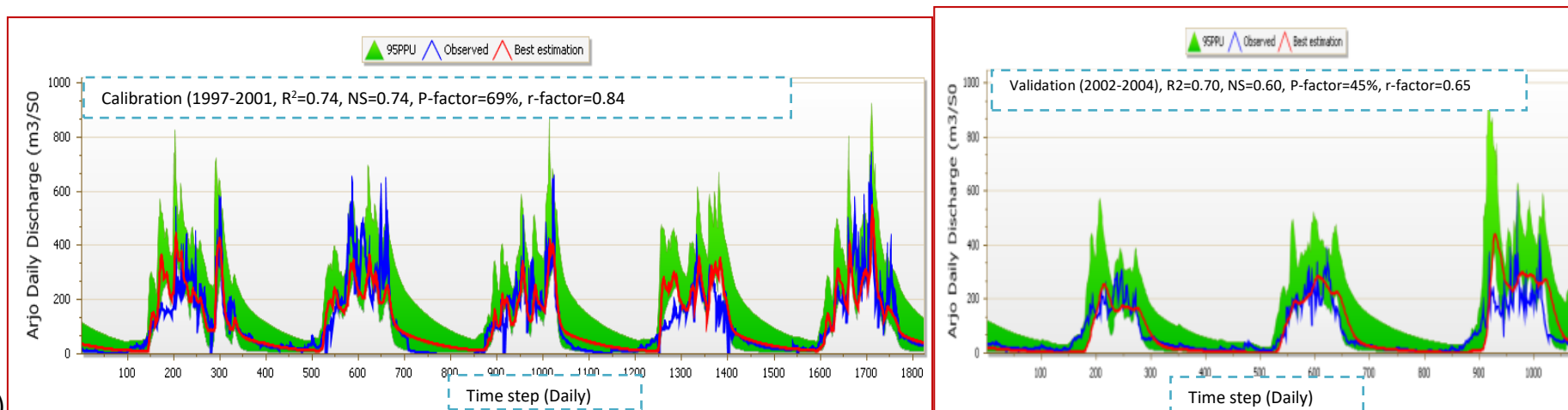
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(a)



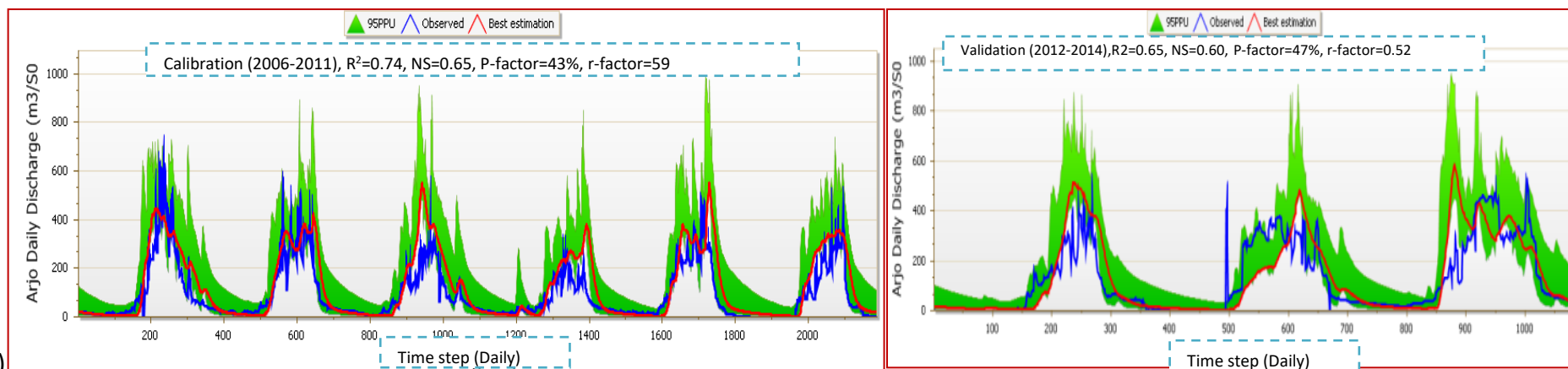
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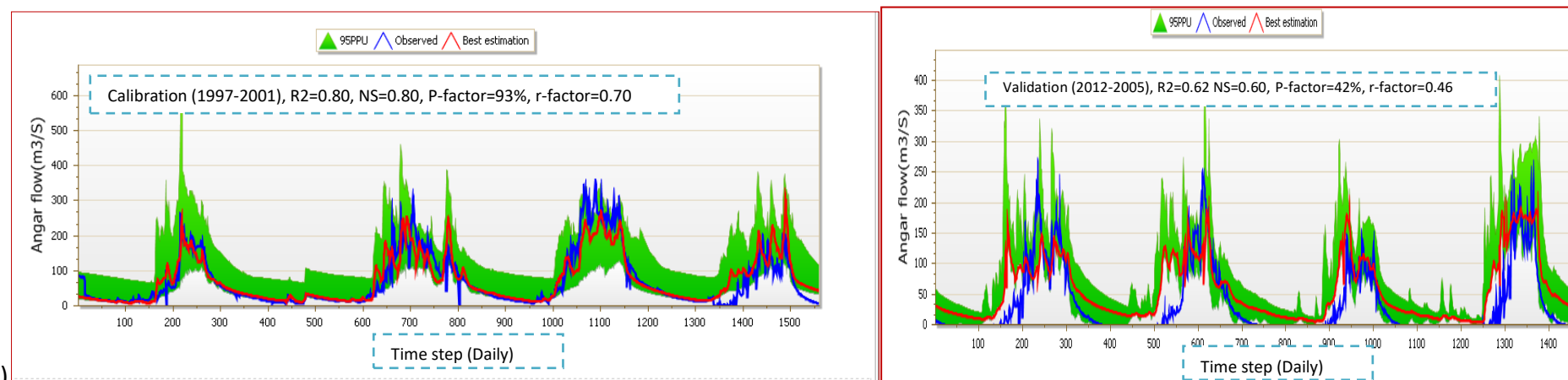
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(c)



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(d)



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Figure 6. Simulated and observed hydrographs at the 3 flow gauging station (a) Dembi,(b, c) Arjo,and (d) Angar at daily time steps

371 After running of SWAT model hydrological water balance is observed and in fact base flow is
 372 overvalued. As it can be seen in table 5 the negative values of PBIAS indicated that over estimating
 373 of flow by the model. This may be due to the reason that the soil shape file applied in this model set
 374 up is single layered soil data. During iteration in in SWAT-CUP algorithm i.e. Sufi-2 calibration
 375 parameters are adjusted to fit observed and simulated hydrograph. Changing of calibration
 376 parameters have significant influence on hydrological component. For this reason, in this study to
 377 decrease Base flow, increase deep percolation (GW_DEEP), Threshold depth of water in the shallow
 378 aquifer required for return flow to occur (GWQMN), increase the groundwater revap-coefficient
 379 (GW_REVAP), and increase the threshold depth of water in shallow aquifer (REVAPMN) was done.
 380 To correct the late shift, the slope (HRU_SLP) increased, and Manning's roughness coefficient
 381 (OV_N) as well as the value of overland flow rate (SLSUBBSN) decreased. SCS runoff curve number
 382 (CN₂) value is getting increased to increase the value of surface run off. Finally, Deep aquifer
 383 percolation fraction (RCHRG_DP) value is also getting larger than that of during Swat run to increase
 384 the amount of water percolated to deep aquifer.

385 3.3. Water Balance of Didessa Sub Basin

386 Water balance analysis of the sub basin is done with a given land use. In addition the calibrated
 387 parameters are reinserted during ArcSWAT run for each watershed to get appropriate balance. This
 388 sort of evaluation has been required to acquire an Understanding of the whole hydrological response
 389 of catchment. At the same time it provides for an essential consideration of the rainfall-runoff
 390 association through a long period of time. The outcome of such investigation considers the general
 391 breakdown of precipitation and their proportions which defining runoff from sub basin, subsurface
 392 and evapotranspiration, and etc. This evaluation involves comparison of input climate data to that of
 393 observed stream flow. The breaking of inputs to the output water balance components could aid in
 394 deciding those possibility sensitivities of the watershed to change in land use or land cover. The
 395 general long term hydrological water balance and hydrological parameters estimated and tabulated
 396 in table 7.

397 **Table 7.**Hydrological water balance ratio and hydrological variables

Hydrology (water balance ratio)	Dembi	Arjo	Dabana	Wama	Angar	L/Didesa
Stream flow/precipitation	0.21	0.2	0.28	0.15	0.20	0.22
Base flow/total flow	0.21	0.28	0.27	0.32	0.39	0.4
Surface run-off/total flow	0.79	0.72	0.73	0.68	0.61	0.60
Percolation/precipitation	0.28	0.34	0.29	0.36	0.35	0.30
Deep recharge/precipitation	0.07	0.13	0.08	0.3	0.15	0.11
ET/precipitation	0.55	0.51	0.50	0.48	0.52	0.52
Hydrological variables (all units in mm)	Dembi	Arjo	Dabana	Wama	Angar	L/Didesa
Evap. and transpiration	964.5	951.7	932.8	961.5	860.3	960.0
Precipitation	1740	1862.9	1861.7	1990.2	1669.9	1850.6
Surface run-off	287.4	264.71	376.41	209.1	204.5	245.8
Lateral flow	3.37	5.2	13.42	98.92	17.66	91.1
Return flow	74.9	95.24	124.6	0.0	115.14	71.9

398 All Units are in (mm)

399 3.4. Flow of major tributaries

400 Average monthly basin rainfall, Evapo-transpiration, surface flow, Potential Evapo-
 401 transpiration and average basin yield is obtained From ArcSWAT output. From this out put one can
 402 understand hydrological situation of the basin in terms of months with high and minimum surface
 403 and base flow. To easily understand the hydrological situation of the sub basin, it is better to classify
 404 the output into two major categories i.e. wet and dry season which is common in Ethiopia. According
 405 to climate of the study area the wet season ranges from May to October where as dry season is when
 406 precipitation is almost negligible and for all months between November and April. Peak flow and
 407 lowermost flow is developed during the month of August and April respectively.

408 The study also extended to evaluate the water yield of the major tributaries to identify the
 409 catchment which contributes maximum annual flow along with their catchment area. The four major
 410 tributaries of Didessa sub basin are: Angar, Dembi or Toba, Wama, and Dabana in their consecutive
 411 order of catchment size.

412 Angar catchment is the largest in case of catchment size (7988.2 Km²), receives annual
 413 precipitation of 1670 mm and takes share water yield of about 4.46×10^9 m³ to Didessa sub basin and
 414 then flows to the Upper Blue Nile River from simulated daily flow.

415 The fourth rank in terms of catchment size (3246.4 km²), tributary of Didessa sub basin is Dabana
 416 catchment. The catchment can receive simulated annual rainfall of 1861.7 mm, and contribute flow
 417 of 2.15×10^9 m³ annually to the upper Blue Nile of Ethiopia.

418 The next tributary of Didessa sub basin is called Wama catchment take third rank in terms of
 419 catchment size (3336.8km²). Wama catchment receives an annual rainfall amount of 1990.2 mm and
 420 donates 2.71×10^9 m³ of annual flow to the upper Blue Nile of Ethiopia.

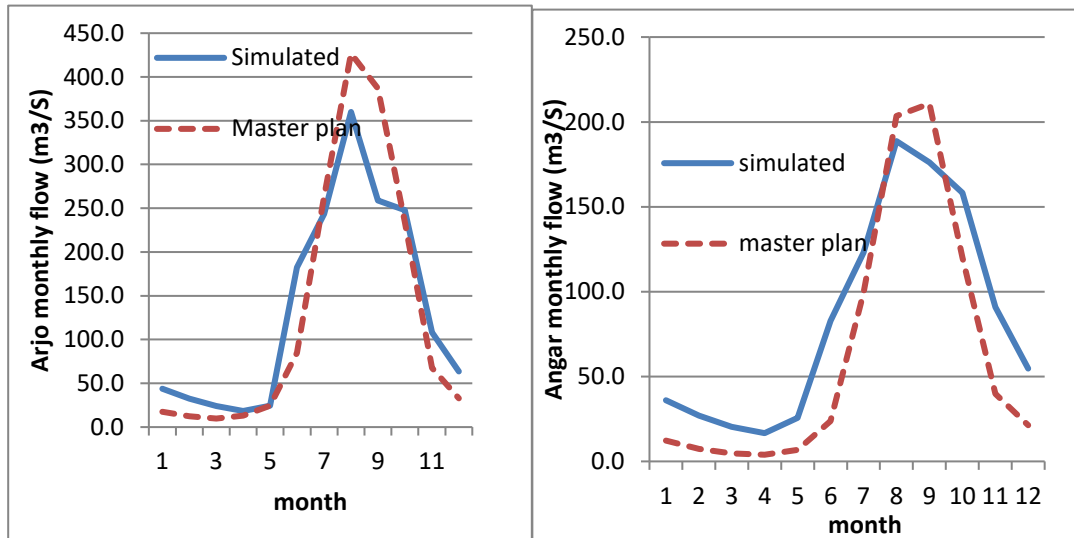
421 The second rank in catchment size (5532.1 km²), and in which the construction of embankment
 422 dam is in progress is Dembi tributary. Dembi receives an annual rainfall of about 1740 mm, and
 423 contribute 3.1×10^9 m³ flow to the upper Blue Nile or Didessa sub basin every year.

424 Generally, Didessa sub basin contributes about quarter of average flow of Abay basin of
 425 Ethiopia. According to simulated output flow volume of about 10.7 Billion meter cube of flow is
 426 annually donated to the Upper Blue Nile of Ethiopia from Didessa sub basin. In percentage it shares
 427 about 26% of Abay basin (54.81×10^9 m³) which is measured at Sudan border [31]. The results are
 428 summarized in table 8.

429 **Table 8.** Average yearly water budget of main tributaries

S.No	Name. of Tributary/ Reach	Precipit ation (mm)	Mean annaul Flow (m ³ /s)	SURQ (mm)	PET (mm)	ET (mm)	Total Annual Water Yield	
							Depth (mm)	Volume (10 ⁹ m ³)
1	Dembi	1740	84.5	287.4	1404.1	964.1	482.6	3.1
2	Dabana	1861.7	67.6	376.4	1456.7	932.4	656.4	2.15
3	Wama	1990.2	86.1	209.1	1195.5	961.0	902.3	2.71
4	Angar	1669.9	147.2	204.4	1359.6	860.2	583	4.6

430 Comparing simulated average monthly flow with Abbay master plan study of 1999, the SWAT
 431 model underestimates the monthly flow during peak for Arjo Didesa and Angar gauging station. The
 432 small difference may be the model used in estimation of flow and the data range used in Master plan
 433 study (1962 to 1992 for Didessa at Arjo, and 1960 to 1992 for Angar at Nekemte) in monthly estimation
 434 as it is shown in figure 8.



435

436

(a)

(b)

437 **Figure 7.** Comparison of simulated vs Abay basin master plan study at gauging stations (a) Arjo Didesa, (b) Angar

438

3.5. Frequency Analysis

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The foremost purpose of frequency analysis to know the relationship between extreme flood with their occurrence through statistical procedures[33]. The end result of this analysis is used in design of hydraulic structures; especially, dams, culverts, diversion structures and etc. frequency analysis are evaluated from simulated flow hydrograph of each major tributary, and can be calculated by using the following formula.

444

$$x_T = \bar{x} + K_T * s , \quad (4)$$

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Where, \bar{x} peak flow mean, s =standard deviation of peak flow and, K_T , magnitude of extreme event at time at return period T.

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The frequency distribution parameters and the distribution type were determined in EasyFit software. EasyFit, which is a data analysis and simulation software was used to fit and simulate statistical distributions of flow data, choose the best model, and use the obtained result of analysis to take better decisions. This analysis method is preferred in cases which one has little or no information about base distributions existing in data, and want to find the general type of distribution [53]. The distribution with the highest frequency was chosen as the best distribution for a particular watershed. [52] The purpose of recognizing distributions for data in essence increases the validity of models, which, in turn, leads to better decisions. The statistical parameters determined by Easyfit like mean, standard deviation, skewnes coefficient are tabulated in Table 9.

456

457

EasyFit shows that all the watershed fit normal and Lognormal model except Dembi and Wama which w found to give good fit with Logpearson-3 and pearson-5 respectively. The values of

458 parameters such as mean (μ), standard deviation (σ) were calculated for each watershed and the result
459 found in Easy fit test was tabulated in (table 9).

460 To decide which distribution is suitable for each major tributaries, the goodness of fit tests,
461 including Kolmogorov- Smirnov (KS), Anderson-Darling (AD), and Chi-square for all the data sets
462 were done for all the watershed at outlet point with 12 years daily data and parameters distribution
463 parameters are found in table 9. Anderson-Darling test is much more sensitive to the tails of
464 distribution, whereas Kolmogorov-Smirnov test is more aware of the center of distribution. To sum
465 up, Anderson-Darling would recommend to use, to get much more powerful test. The Kolmogorov-
466 Smirnov and Anderson-Darling tests were limited to continuous distributions.

467 **Table 9.** Distribution parameters value for each watershed

Test type	Watershed	Distribution type	parameters	Parameters value
Kolmogorov-Smirnov	Dembi	Logpearson(3P)	α, β and γ	1545.9, -0.0388 and 62.83
	Dabana	Weibull	α, β and γ	0.609, 37.51 and 2.495
	Angar	Normal	μ, σ and γ	52.793, 68.0, and 0
	Wama	Lognormal(3P)	μ, σ and γ	0.96, 1.2788 and 0.2718
	Arjo Didessa	Lognormal	μ, σ and γ	4.0523, 1.3164 and 0
Anderson-Darling	Dembi	Lognormal(3P)	μ, σ and γ	2.70, 1.64 and 0.4783
	Dabana	Lognormal(3P)	μ, σ and γ	2.639, 2.0503 and 2.485
	Angar	Normal	μ, σ and γ	52.793, 68.0, and 0
	Wama	Lognormal(3P)	μ, σ and γ	0.96, 1.2788 and 0.2718
	Arjo Didessa	Lognormal	μ, σ and γ	4.0523, 1.3164 and 0
Chi-square	Dembi	Lognormal(3P)	μ, σ and γ	2.70, 1.64 and 0.4783
	Dabana	Lognormal(3P)	μ, σ and γ	2.639, 2.0503 and 2.485
	Angar	Normal	μ, σ and γ	57.793, 68.0, and 0
	Wama	Pearson 5	α, β and γ	1.0816, 1.9942 and 0
	Arjo Didessa	Lognormal	μ, σ and γ	4.0523, 1.3164 and 0

468 Note: these parameters are calculated with 12 year data since EasyFit used cannot handle (>5000)
469 data for trial version.

470 Similarly, using easy fit best distribution is selected as rank of goodness of fit is simulated [35].
471 The basins annual peak flows were extracted from daily flow hydrograph estimated with ArcSWAT
472 model for each watershed. Logarithm of annual peak flow from each watershed at outlet was
473 calculated. The value of average of peak discharge, \bar{x} and standard deviation, σ are calculated from
474 estimated flow hydrograph at each outlet of watershed. And the value of K_T can be calculated for any
475 required return period from equation (6) for lognormal and (7) for logpearson distribution.

476 1. lognormal distribution

477 In lognormal distribution the value of K_T in equation (4) can be found from equation (5) though
478 equation (6) below where (f =probability of exceedence= $1/T$):

$$479 \quad m = \left[\ln \left(\frac{1}{f^2} \right) \right]^{1/2} \quad (0 < f \leq 0.5), \quad (5)$$

480 Then calculating z using the approximation

$$481 \quad y = K_T = m - \frac{2.515517 + 0.802853m + 0.010328m^2}{1 + 1.432788m + 0.189269m^2 + 0.001308m^3}, \quad (6)$$

482 If $f > 0.5$, $(1-f)$ is replaced for f in equation (5)

483 For the lognormal distribution, applied to the logarithms of the variables, and their mean and
484 standard deviation are used in Equation (4). And then antilog of value is taken to get value at
485 treshold.

486 2. **Log Pearson type 3:** to use his method, the first step is to take the logarithms of the hydrologic
487 data, $g = \log(h)$. For log pearson K_T , is approximated from equation (6) using coefficient of
488 skewness (C_s) and studied by [32] as:

$$489 \quad K_T = y + (y^2 - 1)c + \frac{1}{3}(y^3 - 6y)c^2 - (y^2 - 1)c^3 + yc^4 + \frac{1}{3}c^5, \text{ where, } c = \frac{C_s}{6}, \quad (7)$$

490 **Table 10.** Expected probable flood with its frequency of occurrence from model determined annual peak

Threshold Values of (X_T) all units in m^3/s						
Return Period	Angar	Dembi	Dabana	Wama	Didessa@ Confluence of Wama	
2	583.7	458.0	316.8	237.8	657.9	
5	755.5	539.8	396.9	285.7	811.6	
10	864.6	594.0	446.6	317.2	913.4	
20	966.5	646.0	492.2	347.5	1011.0	
25	998.4	662.4	506.4	357.2	1042.0	
50	1095.7	713.2	549.3	387.1	1137.4	
100	1191.1	763.6	590.9	417.3	1232.1	
200	1285.8	813.9	631.7	448.0	1326.5	
1000	1505.4	930.2	725.1	522.3	1545.0	
10000	1825.3	1096.5	858.1	638.4	1857.4	

491 All the values are determined at the confluence of reach

492 According to Arjo Irrigation feasibility of 2007 and 2009 frequency analysis is done for Upper
493 Arjo Didesa at dam site which is approximately same with Dembi watershed. Outlet of Dembi
494 watershed and Dam site is nearer to each other. Hence, Frequency Analysis from simulated flow are
495 compared and found to be similar to that of feasibility study with negligible difference. The estimated
496 100 years flood was $763.6 m^3/s$ and feasibility study estimates as $787 m^3/s$. The difference of the value
497 may be the distribution type I used may be different from that feasibility study.

498 4. Conclusions

499 In this study SWAT model was used as to understand hydrological situation of different
500 watersheds as sources of stream flow has great significance in the effective utilization of water
501 resources. Didessa sub basin has not well studied and understood as compare with the northern sub

502 basins of Blue Nile such as (Tana sub basin). Model assisted hydrological characterization of Didessa
 503 sub basin was handled with different hydrological procedures and methods in this study. First the
 504 observed metrological and hydrological data were statistically tested and found to be consistent and
 505 free of trend. The SWAT Hydrological model for the sub basin was established, calibrated and
 506 validated by means of measured daily and monthly discharge at five gauged places in the study area.
 507 The calibration and validation of the model were measured by the R^2 (coefficient of determination)
 508 and the NS (Nash Sutcliff) model efficiency parameter of at monthly and daily time scale. The values
 509 of R^2 and NS range 0.66 to 0.70 and 0.60 to 0.66, 0.66 to 0.70 and 0.64 to 0.74, 0.62 to 0.80 and 0.61 to
 510 0.8 for Dembi, Arjo Didessa, and Angar tributary watershed respectively, at daily time scale. The
 511 values of R^2 and NS increases at monthly time scale and range 0.80 to 0.86 and 0.72 to 0.83, 0.87 to
 512 0.89 and 0.82 to 0.84, 0.82 to 0.92 and 0.79 to 0.91 for Dembi, Arjo Didessa, and Angar tributary
 513 watershed respectively.

514 Sensitivity analysis realizes parameters those were the most sensitive for the sub basin. And
 515 each parameter is arranged based on their sensitivity rank for each watershed as CN2, GWQMN,
 516 CH_K, ALPHA_BNK and LAT_TIME are the most sensitive parameters in Didessa sub basin.

517 Therefore, the hydrological process in Didessa sub basin is characterized using the simulated
 518 stream flow at five major tributaries watersheds. This characterization include stream flow
 519 hydrograph, identification of sub catchment with highest annual flow, peak flow analysis using fitted
 520 probabilistic distribution and unit.

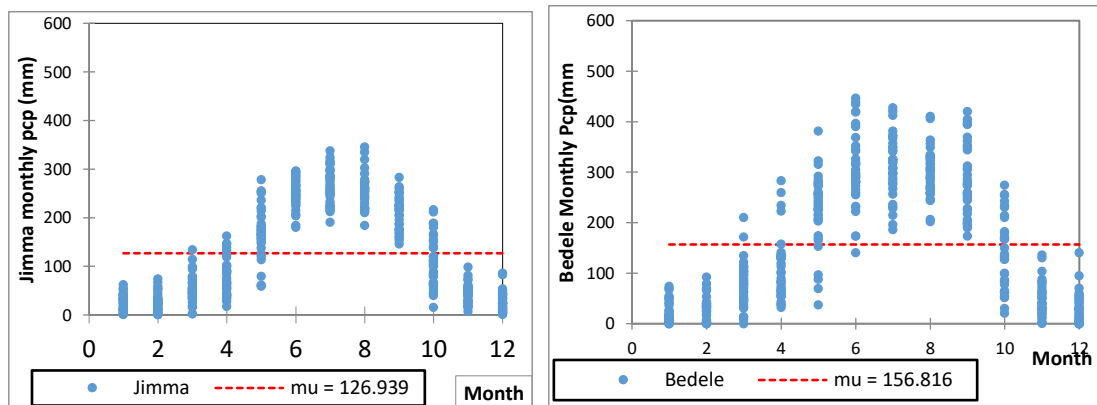
521 Frequency analysis was performed using the simulated annual peak flood series data sets
 522 watersheds of Didesa sub basin. Normal and Lognormal (3P) distribution gives the best fitted
 523 probabilistic distribution for most of watersheds except Dembi and Wama watershed. Using these
 524 fitted probabilistic distribution the peak flow of the watersheds were computed for 2- 10000 years
 525 return period.

526 **Conflicts of Interest:** The author declare no conflict of interest

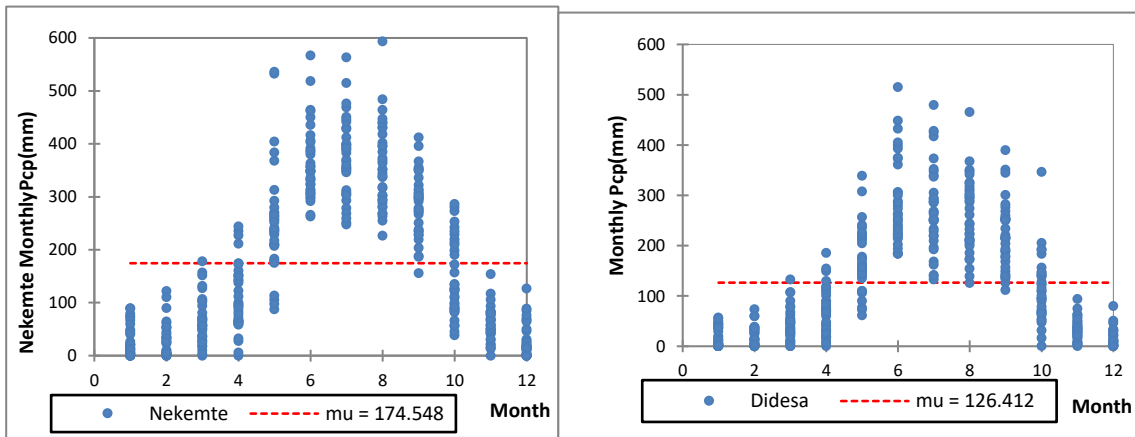
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528 Appendix

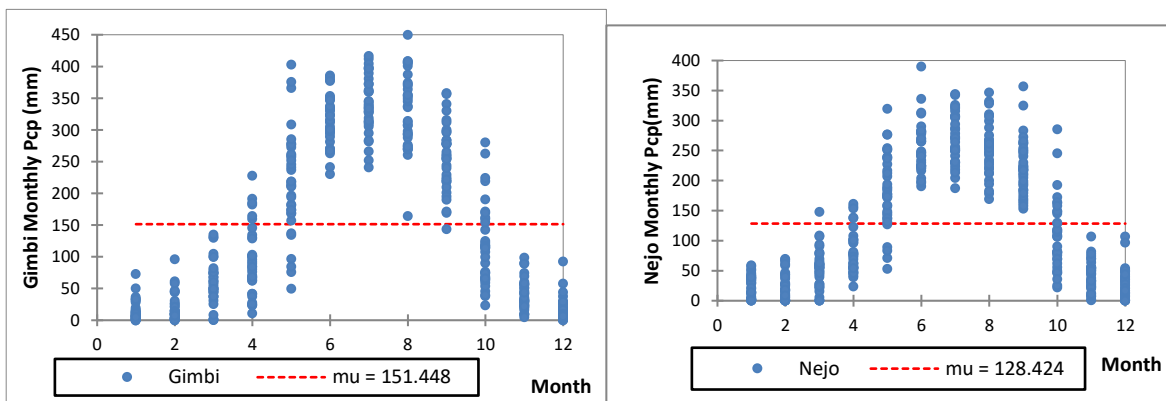
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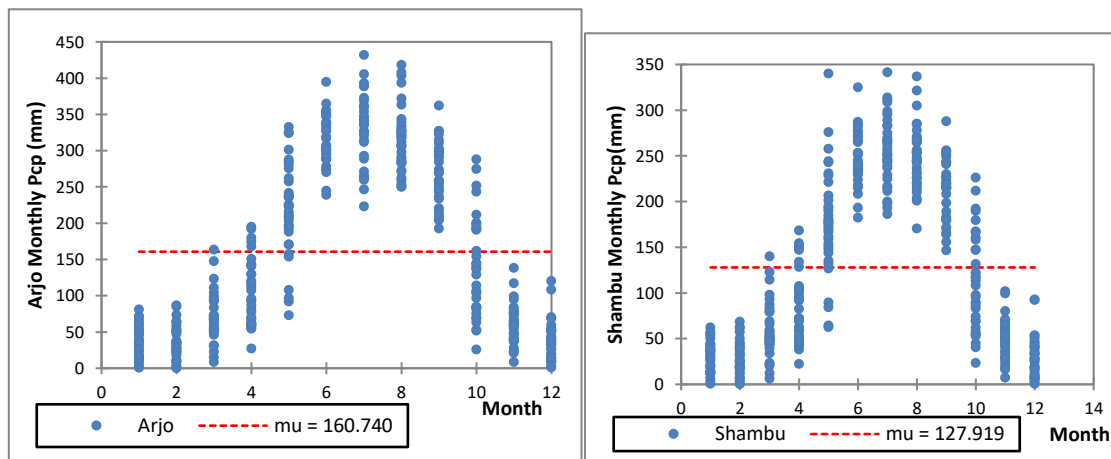
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531



532



533 (mu on graph represents mean monthly rainfall of the station). When mu is represented one horizontal line, the series of
 534 data has one mean and hence homogeneous.

535 **Figure A1.** XLSTAT final output Graphical representation of homogeneity test of rainfall data

536

537

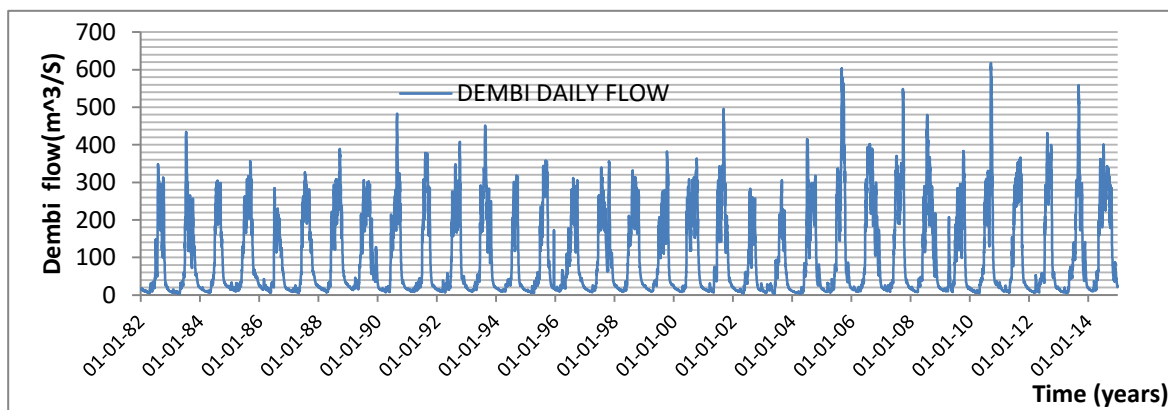
538

539

540 **Table A1.**Parameters' sensitivity Rank at different gauging stations of the sub-basin

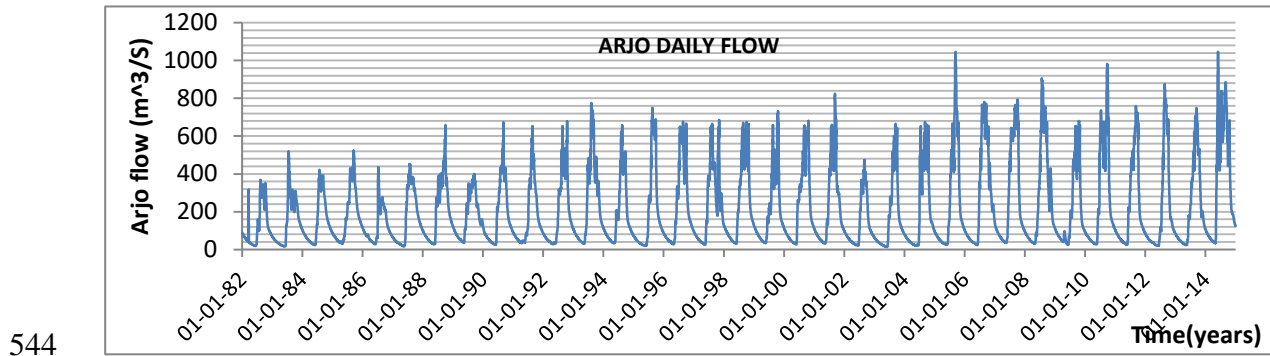
PARAMETER	DEMBI		ARJO		ANGAR		DABANA		WAMA	
	P-value	Rank	P-value	Rank	P-value	Rank	P-value	Rank	P-value	Rank
ALPHA_BF	0.51	8	0.27	4	0.42	6	0.73	14	0.56	10
ALPHA_BNK.	0.03	1	0.47	9	0.84	16	0.89	15	0.28	6
CH_K2	0.53	11	0.15	3	0.35	4	0.27	2	0.33	7
CH_N2	0.13	2	0.88	18	0.35	5	0.58	12	0.51	9
CN2.mgt	0.571	13	0	1	0.32	3	0.51	8	0.09	1
EPCO	0.4	6	0.82	14	0.61	10	0.96	17	0.72	13
ESCO	1	19	0.53	10	1	19	1	19	0.94	17
GW_DELAY	0.82	16	0.87	17	0.55	9	0.5	7	0.79	14
GW_REVAP	0.131	3	0.78	13	0.71	12	0.4	5	0.6	11
GWQMN	0.63	15	0.82	15	0.19	1	0.54	9	0.6	12
HRU_SLP	0.511	9	0.57	11	0.86	18	0.04	1	0.11	2
LAT_TTIME	0.57	12	0.13	2	0.54	8	0.57	11	0.18	3
OV_N	0.52	10	0.95	19	0.3	2	0.56	10	0.21	4
RCHRG_DP	0.61	14	0.41	8	0.72	13	0.92	16	0.61	18
REVAPMN	0.94	17	0.27	5	0.81	14	0.34	3	0.86	15
SLSUBBSN	0.45	7	0.57	12	0.81	15	0.5	6	0.96	19
SOL_AWC	0.96	18	0.33	7	0.84	17	0.6	13	0.24	5
SOL_K	0.38	5	0.85	16	0.51	7	0.37	4	0.91	16
SURLAG	0.3	4	0.28	6	0.69	11	0.99	18	0.34	8

541

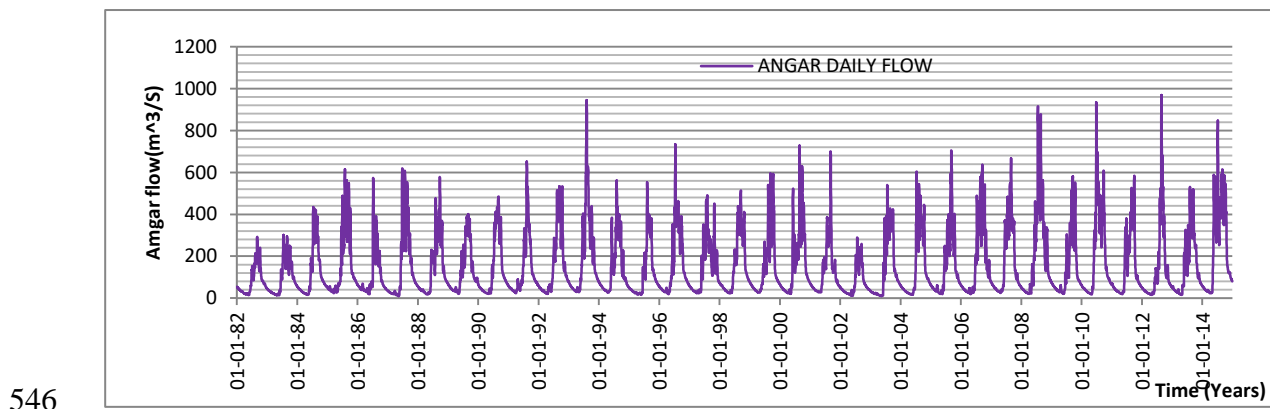


542

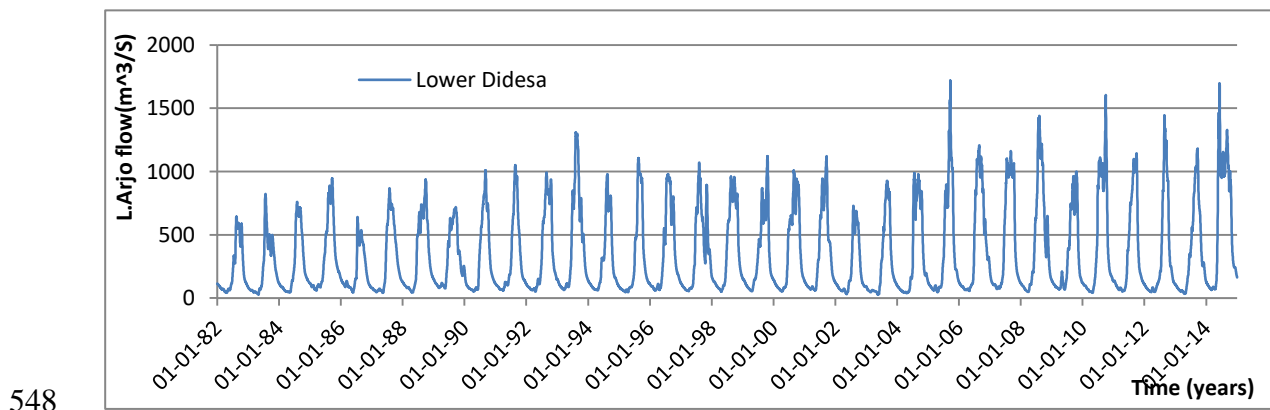
543 (a)



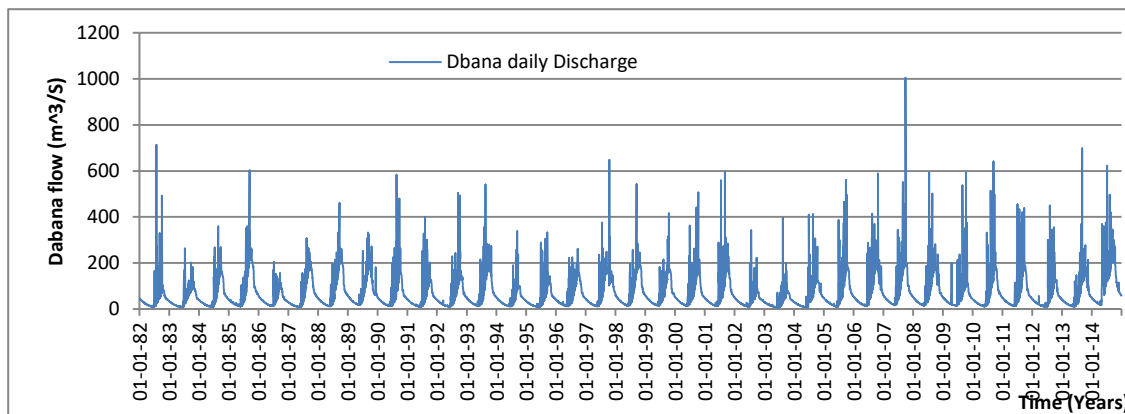
545 (b)



547 (c)



549 (d)



550

551 (e)

552 **Figure A2.** Daily flow Hydrograph: (a) Dembi, (b) Didessa at confluence with Dabana, (c) Angar, (d) Didessa
553 at confluence with Angar, (e) Dabana

554

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