

1 **Preliminary Description of the Origin of a Sedimentary – Exhalative Ore Deposit, in**  
2 **Molango, Hidalgo Mexico**

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10

11 **Abstract**

12 This work shows the preliminary description of the origin of a sedimentary - exhalative outcrop  
13 of Jurassic Lower Pliensbachian. The location of this deposit was achieved by applying an  
14 examination based in the identification of sedimentary transgressions of heterochronies ages and  
15 the identification of a Rift – type mega –structure. According with the methodology, it was  
16 carried out a study of the discordant relationships between two types of sediments: continental  
17 and marine. According the characterization, it was noted the existence of light rare earths, in  
18 values that show positive anomalies in comparison with the distribution of elements in upper  
19 continental crust according to the Clarke [1], reflecting so a felsic affinity of the mineral deposit.  
20 Also, positive anomalies of platinum and Pd, were determined with marginal contents of Au and  
21 Ag; and finally the base metals Zn, Pb and Cu were detected in low contents, which could be due  
22 to the presence of altered shale. According to the sedimentary lithology found, which was of  
23 siliciclastic type; to the exhalative roots observed during the fieldwork; the presence of quartz  
24 minerals such as biotite and muscovite; the presence of minerals of hydrothermal remobilization  
25 like chalcopyrite with some base metals, altered shale, as well as sulfur deficiency; this mineral  
26 reservoir could be defined as a SEDEX – type.

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28 **Keywords:** Rare earths; Platinum; SEDEX ore body; Rift mega structure; transgression

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34 **1. Introduction**

35 Geological-mining exploration has been based on methods that include an orientation survey,  
36 stream sediment geochemistry, sampling gas and, where applicable, a series of drillings in sites  
37 previously located, based on the results obtained in the previous stages. All this, represents both  
38 a great economic investment as well as time.

39 Using the mentioned methodology, until the mid-90s had been located in western Mexico about  
40 60 deposits of Volcanogenic Massive Sulfurs (VMS) [2] mostly, and others sedimentary  
41 exhalative (SEDEX), found in upper Jurassic – Cretaceous rocks and forming clusters. These  
42 deposits were located mainly in the “Terreno Guerrero” (GT) [3], and most of these deposits are  
43 of oceanic affinity with base metals such as Zn – Pb-Cu; of the KUKOKO type like the “Tizapa”  
44 [4] with about 100,000 Tons to 6 MT; in Guanajuato and Baja California with 1 MT having  
45 contents of base metals of Zn-Cu of the kind COPPER KING [2].

46 Conversely, SEDEX deposits with continental affinity lie in the complex Arteaga in Michoacan  
47 and in the metamorphic complex of Xolapa [5], and are of Pb-type with reserves of 2 MT, and in  
48 Zacatecas which is the largest with 36 MT; also it is important to note that both Zacatecas and  
49 Guanajuato [2], have low contents of Ag with some enrichment of Au.

50 The main features of the VMS [4] and SEDEX deposits are scarcely known in Mexico, as well as  
51 other sedimentary deposits that have been studied recently and have been exploited in artisan  
52 way, because its main economic attractiveness is due to the contents of Cu and Ag that they  
53 could have.

54 The fact that Mexico has traditionally been a silver producing country, led the search only in this  
55 metal rich deposits that were mainly epithermal and meso-thermal igneous type from Ternary; so  
56 marine and continental sedimentary deposits of low contents of Ag, have had little attention, and  
57 also these deposits could have base metals such as Cu, Pb, Zn, with contents of Au and Ag in  
58 small quantities [6]. Similarly, the techniques of extraction of ores igneous epi and meso –  
59 thermal having higher silver content, have been the most studied so far, and have had little or no  
60 effectiveness with the synergetic deposits sedimentary considered marginal by the low contents  
61 of Ag, without considering that it may contain other metals accessories that give added value to  
62 this type of deposits. However, advances in the study of new processes for extracting metal  
63 values has achieved effective results in the extraction of Au, Ag and Cu from reservoirs with low

64 values of these metals [7], and considering the large volumes of sedimentary deposits, has  
65 become a profitable exploitation.

66 In eastern Mexico there is a giant syngenetic manganese deposit of marine affinity SEDEX-type  
67 [8] and it is considered as the largest deposit in north America of this kind [9] being a sample of  
68 the potential contents of minerals for this SEDEX-type deposits.

69 In the same way, the first intents to find rare earths in Mexico were related to the prospection of  
70 phosphorites located in eastern Mexico [10] mainly in Tamaulipas, reporting the presence of an  
71 intrusive of Ternary age of nefeline – foidolites and rare earths mineralization. On the other  
72 hand, some geochemical studies were carried out in precambric rocks from Mexico in the  
73 Grenvilliane basement of Huiznopala Gneiss [11], where were obtained positive anomalies of Eu  
74 and low concentrations of Y, Zr, Nb, Th and U, similarly to that reported in the Oaxaqueño  
75 Complex [12] located in the south east of Mexico.

76 As is known, the major occurrence of REE's is in Asia, being almost the 95 % of total  
77 production in China. In this region is located the principal source of REE, which is the Bayan  
78 Obo deposits with high grade and igneous type that corresponds near to 80 % of LREE's [13] but  
79 with low contents of HREE's.

80 It is important to note that actually the tendency is related to the production of HREE, due the  
81 corresponding deposits have more content of these elements, and mining and extractive  
82 processes are comparatively less expensive [14]. Another alternative for recovery of REE's are  
83 the marine nodules, which are located at a depth of 3 to 5 Km [15] in the ocean.

84 According PGE's, its occurrence has been described in Yukon in Canada by Hullbert et al., [16],  
85 on the exploration of platinum from PGE in black slates where were explained deposits of  
86 different origins; the association of V-Cr-PGE's is known in exhalative deposits [17], and also  
87 chromium-spinels associated to Pd minerals have been described in Nairme, Australia [18]  
88 suggesting that PGE's anomalies have its origin in exhalative deposits related to volcanism in  
89 Rift zones, with some presence of organic matter.

90 The traditional methodology used in Mexico to locate mineral deposits such as igneous,  
91 epithermal, meso-thermal and sedimentary, has employed studies based in the primary  
92 delimitation of the area of study, through and exhaustive review of all sources of possible  
93 geological information. So, research can focus the study on the following main aspects: regional

94 geology, the stratigraphy of the area, geochemistry to know principal alterations, tectonics to  
95 describe the structure and finally, the location of a zone to get drillings.

96 With the data obtained in the steps described above, the analysis of the deposits, after carrying  
97 out the study of geological, stratigraphic and structural data, supported by geophysical well logs,  
98 is performed. Indicating the formation (rock) of mining interest perforated, thickness and upper  
99 and lower contacts, which is helpful to delineate the deposit, in time and space.

100 This method of exploration uses both direct and indirect dates that allow in many cases, to  
101 facilitate the work and reduce investment and time. However, this work introduces the study and  
102 analysis of heterochronies transgressions, but in this case only was analyzed the proto Jurassic  
103 transgression of age Pleisbachiano [19], due to the nature of the sedimentary deposits that are  
104 exploited near to the area of study, but these belong to a Kinmerigiano age [9].

105 Finally, this work shows the characterization of a sedimentary - exhalative deposit found in  
106 Molango Hidalgo, Mexico, which have been initially reported in previous work [20], where were  
107 considered the existence of anomalies for REE and precious metals such as Au Pt and Pd, and  
108 the results were quite similar to those found in other deposits [21].

109

## 110 **2. Methods**

### 111 *2.1. Mining exploration to find the ore deposit*

112 The methodology used for the mining exploration to find in this case, the sedimentary ore  
113 deposit, embraces some stages, which are: 1) A recompilation of all geological, geochemical and  
114 geophysical information available for the region of study, which will help to locate targets for the  
115 preliminary exploration. 2) To establish the different types of rocks and determine their relative  
116 ages, marking principally the transgressions and discordances, which is the basis of the proposed  
117 methodology, because this consists in the study of these transgressions, and suggests an  
118 evaluation of the discordant stratigraphic relationships of a marine transgression on rock of  
119 continental origin in a specific area. 3) To determine the main tectonic events in the area of study  
120 where is the transgression, for both compressive and extensive events, and then perform the  
121 tectonic analysis including erosion factors to locate the possible tectonic horst and graven, and  
122 this will link this tectonic event with the site of the transgressive event to find rocks with  
123 potential for the mining exploration. The favorable factors that transgressions offer to find the  
124 site of the mineralization are the following: Physiographic conditions of the graven,

125 physicochemical conditions of the rock, conditions of oxidation – reduction of the mineralized  
126 zone, and dissolution medium. 4) To define the primary and secondary environment of  
127 mineralization, using instrumental and geochemical analysis, in samples taken in the site located  
128 by the previous stages.

129 The prospecting methodology also provides guidelines to define regions with mining potential,  
130 defining the following: The amount of mining horizons producers for each geological formation,  
131 and if is always the same lithological unit or there are more, identification of lateral and vertical  
132 changes in the horizons or formations of minerals found and their physical recognition, defining  
133 the depth and thickness of producing horizons by lithology, and if they are lenses of sand or  
134 sandstone, identifying if their geological – mining behavior is different; and finally defining if  
135 stratigraphic and structural configuration of these horizons of minerals is of great interest.

136

### 137 *2.2. Characterization of mineral*

138 The characterization performed in this work was carried out to obtain accurate data of the  
139 mineralogical phases present, for which was done an analysis of overall phases by thin polished  
140 sections and X – ray Diffraction (XRD) to confirm the mineral phases present.

141 Also, the characterization studies were complemented by Scanning Electron Microscopy (SEM)  
142 to evaluate the texture, particle size and morphology of the phases detected, in the same way the  
143 X-ray mapping and EDS analysis helped us to determine punctual semi – quantitative  
144 composition of some previous phases identified and their distribution in particular zones. Finally,  
145 was conducted an analysis by ICP-MS, which gave the composition of whole rock and the  
146 obtained results were compared with the distribution of elements according to the Clarke [1].

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## 148 **3. Results**

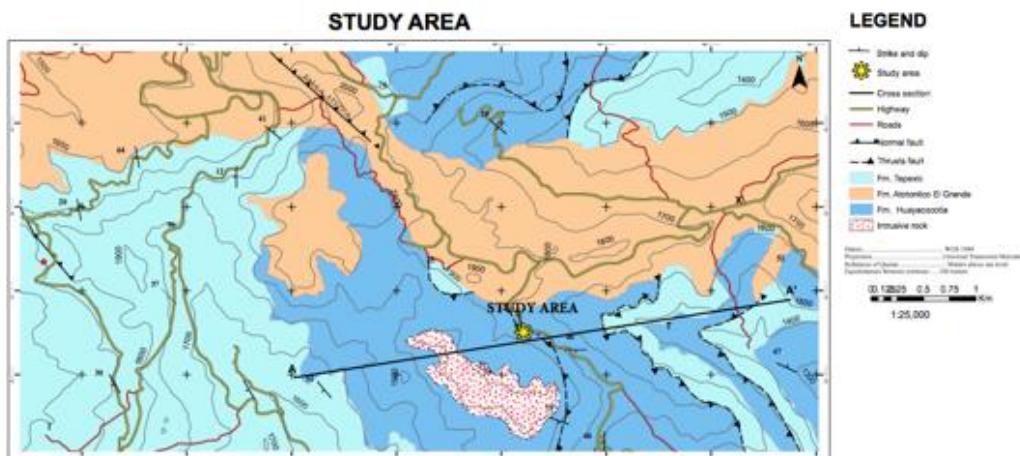
### 149 *3.1. Mining exploration to find the ore deposit*

150 The most important of applying the methodology previously described, is that this was useful to  
151 locate outcrops of sedimentary origin, which could be exhalative. All the raised and described  
152 methodology generated satisfactory results, considering that all elements described in the method  
153 were determined and thus the necessary clues that allowed to find the site of interest were  
154 obtained.

155 The basis of the methodology proposed for the location of minerals was focuses on determining  
156 geological ages, the depths and the upwelling areas. The prospecting steps provided further  
157 guidance to define regions with mining potential, considering that in the area of study was found  
158 only a mining horizon producer, located near the top of the Huayacocotla formation having an  
159 age of Pleinsbachiano – Lower Jurassic.

160 The mineral deposit found (Figure 1) with the above methodology, was located at the following  
161 coordinates; Longitude: -98.700334325 and Latitude: 20.765593 where were observed two main  
162 types of mineralization (Figure 2), which shows the study area found by the indirect method,  
163 where the main transgressions that delimit the location of the site were identified. In the upper  
164 unit, siltstone is observed, but there is a sudden change in the color of the shale into green tones  
165 and with remnants of organic matter, which is indicative again of a continental affinity, and was  
166 also observed strata of sandstone and conglomerate with clasts of less than 3 cm and outcrops of  
167 volcanic rock of intermediate to felsic composition; one located in a Philonian outcrop, having  
168 high silicification with finely disseminated pyrite throughout the matrix and exhalative roots.

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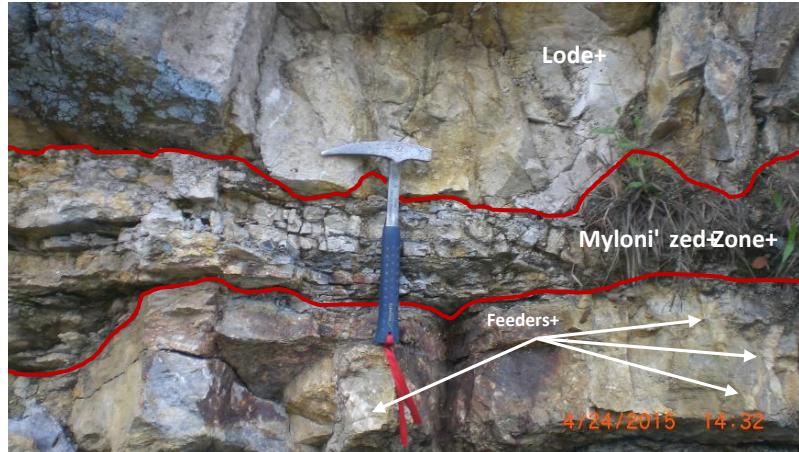
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Figure 1. Geological map showing the study area

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178 The other outcrop is a sequence of sedimentary rocks, showing alteration of minerals with  
179 granoblastic texture, siliciclastic and deformed pyrite. In the first type of outcrop, a hole of 9  
180 meters deep was made on the basis of the observed stock-work structure, in which were detected  
181 the exhalative roots of outcrop and near to the contact, on the top of the transgressive sequence,  
182 as seen in Figure 3.

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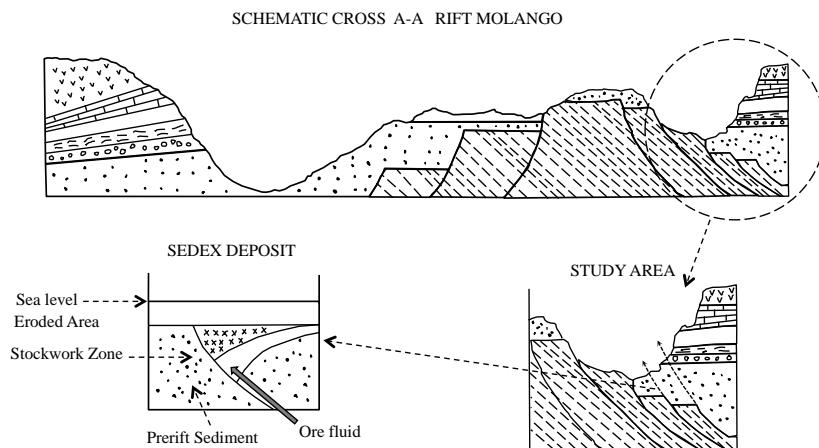
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188 However, in the initial portion of transgressive cycle, were located some exhalative roots such as  
189 seen in Figure 3 that shows the two lithologies of the producer unit. Also the base metal deposits,  
190 were located only in the intermediate unit of the lithology transgressive showing shale and

191 sandstone at rhythmic intervals, which overlies the shale unit with organic matter from the  
 192 Huayacocota Formation. Possibly at the base of the intermediate unit, are the outcropping  
 193 exhalative roots that have not yet been charted. In the Huayacocota Formation also three  
 194 lithological units were recognized; the oldest consists of conglomerate, sandstone, siltstone, and  
 195 shale intervals gap with re-worked fossils, exo-clasts sandstone, volcanic rock fragments and  
 196 smaller proportions of limestone rocks. The size of the measured clasts was 5 to 100 cm in a  
 197 sandy matrix and the size of the conglomerate clasts and clasts, suggest a continental affinity.  
 198 The intermediate unit represents a marine transgression, because in the fieldwork was observed  
 199 that this unit consists of sandstone and shale in a rhythmic sequence. In clusters, the size is not  
 200 greater than 10 cm and the main feature of this unit, is the presence of marine ammonites at the  
 201 base of the same, mainly in the shale. Near the top, the roots of the outcrop exhalative were  
 202 observed.

203 In the area of study, thickness of 350 meters were measured at 900 meters from the  
 204 Huayacocota Formation and thickness of producer horizon surfaced 100 meters, and the upper  
 205 unit shale with organic matter outcrops 250 meters from sedimentary rocks and increases in  
 206 some areas with outcrops of volcanic rocks at 350 meters, the intermediate portion is greater than  
 207 250 meters thickness, while the basal portion was observed at about 450 meters. Figure 4 shows  
 208 the lithological correlation of transgression of the lower Jurassic.

209



210

211 Figure 4. Characterization of the lithology of the studied transgression

212

213 On the other hand, sandstone lenses were observed in the units, and some of them are rich in  
 214 organic matter, indicating also an environment of continental deposit.

215 The mining geological behavior is different in the Huayacocotla training, since the  
216 mineralization is attributed here to the lithology of the intermediate portion and, it infra-lies to  
217 the shale with the organic matter unit. According stratigraphic setting, this is arranged in the  
218 form of bands NWSE direction, while the intermediate unit is structurally associated with high  
219 basement and related to areas of the reverse faults.

220 So, according the obtained results, it was possible to find two geological events of  
221 methodological type that allowed locate this ore deposit, which are due to tectonic environment  
222 and principally by the proto Jurassic Transgression. In the first case and according to megascopic  
223 study, could be observed rests of an antique Rift Triassic – Jurassic in eastern Mexico *op. cit*  
224 [22], and also is fundamental for understanding the mineralization control related in this work.  
225 This tectonic background, leads among others the metal dissolution medium that can be saline  
226 water and the conduction channels due to the topography, resulting in a lithological change from  
227 reductive to oxidant nature [23-25].

228 In the manganese district of Molango, is located the zone containing stratigraphic and structural  
229 vestiges that allowed infer an heterochronic stage of proto and neo Jurassic transgression,  
230 according [26, 27], and that is associated to a SEDEX mineralization located in Cuba [21],  
231 having similar characteristics to the ore deposit found in this work.

232 Furthermore, structurally are observed evidences of a Triassic – Jurassic Rift in the zone of study  
233 [25] and whose base is of age Grenvillian, showing tectonic features similar to those observed in  
234 the North Qaidam Orogen, Western China [28]. Likewise, it follows that the origin of base metal  
235 mineralization and rare earths, perhaps is related to pluvial sediments associated to hydrothermal  
236 activity linked to the Rift that could be the origin of the Gulf of Mexico [25]. This is important,  
237 because was widely accepted that sedimentary deposits (SEDEX) and other sedimentary strati-  
238 form deposits occurred in this type of environment [29, 30].

239

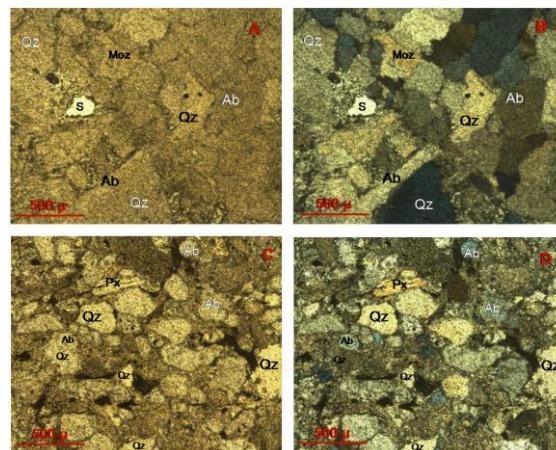
### 240 *3.2. Characterization of the ore deposit*

241 The preliminary characterization of the deposit found, was based on the mineralogical  
242 description by Thin Polished sections, XRD, and morphological identification, some specific  
243 semi-quantitative analysis using SEM-EDS and X – ray mapping to identify the distribution of  
244 elements in samples taken throughout the depth of the blast-hole made from 0 to 9 meters.  
245 Finally, was conducted an analysis of whole rock by fire assay (ICP - MS).

246 *3.2.1. Thin Polished Sections*

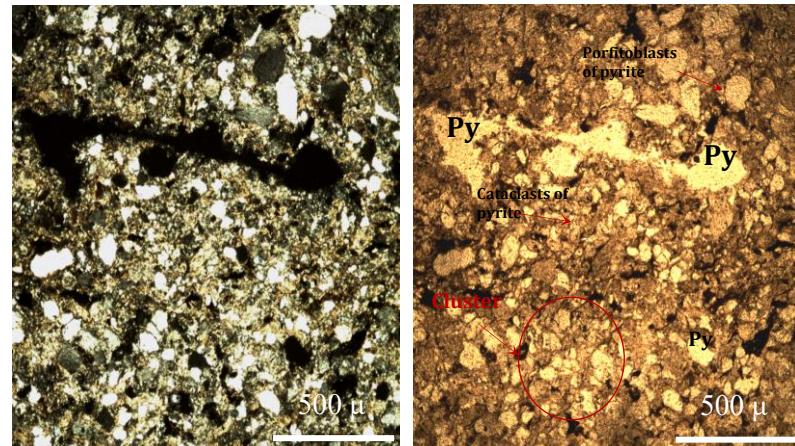
247 Figure 5 shows the results observed in the thin polished section done. Figures 5A and 5B  
 248 correspond to a filonian material, which were examined with parallel light (5A) and cross  
 249 nicols (5B); where from the matrix with angular to sub-angular clasts, was observed a  
 250 replacement by quartz, possibly albite (Ab) and the quartz (Qz) showed in some fragments,  
 251 undulating extinction which may be due to the fact that the material could be subjected to stress.  
 252 In the images, also can be observed sulfur and some fragments of monazite (Moz) imbibed in a  
 253 quartz matrix and the veins with dark minerals are parallel among them, but also are cut among  
 254 themselves in the form of a stock-work.

255 The image 5C was obtained with polarized light and image 5D with cross nicols, and both show  
 256 angular to sub-angular minerals with sedimentary matrix of clasts with an average size of 26  $\mu\text{m}$ ,  
 257 and also were observed quartz (Qz), albite (Ab) and pyroxene (Px), quartz veins with dark  
 258 minerals that partially replace the fluidal-looking clasts. Also can be observed very fine  
 259 fragments of monazite (Moz) near to the quartz veins.



260  
 261 Figure 5. Polished thin sections under: A) parallel light, 4x; B) cross-polarized light 4x (Filonian  
 262 material); C) parallel light, 4x; D) cross-polarized light, 4 x (Sedimentary material). Qy is quartz,  
 263 Ab is Albite, Moz is Monazite, S is Sulphur and Px is pyroxene

264  
 265  
 266  
 267  
 268 Finally, Figure 6 shows another polished section where was observed a collo-form texture that is  
 269 in the center of metamorphic, porfiroblastic and cataclastic zones, found in a fault zone. In  
 270 addition, in this sample the reduction of size by mechanical stresses of the pyrite in grains and  
 271 the disintegration in the form of clusters of pyrite grains in the lower part of image.



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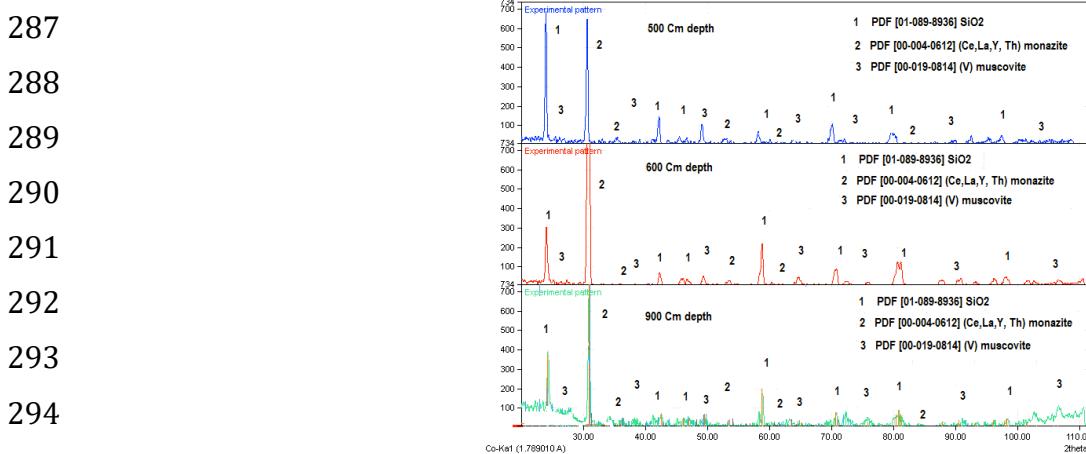
278 Figure 6. Polished thin sections showing the presence of pyrite in the various forms that result  
 279 from the decomposition process, which occurred during sedimentation and thermodynamic  
 280 transformations

281

### 282 3.2.2. X Ray Diffraction

283 The XRD results are mainly shown in Figure 7, where the presence of minerals of high and  
 284 medium temperature such as monazite, quartz type mogonite, muscovite and albite are observed  
 285 for mineral taken at 500 cm depth.

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295

296 Figure 7. XRD spectrum, showing high and medium temperature minerals

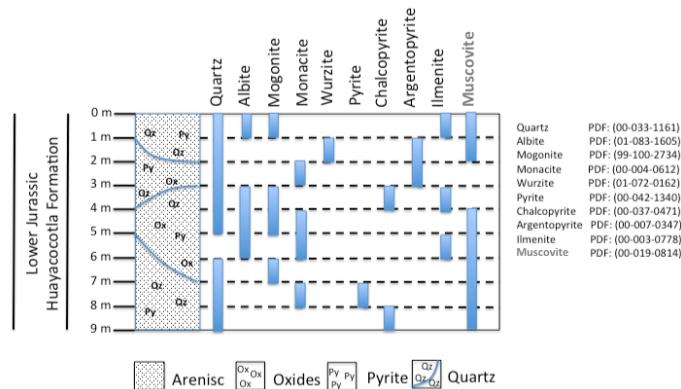
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298 For high temperature minerals, it was noted the presence of illmenite that could be associated  
 299 with the rare earths contents. Also, as depth increases, the amount of quartz decreases and is

300 remarket the presence of minerals like monazite and muscovite that are associated to rare earths,  
 301 platinum, palladium and gold.

302 As a resume, Figure 8 shows the distribution of minerals by XRD from extracted core drilling. It  
 303 can notice that there are minerals of high, medium and low temperature distributed through the  
 304 lithological column, confirming the existence of thermal pulses.

305



306

307 Figure 8. Distribution of mineral phases (by XRD) along the entire length of the drill core

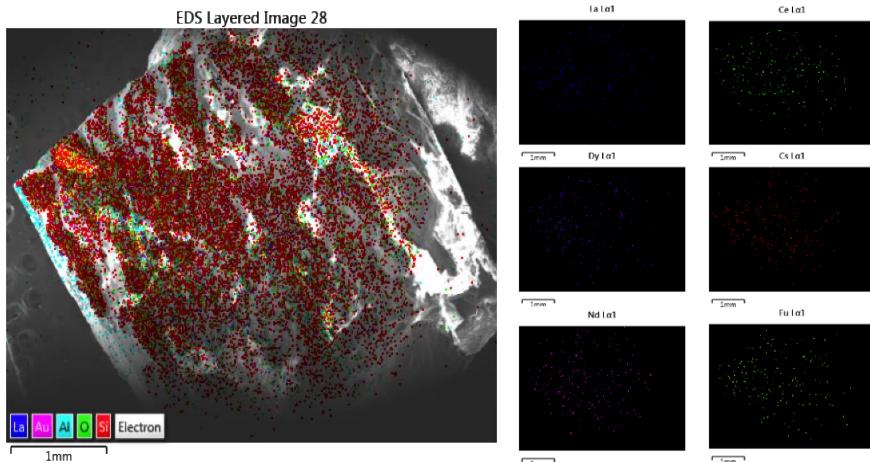
308

### 309 3.2.3. Scanning Electron Microscopy

310 The characterization performed by scanning electron microscopy in conjunction with EDS and  
 311 X-Ray mappings analysis allowed us to establish the presence of REE, Pt, Pd, Au and Ag in the  
 312 most representative particles of the samples analyzed.

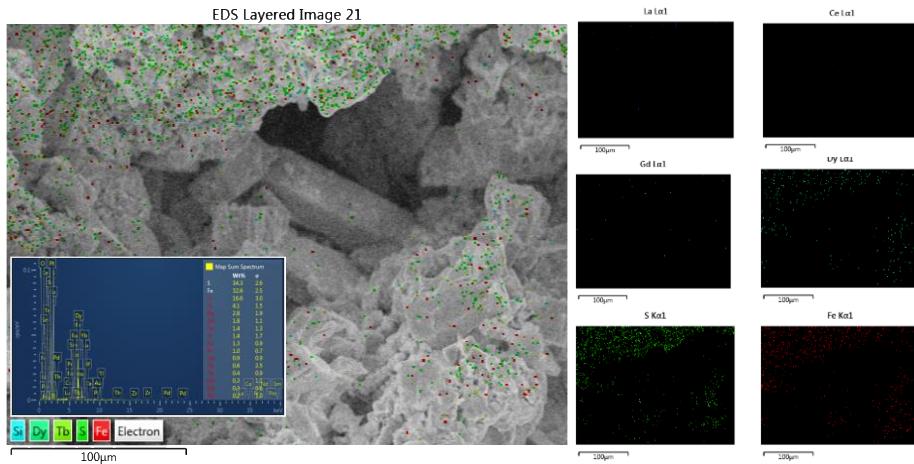
313 Figures 9 and 10 show a semi – detailed images of two particles of mineral, which were observed  
 314 by mapping the overall distribution of the La, Ce, Dy, Cs, Nd, Eu, Gd, S and Fe. In Figure 10,  
 315 can be observed that rare earths elements could be trapped during reduction of pyrite by  
 316 adsorption due to possible production of vacancies in the pyrite botryoidal, and table 1 shows the  
 317 punctual analysis done by EDS of each site showed in Figure 10.

318



319  
320  
321

Figure 9. X-ray mapping (by SEM) of a characteristic particle showing some LREE content



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Figure 10. SEM image showing where EDS analysis was performed, the results of which are reported in Table 1

326

Table 1. General EDS results for each point noted in figure 10

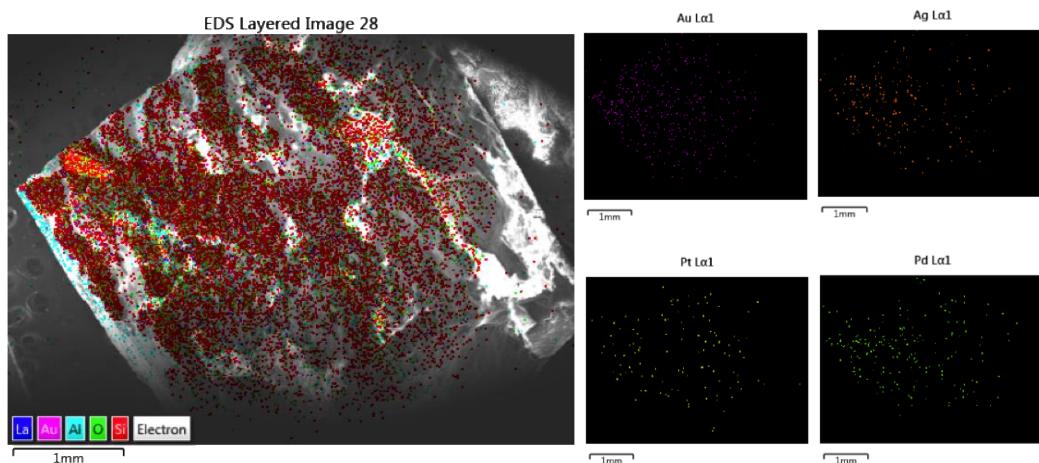
<b>Gd</b>	1.80	-	-	-	-	-	-	-	-	-
<b>Ta</b>	1.40	-	-	-	-	-	-	-	-	-
<b>Tl</b>	1.40	-	-	-	-	-	-	-	-	-
<b>Eu</b>	1.30	-	-	-	-	-	-	-	-	-
<b>Pr</b>	1.00	-	-	-	-	-	-	-	-	-
<b>Sm</b>	0.90	-	-	-	-	-	-	-	-	-
<b>Tb</b>	0.60	-	-	-	-	-	-	-	-	-
<b>Th</b>	0.40	-	-	-	-	-	-	-	-	-
<b>Ho</b>	0.30	-	-	-	-	-	-	-	-	-
<b>Pd</b>	0.30	-	-	-	-	-	-	-	-	-
<b>Yb</b>	0.20	-	-	-	-	-	-	-	-	-

327

328

329 Likewise, Figure 11 shows the mappings made in the same particle of Figure 9, to locate the  
 330 presence and distribution of Pt, Pd, Au and Ag; which could be due to the formation of faults that  
 331 generated channels where occurred a reduction process and then these elements flowed until  
 332 were caught by the shale. The foregoing, together with the results of ICP-MS (whole rock)  
 333 confirm the presence of these strategic elements in the type of deposit described here, which is  
 334 the potential basis for considering this preliminary study, to find an economically attractive  
 335 deposit near to this location or in the same site.

336



337

338 Figure 11. X-ray mapping of Pt, Pd, Au and Ag contained in a mineral particle (SEM)

339 *3.2.3. Analysis by ICP-MS*

340 The results obtained by ICP - MS were compared with the distribution of elements in the upper  
 341 earth crust described by Clarke [1], which could be observed in Table 2. According to the  
 342 obtained results, can be seen that the principal elements detected by ICP – MS have a positive  
 343 anomalies with respect to the Clarke [1]. Tables 3 and 4 show the chemical composition of the  
 344 sample (whole rock) and can be seen that contents of rare earth elements and minor values of Pt,  
 345 Pd and Au are above the average contents according to the Clarke, which may represent an  
 346 economic potential of this ore deposit found using a method of Indirect Mining Exploration,  
 347 once mineral reserves can be determined.

348

349 Table 2. Average distribution of element in the Earth's crust according to Clarke [1]

Range	Elements (% or ppm)
Greater than 10 %	O(46.6), Si(27.7)
1 – 10 %	Al(8.1), Fe(5.0), Ca(3.6), K(2.6), Na(2.8), Mg(2.1)
0.1 – 1 %	C, H, Mn, P, Ti
10 – 100 ppm	Cu, Ce, Co, Ga, La, Li, Nb, Ni, Pb, Sn, Th, Zn, Y
0.1 – 1 ppm	As, B, Br, Cs, Hf, Mo, Sb, Ta, U, W, Lanthanides, Bi, Cd, Y, In, Tl
0.01 – 0.1 ppm	Ag, Pd, Se
0.001 – 0.01 ppm	Au, Ir, Os, Re, Pt, Rh, Ru

350

351

352 Table 3. Elements in samples that are above the average values according to Clarke [1]

Range	Elements (%)
Greater than 10 %	Si(40.8)
1 – 10 %	Fe (36.6), Al (6.2), K (2.9), Mg (1.72), Ca (1.03)
0.1 – 1 %	Ti (0.71), Mn (0.52), P (0.19)

353 Table 4. Distribution of the elements in samples that are above the average values according to  
 354 Clarke (part 2), [1]

Range (ppm)	Elements (ppm)
10 – 100	Pb (1000), Zn (467), Cu (331), Ce (122), Sn (96), La (58), Ni (42), Y (33), Nb (32), Th (25), Ga (21), Li (28), Co (16) As (390), Mo (339), W (181), Sb (129), Cs (122), Nd (46), U (35), Bi (21), Gd (17), B (15), Hf (9), Ta (0.8), Sm (7), Cd (6), Sy (5), Yb (2.1), Eu (2), In (1), Tb (1), Tm (0.4), Ho (0.3), Er (0.1)
0.1 – 1	
0.01 – 0.1	Se (79), Pd (0.05)
0.001 – 0.01	Pt (0.05), Au (0.02)

355

356

357 **4. Discussion**

358 *4.1. Mining Exploration*

359 The method of mining exploration here raised, is focused on prospecting mineralization in  
 360 sedimentary rocks whose minerals preferably are housed in strata or related to these by  
 361 stratigraphic relationships of marine transgressions, on rocks of continental origin and thus find  
 362 and display susceptible traces that could contain mineralization in a given area. Although its  
 363 application is related here only to sedimentary rocks, also it could indirectly be used to define  
 364 other lithologies such as in igneous and metamorphic rocks, which may contain mineralization.  
 365 In addition, it is noteworthy that the sedimentary deposits are widespread and they occur in all  
 366 geological eras, from the Proterozoic to the present, and at different depths.

367 On the other hand, the secular variations in the abundance of sedimentary deposits, for example  
 368 in the study area, are presented during the Lower Jurassic, but the abundance of occurrences also  
 369 could be determined correlating with other mineral deposits of importance already in  
 370 exploitation.

371

372

373

374 *4.2. Characterization of the ore deposit*

375 In the geochemical analysis of trace elements such as Cr and Th, due their low mobility, these  
376 are considered suitable for determining provenance and tectonic setting [31]. As seen from the  
377 results, there is an enrichment of elements large ion such as Cr=80 ppm and Th=25 ppm; all are  
378 higher than the average contents of a Huayacocotla shale barren.

379 On the other hand, the contents (in ppm) of transition trace elements V=89, Co=15.8, Cu=331  
380 and Ni=41.5, are low due to that affinity of the deposit could be felsic or acid.

381 Indeed, while the content of Cr and Ni is low, the content of V could be associated to the  
382 occurrence of metals such as platinum, palladium and gold.

383 Also this deposit has higher concentration of light rare earths (LREE) than the Clarke [1],  
384 because values (in ppm) for heavy rare earths (HREE) are low for Y=32.3 and Ga=21. For this  
385 reason, it could be inferred the presence of associated minerals of REE of the monazite-type [Ce,  
386 La, Nd, Th, (PO<sub>4</sub>)] and bastnaesite minerals.

387 It is important to note that negative anomaly of HREE for Eu and Hf with respect to Clarke [1],  
388 confirms the impoverishment of HREE and reaffirms the felsic affinity of the deposit, and the  
389 possibility that its origin is from a Gneissic protolith, because mineralization has higher  
390 concentrations of LREE. In the same way, a negative anomaly of Ta shows a sedimentary  
391 affinity of the ore deposit. On the other hand, the negative anomaly of sulfur can be due to the  
392 presence of pyrite formed by sulfate reduction in marine environment of lower Jurassic age [32].  
393 Also, the possible contents of organic carbon, the presence of pyrite in botryoidal aggregates,  
394 sulfur clasts, and low contents of base metals, low ratios of Ag/Au, and minor contents of  
395 arsenopyrite and pyrrhotite, could be to the silisticlastic nature that is typical of the Rift  
396 environments [29].

397 Likewise, the study area according to the obtained results could be classified as a Rift, such as  
398 was described by Jowet [33], showing besides a bimodal volcanism, fault-controlled guidelines,  
399 positive gravity, and rapid sedimentation rates over short periods of time that change at slow  
400 rates over long periods. In the same way, it is known [34, 35] that a Rift environment is  
401 conductive to the location of SEDEX deposits of base metals, such as described here.

402 Finally, the Huayacocotla Formation in which this deposit lies, have a great similarity in  
403 lithological, stratigraphic and chronological aspects with another deposit found in the San

404 Cayetano Formation [21], which also was classified as a SEDEX and located in the  
405 Matahambres area of Cuba.

406

## 407 **Conclusions**

408 (1) A sedimentary – exhalative deposit was found in the eastern of the Hidalgo state, in Mexico,  
409 based only in the study of the heterochronic transgressions.

410 (2) Mineralization found is strongly linked to a sedimentary control; besides is close related  
411 with a marine transgression of lower Jurassic age.

412 (3) The mineral outcrop found, consist of two types of mineralization. The first of the Philonian  
413 type, where exhalative roots of the stock-work type were observed at the base. And the  
414 second, of stratiform type formed by a sequence of sandstones and shale mutually  
415 concordant and of marine origin. This lithology, preliminary is the typical one found in  
416 several formations of type SEDEX.

417 (4) According to the characterization carried out by Thin Polished Sections, metamorphism and  
418 cataclastic fragments of mechanically deformed pyrite and quartz were found, which is  
419 indicative of the existence of old tectonism and the movement of transgressions. Likewise,  
420 minerals of reducing environment such as the pyrite that in this case was found in a  
421 botryoidal form, indicating so its marine origin. In the same way, it was possible to identify  
422 micro – veins filled with quartz and disseminated pyrite, remobilization minerals such as  
423 chalcopyrite and finally, monazite; all of them possibly originated in a stock-work.

424 (5) From the analysis performed by ICP, as well as the complementary characterization carried  
425 out by XRD, SEM-EDS; it is concluded that the mineral found contains adequate values of  
426 precious metals such as Au, Ag, Pt and Pd, as well as some rare earths elements, since their  
427 values are above the average classification made by the Clarke [1].

428 (6) All of the above contributes to establish that depending on the results obtained, this outcrop  
429 is a SEDEX type deposit with possibilities to have mining potential. Also, these results  
430 validate the Indirect Method of Mining Exploration to find sedimentary deposits through the  
431 study of transgressions, but could be modified to find any type of deposit, based on its main  
432 lithological characteris.

433

434

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439

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443 Juan Hernández-Ávila, Ma. Isabel Reyes-Valderrama and Rosa A. Vázquez-García, observed the  
444 samples using SEM-EDS and analyzed the data. Eduardo Cerecedo-Sáenz prepared samples in  
445 Thin Polished Sections and analyzed the data. All authors participated in writing the manuscript.

446

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448

449 **References**

450 [1] Levinson, A. A., 1980. Introduction to exploration Geochemistry, Applied Publishing,  
451 Wilmette, Il, USA, Second edition, pp. 924.

452 [2] Miranda, G. M. A., 1995. The volcanogenic massive sulfide and sedimentary exhalative  
453 deposits of the Guerrero Terrane, Mexico, A dissertation submitted to the Faculty of the  
454 Department of Geosciences in the Graduate College in the University of Arizona for the Degree  
455 of Doctor of Philosophy, pp. 298.

456 [3] Campa, M. F. and Coney, P. J. 1983. Tectono-stratigraphic terranes and mineral resource  
457 distributions in Mexico, Canadian Journal Earth Science, 20, pp. 1040 – 1051.

458 [4] Akio, Y. and Koji, K. 2000. Geology of the Tizapa mine in Mexico with special reference to  
459 the present status after development, Shigen Chishitsu, Vol. 50, pp. 1-10.

460 [5] Hermann, U. R., Nelson, B. K. and Ratsbacher, L., 1994. The origin of a terrane: U/Pb zircon  
461 geochronology and tectonic evolution of Xolapa complex (southern Mexico), Tectonics, Vol. 13,  
462 pp. 455 – 474.

463 [6] Ortensen, J. K., Hall, B. V., Bissing, Friedman, R.M., 2008. Age and Paleotectonic Setting of  
464 Volcanogenic Massive Sulfide Deposits in the Guerrero Terrane of Central Mexico: Constraint  
465 from U-Pb Age Isotope Studies, Economic Geology 103, pp. 117 – 140.

466 [7] Salinas, R. E., Hernández, A. J., Rivera, L. I., Cerecedo, S. E., Reyes, V. M. I., Correa, C. M.,  
467 Rubio, M. D., 2016. Leaching of silver contained in mining tailing, using sodium thiosulfate: A  
468 kinetic Study, *Hydrometallurgy*, 160, pp. 6 – 11.

469 [8] Maynard, B. J. and Klein, G. D., 1995. Tectonic Subsidence analysis in the Characterization  
470 of Sedimentary Ore Deposits: Examples from the Witwatersrand (Au), White Pine (Cu), and  
471 Molango (Mn). *Economic Geology*, pp. 37 – 50.

472 [9] Okita, P. M., 1992. Manganese Carbonate Mineralization in the Molango District, Mexico.  
473 *Economic Geology*, V. 87, pp. 1345 – 1366.

474 [10] Herrera, E. M., Rubinovich, K. R., Lozano S. C. R. and Sánchez, Z. J. L., 1991. Nepheline-  
475 rich foidolites and Rare Earths mineralization in the Picacho Tertiary intrusive Complex, Sierra  
476 de Tamaulipas, northeastern Mexico. *Canadian Mineralogist*, Vol. 29, pp. 319 – 336.

477 [11] Lawlor, P. J., Ortega-Gutierrez, F., Cameron, K. L., Ochoa-Camarillo, H., Lopez, R. and  
478 Sampson, D. E., 1999. U-Pb geochronology, geochemistry and provenance of the Grenvillian  
479 Huiznopal Gneiss of eastern Mexico. *Precambrian Research*, V. 94, pp. 73 – 99.

480 [12] Keppie, J. D., Dostal, J., Cameron, K. L., Solari, L., Ortega-Gutierrez, F. and López, R.,  
481 2003. Geochronology and geochemistry of Grenvillian igneous suites in the Northen Oaxacan  
482 Complex, Southern Mexico: tectonic implications. *Precambrian Research*, 120, pp. 365–389.

483 [13] Verplanck, P. L., Van Gosen, B. S., Seal, R. R., McCafferty, A. E., 2014. A deposit model  
484 for carbonatite and peralkaline intrusion – related rare earth element deposits, U.S. Geological  
485 Survey Scientific Investigations Report 2010-5070-J, pp. 58.

486 [14] Kynicky, J., Smith, M. P., Xu, C., 2012. Diversity of rare earth deposits: the key example of  
487 China. *Elements* 8, pp. 361 – 367.

488 [15] Kato, Y., Fujinaga, K., Nakamura, K., Takaya, Y., Kitamura, K., Ohta, J., Toda, R.,  
489 Nakashima, T. and Iwamori, H., 2011. Deep-sea mud in the Pacific Ocean as a potential resource  
490 for rare-earth elements. *Nature Geoscience* 4, pp. 535 – 539.

491 [16] Hulbert, L., Carne, R., Gregoire, C. and Paktunc, D., 1992. Sedimentary nickel, zinc, and  
492 platinum-group-mineralization in the Denovian black shale at the Nickel property, Yukon,  
493 Canada: A new deposit type. *Expl. Min. Geol.* 1, 1, pp. 39 – 62.

494 [17] Goodfellow, W. D., & Lydon, J. W. (2007). Sedimentary exhalative (SEDEX) deposits.  
495 Mineral deposits of Canada: A synthesis of major deposit types, district metallogeny, the

496 evolution of geological provinces, and exploration methods: Geological Association of Canada,  
497 Mineral Deposits Division, Special Publication, (5), 163-183.

498 [18] Pasave, J., 1993. Anoxic sediments – an important environment for PGE; an overview. Ore  
499 Geol. Rev., 8, pp. 425 – 445.

500 [19] Erber, H. K., 1956. El Jurásico Medio y el Calloviano de México: México D.F., Congr.  
501 Geol. Internacional (In Spanish), Vol. 20, monografía, pp. 393.

502 [20] Cerecedo, S. E., Rodríguez, L. V., Andrade, T. P. D., Salinas, R. E., Hernández, A. J.,  
503 Arenas, F. A., 2015. Chemistry and characterization of a mineral deposit mineralogical economic  
504 interest, Mater. Res. Soc. Symp. Proc. Vol. 1766, Materials Research Society.

505 [21] Perez, V. R. G. and Melgarejo, J. C., 1998. El yacimiento Matahambre (Pinar del Río  
506 Cuba): Estructura y mineralogía. Acta Geológica Hispánica (In Spanish), V. 33, # 1-4, pp. 133 –  
507 152.

508 [22] Cerecedo S. E., 2003. La mineralización de cobre-plata del Rift Triásico – Jurásico del  
509 oriente de México, Instituto Politécnico Nacional, Escuela Superior de Ingeniería y Arquitectura.  
510 Master Thesis (In Spanish), pp. 284.

511 [23] Jowet, E. C., 1986. Effects of continental rifting on the location and genesis of stratiform  
512 copper-silver deposits. Geological Association of Canada, Special Paper, 36, pp. 53 – 66.

513 [24] Goodfellow, W. D., Lydon, J. W., and Turner, R. W., 1993. Geology and genesis of  
514 stratiform sediment-hosted (SEDEX) Zn-Pb-Ag sulphide deposits, in Kirkham et al., eds.,  
515 Mineral Deposit Modeling, Geological Association of Canada, Special Paper 40, pp. 201 – 251.

516 [25] Cerecedo, S. E. and Salinas, R. E., 2013. Guía Rápida de Exploración Geológica Minera.  
517 pp. 73, Editorial Académica Española (In Spanish), Saarbrücken, Alemania.

518 [26] Cantú, C. A., 1998. Las transgresiones jurásicas en México, Revista Mexicana de Ciencias  
519 Geológicas (In Spanish), Vol. 15, Num. 1, pp. 25 – 37.

520 [27] Cantú, C. A., 2001. México, margen occidental de la Pangea según evidencias  
521 biogeográficas del Pérmico al Jurásico Inferior, Revista Mexicana del Petróleo (In Spanish), Año  
522 XXVI, Núm. 345, pp. 28 – 35.

523 [28] Zhang, C., Bader, T., Zhang, L., & van Roermund, H. (2015). The multi-stage tectonic  
524 evolution of the Xitieshan terrane, North Qaidam orogen, western China: From Grenville-age  
525 orogeny to early-Paleozoic ultrahigh-pressure metamorphism. Gondwana Research.

526 [29] Sillitoe, R. H., 2002. Rifting, bimodal volcanism, and bonanza gold veins: Society of  
527 Economic Geologist Newsletter, 48, pp. 24 – 27.

528 [30] Goodfellow, W. D., Peter, J. M., Winchester, J. A. and Van Staal, C. R., 2003. Ambient  
529 marine environment and sediment provenance during formation of massive sulfide deposits in  
530 the Bathurst Mining Camp: Importance of reduced bottom waters to sulfide precipitation and  
531 preservation: Economic Geology Monograph, V. 11, pp. 129 – 156

532 [31] Mcleanan, S. M., Hemming, S., McDaniel, D. K. and Hanson, G. H., 1993. Geochemical  
533 approaches to sedimentation, provenance, and tectonics. Johnsson, M. J., Basu, A. (Eds).  
534 Processes controlling the composition of clastic sediments. Geological Society of America.

535 [32] Coleman, M. L., Raiswell, R., Brown, A., Curtis, C. D., Aplin, A. C., Ortoleva, P. J. &  
536 Eglinton, G., 1993. Microbial mineralization of organic matter mechanisms of self-organization  
537 and inferred rates of precipitation of diagenetic minerals (and Discussion), Philosophical  
538 Transactions of the Royal Society of London A: Mathematical, Physical and Engineering  
539 Sciences, 344 (1670), pp. 69 – 87.

540 [33] Jowett, E. C. (1989). Effects of continental rifting on the location and genesis of stratiform  
541 copper-silver deposits. Sediment-Hosted Copper Deposits. Geol Assoc of Canada, Spec Paper,  
542 36, 53-66.

543 [34] Carne, R. C. and Cathro, R. J., 1982. Sedimentary-exhalative (Sedex) Zn-Pb-Ag Deposits,  
544 Northern Canadian Cordillera, Canadian Institute of Mining and Metallurgy, Bulletin, Volume  
545 75, pp. 66 – 78.

546 [35] Goodfellow, W. D., 2004. Geology, genesis and exploration of SEDEX deposits, with  
547 emphasis on the Selwyn Basin, Canada. Attributes and models of some major deposits in India,  
548 Australia and Canada: New Delhi, Narosa Publishing House, pp. 24 -99.