

1 Spatiotemporal Distribution of Satellite-Retrieved 2 Ground-Level PM_{2.5} and Near Real-Time Daily 3 Retrieval Algorithm Development in Sichuan 4 Basin, China

5
6 Chao Gao¹, Xuelei Zhang^{1,2*}, Wenyong Wang^{3*}, Aijun Xiu¹, Daniel Q. Tong² and Weiwei Chen¹

- 7
8 1. Key Laboratory of Wetland Ecology and Environment, Northeast Institute of Geography and
9 Agroecology, Chinese Academy of Sciences, Changchun 130102, China.
10 2. Center for Spatial Information Science and Systems, George Mason University, Fairfax, VA 22030,
11 USA.
12 3. Faculty of Geosciences and Environmental Engineering, Southwest Jiaotong University, Chengdu
13 610031, China.

14
15 * Correspondence: zhangxuelei@neigae.ac.cn (X.Z.); wyyang@home.swjtu.edu.cn (W.W.);
16 Tel.: +86-431-8554-2314 (X.Z.); +86-028-8760-0929 (W.W.)

17 **Abstract:** Satellite-based monitoring can retrieve ground-level PM_{2.5} concentrations with
18 higher-resolution and continuous spatial coverage to assist in making management strategies
19 and estimating health exposures. The Sichuan Basin has a complex terrain and several city
20 clusters that differ from other regions in China: it has an enclosed air basin with a unique
21 planetary boundary layer dynamic which accumulates air pollution. The spatiotemporal
22 distribution of 1-km resolution Aerosol Optical Depth (AOD) in the Sichuan Basin was
23 retrieved using the improved dark pixel method and Moderate Resolution Imaging
24 Spectroradiometer (MODIS) data in this study. The retrieved seasonal AOD reached its
25 highest values in spring and had the lowest values in autumn. The higher correlation ($r = 0.84$,
26 $N = 171$) between the ground-based Lidar AOD and 1-km resolution MODIS AOD indicated
27 that the high-resolution MODIS AOD could be used to retrieve the ground-level PM_{2.5}
28 concentration. The Lidar-measured annual average extinction coefficient increased linearly
29 with the Planetary Boundary Layer Height (PBLH) in the range of 100~670 m, but
30 exponentially decreased between the heights of 670~1800 m. Both the correlation and the
31 variation tendency of simulated PBLH from WRF_SHIN/CALMET were closer to the Lidar
32 observation than that of three other Planetary Boundary Layer (PBL) schemes (the
33 Grenier-Bretherton-McCaa (GBM) scheme, the he Total Energy-Mass Flux (TEMF) scheme
34 and the University of Washington (UW) scheme), which suggested that the simulated PBLH
35 could be used in the vertical correction of retrieval PM_{2.5}. Four seasonal fitting functions were
36 also obtained for further humidity correction. The correlation coefficient between the aerosol
37 extinction coefficient and the fitted surface-level PM_{2.5} concentration at the benchmark station
38 of Southwest Jiao-tong University was enhanced significantly from 0.62 to 0.76 after vertical
39 and humidity corrections during a whole year. During the evaluation of the retrieved
40 ground-level PM_{2.5} with observed values from three cities, Yibin (YB), Dazhou (DZ), and
41 Deyang (DY), our algorithm performed well, resulting in higher correlation coefficients of
42 0.78 ($N = 177$), 0.77 ($N = 178$), and 0.81 ($N = 181$), respectively.

43
44 **Key words:** particulate matter; AOD; planetary vertical corrections; humidity corrections

45 1. Introduction

46 Aerosols are an essential part of atmospheric chemistry and physics, the biosphere,
47 climate, and even human health [1,2]. The climate impact of aerosols is performed as climate
48 forcing, which can be further divided into direct and indirect effects [3]. Moreover,
49 epidemiological studies have also demonstrated that ambient fine particles can seriously
50 deteriorate public health, especially for respiratory and cardiovascular diseases [4,5].

51 With the rapid economic development and urbanization in China, a huge amount of
52 energy consumption and the increased use of motor vehicles have led to the frequent
53 occurrence of serious regional haze pollution. Most of the pollution events are compound
54 secondary pollution, which are caused by fine particles. Previous studies have shown that
55 three primary conditions form regional haze events: the presence of massive emission sources,
56 more static weather conditions, and complex terrain that is unsuitable for pollution diffusion
57 [6]. The spatial distribution of haze pollution in China are further divided into several regions
58 based on emission and pollution characteristics, i.e., the region of Beijing-Tianjin-Tangshan,
59 the region of the Yangtze River Delta, the region of the Pearl River Delta, the region of
60 Northeastern China, and the region of the Sichuan Basin [7].

61 Unlike the geomorphology of the plains and deltas in other regions, the Sichuan Basin is
62 surrounded by mountains and is located in the leeward slope of the Qinghai-Tibet Plateau.
63 The atmospheric layer is stable and it is easy to form an inversion layer. Thus, widespread
64 and long continuous fog and haze pollution frequently occurs over this region. However,
65 there are only limited ground-based monitoring stations located in Sichuan Province, and
66 cannot reflect the spatial and temporal distribution of particulate concentrations during the
67 polluted period, while the remote sensing retrieved particulate concentrations can provide
68 more information to overcome these monitoring deficiencies. Considering the hygroscopicity
69 of particulate matter and vertical distribution, AOD can be reasonably converted to the
70 ground-level particulate concentration via the usage of the boundary layer height and the
71 hygroscopic growth factor. Over recent years, satellite remote sensing retrieved ground level
72 particulate matter concentration has been studied over the world. Hoff and Christopher [8]
73 systematic reviewed the related 30 studies that had addressed on the usage of MODIS or
74 MISA AOD to assess ground-based particulate concentrations during 2001 to 2007. .Gupta et
75 al. [9] obtained a simple linear relationship between the daily mean MODIS AOD and
76 particulate matter mass, and their correlation coefficient was 0.96 over 26 locations in the
77 world. Donkelaar [10] also found significant agreement between satellite-derived and
78 ground-based values outside North America and Europe ($r = 0.9$; $n = 210$). Nicolantonio [11]
79 showed that the correlation coefficient of MODIS AOD and PM concentration was 0.87 by
80 employing the mixing height and relative humidity from the MM5 (Mesoscale Model Version
81 5) in Northern Italy. The relationship between ground-level PM_{2.5} concentrations and the
82 average of MODIS and MISR AOD was further investigated in North America ($r = 0.89$) and
83 elsewhere ($r = 0.91$) [12]. Significant spatial variation of annual mean ground-level PM_{2.5}
84 concentration with remote-sensed PM_{2.5} from MODIS ($r = 0.83$) and MISR ($r = 0.76$) over North
85 America was also reported in Donkelaar [13]. A similar study also reviewed that the MODIS
86 ($r = 0.82$) and MISR ($r = 0.85$) had overall comparable correlations to the ground-level PM_{2.5}
87 concentration in Xi'an, Shaanxi Province [14]. Recently, the major research trends in PM_{2.5}

88 inversion are focused on the usage of high resolution retrievals [15] and/or how spatial resolution
89 impacts AOD/PM_{2.5} relationship [16] and the development of mixed effect models with
90 incorporated land use parameters [17,18]. Donkelaar [17] applied a Geographically Weighted
91 Regression (GWR) to estimate that inversion of the PM_{2.5} concentration were highly consistent
92 ($r=0.9$) with PM_{2.5} concentrations from monitors. Ma also predicted PM_{2.5} concentrations using
93 the GWR model with different meteorological parameters (e.g., boundary level height,
94 temperature, and relative humidity) and land use information with Terra MODIS, Aqua
95 MODIS, and MISR AOD data in China (r ranged from 0.73 to 0.84). Chudnovsky [15]
96 developed a Multi-Angle Implementation of Atmospheric Correction (MAIAC) algorithm to
97 retrieved AOD at 1 km resolution and found that the spatial variability of AOD was depend
98 on PM_{2.5} levels and wind speed, their coefficient of determination ($R^2 = 0.47$) in Boston, it is
99 possible that the wind speed affected the boundary layer structure. Therefore, this paper will
100 use the PBLH and hygroscopic growth factor to eliminate this effect.

101 The atmospheric boundary layer is an important indicator for the diffusion of air
102 pollution in the vertical direction [19]. AOD is described as the integral of extinction
103 coefficients from the ground surface to the top of atmosphere. The extinction coefficient is
104 obtained exactly from numerical equation solving [20], for example, it can be calculated by
105 Mie scattering theory for spherical particles [21] and/or be calculated with the T-matrix
106 method [22] and Discrete Dipole Approximate (DDA) method for non-spherical particles [23].
107 Thus, the first scientific question is to fix the effect of the boundary layer height on satellite
108 remotely retrieved PM_{2.5} concentration. Koelemeijer [24] demonstrated that the correlation
109 coefficient between the PM_{2.5} concentration and AOD could be effectively improved from 0.3
110 to 0.6 by the adaption of boundary layer height and relative humidity extracted from the
111 European Centre for Medium-Range Weather Forecasts (ECWMF) data in Europe. A
112 correlation coefficient was also reported as 0.84 between the PM_{2.5} concentration from
113 Tapered Element Oscillating Microbalance (TEOM) and AOD from ground-based sun
114 photometer observations while considering the Lidar-derived boundary layer height
115 correction [25]. Wang [26] indicated that the correlation between surface level aerosol
116 extinction and AOD was improved as a result of vertical correction, with the coefficient of
117 determination R^2 increasing from 0.35 to 0.56. As above-mentioned, the boundary layer height
118 used for vertical correction was mainly obtained from ground-based Lidar and numerical
119 simulation. At the same time, the exported boundary layer height from the large-scale models
120 (in degrees, e.g., ECWMF) was not enough to satisfy the higher inversion resolution (in
121 kilometers) of satellites over the complex terrain areas. However, to our knowledge, no study
122 had systematically compared the effects of boundary layer height from different data sources
123 (such as the model simulation and ground-based observation) on the retrieved ground-level
124 particulate concentrations. This will be investigated in this study.

125 Relative humidity is another important factor to be corrected during the particulate
126 matter inversion from satellites as the hygroscopic properties of water-soluble substances can
127 increase the extinction coefficient of particles [27]. Several humidity correction methods have
128 been developed to retrieve particulate matter concentration from satellite data such as the
129 empirical regression model [28], the semi-empirical regression model [29], the land-use
130 regression model [30], the multiple regression model [31], the geographically weighted
131 regression model [32], and neural network modeling [33]. In the above-mentioned studies, a

132 hygroscopic growth factor was obtained directly from the empirical formula or regression
133 methods. However, no studies have focused on the establishment of a relationship between
134 the extinction coefficient and the relative humidity, which would offer a new hygroscopic
135 growth factor to correct the humidity effect in the retrieval algorithm.

136 The results presented in this paper focused on comparing the effects of boundary layer
137 height from different data sources on the retrieved ground-level particulate concentrations,
138 and in deriving seasonal hygroscopic growth factors to correct the humidity effect from the
139 exponential fitting method. Finally, a near real-time version of the retrieval algorithm with
140 open-source code was introduced with uncertainty analysis.

141 2. Methodology

142 2.1. Study Area Description

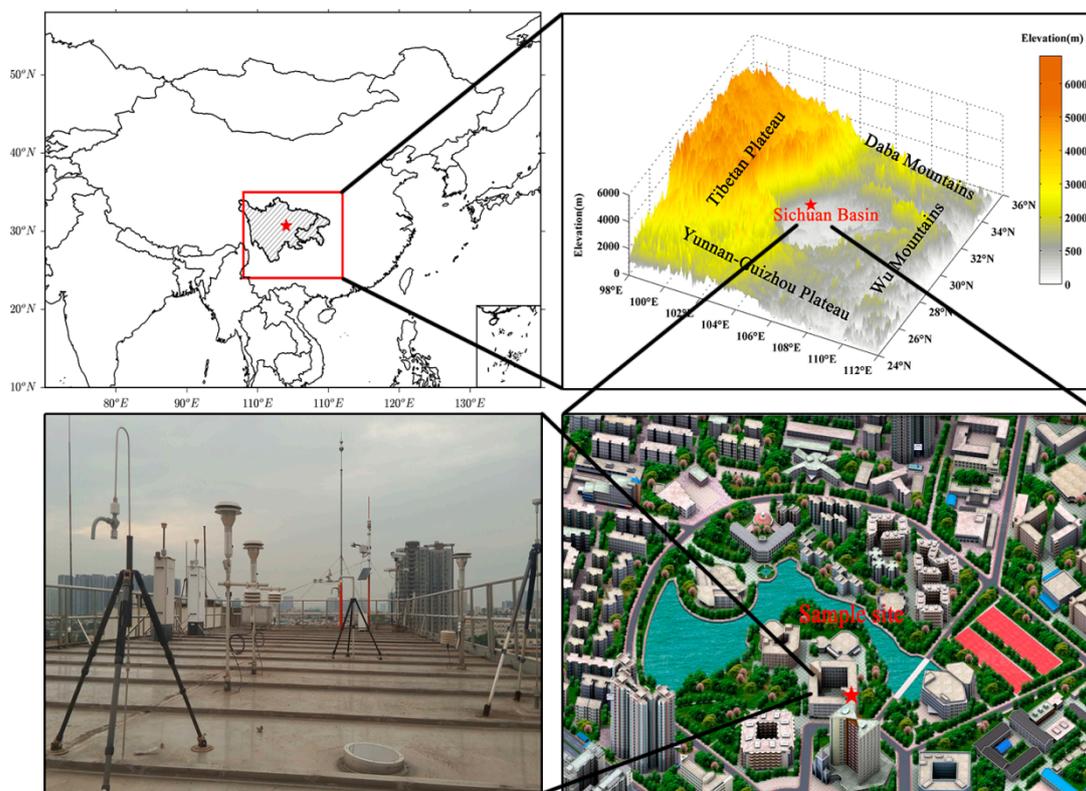
143 The Sichuan Basin (25~35°N, 95°~110°E) covers 260,000 km² of lowland region in
144 Southwestern China, which is surrounded by the Tibetan Plateau to the west, the
145 Yunnan-Guizhou Plateau to the south, the Wu Mountains to the east, and the Daba
146 Mountains to the north (Figure 1). The average altitude of basin is about 500 meters above sea
147 level (ASL) and is much lower than the surrounding mountains (with elevations of 1~3 km
148 ASL). The basin is anchored by the metropolitans of Chengdu-Chongqing, which has a
149 population of approximately 100 million. With accelerated urbanization, intensely emitted
150 atmospheric pollutants and unfavorable diffusion conditions [34] ultimately form regional
151 haze pollution events. The annual PM₁₀ concentration of Sichuan Provincial Environmental
152 Quality Bulletin showed that a period of rapidly declining concentrations from 120 µg/m³ in
153 2004 to 76 µg/m³ in 2008, a period where the concentrations remained stable at around 65
154 µg/m³ from 2009–2011, and a period of the concentrations recovery from 67 µg/m³ in 2012 to
155 80 µg/m³ in 2014 [35]. The ground-level PM_{2.5} concentration was reported publicly after 2013,
156 and is used in this paper.

157 A synergistic suite of monitoring facilities with air quality was situated on a building
158 rooftop at the Southwest Jiao-tong University (30.69°N, 104.05°E; 30 m altitude), which is
159 located in a suburban area of Chengdu. The instruments included a Micro Pulse Lidar (MPL,
160 wavelength at 532 nm; pulse energy at 6~8 µJ; repetition rate at 2500 Hz; Diameter at 178 mm),
161 a Sun-photometer (CIMEL CE-318, wavelength at 532nm), a visibility sensor (FD12P), a semi
162 continuous OC/EC analyzer (SUNSET RT-4), a meteorological station (FRT-X2), and an
163 automatic air quality monitoring station, for analyzing in situ samples for PM_{2.5} and PM₁₀
164 concentrations (Beta Attenuation Particulate Monitor (BAPM), MetOne-1020) and gases
165 (NO_x/NO/NO₂, CO, SO₂, O₃, Hydrocarbons/Total VOC and Benzene).

166 The vertically worked MPL recorded a group of data every 3–4 minutes with a detecting
167 range of 30 km and a resolution of 15 m. Only the AOD, extinction coefficient, and PBLH
168 (which extracted the recorded MPL data and PM_{2.5} concentration from BAPM, and even the
169 observed relative humidity) from the meteorological station were adapted for further analysis
170 in this study.

171 Of all the monitored parameters, the daily PM_{2.5} concentrations and relative humidity
172 were collected from the meteorological station per minute. The vertically worked MPL

173 recorded a group of data every 3–4 minutes with a detecting range of 30 km and a resolution
 174 of 15 m. Only the AOD, extinction coefficient, and PBLH were extracted from the recorded
 175 data for further analysis in this study.



176
 177 **Figure 1.** (a) Locations of the metropolians of Chengdu-Chongqing; (b) Location of the Sichuan
 178 Basin and its elevation; (c, d) Location of the synergistic suite.

179 2.2. Data Sources

180 During the study, the MODIS Level 1 Radiometric and Geolocation data covering the
 181 Sichuan basin were collected through the Level-1 and Atmosphere Archive and Distribution
 182 System (LAADS). Both MODIS Terra and Aqua L1B data (MOD/MYD021KM, Collection 5)
 183 with $1 \text{ km} \times 1 \text{ km}$ resolution were used to retrieve the AOD and water vapor in this study.

184 The Weather Research and Forecasting (WRF) model (version 3.6.1) and the California
 185 Meteorological (CALMET) model (version 6.5.0) were used to calculate the boundary height.
 186 The $1^\circ \times 1^\circ$ of NCEP FNL data were used as WRF input data and were downloaded from the
 187 Computational and Information Systems Laboratory at the National Center for Atmospheric
 188 Research. An additional tool of CALWRF was used to read the WRF model output data and
 189 create a 3D.DAT file for the WRF/CALMET.

190 The extraction and preprocessing of ground-level observed data were conducted
 191 according to the time of satellite passes, i.e., 10:30 AM for MODIS Terra and 1:30 PM for
 192 MODIS Aura. All the above-mentioned data were collected from November 1, 2013 to
 193 October 31, 2014.

194 2.3. Data analysis and retrieval algorithm development

195 Before build the relationship between satellite AOD and ground-based PM_{2.5}
 196 concentrations, the improved dense dark vegetation (DDV) algorithm [36] with calculated
 197 lookup tables (LUTs, built from the radiation transfer model) was adapted to retrieve the
 198 aerosol optical depth over the Sichuan Basin. First, all downloaded Terra/Aqua L1B data were
 199 performed with radiometric correction, geometric correction, and cloud mask. Then, the AOD
 200 was calculated by a linear interpolation between the observed surface reflectance with the
 201 apparent reflectance in the lookup tables. The Second Simulation of the Satellite Signal in the
 202 Solar Spectrum (6S) radiation transmission model was used to calculate the previously
 203 mentioned LUTs [37]. The contents of the H₂O-vapor (g/cm²) and O₃ (cm·atm) are two key
 204 parameters which are required by the 6S model. However, as the column water vapor could
 205 not be directly observed by our meteorological station, it was further calculated with the
 206 MODIS data from the 17, 18, and 19 bands [38]. The acquisition of 3-hourly column ozone
 207 content was collected from the ERA-Interim reanalysis data from the European Centre for
 208 Medium-Range Weather Forecasts (ECMWF). The modified normalized difference water
 209 index (MNDWI) method was used to distinguish water bodies from the land noise [39].

210 In order to retrieval the near-ground particulate concentration, vertical and relative
 211 humidity corrections needed to be further conducted.

212 AOD (τ) is defined as the integral of aerosol extinction along an atmospheric column
 213 from the ground to the top of atmosphere [40], and τ can be calculated by:

$$214 \quad \tau = \int_0^{\infty} \beta_e(z) dz = \int_0^{\infty} \beta_e(0) e^{-\frac{z}{H}} dz = H \beta_e(0) \quad (1)$$

215 where H is the boundary layer height which mainly uses the Lidar and/or numerical
 216 simulation computation during the vertical correction.

217 The blind zone and transition zone height of incomplete after pulse correction (overlap
 218 area) is approximately 200 m [41,42]. An extinction coefficient at a height of 200 to 800 m was
 219 taken as the measured value.

220 WRF mainly simulates the boundary layer height through the four boundary layer
 221 parameterization schemes of GBM [43], UW [44], TEMF [45], and Shin-Hong (SHIN) [46].
 222 CALMET uses the objective analysis method to adjust for kinematic effects of terrain, slope
 223 flows, and terrain blocking effects, and calculates the boundary layer height. The column
 224 extinction coefficient was obtained from the aerosol optical thickness divided by the
 225 boundary layer height.

226 The aerosol chemical components chiefly consist of water-soluble substances such as
 227 sulfates, nitrates, and so on. Absorbing moisture from the atmosphere, fine particles easily
 228 reach the deliquescence point, and eventually the hygroscopic properties affect the optical
 229 properties [47], and the ratio of the water-soluble substances to the particulate mass
 230 concentration was 50.7% in Chengdu [48]. However, the extinction coefficient and relative
 231 humidity from ground monitoring were used as the estimated hygroscopic growth factor; the
 232 satellite inversion of particulate matter concentration is rarely used.

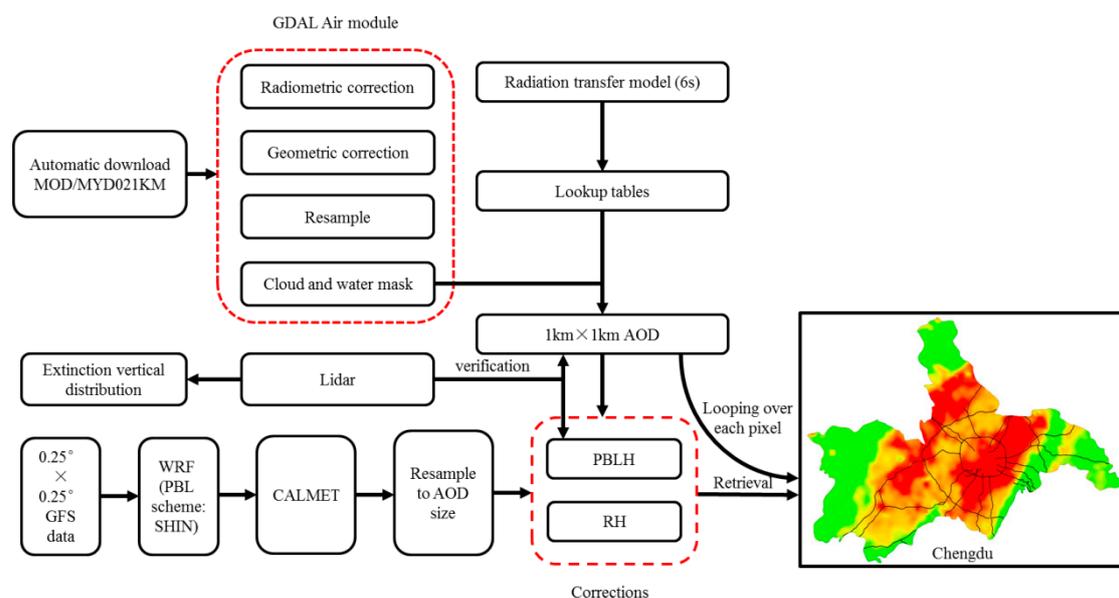
233 By means of the extinction coefficient and relative humidity, a method for calculating the
 234 hygroscopic growth factor in this study using the following formula [49,50]:

$$f(RH) = \frac{\sigma_{ext}(RH)}{\sigma_{ext}(RH_{ref})} = \left(\frac{1-RH}{1-RH_{ref}} \right)^{-\lambda} \quad (2)$$

where the extinction coefficient σ_{ext} depends on the relative humidity RH. λ is the Hänel growth coefficient, which is dependent on the aerosol property. RH_{ref} is the average value of relative humidity < 40% for dry conditions [51].

Finally, we developed the retrieval algorithm with the following three steps to enhance the ability of diagnosis and consultation for local environmental protection agencies, and further improved our retrieval algorithm to a near real-time version with on-line produced visualization products. The workflows are described as follows:

- (1) A script was used to automatically download the near real-time product from NASA. The original downloaded high-resolution MODIS product was first preprocessed via the GDAL_Air module (including the radiometric and geometric corrections, resampling, and masking), then produced a high-resolution AOD product with the application of Lookup tables from the 6S radiation transmission model. At the same time, this AOD are verified by the Lidar AOD
- (2) Before conducting the retrieval of $PM_{2.5}$, the regional WRF model with downscaling to 1 km resolution should be simulated with the GFS initial forecast fields and the SHIN PBL scheme (the WRF model was set to automatically run at 6:00 AM of every day). The wrfout file was further used in the CALMET model, and the final output results of the gridded PBLH and RH should be resampled to the grids of the inversed high-resolution AOD product.
- (3) The PBLH from WRF_SHIN/CALMET was extracted and fitted with the vertical correction function on each pixel of the inversed AOD. The seasonal humidity correction function was also selected at the same time according to the specific date. Finally, the regional gridded $PM_{2.5}$ concentrations were retrieved with vertical and humidity corrections after looping calculations over each pixel of the high-resolution AOD product.



263
264

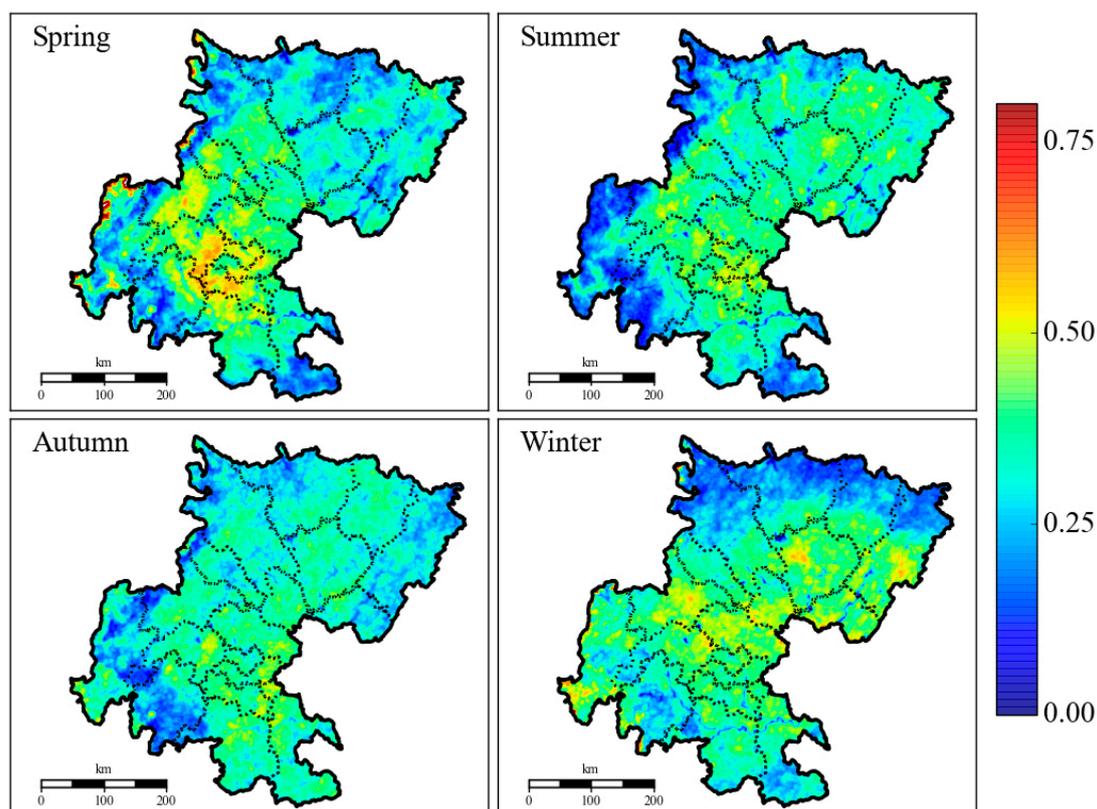
Figure 2. Flowchart of near real-time retrieval algorithm for $PM_{2.5}$ concentration over the Sichuan Basin.

265 Note that if this near real-time algorithm were to be used in other East Asian regions, we
266 recommend the use of the humidity correction function from Song [52], which has been
267 further verified by Tao [53].

268 3. Results and Discussion

269 3.1. Spatial and Temporal Distribution of Retrieved AOD

270 The retrieved AOD distributions (1 km x 1 km) based on the DDV algorithm over the
271 Sichuan Basin in four different seasons are explicitly illustrated in Figure 3. The retrieved
272 results also showed that higher AOD values were distributed in the central and northeastern
273 parts of the Sichuan Basin in winter and spring, especially for the three city groups of
274 Chengdu-Meishan, Bazhong-Dazhou, and Zigong-Yibin-Luzhou, but had homogeneous and
275 lower AOD distributions depicted in summer and autumn. The pattern of spatial distribution
276 for AOD was consistent with the emission distributions of anthropogenic particulates from
277 the Multi-resolution Emission Inventory for China (MEIC) [54], which is mainly related to the
278 urban agglomeration of the Chengdu plain and the northeast agglomeration of industry and
279 transportation development. The highest relative humidity and temperature were reported in
280 July and August in the Sichuan Basin, which would enhance the formation and conversion of
281 secondary organic aerosols (SOAs), and the AOD values presented as higher values when
282 compared to other regions in China (e.g., Northeast China) [55,56]. Regional statistics of
283 annual AOD over the Sichuan Basin from March 2000 to February 2010 was reported as 0.848
284 by MODIS Atmosphere level2 Aerosol Product with resolution $1^{\circ}\times 1^{\circ}$ [55]; however, the latest
285 study from Liu [56] showed that the annual AOD values were 0.33 for the Sichuan Basin and
286 0.65 for the city of Chengdu with the MODIS AOD (3x3km). The revised results from
287 high-resolution MODIS images presented as 0.29 for the Sichuan Basin and 0.57 for Chengdu,
288 respectively.



289

290

291

Figure 3. Spatial and temporal distributions of retrieved AOD from 1 km MODIS over the Sichuan Basin.

292

293

294

295

296

297

298

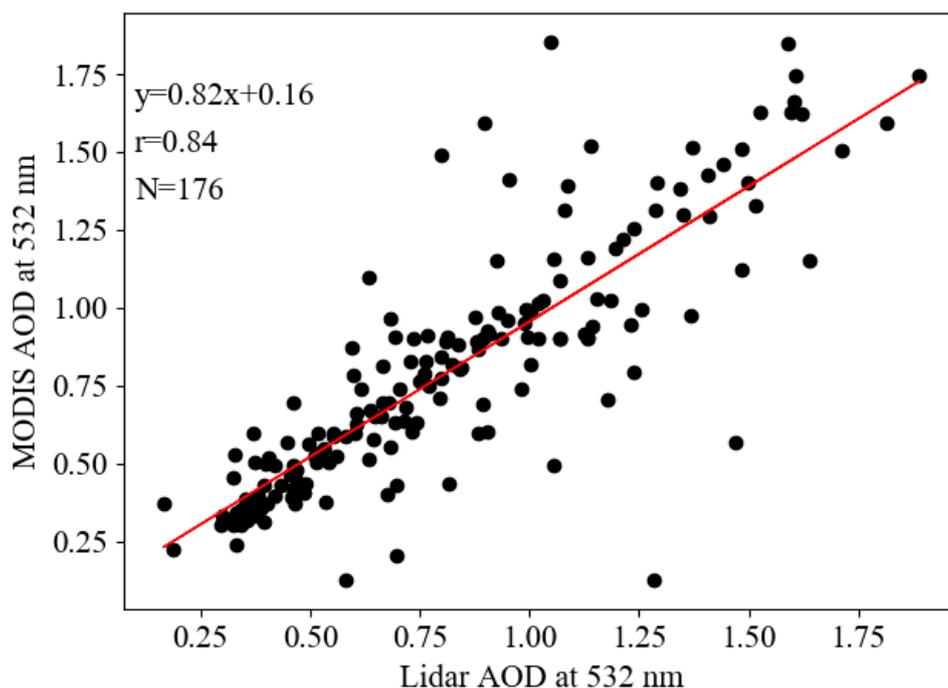
299

300

301

302

To further verify the accuracy of the revised high-resolution AOD, the observed AOD values of ground-based scanning Lidar were collected from the super-monitoring station of Southwest Jiaotong University. As the station is located in the overland part of a suburban area, the observed Lidar AOD was dramatically varied due to the changes in meteorological factors. Thus, the ground-based Lidar data from 176 days were specifically extracted according to the satellite transit time, and the correlation between the MODIS AOD and Lidar AOD is illustrated in Figure 4. The highest correlation ($r = 0.84$) under the condition of 95% confidence indicated that the spatial and temporal distribution of high-resolution AOD could be applied to further retrieve the $PM_{2.5}$ concentrations. A similar coefficient was also reported as 0.78 when compared to the Lidar AOD at a 1 km resolution MODIS AOD taken at the Hong Kong International Airport from January 2006 to September 2007 [57].



303

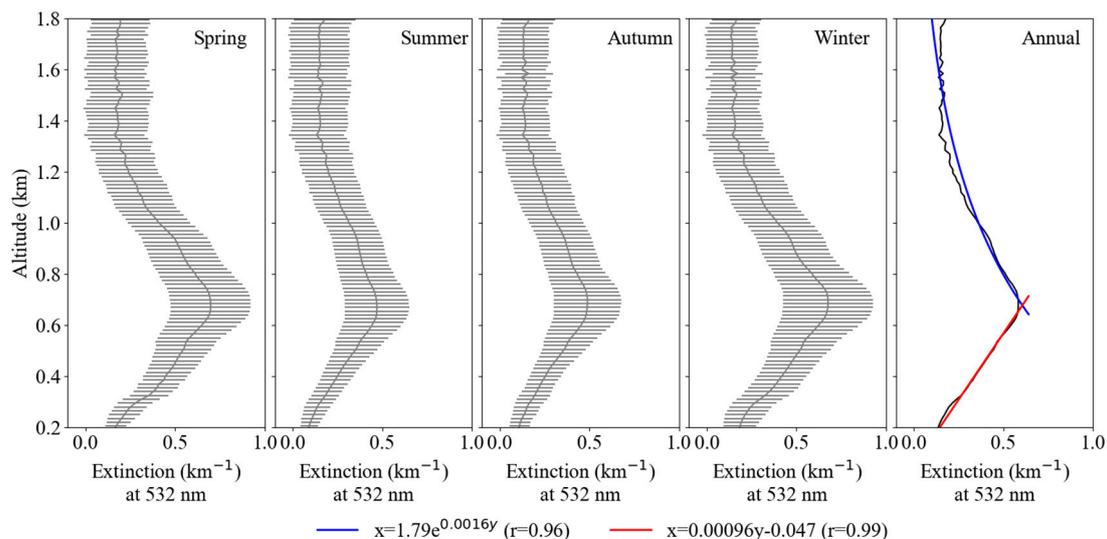
304 **Figure 4.** Correlation between the ground-based Lidar AOD and 1 km resolution MODIS AOD at the
305 site of Southwest Jiao-tong University.

306

3.2. Vertical Correction of Extinction Coefficient

307

308 The vertical distribution of aerosol extinction coefficient (km^{-1}) can be obtained from the
309 back scattered signals received by Lidar, and the atmospheric boundary layer height can be
310 derived using the algorithm presented by He and Mao [58]. Figure 4 shows the seasonal and
311 annual averaged vertical profiles of the aerosol extinction coefficients retrieved from the Lidar
312 measurements at the site of Southwest Jiao-tong University. Strong aerosol extinction was
313 observed at the near ground surface due to the locally emitted aerosols (The horizontal black
314 error bars indicate the maximum and minimum values of the Lidar-derived aerosol extinction
315 coefficient at 0.2 ~ 1.8 km). Then, the coefficient dramatically decreased from 0.7 to 0.02 at a
316 height of 100 m under turbulent flow at vertical direction. The extinction coefficient gradually
317 increased and reached the major peak of 0.58 at the height of 670 m, and then decreased
318 exponentially to about 0.12 at a height of 1800 m. This pattern was obviously different to the
319 observed results with monotonically decreasing trends at Lanzhou and Taiwan in China, and
320 Gwangju in Korea [59,61], but similar to the pattern observed from regional haze events and
dramatic biomass burning in the North China Plain and Singapore [62,63].



321

322 **Figure 5.** Seasonal and yearly mean aerosol vertical extinction profiles at the site of Southwest Jiao-tong
 323 University.

324

325 The time series variations between the monitored $PM_{2.5}$ concentration from the ground
 326 site and the vertical profile of particle extinction are depicted in Figure 6. With the Lidar
 327 measurements, it was found that when the concentration of particulate matter was high, the
 328 high extinction coefficient was mainly concentrated at about 500–800 m in winter and spring.
 329 This can be explained by the shallow PBLH severely limiting the diffusion of pollutants in the
 330 vertical. However, the higher values of particle extinction frequently fluctuated between the
 331 heights of 200 m to 2000 m in summer and autumn. Moreover, it was identified that the
 332 distribution of profiles in different seasons had a similar variation trend, thus a yearly fitting
 333 was conducted to obtain the linear and exponential functions for the vertical correction in
 334 retrieval ground-level $PM_{2.5}$ concentration.

335

336

337

338

339

340

341

342

343

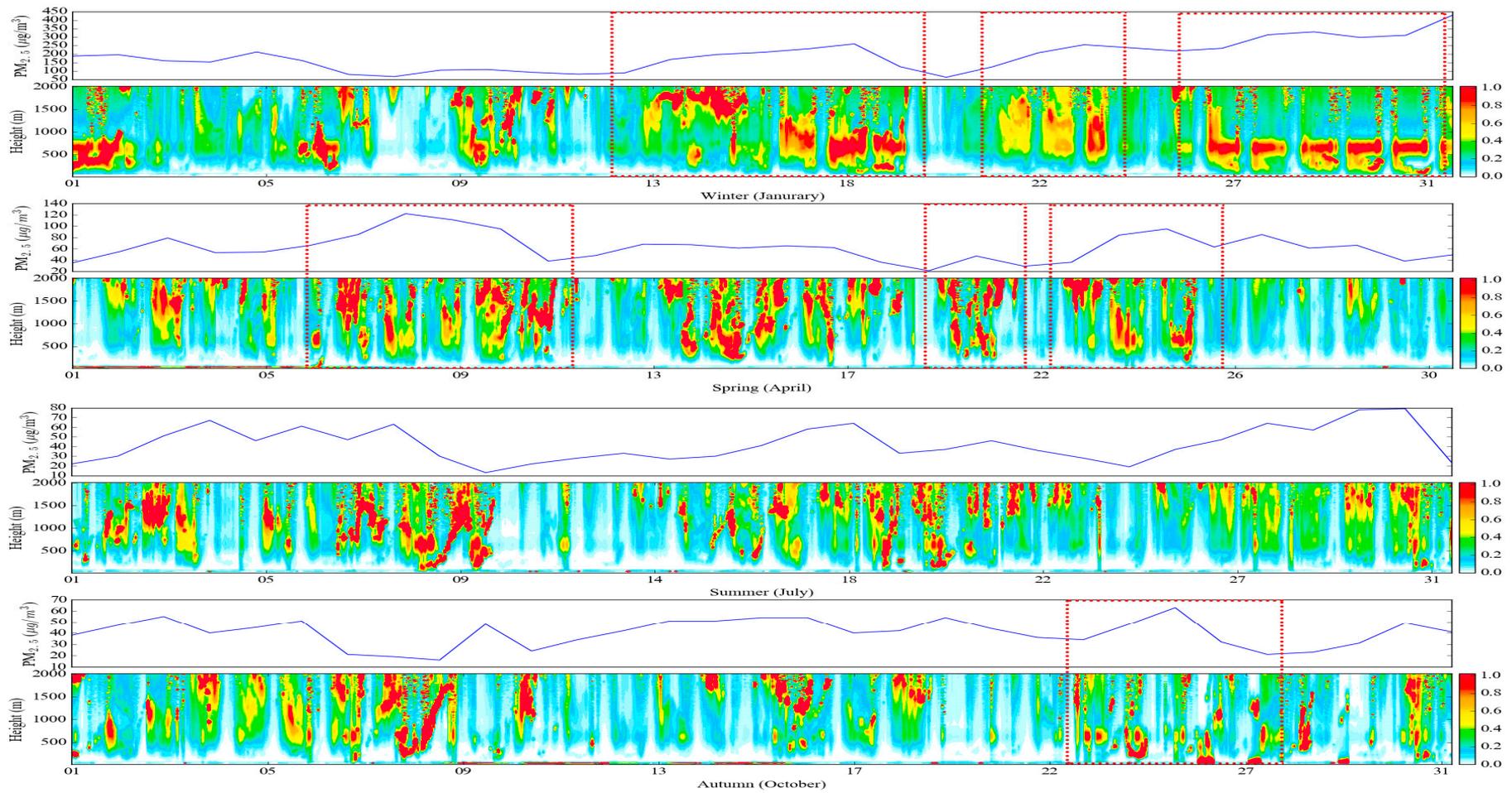
344

345

346

347

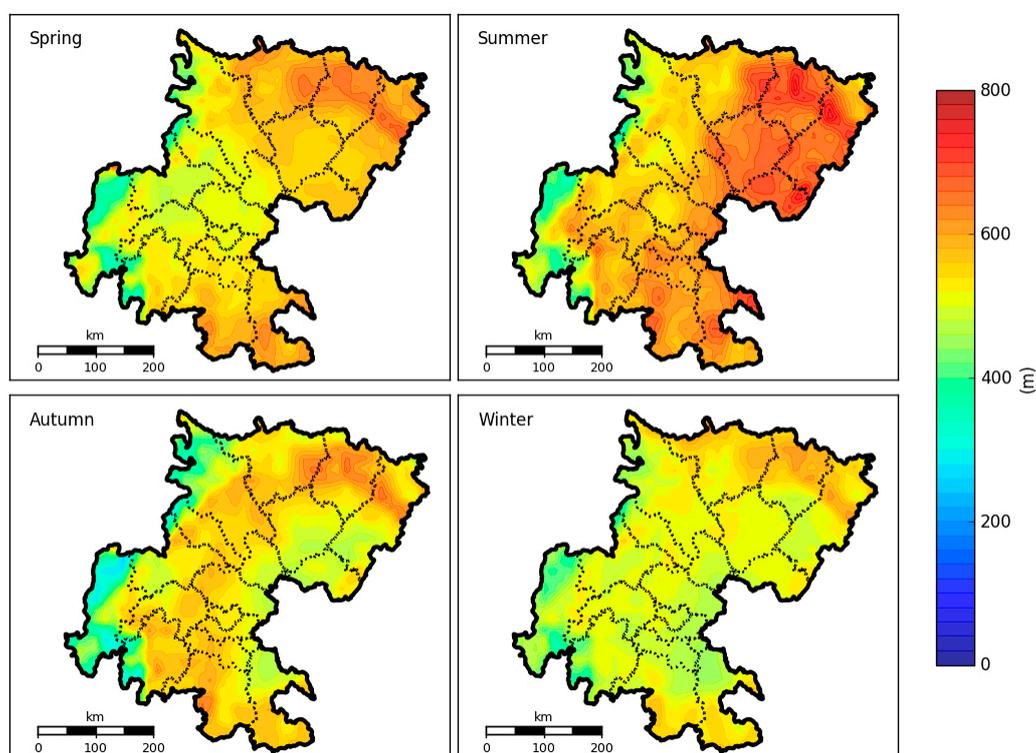
348



349
350

Figure 6. The time series variation of ground-based $PM_{2.5}$ concentrations and vertical profiles of particle extinction over the Sichuan Basin in 2014.

351 The increasingly available Lidar-based aerosol extinction profiles provide insights into
352 the boundary layer as well as the residual above it, but only very limited numbers of the
353 instrument are located in the study regions. To obtain the spatial distribution of PBLH over
354 the whole Sichuan Basin, we simulated the PBLH using the numerical models of WRF and
355 WRF/CALMET, and further compared them with the data from ground-based Lidar. The
356 simulated results for four PBLH parameterization schemes (GBM, SHIN, TEMF, UW) from
357 WRF, along with the result from WRF/CALMET over the Sichuan Basin are illustrated in
358 Figure 7. The results showed that the height of PBLH in winter was the lowest (346 m),
359 followed by autumn (467 m), spring (496 m), and summer (585 m). It also clearly showed that
360 the heights of PBLH were higher in the northeastern part of the Sichuan Basin than in the
361 western part that adjoins to the steep edge of the Tibetan Plateau.



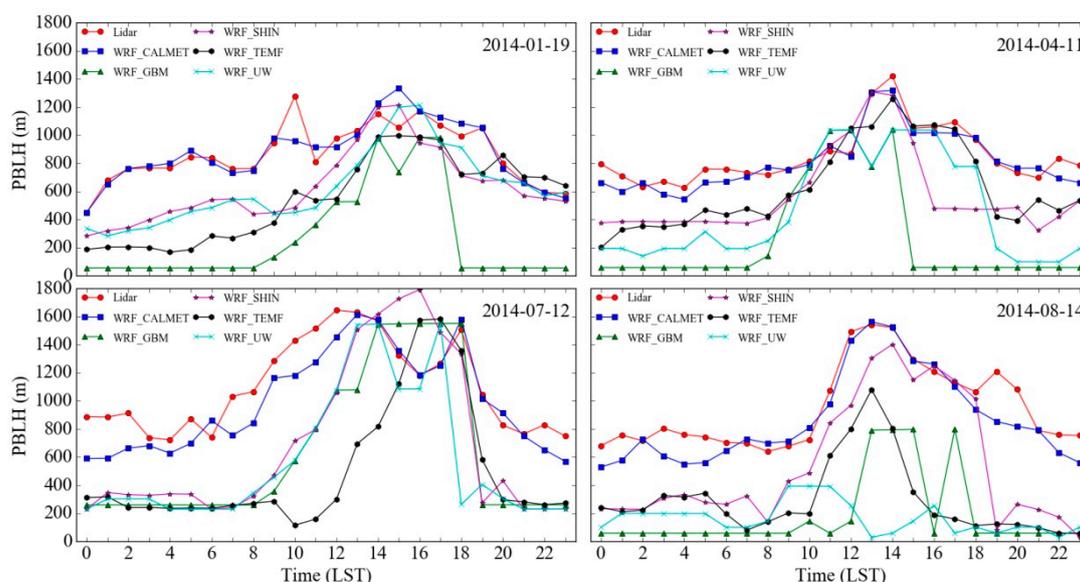
362

363 **Figure 7.** The simulated spatial distribution of planetary boundary layer height over the Sichuan Basin
364 during November 1, 2013 to October 31, 2014.

365

366 In Figure 8, diurnal Lidar observations were selected as the benchmark data, then
367 compared with the extracted WRF and WRF/CALMET simulation results for the site of
368 Southwest Jiaotong University on randomly selected dates (7 January, 21 April, 20 July, and 11
369 August, 2014). The height of the boundary layer observed by Lidar was in good agreement
370 with the simulated boundary layer by the SHIN scheme than with the other three schemes in
371 WRF. The simulation result from WRF with the SHIN scheme was input as an initial-guess
372 field in the CALMET model to further simulate the wind field and PBLH with finer resolution.
373 However, both the correlation and the variation tendency of WRF/CALMET were closest to
374 the observation than that of the four WRF schemes. It is also verified that the WRF/CALMET
375 with surface wind field diagnosis was more flexible than WRF when simulating atmospheric

376 boundary height in complex terrain such as in basin and mountain areas [64].



377

378 **Figure 8.** The evaluation of simulated diurnal variation of PBLH from four WRF PBL schemes and
 379 WRF/CALMET with the observation from ground-based Lidar.

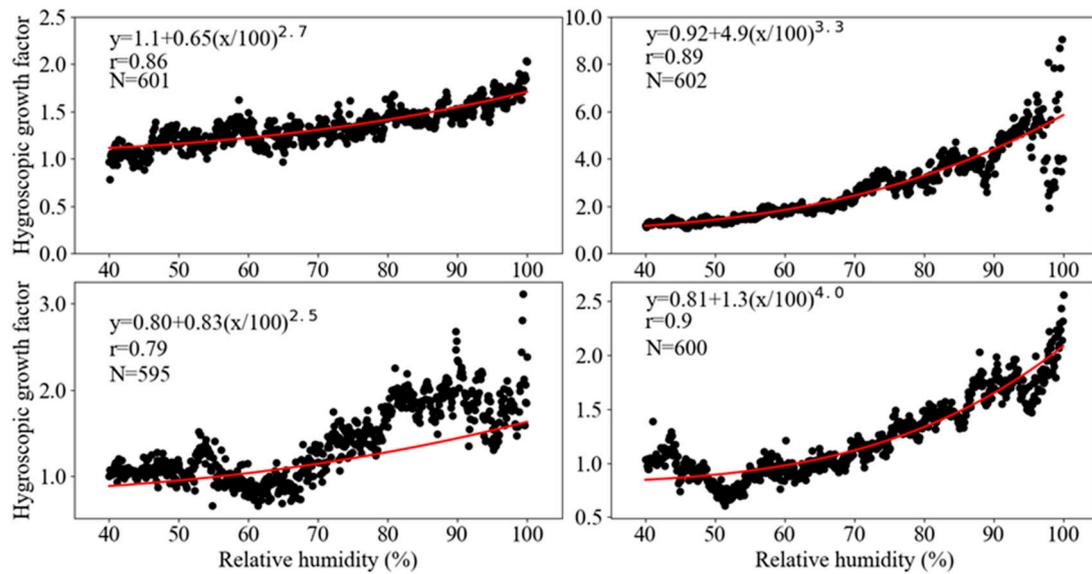
380

381 The simulated vertical distribution of PBLH from WRF_SHIN/CALMAT on the satellite
 382 transit time was adapted to conduct daily vertical correction looping over each grid of the
 383 MODIS images to retrieve the $PM_{2.5}$ concentrations. The NetCDF format results output from
 384 WRF_SHIN/CALMAT were converted to a GeoTIFF format and resampled to meet the
 385 MODIS grids via Python programs, which are available open-source as supplied in the
 386 appendix materials. The vertical optical thickness was corrected by the above obtained
 387 boundary layer height, and the extinction coefficient was used to fit the $PM_{2.5}$ concentration.
 388 The correlation between the extinction coefficient and retrieved $PM_{2.5}$ concentration for the
 389 site of Southwest Jiao-tong University are shown in Figure 10.

390 The results showed that the correlation coefficient r was presented as 0.62 (Figure 10a),
 391 and the reason for this poor correlation was that the extinction coefficient needed to have
 392 further humidity correction conducted.

393 3.3. Humidity Correction of Extinction Coefficient

394 To improve the correlation after the inversion correction, the hygroscopic growth factors
 395 of particles were calculated with the mass extinction efficiency with Equation (2) in Section
 396 2.4; then, an exponential method was used to fit the relationship between the measured
 397 relative humidity and hygroscopic growth factors for the extinction coefficient. The fitting
 398 results for spring, summer, autumn, and winter are compiled in Figure 9. These revealed that
 399 the average mass extinction efficiency of the particulates continuously increased with relative
 400 humidity. The increase trend was moderate when the relative humidity increased from 40–
 401 80%, whereas the growth factor increased rapidly when relative humidity was higher than
 402 80%.



403

404

Figure 9. Correlation between the hygroscopic growth factor and the relative humidity across four seasons over the Sichuan Basin.

405

406

407

408

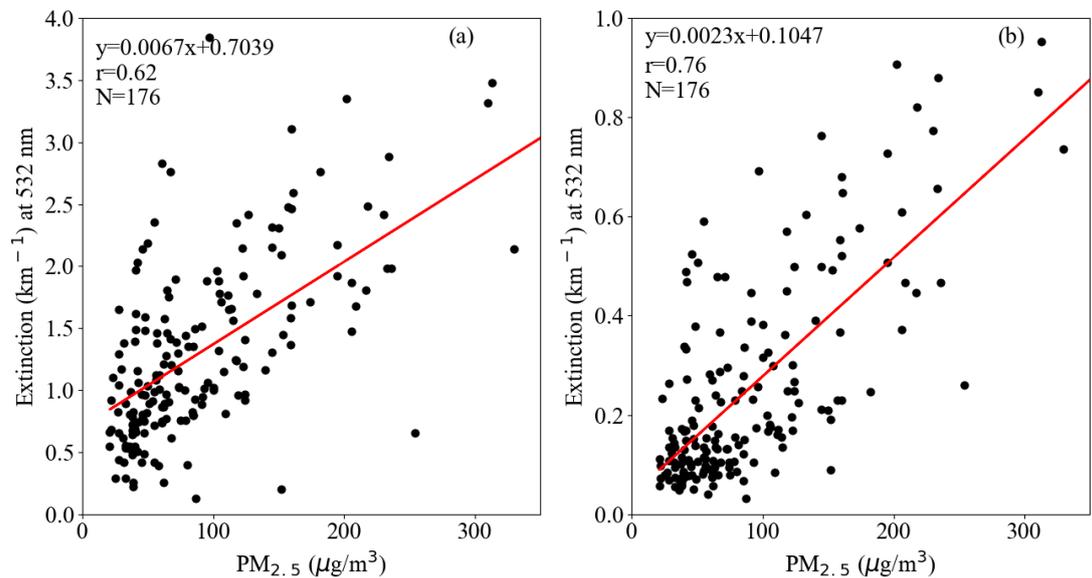
409

410

411

412

The average moisture absorption growth factor was 2.79 in summer, followed by 1.48 in the autumn, 1.23 in spring, and the minimum was 1.22 in winter. The fitting exponential functions had the highest correlation coefficient of 0.9 in winter, followed by 0.79 in summer, 0.86 in spring, and 0.79 in autumn. The fitted exponential formulas were used for humidity correction, and this led to the correlation coefficient between the extinction coefficient and the retrieved $PM_{2.5}$ concentration increasing significantly from 0.62 to 0.76 (Figure 10b).



413

414

Figure 10. Correlation analysis between the aerosol extinction coefficient and the fitted surface-level $PM_{2.5}$ concentration for the site of Southwest Jiao-tong University. (a) Results from vertical correction; and (b) Results from the vertical and humidity corrections.

415

416

417

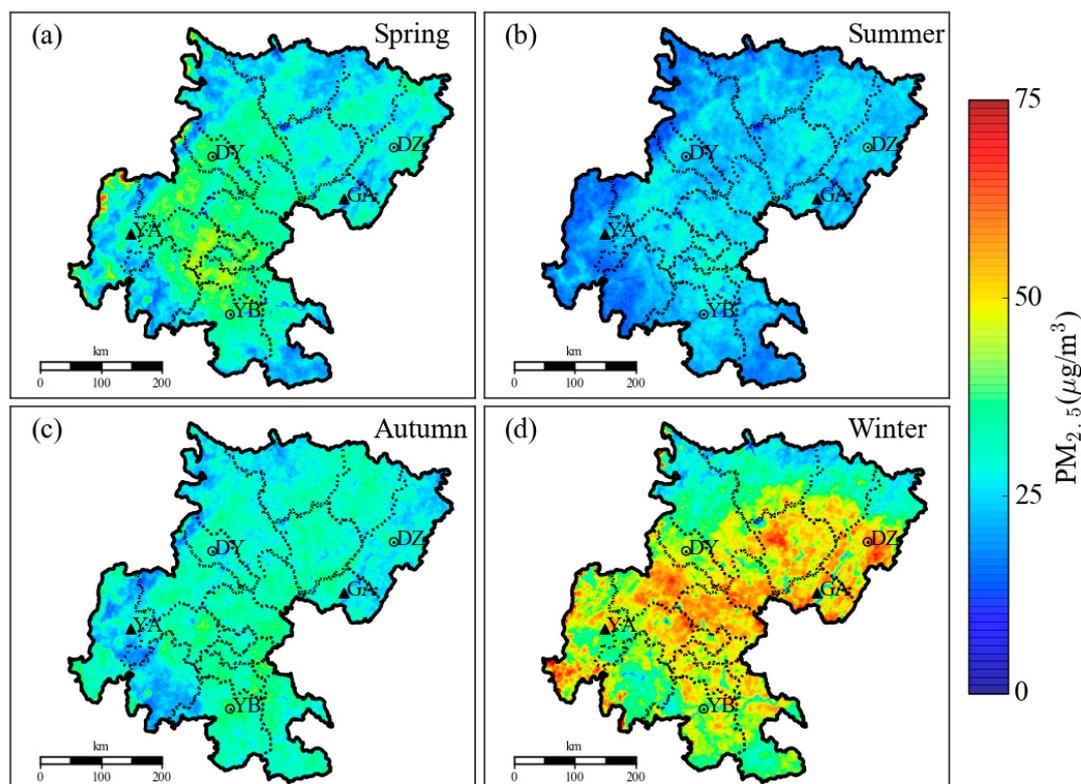
418

419

420

Assuming that relative humidity varies smoothly, the spatial distribution of relative humidity can be used to retrieve the temporal and spatial distribution of ground-level $PM_{2.5}$ concentrations (Figure 11). First, we calculate the dry extinction coefficient based on the fitted

421 four-season hygroscopic growth factors, and then establish the relationship between the dry
 422 extinction coefficient and the particulate matter concentration. Second, the relative humidity
 423 in the Sichuan Basin was calculated by WRF SHIN / CALMET, and then the hygroscopic
 424 growth factor for each grid of $1\text{km} \times 1\text{km}$ (matching the inversion AOD) was computed using
 425 the formula (Eq. 2) in section 2.3. Finally, use the vertical correction extinction coefficient
 426 divided by hygroscopic growth factor is the concentration of particulate matter. It was clear
 427 that the highest retrieved $\text{PM}_{2.5}$ concentrations presented at the bottom zone of the Sichuan
 428 Basin in winter, especially for the city groups of Chengdu-Meishan, Bazhong-Dazhou, and
 429 Zigong-Yibin-Luzhou (Figure 11a). Throughout the remaining cities of the Basin, the
 430 contribution to the $\text{PM}_{2.5}$ emission was also larger, suggesting that the control of haze in the
 431 Sichuan Basin demands further regional joint prevention and controls. $\text{PM}_{2.5}$ pollution in
 432 spring was still serious when excluding the northern city group of Bazhong-Dazhou (Figure
 433 11b), but the air quality in most cities was generally good in summer and autumn (Figures 11c
 434 and 11d). The seasonal variation and distribution characteristics of fine particulates could be
 435 well identified via the retrieval of high-resolution MODIS images, which was consistent with
 436 the monitoring result that the worst air quality occurred at the bottom zone (11 cities) of the
 437 Basin in winter, and the second highest pollution occurred across the Basin excluding Yaan
 438 (YA) and Guangan (GA) in spring [51].



439
 440 **Figure 11.** The temporal and spatial distribution of retrieved high-resolution $\text{PM}_{2.5}$ concentrations across
 441 the Sichuan Basin during November 1, 2013 to October 31, 2014.

442

443 To evaluate the rationality of the spatial distribution for the retrieved $\text{PM}_{2.5}$ concentration,
 444 we compared the retrieved values with the ground-observed $\text{PM}_{2.5}$ concentration at three
 445 other stations in the Sichuan Basin from January to October, 2014 (Figure 12). Three sites were

446 selected by considering the discrete distribution around the basin to evaluate the retrieval
447 accuracy, i.e., Yibin (YB), Dazhou (DZ), and Deyang (DY) (circle marked sites in Figure 11).
448 The results indicated that the retrieved PM_{2.5} concentrations were slightly overestimated for
449 severe pollution events in winter, but were slightly underestimated in spring at the site of DY.
450 The correlation coefficients for the evaluation during a whole year were 0.78, 0.77, and 0.81
451 for the sites of YB, DZ, and DY, respectively. The retrieved seasonal distribution of PM_{2.5}
452 concentration was closer to the observed results in DY than in YB and DZ.

453

454

455

456

457

458

459

460

461

462

463

464

465

466

467

468

469

470

471

472

473

474

475

476

477

478

479

480

481

482

483

484

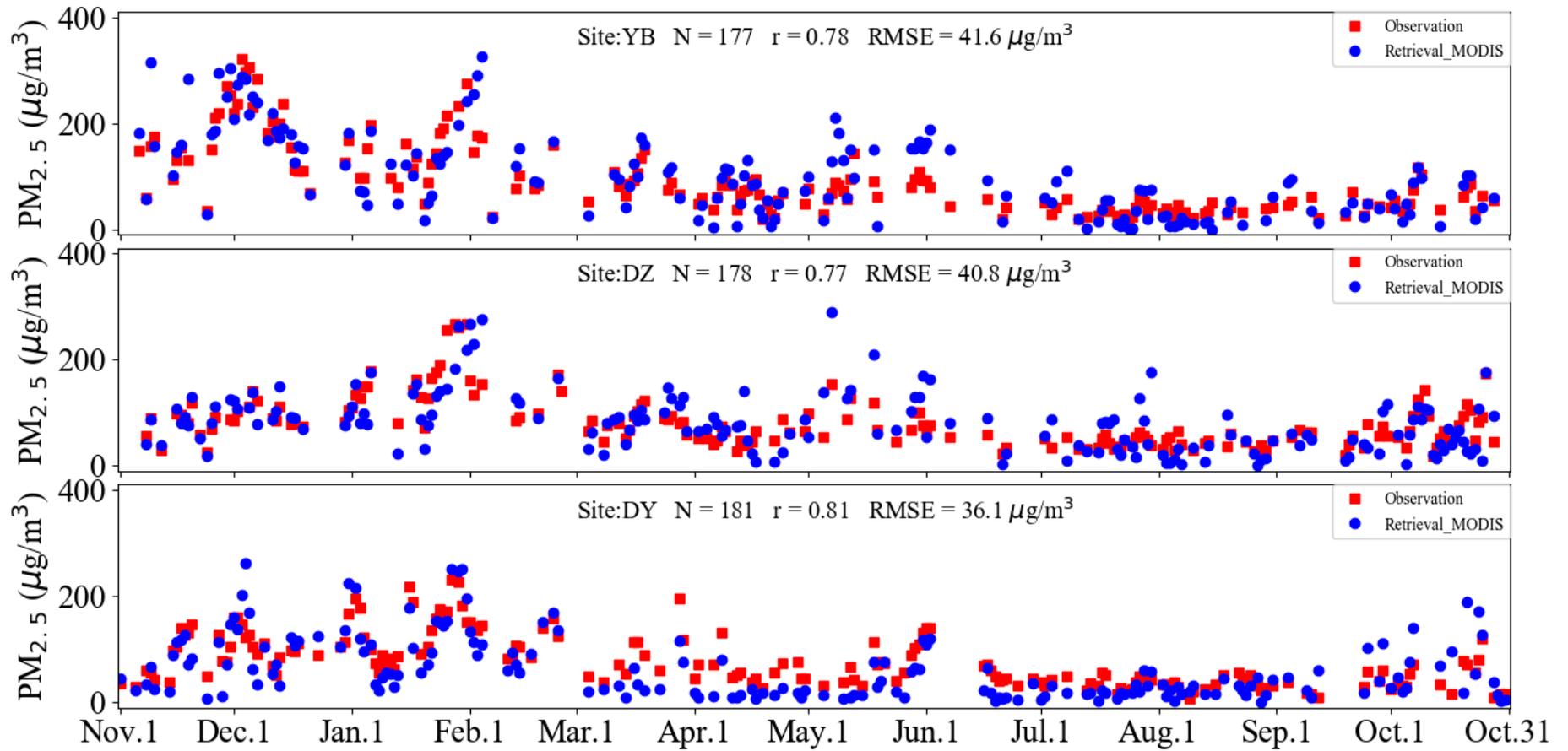
485

486

487

488

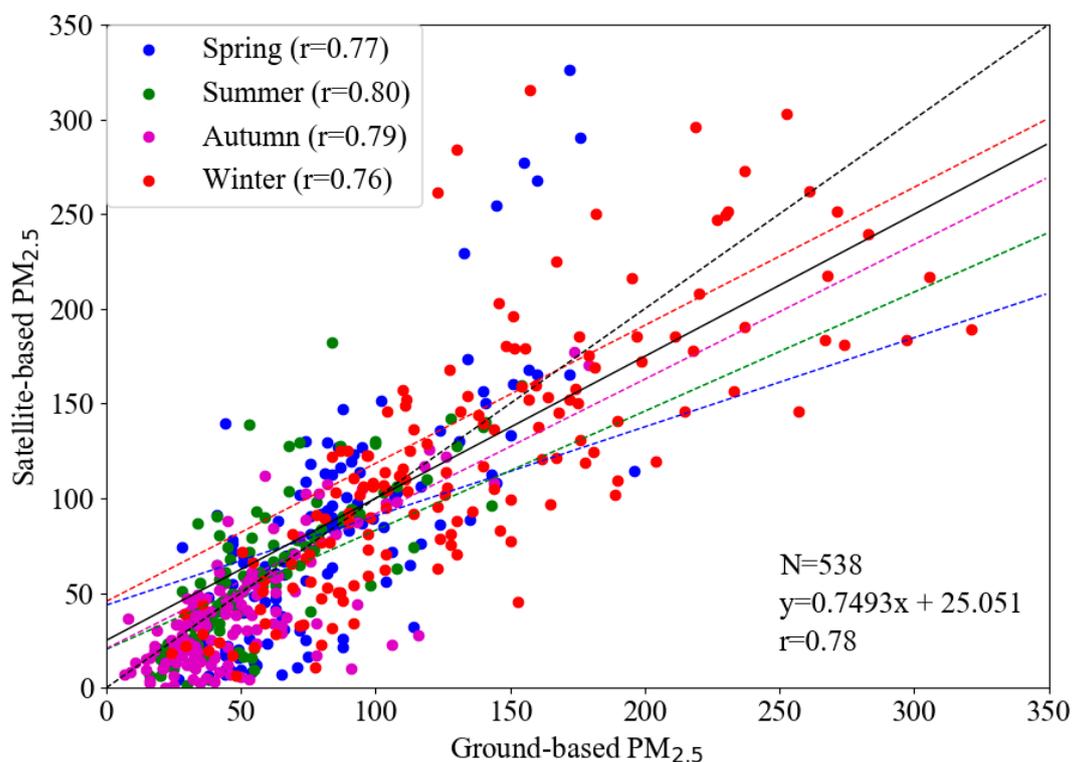
489



490
491 **Figure 12.** Time series evaluation of retrieved ground-level $PM_{2.5}$ with surface measurements at the sites of Yibin, Dazhou, and Deyang during November 1, 2013 to
492
493 October 31, 2014.

490
491
492
493

494 Annual average MODIS retrieved $PM_{2.5}$ concentration over all three sites (80.80 ± 50.68
 495 $\mu\text{g}/\text{m}^3$) is approximately 2.61% lower than the ground-based measurements (82.96 ± 55.63
 496 $\mu\text{g}/\text{m}^3$) with a similar variation dynamic range. A similar study, Liu [65] found that the
 497 annual average over with spatial scale of 17.6 km MISR retrieved $PM_{2.5}$ concentration over
 498 the United States is approximately 10% lower than the EPA measurements, which further
 499 suggested that the application of higher spatial resolution AOD data could significantly
 500 improve the retrieval agreement. The MODIS retrieved $PM_{2.5}$ concentration are generally
 501 higher than the ground-based measurements in YB (7.28%) and DZ (1.05%) which located in
 502 eastern Sichuan Basin, but lower in the western part, such as DY (6.69%). A scatterplot in
 503 Figure 13 illustrated that annual average MODIS retrieved $PM_{2.5}$ concentrations have a good
 504 linear relationship with ground-based measurements ($r = 0.78$, linear regression slope =
 505 0.7493).



506

507 **Figure 13.** Comparison of daily averaged $PM_{2.5}$ retrieved from MODIS and ground-observed (site: YB, DZ
 508 and DY) in different seasons.

509

510 Seasonally, the difference between MODIS $PM_{2.5}$ and ground-based $PM_{2.5}$ measurements
 511 is the largest in the summer ($64.94 \mu\text{g}/\text{m}^3$ versus $57.20 \mu\text{g}/\text{m}^3$, 13.54% difference) and
 512 smallest in the spring ($81.36 \mu\text{g}/\text{m}^3$ versus $82.41 \mu\text{g}/\text{m}^3$, -1.27% difference). The MODIS $PM_{2.5}$
 513 concentrations are generally lower than ground-based measurements except during the
 514 summer (13.54% higher).

515 We also compared the daily averaged $PM_{2.5}$ retrieved from MODIS with the
 516 ground-observed values from the sites of YB, DZ and DY in different seasons (Figure 13).
 517 Good agreements are shown with $r = 0.78$ ($N = 538$) for $PM_{2.5}$ throughout the whole year and
 518 $r = 0.77$, 0.80, 0.79 and 0.76 in spring, summer, autumn and winter, respectively. More dust
 519 storms and biomass burning occur in spring and winter could result in larger variations in

520 aerosol types, which could decrease the correlation coefficient [66]. The vegetated land
521 surface is facilitated by satellite retrievals in the summer [67].

522 In this analysis, it is obviously that the capability of 1km MODIS AOD to predict
523 ground-based PM_{2.5} concentrations can be substantially enhanced by vertical and humidity
524 corrections with model simulated parameters. A database of boundary layer height and
525 humidity growth factor for the Sichuan Basin region was built for use in retrieving the
526 high-resolution ground-level PM_{2.5} concentrations from the 1 km MODIS AOD products.
527 The uncertainties were mainly produced from assumptions such as the fitted seasonal
528 exponential function of humidity growth factor, and the GFS meteorological initial forecast
529 fields, which are used in daily retrieval works. Moreover, during the inversion progress of
530 aerosol optical depth, only eight fixed aerosol modes are provided in the 6S radiative
531 transfer model, which cannot fully reflect the complex optical properties of the entire
532 Sichuan Basin. The model simulated PBLH still has bias with comparison the actual height
533 and this would lead to uncertainty of vertical correction. Moreover, another important
534 uncertainty came from the usage of model simulated surface-level humidity, however, the
535 columnar integrated humidity should be applied. Ground-based observations from a
536 sun-photometer should be conducted to obtain more detailed optical parameters for future
537 works. In addition, only one super-monitoring station was selected as the benchmark and
538 three ground-based observation sites were adopted for daily product evaluation, surface
539 observations from more sites, and longer periods would be useful to further verify the
540 applicability of AOD-derived PM_{2.5}.

541 4. Conclusions

542 Although 140 ground-level monitoring stations have been built across 21 cities of the
543 Sichuan Province since 2014, they still have limited spatial resolution and coverage. Recently,
544 multi-covariates model-based approaches have been reported to retrieve ground-level PM_{2.5}
545 concentrations from high-resolution MODIS AOD products. In this study, the ground
546 observed data from a super-monitoring station were selected as the benchmark; and we
547 combined the WRF_SHIN/CALMET simulated PBLH data and meteorological data to
548 account for the vertical correction and the integrated seasonal humidity effect in the retrieval
549 of PM_{2.5} concentrations. An algorithm with an open-source code was developed to estimate
550 PM_{2.5} concentration over the Sichuan Basin based on the near real-time obtained MODIS AOD
551 at 1 km resolution. The correlation coefficient between the aerosol extinction coefficient and
552 the fitted surface-level PM_{2.5} concentration at the benchmark station was significantly
553 enhanced from 0.62 to 0.76 after vertical and humidity corrections during a whole year period.
554 Further evaluation of the retrieved ground-level PM_{2.5} with the observed values in three cities,
555 Yibin (YB), Dazhou (DZ), and Deyang (DY), showed generally better agreements, with the
556 correlation coefficients of 0.78 (N = 177), 0.77 (N=178), and 0.81 (N = 181), respectively. The
557 results showed that the near real-time algorithm had the capacity to identify PM_{2.5}
558 spatiotemporal distribution on regional and urban scales with complex basin terrain, and to
559 provide helpful information for the diagnosis, consultation, and control of heavy haze events
560 in environmental protection agencies.

561

562 **Acknowledgments:** This work was financially supported by the National Natural Science Foundation
563 of China (NSFC) (No. 41771071 and 21407148), the National Key R&D Plan (No. 2017YFC0212304),
564 Frontier Science Research Plan of CAS (No. QYZDB-SSW-DQC045), and the Youth Innovation
565 Promotion Association of CAS (No. 2017275). The authors would like to thank Qingqing Tong and
566 Xinyuan Cao for their recommendations for improving this manuscript.

567

568 **Author Contributions:** Chao Gao and Wenyong Wang carried out the data collection and related
569 analysis; Chao Gao, Xuelei Zhang, and Weiwei Chen plotted the figures and wrote this manuscript;
570 Chao Gao wrote the open-source codes; and Aijun Xiu and Danile Q. Tong gave modification
571 suggestions for this manuscript.

572

573 **Code availability:** The open-source codes for the retrieval algorithm can be accessed in the
574 supplementary file.

575

576 **Conflicts of Interest:** The authors declare no conflict of interest.

577

578 **References**

579

- 580 1. Pöschl, U. Atmospheric aerosols: composition, transformation, climate and health
581 effects. *Angew. Chem. Int. Edit.* **2005**, *44*, 7520-7540.
- 582 2. Mahowald, N.; Ward, D.S.; Kloster, S.; Flanner, M.G.; Heald, C.L.; Heavens, N.G.; Hess,
583 P.G.; Lamarque, J.F.; Chuang, P.Y. Aerosol impacts on climate and biogeochemistry.
584 *Annu. Rev. Env. Resour.* **2011**, *36*, 45-74.
- 585 3. Lohmann, U.; Feichter, J. Global indirect aerosol effects: a review. *Atmos. Chem. Phys.*
586 **2005**, *5*, 715-737.
- 587 4. Rd, P.C.; Burnett, R.T.; Thun, M.J.; Calle, E.E.; Krewski, D.; Ito, K.; Thurston, G.D. Lung
588 cancer, cardiopulmonary mortality, and long-term exposure to fine particulate air
589 pollution. *Jama.* **2002**, *287*, 1132.
- 590 5. Chen, L.C.; Lippmann, M. Effects of metals within ambient air particulate matter (PM)
591 on human health. *Inhal. Toxicol.* **2009**, *21*, 1-31.
- 592 6. Huang, R.J.; Zhang, Y.; Bozzetti, C.; Ho, K.F.; Cao, J.J.; Han, Y.; Daellenbach, K.R.;
593 Slowik, J.G.; Platt, S.M.; Canonaco, F.; Zotte, P.; Wolf, R.; Pieber, S.M.; Bruns, E.A.;
594 Crippa, M.; Ciarelli, G.; Piazzalunga, A.; Schwikowski, M.; Abbaszade, G.;
595 Schnelle-Kreis, J.; Zimmermann, R.; An, Z.; Szidat, S.; Baltensperger, U.; El Haddad, I.;
596 Prévôt, A.S. High secondary aerosol contribution to particulate pollution during haze
597 events in China. *Nature.* **2014**, *514*, 218.
- 598 7. Zhang, X.Y.; Wang, Y.Q.; Niu, T.; Zhang, X.C.; Gong, S.L.; Zhang, Y.M.; Sun, J.Y.;
599 Brandt, J. Atmospheric aerosol compositions in China: spatial/temporal variability,
600 chemical signature, regional haze distribution and comparisons with global aerosols.
601 *Atmos. Chem. Phys.* **2012**, *12*, 779-799.
- 602 8. Hoff, R.M.; Christopher, S.A. Remote sensing of particulate pollution from space: have
603 we reached the promised land?. *Japca. J. Air. Waste. Ma.* **2009**, *59*, 645-675.
- 604 9. Gupta, P.; Christopher, S.A.; Wang, J.; Gehrig, R.; Lee, Y.; Kumar, N. Satellite remote
605 sensing of particulate matter and air quality assessment over global cities. *Atmos.*

- 606 *Environ.* **2006**, *40*, 5880-5892.
- 607 10. Van Donkelaar, A.; Martin, R.V.; Brauer, M.; Boys, L.B. Global Fine Particulate Matter
608 Concentrations from Satellite for Long-Term Exposure 2 Assessment 3. *Assessment*.
609 **2015**, *3*, 4.
- 610 11. Nicolantonio, W.D.; Cacciari, A.; Bolzacchini, E.; Ferrero, L.; Volta, M.; Pisoni, E.
611 MODIS aerosol optical properties over North Italy for estimating surface-level PM_{2.5}.
612 *Proceedings of Envisat Symposium*. **2007**, *636*, 23-27
- 613 12. Van Donkelaar, A.; Martin, R.V.; Brauer, M.; Kahn, R.; Levy, R.; Verduzco, C.;
614 Villeneuve, P.J. Global estimates of ambient fine particulate matter concentrations from
615 satellite-based aerosol optical depth: development and application. *Environ. Health*.
616 *Persp.* **2010**, *118*, 847.
- 617 13. Van Donkelaar, A.; Martin, R.V.; Park, R.J. Estimating ground-level PM_{2.5} using aerosol
618 optical depth determined from satellite remote sensing. *J. Geophys. Res-atmos.* **2006**, *111*,
619 5049-5066.
- 620 14. You, W.; Zang, Z.; Pan, X.; Zhang, L.; Chen D. Estimating PM_{2.5} in Xi'an, China using
621 aerosol optical depth: A comparison between the MODIS and MISR retrieval models.
622 *Sci. Total. Environ.* **2015**, *505*, 1156-1165.
- 623 15. Chudnovsky, A.A.; Kostinski, A.; Lyapustin, A.; Koutrakisa, P. Spatial scales of
624 pollution from variable resolution satellite imaging. *Environ. Pollut.* **2013**, *172*, 131-138.
- 625 16. Lyapustin, A.I.; Wang, Y.; Laszlo, I.; Hilker, T.; Hall, F.G.; Sellers, P.J.; Tucker, C.J.;
626 Korkin, S.V. Multi-angle implementation of atmospheric correction for MODIS
627 (MAIAC): 3. Atmospheric correction. *Remote. Sens. Environ.* **2012**, *127*, 385-393.
- 628 17. Van Donkelaar, A.; Martin, R.V.; Brauer, M.; et al. Hsu, N.C.; Kahn, R.A.; Levy, R.C.;
629 Lyapustin, A.; Winker, D.M. Global estimates of fine particulate matter using a
630 combined geophysical-statistical method with information from satellites, models, and
631 monitors. *Environ. Sci. Technol.* **2016**, *50*, 3762-3772.
- 632 18. Ma, Z.; Hu, X.; Huang, L.; Bi, J.; Liu, Y. Estimating ground-level PM_{2.5} in China using
633 satellite remote sensing. *Environ. Sci. Technol.* **2014**, *48*, 7436-7444.
- 634 19. Wilson, J.D.; Sawford, B.L. Review of Lagrangian stochastic models for trajectories in
635 the turbulent atmosphere. *Bound-lay. Meteorol.* **1996**, *78*, 191-210
- 636 20. Hauschildt, P.H.; Baron, E. Numerical solution of the expanding stellar atmosphere
637 problem. *J. Comput. Appl. Math.* **1999**, *109*, 41-63.
- 638 21. Hess, M.; Koepke, P.; Schult, I. Optical properties of aerosols and clouds: The software
639 package OPAC. *B. Am. Meteorol. Soc.* **1998**, *79*, 831-844.
- 640 22. Mishchenko, M.I.; Travis, L.D.; Mackowski, D.W. T-matrix computations of light
641 scattering by nonspherical particles: a review. *J. Quant. Spectrosc. Ra.* **1996**, *55*, 535-575.
- 642 23. Draine, B.T.; Flatau, P.J. Discrete-dipole approximation for scattering calculations.
643 *JOSAA*. **1994**, *11*, 1491-1499.
- 644 24. Koelemeijer, R.B.A.; Homan, C.D. Matthijsen, J. Comparison of spatial and temporal
645 variations of aerosol optical thickness and particulate matter over Europe. *Atmos*.
646 *Environ.* **2006**, *40*, 5304-5315.
- 647 25. Boyouk, N.; Léon, J.F.; Delbarre, H.; Podvin, T.; Deroo, C. Impact of the mixing
648 boundary layer on the relationship between PM_{2.5} and aerosol optical thickness.
649 *Atmos. Environ.* **2010**, *44*, 271-277.

- 650 26. Wang, Z.; Chen, L.; Tao, J.; Zhang, Y.; Su, L. Satellite-based estimation of regional
651 particulate matter (PM) in Beijing using vertical-and-RH correcting method. *Remote.*
652 *Sens. Environ.* **2010**, *114*, 50-63.
- 653 27. Flores, J. Michel; Bar-Or, R.Z.; Bluvshstein, N.; Abo-Riziq, A.; Kostinski, A.; Borrmann,
654 S.; Koren, I.; Koren, I.; Rudich, Y. Absorbing aerosols at high relative humidity: linking
655 hygroscopic growth to optical properties. *Atmos. Chem. Phys.* **2012**, *12*, 5511-5521.
- 656 28. Kumar, N.; Chu, A.; Foster, A. An empirical relationship between PM_{2.5} and aerosol
657 optical depth in Delhi Metropolitan. *Atmos. Environ.* **2007**, *41*, 4492-4503.
- 658 29. Tian, J.; Chen, D. A semi-empirical model for predicting hourly ground-level fine
659 particulate matter (PM_{2.5}) concentration in southern Ontario from satellite remote
660 sensing and ground-based meteorological measurements. *Remote. Sens. Environ.* **2010**,
661 *114*, 221-229.
- 662 30. Mao, L.; Qiu, Y.; Kusano, C.; Xu, L. Predicting regional space-time variation of PM_{2.5}
663 with land-use regression model and MODIS data. *Environ. Sci. Pollut. R.* **2012**, *19*,
664 128-138.
- 665 31. Gupta, P.; Christopher, S.A. Particulate matter air quality assessment using integrated
666 surface, satellite, and meteorological products: Multiple regression approach. *J.*
667 *Geophys. Res-atmos.* **2009a**, *114*, 311-311.
- 668 32. Song, W.; Jia, H.; Huang, J.; Zhang, Y. A satellite-based geographically weighted
669 regression model for regional PM_{2.5} estimation over the Pearl River Delta region in
670 China. *Remote. Sens. Environ.* **2014**, *154*, 1-7.
- 671 33. Gupta, P.; Christopher, S.A. Particulate matter air quality assessment using integrated
672 surface, satellite, and meteorological products: 2. A neural network approach. *J.*
673 *Geophys. Res-atmos.* **2009b**, *114*, 311-311.
- 674 34. Chen, Y.; Xie, S. Temporal and spatial visibility trends in the Sichuan Basin, China,
675 1973 to 2010. *Atmos. Res.* **2012**, *112*, 25-34.
- 676 35. Environmental Quality Report of Sichuan Province in 2015. Available online:
677 <http://www.schj.gov.cn/hjgl/hjjcydc/hjzkgb/201506/P020170804637398461312.pdf>
678 (accessed on 1 May 2015) (In Chinese)
- 679 36. Levy, R.C.; Remer, L.A.; Dubovik, O. Global aerosol optical properties and application
680 to Moderate Resolution Imaging Spectroradiometer aerosol retrieval over land. *J.*
681 *Geophys. Res-atmos.* **2007**, *112*, 3710-3711.
- 682 37. Kotchenova, S.Y.; Vermote, E.F.; Matarrese, R.; Klemm, F.J. Validation of a vector
683 version of the 6S radiative transfer code for atmospheric correction of satellite data.
684 Part I: Path radiance. *Appl. Optics.* **2006**, *45*, 6762-6774.
- 685 38. Gao, B.C.; Kaufman, Y.J. Water vapor retrievals using Moderate Resolution Imaging
686 Spectroradiometer (MODIS) near-infrared channels. *J. Geophys. res-atmos.* **2003**, *108*,
687 1007-1021.
- 688 39. Xu, H.Q. Modification of normalized difference water index (NDWI) to enhance open
689 water features in remotely sensed imagery. *Int. J. Remote. Sens.* **2006**, *27*, 3025-3033.
- 690 40. Liu, Y.; Sarnat, J.A.; Kilaru, V.; Jacob, D.J.; Koutrakis, P. Estimating ground-level PM_{2.5}
691 in the eastern United States using satellite remote sensing. *Environ. Sci. Technol.* **2005**,
692 *39*, 3269-3278.
- 693 41. Yang, D.; Li, C.; Lau, A. K. H.; Li, Y. Long-term measurement of daytime atmospheric

- 694 mixing layer height over Hong Kong. *J. Geophys. Res-atmos.* **2013**, *118*, 2422-2433.
- 695 42. Qin, K.; Wang, L.Y.; Wu, L.X.; Xu, J.; Rao, L.L.; Letu, H.; Shi, T.W.; Wang, R.F. A
696 campaign for investigating aerosol optical properties during winter hazes over
697 Shijiazhuang, China. *Atmos. Res.* **2017**, *198*, 113-122.
- 698 43. Grenier, H.; Bretherton, C.S. A Moist PBL Parameterization for Large-Scale Models and
699 Its Application to Subtropical Cloud-Topped Marine Boundary Layers. *Mon. Weather.*
700 *Rev.* **2001**, *129*, 357.
- 701 44. Bretherton, C.S.; Park, S.S. A New Moist Turbulence Parameterization in the
702 Community Atmosphere Model. *J. Climate.* **2009**, *22*, 3422-3448.
- 703 45. Angevine, W.M.; Jiang, H.; Mauritsen, T. Performance of an Eddy Diffusivity-Mass
704 Flux Scheme for Shallow Cumulus Boundary Layers. *Mon. Weather. Rev.* **2010**, *138*,
705 2895-2912.
- 706 46. Shin, H.H.; Hong, S.Y. Representation of the Subgrid-Scale Turbulent Transport in
707 Convective Boundary Layers at Gray-Zone Resolutions. *Mon. Weather. Rev.* **2015**, *143*,
708 250-271.
- 709 47. Tang, I.N.; Chemical and size effects of hygroscopic aerosols on light scattering
710 coefficients. *J. Geophys. Res-atmos.* **1996**, *101*, 19245-19250.
- 711 48. Wang, Q.Y.; Tao, J.; Ren, P.K. Shen, Z.X.; Luo, L.; Tang, X.Y. Chemical composition of
712 PM_{2.5} and its impact on visibility at Chengdu in 2010 winter. *J. Earth. Environ.* **2012**, *5*,
713 1104-1108.
- 714 49. Zieger P.; Fierz S.R.; Gysel M.; Str m, J.; Henne. S.; Yttri, K.E.; Baltensperger, U.;
715 Weingartner, E. Effects of relative humidity on aerosol light scattering in the Arctic.
716 *Atmos. Chem. Phys.* **2010**, *10*, 3875-3890.
- 717 50. Titos, G.; Lyamani, H.; Cazorla A.; Sorribas, Mar.; Inmaculada, F.M.; Alfred, W. Study
718 of the relative humidity dependence of aerosol light-scattering in southern Spain.
719 *Tellus. B.* **2014**, *66*, 136-140.
- 720 51. Lin, C.Q.; Li, Y.; Yuan, Z.B.; Lau, A.K.H.; Li, C.C.; Fung, J.C.H. Using satellite remote
721 sensing data to estimate the high-resolution distribution of ground-level PM_{2.5}. *Remote.*
722 *Sens. Environ.* **2015**, *156*, 117-128.
- 723 52. Song, C.H.; Park, M.E.; Lee, K.H.; Ahn, H. Lee, J.; Kim, J.Y.; Han, K.M.; Kim, J. Ghim,
724 Y.S.; Kim Y.J. An investigation into seasonal and regional aerosol characteristics in
725 East Asia using model-predicted and remotely-sensed aerosol properties. *Atmos. Chem.*
726 *Phys.* **2008**, *22*, 6627-6654.
- 727 53. Tao, J.H.; Wang, Z.F.; Xu, Q.; Li, L.J.; Fan, M.; Tao, M.H.; Su, L.; Chen, L.F. Particulate
728 matter mass extinction hygroscopic growth model in Beijing. *J. Remote. Sens.* **2015**, *1*,
729 12-24.
- 730 54. Ning, G.C.; Wang, S.G.; Ma, M.J.; Ni, C.J.; Shang, Z.W.; Wang, J.X.; Li, J.X.
731 Characteristics of air pollution in different zones of Sichuan Basin, China. *Sci. Total.*
732 *Environ.* **2018**, *612*, 975-984.
- 733 55. Luo, Y.X.; Zheng, X.B. Zhao, T.L. Chen, Juan. A climatology of aerosol optical depth
734 over China from recent 10 years of MODIS remote sensing data. *International Journal of*
735 *Climatology.* **2014**, *3*, 863-870.
- 736 56. Liu, X.; Chen, Q.; Che, H.; Zhang, R.; Gui, K.; Zhang, H.; Zhao, T. Spatial distribution
737 and temporal variation of aerosol optical depth in the Sichuan basin, China, the recent

- 738 ten years. *Atmos. Environ.* **2016**, *147*, 434-445.
- 739 57. Chan, P.W. Comparison of Aerosol Optical Depth (AOD) Derived from Ground-Based
740 Lidar and MODIS. *Open. Atmos. Sci. J.* **2009**, *3*, 131-137.
- 741 58. He, Q.S.; Mao, J.T. Observation of urban mixed layer at Beijing using a micro pulse
742 Lidar. *Acta. Meteorol. Sin.* **2005**, *3*, 374-384.
- 743 59. Woo, K.S.; Yoon, S.C.; Kim, J.Y.; Kim, S.Y. Seasonal and monthly variations of
744 columnar aerosol optical properties over East Asia determined from multi-year
745 MODIS, Lidar, and AERONET Sun/sky radiometer measurements. *Atmos. Environ.*
746 **2007**, *8*, 1634-1651.
- 747 60. Allen, C.D.; Tsai, T.C.; Chen, J.P.; Chang, S.C.; Jeng, Y.J.; Chiang W.L.; Lin N.H.
748 Interpreting aerosol Lidar profiles to better estimate surface PM_{2.5} for columnar AOD
749 measurements. *Atmos. Environ.* **2013**, *79*, 172-187.
- 750 61. Cao, X.; Wang, Z.; Tian, P.; Wang, J.; Zhang, L.; Quan, X. Statistics of aerosol extinction
751 coefficient profiles and optical depth using Lidar measurement over Lanzhou, China
752 since 2005-2008. *J. Quant. Spectrosc. Ra.* **2013**, *122*, 150-154.
- 753 62. Zhao, X.J.; Zhao, P.S.; Xu, J.; Meng W.; Pu W.W.; Dong, F.; He, D.; Shi, Q. F. Analysis of
754 a winter regional haze event and its formation mechanism in the North China Plain.
755 *Atmos. Chem. Phys.* **2013**, *11*, 5685-5696.
- 756 63. Salinas, S.V.; Chew, B.N.; Jukka, M.; Campbell, J.R.; Welton, E.J.; Reid, J.S.; Yu L.E.;
757 Liew, S.C. Physical and optical characteristics of the October 2010 haze event over
758 Singapore: A photometric and Lidar analysis. *Atmos. Res.* **2013**, *122*, 555-570.
- 759 64. Wang, C.H.; Hu, J.; Feng, S.L.; Jin, S.L.; Zhang, F.M.; Liu, C. Comparing different
760 boundary layer schemes of WRF by simulation the low-level wind over complex
761 terrain. *Aimsec.* **2011**, *2*, 6183-6188.
- 762 65. Liu, Y.; Park, R.J.; Jacob, D.J.; Li, Q.B.; Kilaru, V.; Sarnat, J.A. Mapping annual mean
763 ground-level PM_{2.5} concentrations using Multiangle Imaging Spectroradiometer
764 aerosol optical thickness over the contiguous United States. *J. Geophys. Res-Atmos.*
765 **2004**, *109*, 22.
- 766 66. Tao, J.; Zhang, L.; Engling, G.; Zhang, R.J.; Yang, Y.H.; Cao, J.J.; Zhu, C.S.; Wang, Q.Y.;
767 Luo, L. Chemical composition of PM_{2.5} in an urban environment in Chengdu, China:
768 importance of springtime dust storms and biomass burning. *Atmos. Res.* **2013**, *122*,
769 270-283.
- 770 67. Li, C.; Hsu, N.C.; Tsay, S.C. A study on the potential applications of satellite data in air
771 quality monitoring and forecasting. *Atmos. Environ.* **2011**, *45*, 3663-3675
- 772