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Article

Reconstruction of Ancient Volcanic Edifices Using Magnetotellurics: A Case Study from Carboniferous Zhibo Iron Deposit in Western Tianshan Metallogenic Belt in China

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Abstract

Volcanoes are the main pathway to the surface for the heat and material within the Earth. It is the window into the Earth because they provide and record many clues to what is going on in the subsurface. Active and potentially active volcanoes greatly affected the course of human history and are closely related to the present tectonics of the Earth. As a result, a great number of active volcanoes have experienced comprehensive geophysical studies. However, understanding to the internal structure of the ancient volcano complexes remains very poor. We propose a new approach to study the structure of the ancient volcano complexes based on the comparison of the magnetotelluric (MT) observation from Zhibo (ZB) ancient volcano and the active mid-oceanic ridges volcanoes from Iceland and intracontinental volcanoes from north China. MT responses of magma chambers from active volcanoes shows very close resistivity values. Based on these observations, we assumed that the ancient active volcano chambers have similar resistivity as current active volcanoes. We then reconstruct the ancient Carboniferous volcanoes complex in ZB using the ratio of the lower portion of the MT responses from ZB and active volcanoes. The results imply the existence of the fossil magma chambers at some 5 km in depth. It serves as the critical indication for a volcanic center and confirms the magmatic origin of the ZB volcanic rock-hosted iron deposits.

Keywords: crustal structure; northern Tibetan Plateau; magnetotelluric; Qimantagh; metallogenic belt

1. Introduction

Volcanoes are the main pathway to the surface for heat and material within the Earth, most of them are associated with the boundaries of tectonic plates [1,2]. More than 1400 volcanoes around the world have ever been active, and more than 1500 are active or potentially active in the last 10,000 years, some 1200 volcanoes are thought to have been active during the Pleistocene, but only tens of these are well documented or studied [3,4]. Active and potentially active volcanoes greatly affected the course of human history and are closely related to the present tectonics of the Earth [5,6]. A great number of active volcanoes have experienced comprehensive geophysical studies [7–17]. However, very few geophysical observations have been conducted to the ancient volcanoes, although they have significantly contributed to the structure, surface, composition, biosphere, and evolution of the Earth.

The presence of fluids and partial melt in the magma and mush storage reservoirs of active volcanoes results in lower electrical resistivity compared to their hosting geological units, and geo-electrical methods are frequently employed for volcanoes studying [16,18,19]. Among them, Magnetotellurics (MT) are widely used due to the great penetration depth and sensitivity to temperature, water content and metallic mineral content [9,10,12,21,22]. MT has revealed two conductive layered structure in many active volcanoes system, an upper layer reflects hydrothermal system and a deeper layer or mass means the magma or partial melting source[23]. Low-velocity anomalies have been imaged in magma systems as well [24]. The energy, volatile materials, and space within a volcano are thought to be responsible for the formation of volcano-hosted or related ore deposits. As a result, ancient volcanic edifices are important in the formation and storage of mineral, oil, and gas deposits. How to locate and map ancient volcanic edifices using MT, whether the ancient volcano has relative low-resistivity anomalies, remains unclear.

We conducted a broadband MT (BMT, which combined audio magnetotelluric and MT) exploration of the ZB submarine volcano rock-hosted iron deposits (MVRI). To develop a new approach for the reconstruction of the ancient volcanic complexes, we compare the MT observation among the active volcanoes and ZB ancient Carboniferous volcanoes complex. We calculate a ration between the average MT response of the active volcano and those in ZB. Based on the laboratory measurements, modeling, and inversion, we reconstruct the ZB ancient Carboniferous volcano using the ratio of the observational apparent resistivity in the clower frequency portions of the MT responses from ZB and the Iceland active volcanoes. Our BMT results show the response of the residual paleo-volcanic edifice, refine the understanding of volcano mineral control system in ZB.

2. Geological Setting

ZB iron deposit is location in the eastern section of Western Tianshan, northwest China. Western Tianshan is in the southwestern part of the Central Asian Orogenic Belt (Figure 1), one of the largest Phanerozoic accretionary orogenic systems in the world [25]. It is a Paleozoic orogenic collage formed by multiple subduction events and the accretion of several terranes between the Junggar Terrane and the Tarim Craton [25,26]. Moreover, it controls the tectonics of Central Asia[26,27]. The Western Tianshan Metallogenic Belt (WTMB) is a typical submarine volcano rock-hosted iron deposits (MVRI) in China[28]. It has several large to medium deposits of submarine Carboniferous volcanic-hosted iron (both hematite and magnetite). More than 1.2 billion tons of iron ore resources with an average grade of 40% (up to >60%) have been found in the Awulale iron deposits belt, eastern section of the WTMB, which is thought to have excellent prospecting potential for high-grade iron ore deposits [29,30]. All the iron ore deposits are commonly hosted in the widespread Early Carboniferous submarine volcano-related rocks[29]. Some of these deposits are characterized by well-developed skarn assemblages; others are considered to have been directly controlled by the ancient volcanic edifices. Surface and mining geological, petrological, geochemical, and surface and aeromagnetic analyses have been employed to explore the mineral deposits and mineral systems, making many new discoveries about the deposits and the structure and origin of the mineral systems. Central type Carboniferous volcanoes are speculated to control the distribution of the iron deposits in the Awulale Iron Metallogenic Belt in the eastern portion of WTMB. The ZB iron deposit is one of the largest volcanic-hosted magnetite deposits in the WTMB (Figure 1c). It has attracted widespread interest from geologists and mineralogists for its large size and high grade, as well as the controlling ancient volcanic edifices. Previous geochemical and geochronological studies determined that the iron mineralization at ZB was originated from the melting of a mantle wedge fertilized during subduction with subordinate hydrothermal overprinting [30–32] or hydrothermal origin[33]. The nature and tectonic affinity of the volcanic host rock is still controversial.

Remote sensing and geological field observations confirm the existence of an ancient volcano edifice. A huge Aikendaban oval caldera was recognized by remote sensing [33]. It is in the eastern Awulale belt where the ZB iron deposit is seated. There are several other large-scale MVRI, which are considered to be controlled by this volcanic edifices [29,33]. Field work in ZB has found some

typical characteristics of volcano activity, such as the stomata and almond-shaped structures in the ore deposit [34]. Shen et al. [33] reported new evidence of an ancient volcano edifice different from the huge Aikendaban oval caldera based on field observations of the ZB deposit. They found andesites in volcanic overflow facies as well as tuff and breccia-bearing tuff in volcanic sedimentary facies in the eastern mining area of ZB. Besides, outcrops of andesitic volcanic breccia and agglomerate rocks have been found in the eastern mining area and in some drill cores. These indicate that there were two stages of volcanic activity: the eruptive and overflow volcanic stages. An arc-shaped fault system and formations as well as the occurrence of the ore body, which were controlled by the ring fault system, imply the existence of a caldera and weak erosion in the ZB iron deposit.

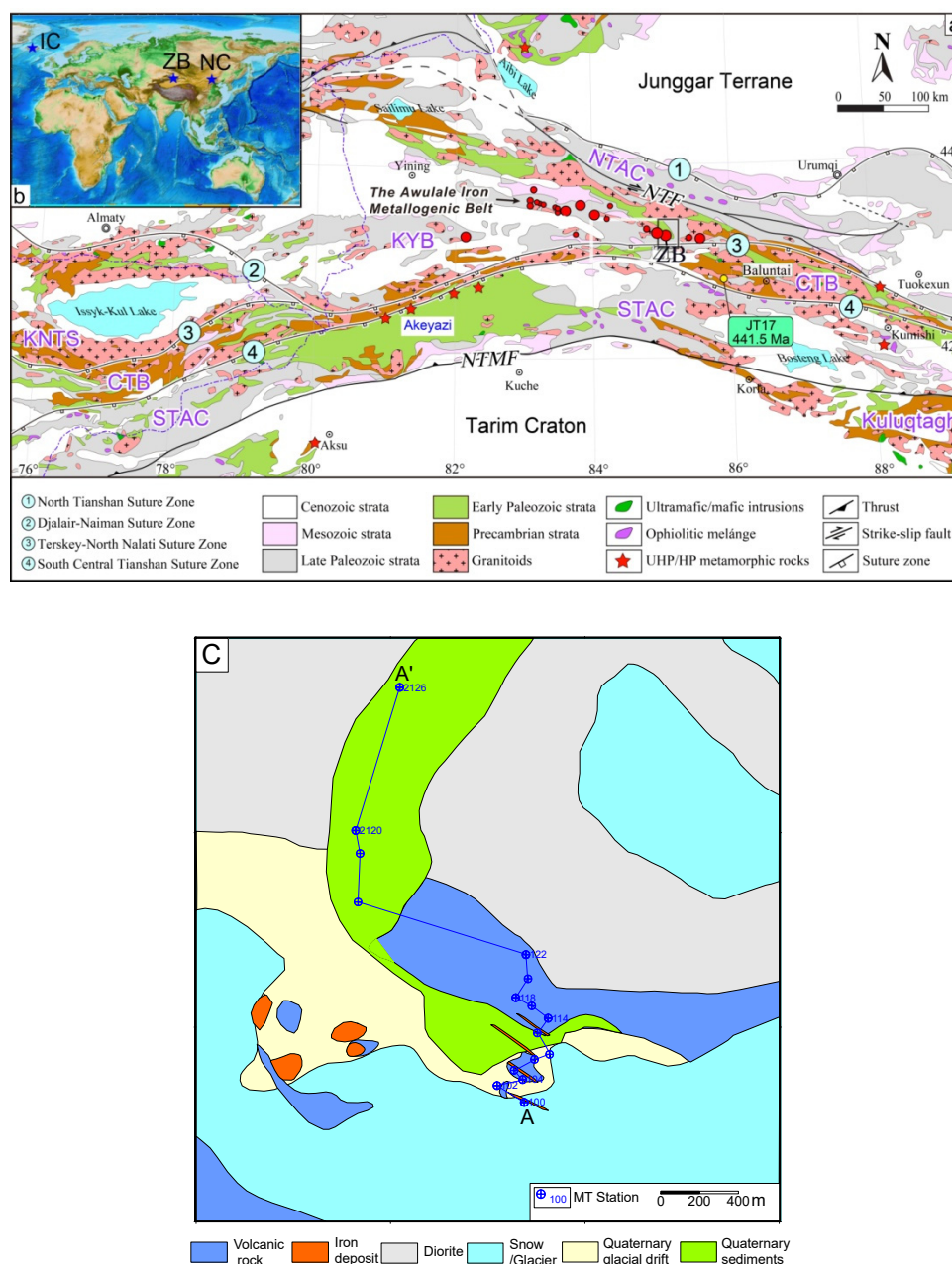


Figure 1. Tectonic setting of ZB iron deposit [26]. (a) is the tectonic outline of the Central Asian Orogenic Belt (CAOB) and environs. (b) is the simplified tectonic divisions of the Western Tianshan Orogen and adjacent. (c) shows the geological map of the ZB iron deposit and the locations of MT stations. NTAC–North Tianshan Accretionary Complex, KYB–Kazakhstan–Yili Block, KNTS–Kyrgyzstan North Tianshan Block, CTB–Central Tianshan Block, STAC–South Tianshan Accretionary Complex. Major faults: NTF = North Tianshan Fault; NTMF = Northern Tarim Margin Fault [26].

Petrological and geochemical analyses provide plenty of evidence for the volcanic and magmatic origin of the ZB iron deposit. It is reported that this deposit is close to the center of a volcano. The U-Pb zircon age of the host rocks ranges from 350 to around 300 Ma in three stages [35]. Based on their tholeiitic-calc-alkaline affinities, most of these volcanic rocks in ZB are considered to have formed an active continental margin arc. Primitive mantle-normalized incompatible element diagrams, their relatively complete rock assemblages, and Th/Yb versus Nb/Yb diagrams indicate that they have remarkable negative Nb, Ta, and Ti anomalies [35]. Electron probe micro-analyzer data show that the average Co content of pyrite from the magmatic phase in ZB is some six times higher than that of the hydrothermal phase. The Co/Ni ratio is about 20 times higher. Moreover, $\delta^{34}\text{S}$ for the pyrite indicates that the magmatic-origin sulfur is characterized by a deep source [31,36]. A rock geochemical analysis indicated that the ZB iron deposit was mainly formed by the affection of magmatism associated with volcanic activity and hydrothermal metasomatism of magmatic-hydrothermal fluid [29,31]. An ore-forming fluid analysis showed that the metallogenic temperatures during the magmatic mineralization in ZB were about 650–700°C and that the composition of the fluid in the inclusions of magnetite are of the $\text{Na}^+\text{-Cl}^{\text{-}}(\text{SO}_4^{2-})$ type. This indicates that the metallogenic fluids activated during magmatic mineralization are Fe-rich melts [31].

3. Data and Methods

We carried out an AMT and MT study in the ZB iron deposit and its vicinity, the AMT and MT results are combined as broadband MT (BMT). Data from 19 of the stations were acquired (Figure 2) and 8 of among them were used for the further processing of reconstruction for considering of the data quality. The working area is of rough mountainous terrain, these make it is very difficult to find suitable places to set up a data acquisition system and the survey line could not go straight. The common station spacing was 100 m. We collected six MT station data sets from mid-ocean ridge (Iceland) and inland (north China) to study the MT response of the different type active volcanoes. The naturally occurring magnetotelluric fields are used as the field source of MT and AMT. The natural field contains electromagnetic waves with different frequency component, they are source used to investigate the underground electrical conductivity of the Earth in different depth[37]. Frequency components in band of 10 Hz to 10 kHz are employed for AMT sounding, those of several hundred hertz to less than 0.001 Hz are used for MT sounding. AMT and MT measured the orthogonal electric (E) and magnetic (H or B) fields behave like plane waves on the Earth surface[38,39]. Frequency-based impedance results could be extracted from the amplitude, phase, and directional relations between the surface electric and magnetic fields. The subsurface distribution of electrical conductivity are then converted from the impedance results [40,41].

We use three MT receiver units for data acquisition. A detailed field operation have been described in He et al.[39]. Some stations lie close to an open operational mining pit, and data from these stations were affected by artificial electromagnetic and vibration noise at different levels. Data with a high noise level are rejected in further processing. We utilize electromagnetic array profile filtering[42] as a pre-processing step to correct for topographic static shifts and artificial noise effects in the data along the survey line direction (XY), the data are inverted using Bostick conversion with space filtering for the selected sections [43,44].

The reconstruction approach of ancient volcanic edifices using magnetotellurics are based on the assumption that the ancient active magmatic chamber has similar resistivity to update active magmatic chamber, for the resistivity of the magmatic chamber are mainly controlled by the temperature and water content. With the cooling and degassing of the magmatic chamber, its resistivity increase. However, the resistivity is also employed for determine the buried depth of magmatic chamber. , Even the active and ancient magmatic champers are seated in the same depth, the conversion depth from frequency to space domain of an active and ancient magmatic chamber are different because the conversion depth are determined by the resistivity of the magmatic chamber. We propose a reconstruction approach by using the ration between the average apparent resistivity of active volcano and ancient volcano system. The conversion depth and the resistivity of

the ancient volcanic edifices are then reconstructed based on this ratio. The details of why and how to determine of the ratio is discussed in Section 5.1.

4. Results

We combined the observational results from the same AMT and MT station into one BMT curve with a broad frequency range from greater than 10,040 to some 0.001 Hz. The location of the stations is shown in Figure 1c and the observational curve shows in Figure 2. The resistivity curve from 10 among the 12 stations have an obvious fine-layer high (high resistivity)-low (low resistivity)-high-low-high (H-L-H-L-H) electrical structure. It is hard to identify structures in the data from the other 2 stations due to the artificial noise. Figure 3 shows the typical resistivity curves of 12 stations from ZB deposits (Figure 3a) in the ancient Carboniferous volcanoes, North Iceland (IC) as a representative of an active mid-oceanic ridge volcano (Figure 3b), north China (NC) with active intraplate volcanoes (Figure 3c), and the vicinity of ZB, an non-volcanic area (Figure 3d). The ZB deposits and the IC active volcano have very similar geo-electrical structure, characterized as H-L-H-L-H, but the resistivity value at ZB is more than 10 times higher than at IC. The intraplate active volcanoes in north China (NC) have H-L-H-L structure. There was no obvious low-resistivity layer in the non-volcanic area (Figure 3d).

Figure 3e shows the logarithm of the resistivity as a function of frequency averaged from three selected stations at ZB, IC, and NC. It confirms the similar geo-electrical structures at IC and ZB. The average resistivity at ZB was around 1000 Ω.m for frequencies from 10,000 to 0.001 Hz. There are two relative resistivity lows around 10 and 0.01 Hz. The average resistivity for IC varies from 100 to tens of Ω.m with two relative lows centered at 1.72 and 0.046 Hz. The resistivity for NC varies from high to low to high as the frequency decreased from 320 to 0.094 Hz. The resistivity low is at 15 Hz. The resistivity decreases with the frequency decreasing from 0.094 to 0.0011 Hz.

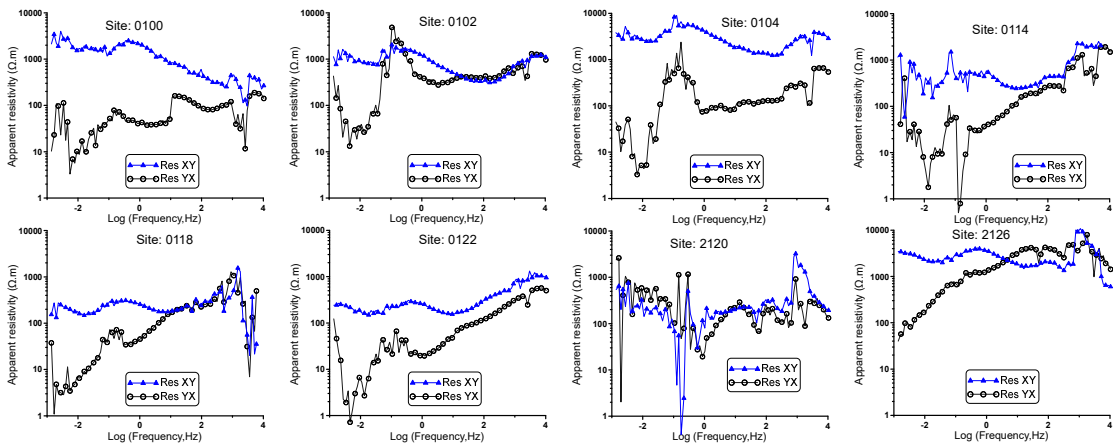


Figure 2. The apparent resistivity versus frequency curves of 12 BMT stations show general geoelectrical variation with depth, the low frequencies refer to the deep and the higher refer to shallow. The relative low resistivity layer in Rxy direction which might reflect the magma chamber remains could be identified from data of 10 stations.

Results from IC and NC indicate that the mid-oceanic ridge and intraplate active volcanoes have very similar average resistivity at the low frequency portion. Figure 3f shows the comparison result of the 1-D MT models (using the algorithm of Whittall and Oldenburg[45]) between the active and ancient volcano. The model of the active volcano was based on the Bostick conversion of the average resistivity curves for IC, as shown in Figure 3f. The 1-D geoelectrical structure of the ancient volcano model is the same to those of the active volcano, but the resistivity is 30 times greater in value than those of the active model. Overall, the average resistivity of the ancient volcano is 21.88 times greater than that of the active volcano. Their ratio of the lowest resistivity in the lower frequency band is 30.

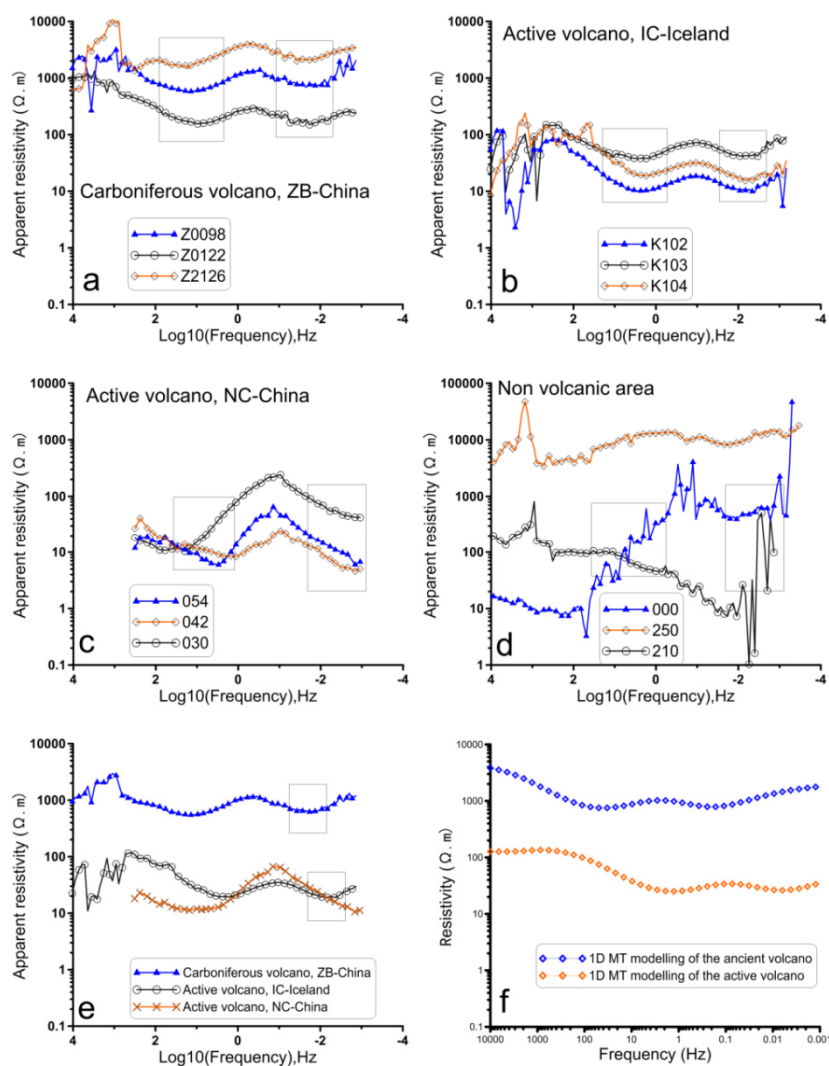


Figure 3. The typical MT responses in term of the resistivity curves, 3a shows the ZB deposits in the ancient Carboniferous volcanoes, 3b are acquired from the active mid-oceanic ridge volcano in Iceland, 3c from north China with active intraplate volcanoes, and 3d the vicinity of ZB, a non-volcanic area. 3e is the logarithm of the resistivity averaged for three selected stations at ZB, IC, and NC. 3f shows the comparison of 1-D MT models of the active and ancient volcanoes. The model of the active volcano used data from the Bostick conversion from the average resistivity curves of IC 3b. The ratio of the resistivity of the ancient and active volcano models for each layer is 30.

The effect of water content on the resistivity of the rock sample was measured by comparing the results before and after sample absorbed water in room conditions (Figure 4). The initial resistivity of leucogranite (measured time: 02-04-2022,13:26) increases to a peak values at 731071 $\Omega.m$ from 10,000 to 1.5 Hz, then decrease slowly down to 241558 $\Omega.m$ from 1.5 to 0.0032 Hz, and then decrease roughly down to 803 $\Omega.m$ in from 0.0032 to 0.001 Hz. Measurement results show very little variation with frequency in next three days. The peak resistivity of sample decrease from 731071 to 296 $\Omega.m$, the ratio of highest to lowest resistivity is some 0.04%, after some 0.47% weight percent water was absorbed by the sample. The result indicates that the water content dominate the resistivity of the rock sample.

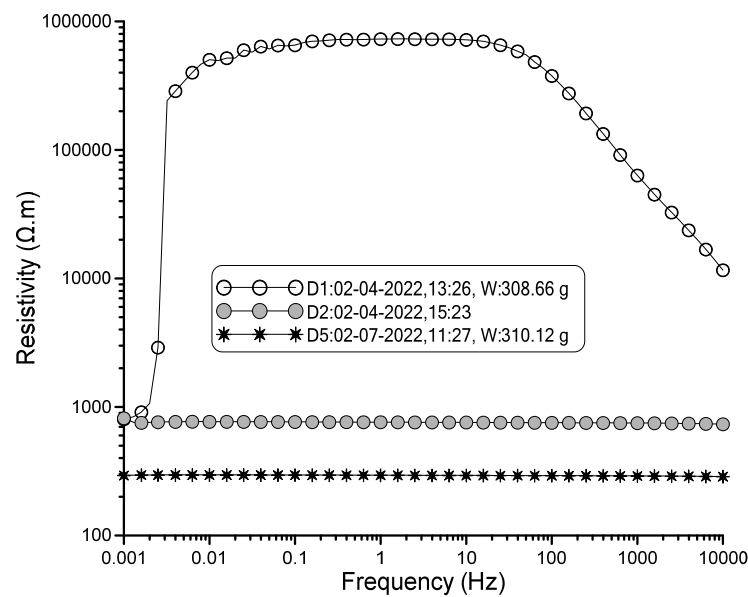


Figure 4. Time-lapse complex resistivity of the rock sample, pre- and post- absorbing water. The water in the rock samples was absorbed from the wet sample hold during measurement with a weight (W) variation of 1.46g around 0.47%. The duration of measurement is some three days.

Figure 5 shows the BMT conversion results for the observational and reconstructed data, the reconstruction correction ratio is 0.0315 (Table 1), how to determine the correction ratio are discussed in next section. The result converted directly from the observational data (Figure 5a) show four distinct layers apart from the top resistant layer: (1) an top layer with a spatially variable resistivity that contains a number of discrete low-resistivity zones from the surface to 10 km below the sea level, (2) an upper high-resistivity layer with a conversion resistivity greater than 1000 Ω.m, (3) an intermediate layer with relatively lower resistivity from 100 to 500 Ω.m seated around 35 to 40 km below sea level, and (4) a bottom high-resistivity layer with a conversion resistivity greater than 500 Ω.m. The uppermost low-resistivity layer may reflect the aqueous fluids and shallow iron deposits in ZB. The top high-resistivity layer above it is glacial debris. Figure 6b and 6c show the conversion result of the reconstructed data in different depth scale. There are two distinct conducting layers with resistivity from 5 to 10 Ω.m. The upper conducting layer is around 2 km below the surface. It might reflect the mainly iron deposits and their host rocks. The lower conducting layer is speculated to be the reflection of the active magma chamber in depth of 4 to 8 km. A layer with a relatively high resistivity more than 15 Ω.m separates the two conducting layers.

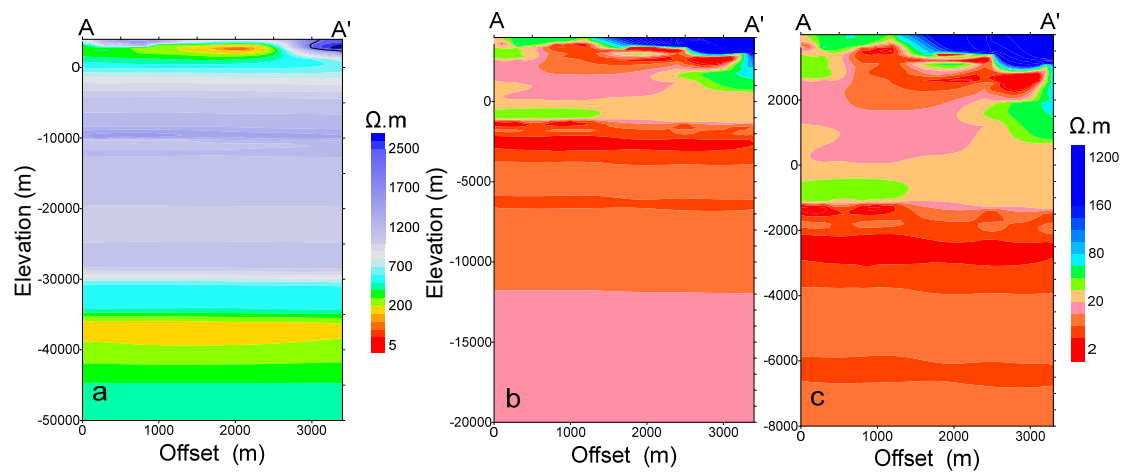


Figure 5. Resistivity conversion result of the cross-section from BMT results trend from south (A) to north (A') of the ZB iron deposit. The reconstruction are conducted by correct the observed BMT data with a reconstruction

correction ratio of 0.0315. (a) Data converted from the observational raw data (pre reconstruction), with mainly high resistivity and a relatively low resistivity layer around 25 to 30 km below sea level. (b) The geo-electrical structure of post-reconstruction. After the correction, the lower deep resistivity layer is seated around 1 to 2 km below sea level. (c) An enlargement of the upper part of (b).

5. Discussion

5.1. Determining the Reconstruction Correction Ratio

The magma and melting in the active volcanoes feature high temperature and contain dissolved water in their composition that reduces its resistivity and result in the strong resistivity contrasts between the active magma chamber and hosting units [46]. This contrast represents an appropriate framework for the application of electromagnetic methods, such as MT, to study the property and the volcano structure. The active magma chambers are always characterized by deep buried low resistivity layer and the location could be infer by the MT conversion results. Understanding to the buried depth and the electrical property of the fossil magma chamber remain very poor. Once a volcano went bone-dry and cooled, we can consider the fossil magma chambers are immobile at the certain buried depth apart from the later tectonic uplift and erosion. Several Carboniferous caldera are preserved in the Western Tianshan, indicate that this area has experienced very weak erosion.

The buried depth of geological units infers from the observational MT data are basically determined by two parameters: the impedance, which is always in terms of resistivity and phase, and the corresponding frequencies range. A simple estimation is that the buried depth is proportional to the square root of the resistivity at a given frequency. During the evolution from active to fossil, the magma chambers lose its high temperature and water content and then turn from conductive to resistant at the same buried depth. However, the buried depth would be overestimated by MT conversion due to their increased observational resistivity. We attempt to reconstruct the ancient volcanoes by correcting the observational resistivity of the fossil magma chamber to those of the active magma chamber.

The low resistivity anomalies in the active volcanoes are mostly plausible interpreted as melt, saline fluids and the high temperature, although it difficult to differentiate between fluids and melt with MT alone due to the inherent non-uniqueness of the MT problem and the tendency for conductive bodies to smear vertically in regularized inversion[47]. The temperature will weakly affect the resistivity of chamber system of the active volcanoes due to the presence of fluid phase [48]. Modeling result of bulk resistivity as a function of fluid fraction by Yang et al. [47] shows that the bulk resistivity of the melt system would decrease 2~3 time when the fluid fraction range from 7 to 26 percent. Our measurement result of the rock sample the absorbed water could decrease the samples resistivity from 731071 to 296 $\Omega\cdot\text{m}$ (2467 times) in the first two hours, but only decrease 2.56 times in the next two days (Figure 3). These indicate that the fluid fraction will also weakly affect the resistivity of chamber system after the conductive fluid are connected [48–50]. The above discussions infer that most of the active melt chamber might have similar resistivity even they feature as different temperature and fluid content. The observed MT result from the mid-oceanic ridge (IC) and intraplate (NC) active volcanoes (Figure 3e) have potential provided illustrations to our inference.

We assume that the magma systems of ancient active volcanoes feature resistivity close to those of present active volcanoes. We use the average apparent resistivity acquired from several stations of currently active volcanoes and select the lowest portion from the apparent resistivity curve in the low-frequency band. The average of nine adjacent frequencies is used to represent the apparent resistivity of the active volcano. In the same way, the apparent resistivity of an ancient volcano provides values about its fossil chamber. The ratio of the average apparent resistivity of an active volcano and an ancient volcano is employed as the reconstruction correction (Table 1). The ratio value at 0.0315 was used to recalculate the apparent resistivity for all BMT data sets from each station in the ZB iron deposit. Although it is immature, this is the important way we can make for the reconstruction of the ancient volcanic magma chambers so far.

5.2. Refined Understanding Marine Volcanic-Hosted ZB Iron Deposit

The ZB iron deposit represents a typical volcanic-hosted magnetite deposit in the Awulale belt, eastern part of the Western Tianshan, the tectonic setting is still controversial [29,30]. A huge Aikendaban oval caldera was recognized in the eastern Awulale belt where the ZB iron deposit is seated. Apart from ZB, several other large-scale MVRI, which are considered to be controlled by this volcanic edifices [29,33], have been discovered. Petrological and geochemical analyses have provided a great many of observations for understanding the volcanic and magmatic origin of the ZB iron deposit. It is reported that this deposit is close to the center of a volcano[33].

The MT observations from ZB and an active IC volcano have resistivity curves with similar shapes (Figure 3). These indicate the existence of a volcano system and potentially imply the key information for understanding volcanic structure, but the resistivity is obviously greater to those of the active volcano system. We reconstructed the ancient volcano system (Figure 6) using MT data based on the assumption that the active magma chambers feature almost same resistivity in value to the ancient chambers before they are cooled and dried. The result shows a five-layer geo-electrical structure. A comprehensive model are in interpreted as shown in Figure 6 based on the rock physics analysis, geochemical study, drilling, geological and this MT observation. The resistive top layer in ZB is mainly composed of glacial and frozen detritus, the upper conductive layer reflects the iron deposits and their related mineralization and alteration deposits. Inferred from the understanding of active volcanoes, the deeper conductive layer is interpreted as the fossil magma chamber. Before reconstruction, our MT result reflects the seated depth of the fossil magma chamber locates at some 30 km (Figure 5a). The depth is corrected to around 5 km after reconstruction (Figure 5b). This magma chamber provides important space for magmatic mixing and iron enrichment, and it eventually contributed to the metallogenesis of the ZB iron deposit. Our MT study has refined the understanding of the volcano system in the ZB iron deposit.

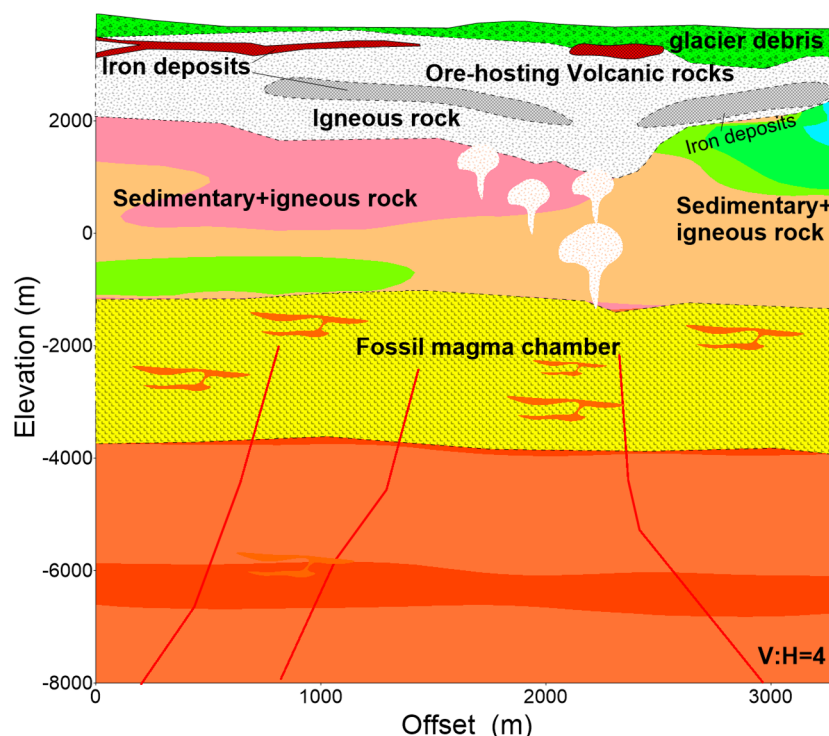


Figure 6. Refined and simplified model of the ZB iron deposit mineral system by reconstructing ancient Carboniferous volcanoes based on MT observations. It is a typical volcano-hosted iron deposit mineral system. A near surface volcanic iron deposit layer serves the main ore-bearing layer, its top portion has been verified by drilling borehole. A fossil magma chamber is seated in upper crust beneath the ZB iron deposit. The red polylines reflect the pathways for heat, magma and/or fluid. The base map is geo-electrical section of post-reconstruction.

6. Conclusions

Broadband MT observations reveal similar geo-electrical structures in terms of apparent resistivity curves type in ZB ancient Carboniferous volcanoes and Iceland active volcano except different value. Close average apparent resistivity in the corresponding frequency band, which reflects the magma chamber, are also found among the active mid-oceanic ridge and intra-continental volcanoes. Laboratory results indicate that the active volcanoes chambers might feature almost same conductivity in the state that the partial melt and/or fluid exist. Based on the assumption that the magma systems of ancient active volcanoes feature proximate resistivity to those of present active volcanoes, we acquired the reconstruction correction from MT responses. We have reconstructed the ZB ancient Carboniferous volcanoes using the correction ratio of the MT observed resistivity between the ZB ancient volcano and IC active volcano. The understanding to mineral system has refined using the reconstructed model. A fossil magma chamber in the upper crust, which acts as the critical role to mineral system, is interpreted. We made a new approach to reconstruct ancient volcanic edifices based on MT study.

Author Contributions: Conceptualization, Lanfang He; methodology, validation, investigation. Lanfang He; Ping Shen, resources. Lanfang He; Zhongxing Wang; writing—original draft preparation. Lanfang He; Ping Shen; Xi Zhang; writing—review and editing. Lanfang He; Ping Shen; Zhongxing Wang; project administration. All authors have read and agreed to the published version of the manuscript.

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Data Availability Statement: The data used to support the findings of this study are available from the corresponding author upon request.

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Conflicts of Interest: The authors declare no conflict of interest.

Appendix A

Table 1. Apparent resistivity value used for calculating the reconstruction correction shift coefficient.

Number	Rxy-IC	Rxy-ZB	Ratio
1	19.348	643.622	0.0300
2	19.701	681.674	0.0290
3	18.936	647.165	0.0290
4	19.843	647.734	0.0306
5	18.944	622.475	0.0304
6	19.186	632.86	0.0303
7	20.876	637.369	0.0328
8	23.785	654.074	0.0364
9	24.384	695.172	0.0351
Average	20.556	651.349	0.0315

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