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Article

Discussion on Genetic Types and Sources of Natural Gas in Y5 Well Area on the North Slope of Baiyun Sag

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Abstract: In recent years, the genesis of natural gas and the definition of major gas source rocks in the Baiyun Sag of the Pearl River Mouth Basin have attracted extensive attention from scholars. There is still considerable controversy over the genesis and the main gas source rocks of natural gas. In this study, the chemical kinetics method of natural gas generation and carbon isotope fractionation was used to study the generation history and isotope evolution history of Paleogene natural gas in the Y5 well area located on the north slope of Baiyun Sag. And other geochemical discriminant methods were applied simultaneously to determine the type and source rocks of natural gas. The results show that the natural gas in the Y5 well area on the northern slope of Baiyun Sag is mainly kerogen degradation gas, mixed with a small amount of oil cracking gas. The mixture of kerogen degradation gas and a small amount of crude oil cracking gas caused the inversion of the carbon isotope sequence of ethane and propane. The natural gas in the Y5 well area mainly comes from delta-shallow lacustrine facies source rock. The high-quality reservoirs in the shallower strata accumulate natural gas generated in the early stage, and the tight reservoirs in the deeper strata mainly accumulate natural gas generated in the late stage.

Keywords: Baiyun Sag; thermal simulation experiment; hydrocarbon generation kinetics; carbon isotope fractionation; natural gas genesis

1. Introduction

The Baiyun Sag is located in the deep-water area of the northern continental slope of the South China Sea, and it is a secondary structural unit inside the Zhu II Depression of the Pearl River Mouth Basin (Figure 1). The water depth of the Baiyun Sag varies greatly, and the Baiyun Sag can be divided into four sub-sags: the Baiyun main sub-sag, the west sub-sag, the east sub-sag and the south sub-sag [1]. The Baiyun Sag has the most complete Mesozoic and Cenozoic strata in the Pearl River Mouth Basin. The deep-water area of Baiyun Sag has been proven to be a gas-rich sag [2], which has the characteristics of concurrent oil and gas generation and is dominated by kerogen degradation gas. Among them, commercial gas reservoirs in Baiyun Sag are mainly found in the northern part of Baiyun main sag to Panyu low uplift and the W3-4 to H29 nose-shaped structural belt in the eastern part of Baiyun main sag, all of which are located in the uplift areas. The gas-rich strata are concentrated in the lower Zhujiang Formation, and only at the H19 area are there shallow natural gas accumulations in the Hanjiang Formation–Yuehai Formation. From the northern part of Baiyun main sag to the Panyu low uplift, condensate gas reservoirs with high gas–oil ratios are predominant, while the gas–oil ratios of W3-4, H28, and H29 condensate gas reservoirs in the eastern part of Baiyun main

sag is relatively lower. It can be seen that the maturity of natural gas in the condensate gas reservoirs in the two zones is different.

There is a lack of research on the gas generation characteristics of the main gas source rocks in this area and the contribution of gas sources to accumulation. Multiple sets of source rocks in Wenchang–Enping Formation have reached the high–over mature stage under the “hot basin” system with high heat flow and high geothermal gradient in Baiyun Sag [3], and have a complete oil and gas generation sequence of light oil, kerogen degradation gas, and crude oil cracking gas. Scholars have great disputes about the contribution to accumulation of different gas source rocks in the study area [4,5]. In general, scholars have different opinions mainly because relatively few exploratory wells have encountered Paleogene mudstone in this area, and there is a lack of actual drilling data on source rocks in this area. Besides, the study area is located in the ocean–continent transition zone, with various types of coal measures and lacustrine source rocks, and there is a lack of systematic research on the gas generation characteristics of different types of source rocks.

The chemical kinetic theory was introduced by scholars such as Maier [6]and Allred [7], and it was applied to the dry distillation process of shale oil and coal in the early stage. Since the 1970s, it has been widely used in the study of the hydrocarbon generation process of kerogen in source rocks [8]. The research on hydrocarbon generation kinetics of OM has experienced three stages: basic understanding, rapid development, and in–depth understanding. The characteristics and contents of the research range from the establishment of the initial model and the optimization of parameters to the wide application in the evaluation of oil and gas resources, and then to the study of the hydrocarbon generation kinetics of single compounds and the influence of the uncertainty of kinetic parameters on the results of geological applications [9]. At present, there are four kinds of common kinetic models, including the overall reaction model [7], the Friedman reaction model [10–12], the sequential reaction model [13,14], and the parallel reaction model [15,16]. This study comprehensively considers various factors such as model principle, model accuracy, and application difficulty, and selected the parallel first–order reaction model. Scholars have demonstrated the applicability and practicability of a limited number of parallel first–order reaction models [17]. The isotope fractionation kinetic model is based on the hydrocarbon generation kinetic model, combined with the geological background, the natural gas accumulation model and gas source can be clarified [18–21].

Regarding the above problems, this study took the Y5 well area as an example to carry out the study on the kinetics of natural gas generation and carbon isotope fractionation, and the process of natural gas generation and carbon isotope fractionation was characterized quantitatively. The genetic type and source of natural gas were clarified by combining the thermal simulation results and geochemical characteristics of natural gas in the study area.

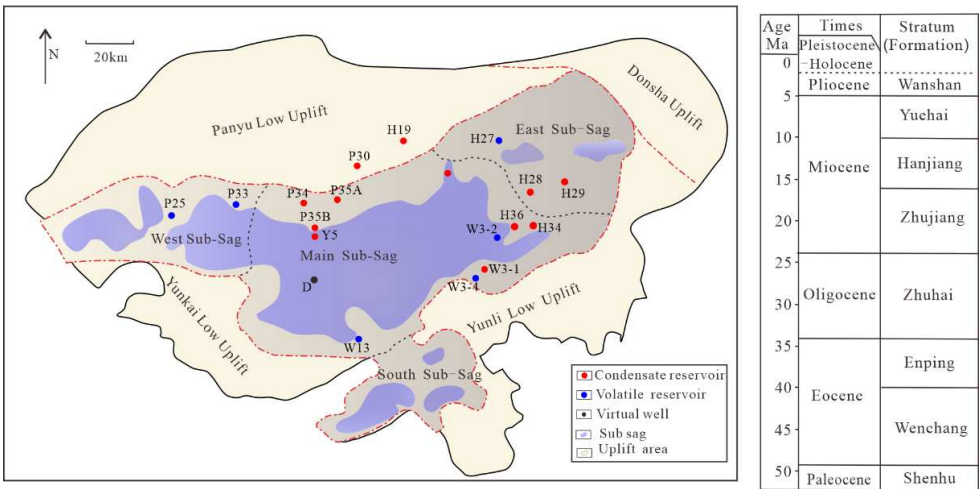


Figure 1. Regional geological setting and composite stratigraphic column of Baiyun Sag.

2. Materials and Methods

2.1. Sample and experiment

7 samples, including 4 source rock samples (Table 1) and 3 crude oil samples, from the Pearl River Mouth Basin were selected for hydrocarbon generation experiments. The organic matter (OM) type of two source rock samples from the W well area is type II₁, and the gas produced is a mixture of kerogen degradation gas and crude oil cracking gas. The OM types of source rocks from the P well area and H well area is type II₂ and type II₁–II₂ respectively, and the gas produced is mainly kerogen degradation gas. The selection of different types of gas source rocks can provide a basis for the establishment of a chart for determining the genesis of natural gas. A normal crude oil sample F14 from the semi–deep lacustrine facies of the Wenchang Formation, a heavy oil sample F7 from the shallow lacustrine facies to semi–deep lacustrine facies of the Zhujiang Formation, and a condensate oil sample Y25 from shallow lacustrine facies of the Zhujiang Formation were selected, respectively.

Table 1. Basic geochemical parameters of the samples.

Well	Sample type	Formation	R _o (%)	T _{max} (°C)	S ₂ (mg/g)	TOC (%)	HI (mg/g)	Type of OM
W13	Coal from delta facies	Enping	0.7	428	156.94	48.32	325	II ₁
W3-4	Mudstone from shallow lake facies	Wenchang	0.8	447	3.11	1.32	275	II ₁
P33	Mudstone from delta facies	Enping	1.2	438	1.53	1.65	93	II ₁ –II ₂
H36	Mudstone from delta facies	Wenchang	0.9	428	1.82	1.12	163	II ₂

This study designed and carried out a gold tube thermal simulation experiment with variable pressure to reflect the true evolution characteristics of hydrocarbon generation under geological conditions. A total of 12 pressure points (60MPa–115MPa, with a pressure setting interval of about 5MPa) were set in this experiment, and the temperature was raised from 200°C to 600°C at a heating rate of 2°C/h and 20°C/h, respectively. One autoclave would be taken out at every object temperature point with a temperature interval of about 25°C, and the gold tube in it was taken out while it was cooling. Then the gold tube was placed in a special gas collection system for the GC analysis and the isotope analysis.

2.2. Establishment and calibration of natural gas generation and carbon isotope kinetic models

2.2.1. Hydrocarbon generation kinetic model

The parallel first–order reaction model assumes that the hydrocarbon generation process of OM is composed of N parallel first–order reactions [17]. The activation energy, pre–exponential factor and the original hydrocarbon generation potential of OM corresponding to each reaction is E_i , A_i , and X_{i0} , $i=1, 2, 3, \dots, N$. When a certain reaction time t is reached, the hydrocarbon generation amount of the i –th reaction is X_i , then:

$$\frac{d(X_i)}{d(t)} = K_i (X_{i0} - X_i) \tag{1}$$

In formula (1), K_i is the reaction rate constant of the i –th hydrocarbon generation reaction, which can be obtained according to the Arrhenius formula:

$$K_i = A_i \exp\left(\frac{-E_i}{RT}\right) \tag{2}$$

Since the pyrolysis experiment of OM under laboratory conditions adopts a constant rate of temperature rise (assuming that the rate of temperature increase is D , °C·min^{–1}), it can be obtained:

$$\frac{d(T)}{d(t)} = D \quad (3)$$

Combining the above formulas, the hydrocarbon generation amount of the i -th reaction can be obtained.

$$X_i = X_{i0} \left(1 - \exp \left(- \int_{T_0}^T \left(\frac{A_i}{D} \exp \left(- \frac{E_i}{RT} \right) \right) dT \right) \right) \quad (4)$$

The total hydrocarbon generation amount of N parallel reactions is:

$$X = \sum_{i=1}^N X_i = \sum_{i=1}^N \left(X_{i0} \left(1 - \exp \left(- \int_{T_0}^T \left(\frac{A_i}{D} \exp \left(- \frac{E_i}{RT} \right) \right) dT \right) \right) \right) \quad (5)$$

In the formula (5), X is the total amount of hydrocarbon generation of all N parallel reactions; N is the number of parallel first-order reactions.

2.2.2. Kinetic model of carbon isotope fractionation

Scholars systematically compared the current isotope kinetic models [22], and it has been confirmed that the isotope kinetic model based on the Cramer III model has an ideal simulation effect for the quantitative description and geological application of isotope fractionation. In this isotope kinetic model, $^{12}\text{CH}_4$ and $^{13}\text{CH}_4$ are regarded as two different products, and their kinetic parameters are calculated respectively. In the Gramer III model, the formation of methane is regarded as the result of n first-order reactions. For each reaction, the generating rates of $^{12}\text{CH}_4$ and $^{13}\text{CH}_4$ are different:

$$\frac{k_{12c}}{k_{13c}} = \frac{A_{12c} \times \exp \left(- \frac{E_{12c}}{RT} \right)}{A_{13c} \times \exp \left(- \frac{E_{13c}}{RT} \right)} \quad (6)$$

The carbon isotope ratio model of cumulatively produced methane is:

$$R_{\text{accu}} = \frac{\sum_{i=1}^n c_{i13c}(t)}{\sum_{i=1}^n c_{i12c}(t)} = \frac{\sum_{i=1}^n f_{i13c}^0 - f_{i13c}^0 \times \exp \left(- \int_0^t k_{i13c}(T(t)) dt \right)}{\sum_{i=1}^n f_{i12c}^0 - f_{i12c}^0 \times \exp \left(- \int_0^t k_{i12c}(T(t)) dt \right)} \quad (7)$$

Among them, k_{i12c} and k_{i13c} are the generation rate coefficient of $^{12}\text{CH}_4$ and $^{13}\text{CH}_4$ of the i -th reaction; c_{i12c} and c_{i13c} is the cumulative $^{12}\text{CH}_4$ and $^{13}\text{CH}_4$ production at the time t of the i -th reaction; f_{i12c}^0 and f_{i13c}^0 is the hydrocarbon generation potential of $^{12}\text{CH}_4$ and $^{13}\text{CH}_4$ of the i -th reaction, which is equal to the product of the corresponding reaction fraction and the total hydrocarbon generation potential.

3. Results

Figures 2 and 3 show the conversion rate curves of methane and total gas (C_{1-5}) and the carbon isotope value evolution curve of methane. As the degree of thermal evolution increases, the conversion rate of OM to methane and total gas (C_{1-5}) gradually increases. The carbon isotope values of methane show an overall upward trend, but there is a weak downward trend in the early stage of thermal evolution of OM.

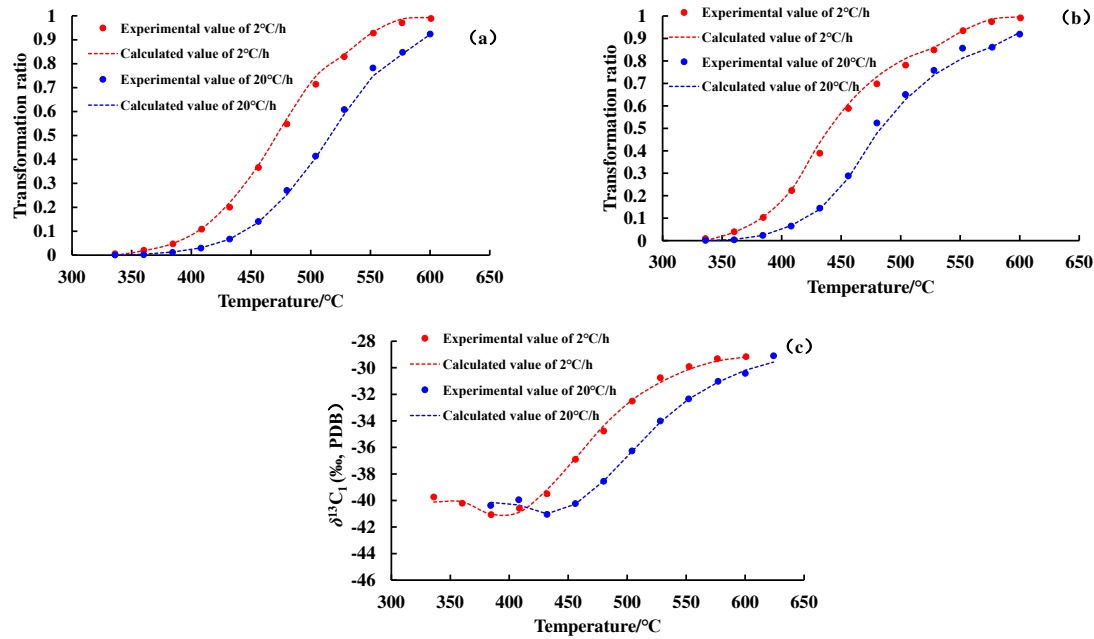


Figure 2. Transformation ratio and carbon isotope value curve of pyrolysis gas from coal measure source rock in Well W13. (a) The transformation ratio of methane; (b) The transformation ratio of C₁₋₅; (c) Carbon isotope value curve of methane.

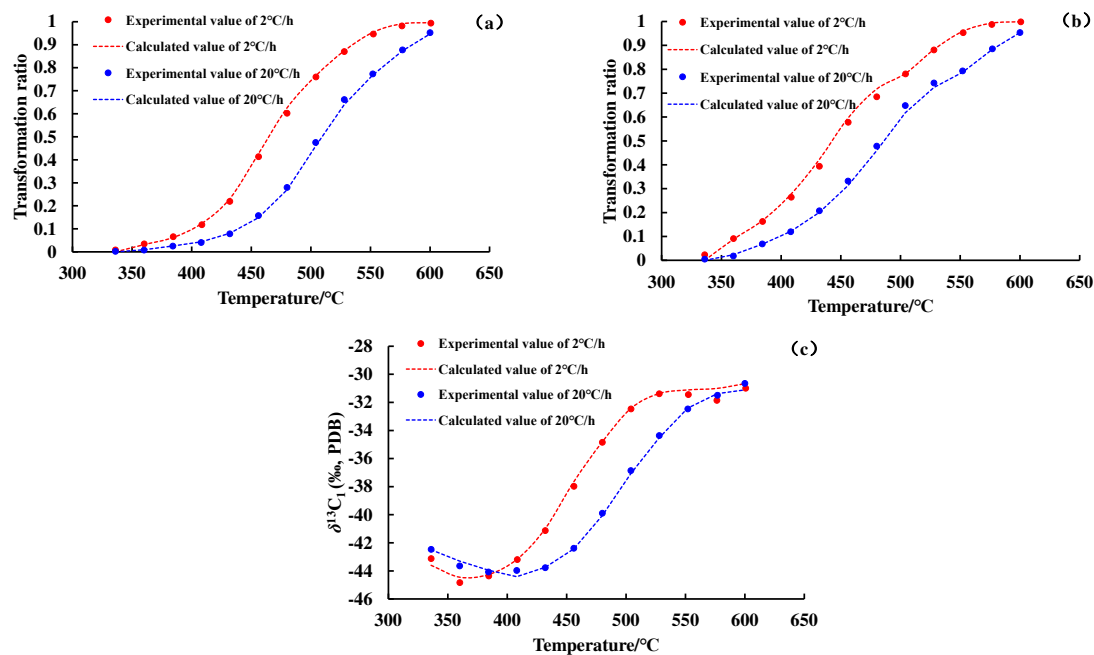


Figure 3. Transformation ratio and carbon isotope value curve of pyrolysis gas from mudstone in Well W3-4. (a) The transformation ratio of methane; (b) The transformation ratio of C₁₋₅; (c) Carbon isotope value curve of methane.

The transformation ratio curves and the carbon isotope value evolution curves of methane and C₁₋₅ calculated by the hydrocarbon generation kinetic model and the Cramer III model are shown in Figures 2 and 3. A nice match for the experiment data reflects the feasibility of the hydrocarbon generation kinetic model and isotope fractionation kinetic model, and lays a foundation for subsequent geologic applications. The kinetic parameters of $^{12}CH_4$, $^{13}CH_4$, and C₁₋₅ generated from the pyrolysis of different samples are shown in Figures 4 and 5.

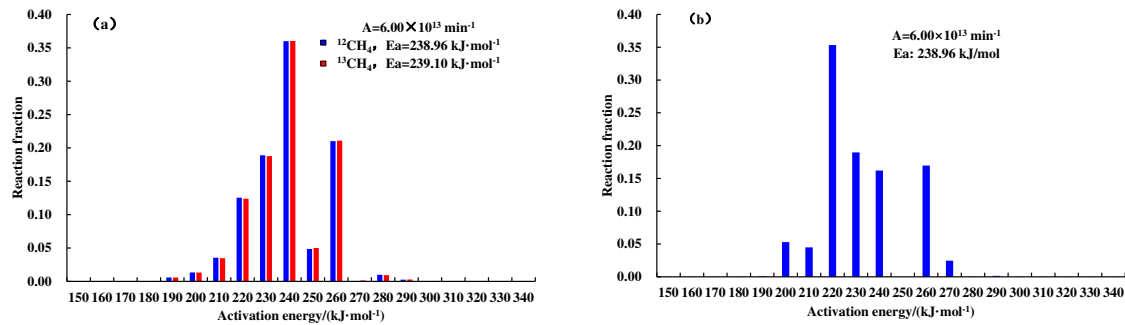


Figure 4. Hydrocarbon generation activation energy distribution of coal from Well W13. (a) Methane; (b) C₁₋₅.

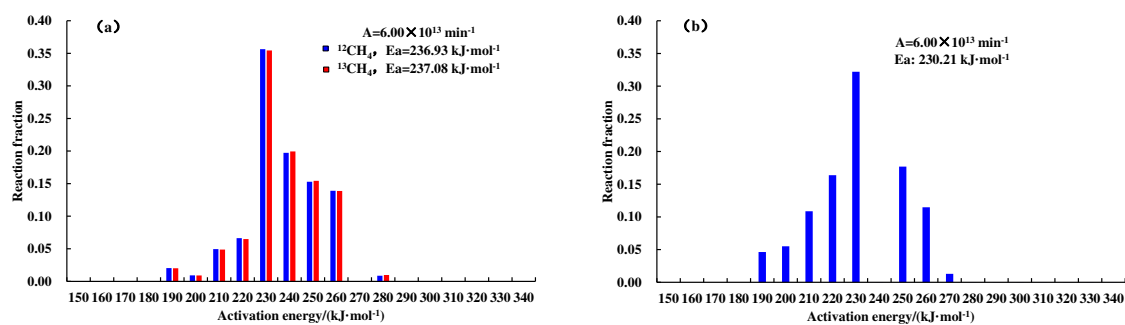


Figure 5. Hydrocarbon generation activation energy distribution of mudstone from Well W3-4. (a) Methane; (b) C₁₋₅.

4. Discussion

4.1. Identification of natural gas genetic types

4.1.1. Research based on the criteria for determining the genesis of natural gas

Prinzhofer et al. proposed to identify the genesis of natural gas based on the ratio of ethane to propane content and their difference of carbon isotope [23]. In this study, based on the thermal simulation experiment of kerogen degradation gas and crude oil cracking gas, the discriminant chart of kerogen degradation gas and crude oil cracking gas was revised. As shown in Figure 6a, the thermal simulation gas from source rocks (mixed gas produced by primary degradation of kerogen and secondary cracking of oil) has the characteristics of a small carbon isotopic gap ($-6\text{‰} < \delta^{13}\text{C}_2 - \delta^{13}\text{C}_3 < 0\text{‰}$), and a high content ratio ($\ln(\text{C}_2/\text{C}_3) > 0.5$) between ethane and propane. The carbon isotopic gap between ethane and propane remained stable with the increase of content ratio on the whole. Correspondingly, crude oil cracking gas has the characteristics of a large carbon isotopic gap ($\delta^{13}\text{C}_2 - \delta^{13}\text{C}_3 < -6\text{‰}$), and a low content ratio ($\ln(\text{C}_2/\text{C}_3) < 0.6$) between ethane and propane. The carbon isotopic gap between ethane and propane decreases with the increase of content ratio on the whole. The geochemical characteristics of natural gas in the Y5 well area is generally consistent with that of the gas generated from the source rock thermal simulation, showing a mixture of kerogen degradation gas and crude oil cracking gas.

The natural gas genesis discrimination chart based on the relationship between $\text{C}_1/(\text{C}_2+\text{C}_3)$ and $\delta^{13}\text{C}_1$ [24] can also be used to discriminate the genesis of natural gas in the Y5 well area. As shown in Figure 6b, the data points of crude oil cracking gas are all located in the lower left corner of thermogenic gas, with a low $\text{C}_1/(\text{C}_2+\text{C}_3)$ ratio (< 50), and a light carbon isotopic value of methane ($-35\text{‰} < \delta^{13}\text{C}_1 < -50\text{‰}$). The kerogen degradation gas has the same characteristics as the crude oil cracking gas in the low maturity stage. But in the high maturity stage, the kerogen degradation gas

shows the characteristics of a high $C_1/(C_2+C_3)$ ratio and heavy methane carbon isotopic value, which are clearly different from those of crude oil cracking gas. The characteristics of both kerogen degradation gas in the high mature stage and the natural gas in the Y5 well area is between the characteristics of typical Type II and Type III kerogen degradation gas.

According to the above, it can be seen that the natural gas in the Y5 well area is a mixture of kerogen degradation gas and crude oil cracking gas, and it needs to be further confirmed which is the main kind. Li et al. discovered that affected by the differences in structure and activation energy of gas generation between crude oil and kerogen, the C_2/C_3 ratio of crude oil cracking gas is lower than that of kerogen degradation gas at the same stage of evolution, and the C_1/C_2 ratio of crude oil cracking gas at low/medium maturity ($R_o < 1.0\%$) and over maturity ($R_o > 2.0\%$) stages is smaller than that of kerogen degradation gas [25]. It can be seen from Figure 6c that the natural gas in the Y5 well area is mainly kerogen degradation gas.

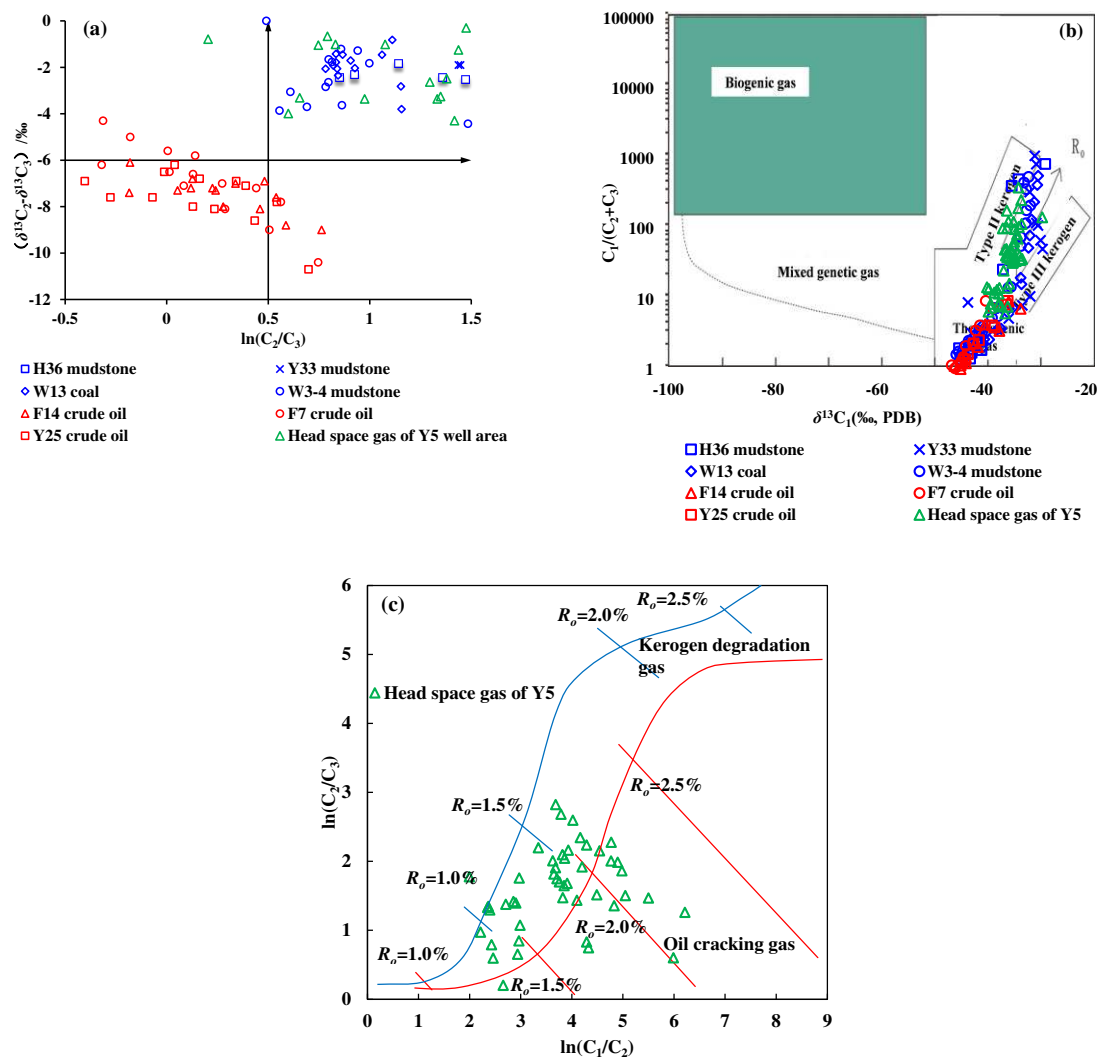


Figure 6. Discrimination of genetic types of natural gas. (a) $\ln(C_2/C_3)$ vs $(\delta^{13}C_2 - \delta^{13}C_3)$; (b) $\delta^{13}C_1$ vs $C_1/(C_2+C_3)$; (c) $\ln(C_1/C_2)$ vs $\ln(C_2/C_3)$.

Dai established the relationship between the carbon isotope value of methane and maturity based on statistical data [26]. Oil-type gas and coal-type gas follow the formula $\delta^{13}C_1 = 15.80 \log R_o - 42.20$ and $\delta^{13}C_1 = 14.12 \log R_o - 34.39$, respectively. Using these formulas, the maturity of source rocks can be calculated from the carbon isotope value of methane. It can be seen from Figure 7 that the R_o values calculated by using the coal-type gas formula are basically consistent with or slightly higher than those of the source rock (some natural gas originates from the source rock with higher maturity in the center of the sag). However, the R_o values calculated by the oil-type gas formula is higher than

those of the source rock, and the R_o values calculated below the ZH440 layer are between 4% and 8%, which is inconsistent with the geological condition of the study area. However, ultraviolet (UV) fluorescence observation revealed the presence of cracking residual asphalt in the reservoir pores of in Y5 well area, which proved that a certain amount of crude oil was charged in the early stage and cracked into gas in the later stage. In summary, the natural gas in the Y5 well area on the northern slope of Baiyun Sag is mainly kerogen degradation gas, and may contain a small amount of crude oil cracking gas.

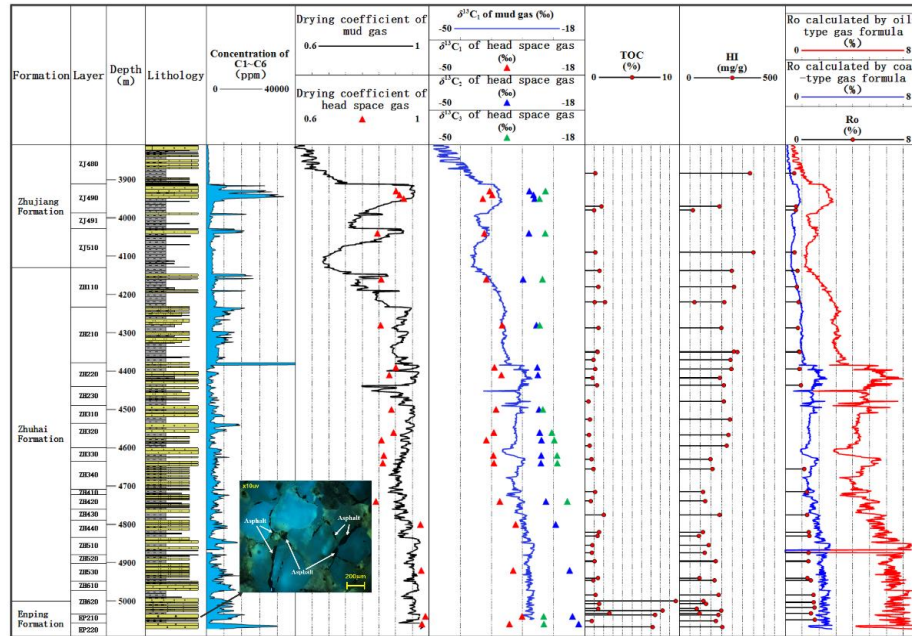
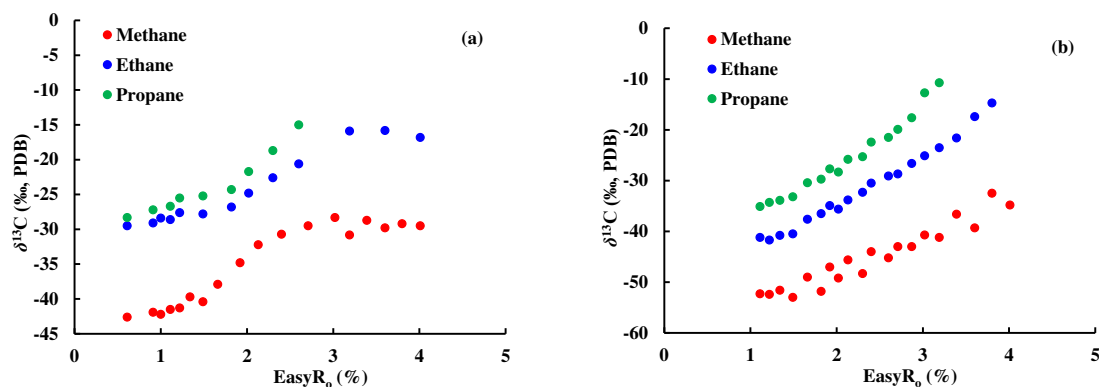


Figure 7. Comprehensive histogram of geochemical characteristics of Well Y5.

4.1.2. Discrimination based on isotope sequence reversal of natural gas

Part of the natural gas in the Y5 well area has the characteristics of $\delta^{13}C_2 > \delta^{13}C_3$ (Figure 7), while the gas in the thermal simulation experiment shows a normal carbon isotope sequence (Figure 8a,b). Studies by Dai have shown that mixed sources are the main reason for the reversal of natural gas isotopic sequences [26]. This study calculated the isotopic composition of gas mixture of Y33 source rock pyrolysis gas (representing kerogen degradation gas) and F14 crude oil cracking gas. The results showed that the mixing of kerogen degradation gas and crude oil cracking gas can cause the reversal of the isotopic sequence of ethane and propane but does not affect the isotopic sequence of methane and ethane. It was also found that the higher the maturity of gas, the lower the proportion of crude oil cracking gas that needs to be mixed for reversal. When the R_o is 1.22%, the mixing ratio of kerogen degradation gas to crude oil cracking gas is 2:1 corresponding to the reversal of ethane and propane isotope sequence, and the corresponding ratio is 15:1 when R_o is 2.30% (Figure 8c).



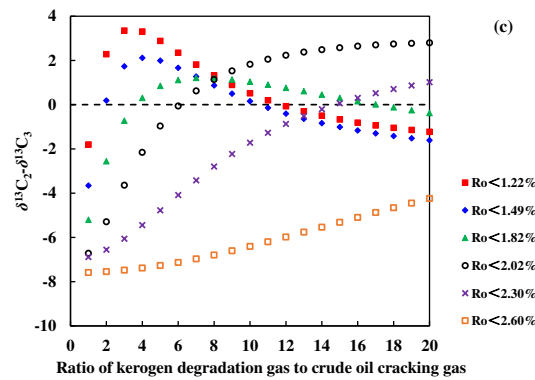


Figure 8. Isotope characteristics of single source thermal simulation gas and mixed thermal simulation gas. (a) The carbon isotope sequence of the kerogen degradation gas; (b) The carbon isotope sequence of the oil cracking gas; (c) The reversals of ethane and propane isotope sequence caused by the mixing of kerogen degradation gases and crude oil cracking gases with different maturities.

4.2. Discussion on the gas source

Combining the carbon isotope kinetic parameters of methane with the thermal history information of the Well Y5 and Well D (Figure 9), the carbon isotope value curves of natural gas generated accumulatively from different geologic times to the present were acquired (Figure 10). The homogenization temperature of the saline inclusions associated with the gas-washed oil inclusions with blue fluorescent in the Y5 well area is 170 °C–180 °C, which corresponds to the start time of natural gas charging is 11Ma (Figure 9a). It can be seen from Figure 10 that during the thermal evolution of OM, the carbon isotope of natural gas gradually becomes heavier. Since 11 Ma, the carbon isotope value of methane generated from the source rocks in Enping Formation and Wenchang Formation ranges from -42‰ to -28‰ , which is consistent with the carbon isotope value of methane in Well Y5. It shows that the natural gas in the Y5 well area mainly comes from delta to shallow lacustrine facies source rocks. From shallow to deep, the drying coefficient of natural gas increases gradually, and the carbon isotope value gradually becomes heavier, reflecting that the high-quality reservoirs in the shallower layer accumulated natural gas generated in the early stage, and the tight reservoirs in the deeper layer mainly accumulated natural gas generated in the late stage.

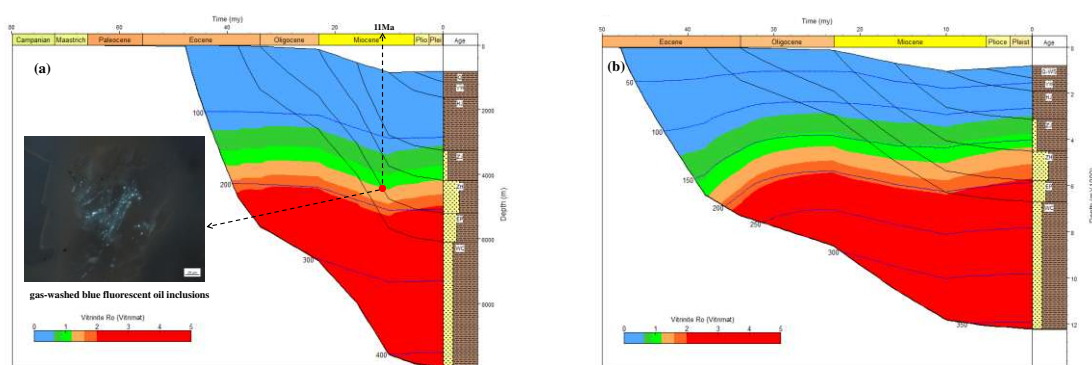


Figure 9. Sedimentary burial history and thermal history of north slope of Baiyun Sag. (a) Well Y5; (b) Well D, a virtual well located on the center of the main sub-sag.

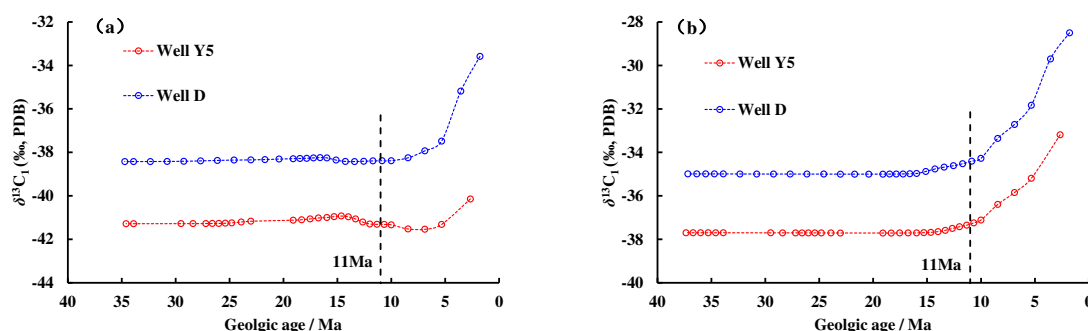


Figure 10. Carbon isotope value curves of methane generated from different geologic time to present. (a) Methane generated from the coal measures in the Enping Formation; (b) Methane generated from the lacustrine source rock in the Wenchang Formation.

5. Conclusions

Based on the data from thermal simulation experiment, this study revised the discriminant chart of the genesis of natural gas, the genetic type and source of natural gas in the Y5 well area on the northern slope of Baiyun Sag were clarified by combining the thermal simulation results and the method of carbon isotope kinetic. This research shows that:

The natural gas in the Y5 well area on the northern slope of Baiyun Sag is mainly kerogen degradation gas, mixed with a small amount of oil cracking gas. The mixing of these two types of natural gas results in the reversal of the isotopic sequences of ethane and propane, and the higher the maturity of the natural gas, the lower the proportion of cracking gas from crude oil required for the reversal.

The exist of blue fluorescent oil inclusions in Y5 well area indicated gas washing occurred at 11Ma. The isotopic values of gas generated since 11Ma from the Enping Formation coal measures and Wenchang Formation shallow lake mudstone range from -44.3‰ to -29‰, which is basically consistent with the actual isotopic composition of natural gas in the Y5 well area. The high-quality reservoirs in the shallower strata accumulated natural gas generated in the early stage, and the tight reservoirs in the deeper strata mainly accumulated natural gas generated in the late stage.

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