Post-Collisional Cu-Au Porphyry and Associated Epithermal Mineralisation in the Eastern Mount Isa Block: A New Exploration Paradigm for NW Queensland

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Article

Post-Collisional Cu-Au Porphyry and Associated Epithermal Mineralisation in the Eastern Mount Isa Block: A New Exploration Paradigm for NW Queensland

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Abstract

Post-collisional Cu-Au-Ni-Co-Pt-Pd-Sc porphyry, [Duck Creek porphyry system (DCPS)], with overlying Au-Te-Bi-W-HRE epithermal mineralisation, [Highway epithermal system (HES)] has been discovered in the core of the Mitakoodi anticline, southwest of Cloncurry. Xenotime and monazite geochronology indicates mineralisation occurred between ~1490 and 1530 Ma. Host rock lithologies show widespread potassic and/or propylitic to phyllic alteration. Paragenesis of porphyry sulphides indicate early crystallisation of pyrite, followed by chalcopyrite, with bornite forming by hydrothermal alteration chalcopyrite. Cu sulphides also show the effect of supergene oxidation alteration with rims of covellite, digenite and chalcocite. Redox conditions deduced from V/Sc systematics indicate that the DCPS contains both highly oxidized (typical of porphyries) and reduced lithologies, typical of plume generated tholeiitic and alkaline suites. Ni/Te and Cu/Te systematics plot within the fields defined by epithermal and porphyry deposits. Duck Creek chalcophile and highly siderophile element (Cu, MgO and Pd) systematics resemble data from porphyry mineral systems, at Cadia, Bingham Canyon, Grasberg, Skouries, Kalmakyr, Elaisite, Assarel and Medet. SAM geophysical inversion models suggest the presence of an extensive porphyry system below the HES. A progressive increase in molar Cu/Au ratios with depth from the HES to the DCPS, supports this conclusion. Three metal sources contributed to the DCPS-HES viz., tholeiitic ferrogabbro, potassic ultramafic to mafic system and a Fe and Ca-rich alkaline system. The latter two imparted non-crustal superchondritic Nb/Ta ratios that are characteristic of many deposits in the eastern Mount Isa Block. The associated tholeiite and alkaline magmatism reflect mantle plume upwelling through a palaeo-slab window that had accreted below the eastern flank of the North Australian craton following west verging collision by the Numil Terrane. Discovery of this linked mineral system provides a new paradigm for mineral exploration in the region.

Keywords: post-collisional porphyry; epithermal Au; North Australian craton; tholeiite and alkaline magmatism; superchondritic Nb/Ta ratios; slab window; exploration paradigm

1. Introduction

Supply chain models indicate that during the next two decades, the demand for copper will exceed the total tonnage that has ever been mined throughout earth history. This supply deficit is exacerbated by reduced production from major mines in Chile, Peru and Indonesia, a situation that has been compounded by a dearth of discoveries and by a reduction in the size and grade of these deposits. These factors provide the rationale for exploration and discovery of new copper-rich systems, in geopolitically safe jurisdictions. Cu-Au bearing porphyry deposits are attractive

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exploration targets because, although they have low copper and gold grades, e.g., 0.5 to 1.5 wt% Cu and 0.1 to 1.5 g/t Au, they are high tonnage systems, ranging from 10^6 to 10^9 tonnes [1,2].

Porphyry deposits occur in continental crust above subduction zones, where oceanic slabs release volatiles during dehydration, metasomatizing the overlying mantle wedge, facilitating partial melting and forming water-rich, metal-bearing oxidized juvenile mafic to felsic magmas [3–5]. However, Porphyry deposits also occur in post-subduction continental collision zones, associated with less hydrous magmas [6,7]. Metals are derived from several sources [2,8–10]. Copper is largely derived from hydrated subducted oceanic lithosphere and is introduced into the overlying mantle wedge by hydrous fluids. Metals are also derived from alkaline and tholeiitic plume magmas that enter the mantle wedge via slab tears, some of which may be "fossil" tears [8,10]. Porphyry systems containing a significant plume component are also generally enriched in Au and PGEs [11–16].

In water-rich sulphate bearing oxidized magmas that generate porphyry Cu-Au deposits [17], sulphur controls the behaviour of Cu and other chalcophile elements, due to the high partition coefficients that exist between these elements in sulphides and silicate melts. Cu-rich sulphides in parental magmas are destroyed by oxidation increasing the chalcophile element concentration of the resulting sulphur-undersaturated magmas [18]. In these sulphide undersaturated magmas Cu and Au both behave like incompatible elements [17], enabling chalcophile element concentrations to increase in evolving sulphate-bearing magmas. Precipitation of metals in porphyry systems is finally driven by sulphate reduction, crystallization of magnetite, decreasing pH and degassing of oxidized gases [23]. Elements liberated from sulphides by sulphate reduction are scavenged by aqueous fluids. These fluids transport metals to shallower depths, into epithermal systems above the porphyry stockwork mineralisation [20]. Such epithermal vein and breccia precious metal-rich systems form from silica- and bicarbonate-rich hot springs at palaeo-depths of less than 1 km. [21]. forming a distinctive array of textures, that are valuable exploration vectors [22].

Porphyry deposits are mainly preserved in less eroded Phanerozoic upper crust at depths between 2–4 km [23,24]. However, a few examples of Precambrian porphyry-epithermal mineralisation have been reported. For example, in the Archaean (2.83–2.82 Ga) Kolmozero–Voronya greenstone belt in the Kola Peninsula [25], in the Palaeo-to-Meso-Archean North Pilbara Terrane [26] and in Neoarchean crust at Boddington in the Southwest Yilgarn craton [27]. Porphyry mineralisation of Palaeoproterozoic age has also been described from the Tapajós Mineral Province in Amazonia, Brazil [28] and the Great Bear magmatic zone in Canada [29]. In fact, in the Great Bear magmatic zone, there is a continuum of deposits styles ranging between iron oxide coper gold IOCG mineralisation and arc-generated porphyry Cu-Au deposits [30]. Temporally and spatially associated IOCG and porphyry mineralisation of Mesoproterozoic age also occurs in the Gawler Craton, South Australia [31]. These provide a further example of the close spatial association between IOCG and porphyry deposits reported from Northern Chile [32]. According to Sillitoe [33] this close association could reflect variation in level of exposure, with IOCG magnetite-Cu-Au deposits occurring at depth and oxidized porphyry style hematite Cu-Au mineralisation forming at shallower crustal levels.

In this paper, we report the discovery of Mesoproterozoic post-collisional porphyry and epithermal mineralisation in the eastern segment of the Mount Isa Block. It is preserved in a zone of low strain in the core of the Mitakoodi anticline, approximately 25 km southwest of Cloncurry. The mineral system is hosted by a suite of post-tectonic intrusive lithologies, that include tholeiitic ferrogabbro, leuco-gabbronorite and anorthosite, as well as an alkalic suite that includes websterite, ultramafic lamprophyre, melilite-bearing gabbro, and monzodiorite. These units crop out as roof pendants that intrude Marraba Volcanic meta-basalt, meta-andesite, tuffaceous units. Occurring above the porphyry mineralization, at a higher palaeo-crustal level, are sheets of carbothermal lithologies and associated siliceous crackle breccias that preserve distinctive epithermal textures indicative of deposition from fumarolic CO_2-rich fluids. The epithermal vein and breccia system has a strike length of > 20 km and is structurally controlled by axial plane shear zones of parasitic folds on the regional Mitakoodi anticline.

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We review key lithogeochemical features of the Highway-Duck Creek mineral system using an extremely large petrological geochemical data base. Using pivotal geochemical vectors that facilitated the discovery, we present an alternative metallogenic model for the Cloncurry region that better accommodates the geodynamic model for the tectonic evolution of the eastern flank of the North Australian Craton. Given the coherence of the lithogeochemical data, this model may be applicable to other deposits in the region, particularly, to the so-called, "within craton IOCG deposits" that have recently been classified as "intracratonic copper-gold deposits" by Brauhart and Groves [34].

The Mount Isa block, also referred to as the NWMP (Figure 1), is spectacularly enriched in a wide range of metals, including the Tier 1 sediment-hosted Mount Isa copper deposit, and the Ernest Henry iron–oxide–copper–gold (IOCG) deposit [35,36]. However, since discovery of the Ernest Henry deposit more than 30 years ago, the region remains remarkably under explored and few significant new discoveries have been made. In view of the significant metal endowment of the province, this limited exploration success is an enigma. It may reflect the fact that current exploration models do not fully accommodate the possible existence of different mineral systems. Alternatively, it may simply reflect the fact that "barren" magnetic targets have been drilled rather than "fertile" non-magnetic hydrothermally altered targets.

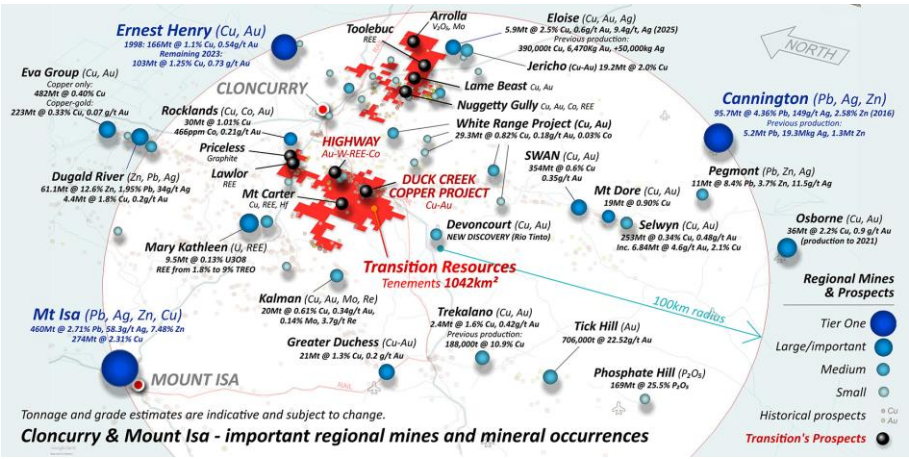

Figure 1. Map showing the location of the Highway-Duck Creek systems in relation to principle mineral deposits in the Northwest Minerals Province.

2. Geological Background

Exploration in the in the Cloncurry segment of the NWMP has largely focused on the post-Isan Orogeny Cu-Au-Co-REE-Re-Mo-U IOCG, or iron sulphide-rich (ISCG) mineral systems. Metals are interpreted to have been derived either from a granitic source, such as the Williams-Naraku Batholith; [37], or by mixing between magmatic fluids with basin brines [38–40]. Fluid inclusion studies, and both stable and rare gas isotopes indicate that, although fluids largely display crustal isotopic signatures, there is a cryptic mantle signature [41]. However, neither the granitic, or the brine source, accounts for the wide range of chalcophile (Ag, As, Bi, Cd, Cu, Hg, In, Pb, S, Sb, Se, Te, Tl, and Zn), siderophile (Co, Ni, Mo, W, Re, Os, Pt, Pd and Au) or lithophile (Sc, V, Cr, Y, Zr, Nb, REE, Th and U) abundances reported in these deposits. The multi-element association of Co-Ni-V-Sc-PGE with high Nb/Ta ratios reported in many Cloncurry deposits, is clearly inconsistent with derivation from a crustal granitic source and requires the involvement of an ultramafic/mafic mantle-plume component.

Such an alkaline ultramafic suite occurs at Mount Cobalt, ~90 km south of Cloncurry [42]. These lithologies include olivine-bearing pyroxenites, with a superchondritic average Nb/Ta ratio of 25. As discussed later, elevated Nb/Ta ratios are clearly inconsistent with a crustal metal source, indicating Nb-Ta fractionation in either a mantle plume alkaline magmatic system [43], in metasomatized continental lithospheric mantle [44] or in deep arc eclogitic cumulates stored between

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400 and 650 km in the transition zone [45]. The Mount Cobalt lithologies are extremely enriched in Co 10 wt.%, Ni 3117 ppm, Sc 278 ppm, Au 5 ppm, Pd 8 ppb, Pt 3.7 ppb, and they have significant concentrations of heavy rare earth elements, Dy up to 315 ppm and Tb up to 90 ppm. Thus, these lithologies provide direct evidence of a viable ultramafic source that accounts for the anomalous highly siderophile elements (Co, Ni, Mo, W, Re, Os, Pt, Pd) and Au in Cloncurry district mineral systems.

Furthermore, Ca-rich pyroxenites, melilite-bearing gabbroic lithologies and ultramafic lamprophyres with elevated superchondritic Nb/Ta ratios, up to 202, with an average of 48.6 (SD 29), are also present at the Elaine Dorothy copper prospect, near Mary Kathleen uranium mine (Figure 1). They define a significant and continuous depth interval between 232m and 871m. Mafic lithologies with superchondritic Nb/Ta ratios also present at Lawlor (Figure 2), ~10 km northeast of Highway, where values range up to 390 an average of 54 (SD 81). These examples clearly provide additional evidence for the role of mafic alkaline magmatism in the area [46].

Such cryptic evidence for the role of plume magmatism within the eastern segment of the North Australian Craton, and the metal association of high-K calc alkaline hosted porphyry mineralisation in the eastern NWMP is consistent with the inferred geodynamic evolution of the region. The eastern segment of the North Australian Craton is interpreted to be an early Proterozoic convergent Andean margin that developed above a west dipping subduction zone [47]. This interpretation was subsequently confirmed by reflection seismology, with the identification of a prominent west dipping feature, the Gidyea Suture, that extended below the Mount Isa Block that juxtaposed the outboard suspect Numil terrane with the North Australian Craton [48,49]. Thus, the Mesoproterozoic metallogenic endowment of the eastern NWMP, appears to reflect the geodynamic interplay between subduction zone magmatism, back arc basin rift magmatism, slab tear facilitated tholeiitic and alkaline plume magmatism, collisional orogenesis and finally crustal anatectic magmatism that produced the regionally widely distributed alkaline and peraluminous Williams Naraku Batholith intrusions.

The involvement of Mesoproterozoic intraplate tholeiitic and alkaline plume magmatism in the NWMP is consistent with the geodynamic model that explains the breakup of the Columbia supercontinent, when the Paleo-Australian components of Columbia, i.e., the Gawler, Curnamona and North Australian Cratons traversed an upwelling mantle plume [50,51]. The eastern margin of the North Australian Craton in the Cloncurry area is a collisional orogenic belt, comprising an accreted collage of subduction generated Palaeoproterozoic crust and telescoped back arc basins [36,52]. As collisional orogens are commonly underlain by slab tears [53,54], these tears provide a window for penetration by an upwelling plume into either the mantle wedge or the sub-continental lithospheric mantle. The slab tear window most likely accreted to the base of the sub-continental lithospheric mantle where it provided a conduit for ingress by upwelling plume magmas. Slab tears are a common feature of continent-continent collisional environments [54]. Metals in Mesoproterozoic NWMP deposits are interpreted to have been derived from two sources; Cu from altered subducted oceanic crust that was transported via hydrous fluids into the mantle wedge, and Cu, Au and platinum group elements introduced into continental

lithosphere by tholeiitic and alkalic plume [55,56].

2.1. The Highway - Duck Creek Mineral System

Epithermal Au-Te-W-REE mineralisation, and the underlying porphyry Cu-Au-Co-Ni-Sc system, discovered southwest of Cloncurry (Figures 1 & 2) is a greenfield discovery as there no evidence in the field to indicate any previous mining activity in the area. Although the area contains a myriad of historic supergene copper workings, it had been overlooked for modern exploration for almost 100 years. The discovery was made using p-XRF analyses that showed elevated tungsten (W) gold (Au) and yttrium (Y) in silica-rich breccia that exhibited epithermal quartz textures. The name "Highway" was coined to acknowledge the anomalously ***high*** concentrations of ***W***, ***Au*** and ***Y*** in these breccias.

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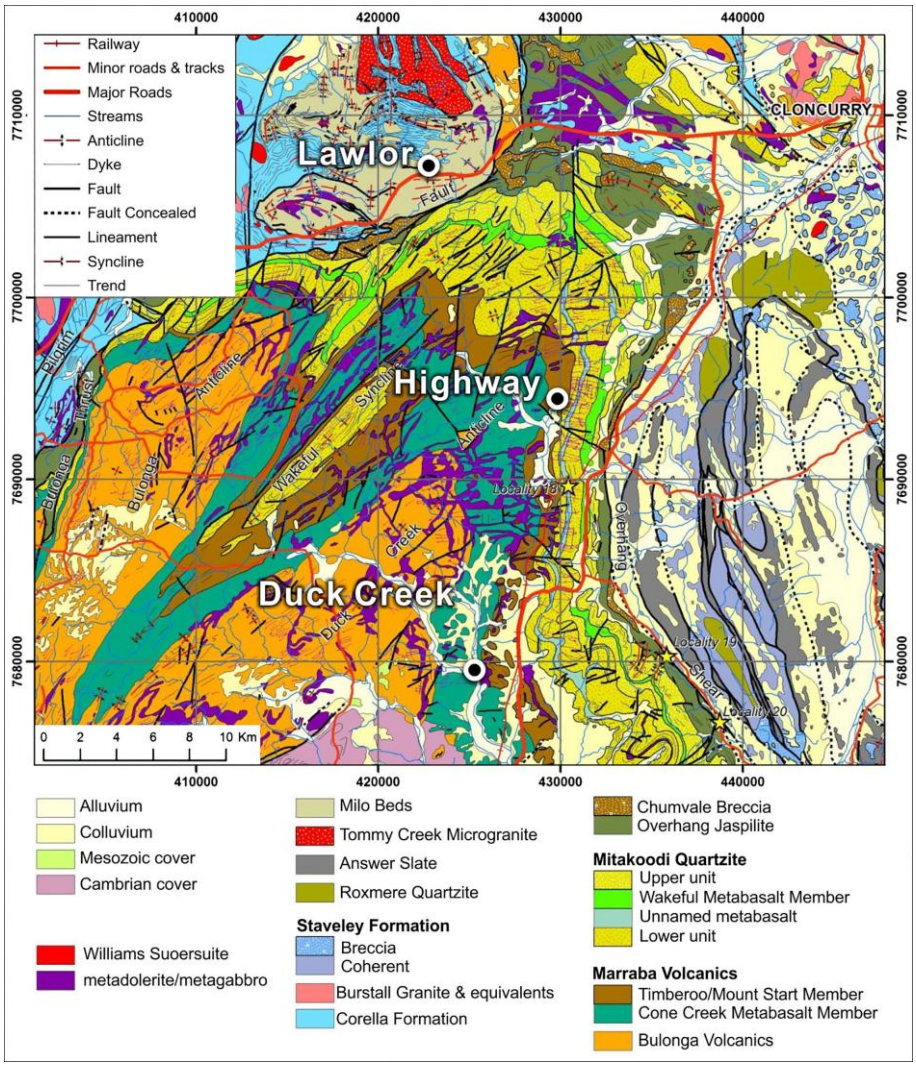

Figure 2. Geological map of the Mitakoodi anticline showing locations of the Highway, Duck Creek and Lawler prospects.

The silicic breccias and associated carbothermal veins, that contain elevated Au and tellurium Te have been traced for more than 20 km along the eastern limb of the Mitakoodi Anticline [57,58] (Figure 2). They are interpreted to have crystallized from boiling silica and CO_2 - rich hydrothermal and carbothermal fluids at shallow crustal depths. These fluids accessed the axial plane shear zones of parasitic folds around the Mitakoodi anticline where they precipitated.

The Duck Creek porphyry system occurs at a deeper level of crustal exposure in the core of the Mitakoodi Anticline, west of the Highway epithermal mineralisation. Litho-geochemical data for drill core indicate that mineralisation occurs in a post-tectonic potassic suite that includes olivine websterite, pyroxenite, gabbro, leucogabbro, anorthosite and monzodiorite and monzonites. In view of the common association between potassic magmatism and large, high-grade Cu-Au porphyry deposits (Müller and Groves 2019), recognition of the involvement of post-tectonic potassic magmatism in the eastern NWMP has important implications for the resource potential of the Highway-Duck Creek discovery. The Highway-Duck Creek epithermal-porphyry mineralization post-dates the ~1650 Ma) thermotectonism in the Mount Isa Block.

The discovery was facilitated through application of a new exploration model for the Cloncurry region based on results reported by Collerson in [42]. This model specifically addressed apparent deficiencies with the IOCG model that specifically focused on the role played by post-tectonic Williams Batholith granitoids, whose felsic compositions failed to explain the distinctive element association viz., Cu-Au-Ni-Co-Pt-Pd-Sc (±HREE+Y) and the superchondritic Nb/Ta recorded in the

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Highway-Duck Creek mineralisation. The model provided an alternative explanation for the spatial association and source of metals in iron-oxide-copper-gold (IOCG) and iron-sulphide-copper-gold (ISCG) mineralisation in the Cloncurry area cf., Skirrow [59], that have recently been assigned to be members of a new clan of intracratonic ore deposits by Brauhart and Groves [34]. By accommodating spatial and geochemical relationships of mineralisation, within a coherent narrative that is consistent with models for the geodynamic evolution of eastern margin of the North Australian Craton, the new model clearly provides a plausible exploration basis for future discoveries.

2.2. Geological Setting of the Mitakoodi Domain

The megascopic scale Mitakoodi anticline (Figure 2) [57,58] and an associated west-verging thrust system, formed during the Isan Orogeny at ~ 1650 Ma [57,58,60,61]. The core of the Mitakoodi anticline is a region of low finite strain that preserves a varied sequence of Marraba Volcanic lithologies including the Bulonga Volcanics, the Cone Creek and the Timberoo Members. The Bulonga Volcanics are conformably overlain by interbedded mafic volcanics and fine clastic units of Marraba Volcanics and by a thick (up to 2000 m) succession of quartzo-feldspar arenite, termed the Mitakoodi Quartzite. The Mitakoodi Quartzite contains minor interbedded lenses of siltstone and metabasalt as well as a unit of thinly bedded tuff, near the top of the sequence.

The top of the quartzite contains a unit of porphyritic rhyolite which has been mapped for a strike length of >3 km. As rhyolite is a high viscosity magma, that forms lava domes, not extensive flows, the rhyolite sequence is possibly better interpreted as a pyroclastic ignimbrite sheet or tuff. A U-Pb SHRIMP zircon age of 1755 ± 4 Ma for the "rhyolite" [58] is within error of the age of the felsic Bulonga volcanic unit dated by Neumann et al., [62] in the core of the Mitakoodi anticline. The entire package of lithologies that crop out in the Mitakoodi Culmination represent a previously unrecognized intraplate suite, that was possibly deposited in a back-arc rift environment. Some mafic units are quite massive and most likely represent high-level ultramafic to mafic plutonic bodies.

The Mitakoodi Domain was subsequently intruded by phases of the Wimberu Granite member of the Eureka Supersuite of plutons within the Williams Batholith. The Wimberu Granite yields LA-ICP-MS zircon U/Pb ages of 1518 ± 4 and 1483 ± 4 Ma [63]. This younger Williams - Naraku Batholith magmatism between ~1530 – 1490 Ma, resulted in the formation of a complex suite of granitoids ranging in composition from dioritic through monzonitic, syenitic to granitic compositions with A-type geochemical signatures. Highway lithologies intrude regionally extensive rhyodacitic and rhyolitic lithologies of the 1750±4 Ma Bulonga Volcanics [62] in the core of the Mitakoodi anticlinorium (Figure 2). This has been confirmed by laser ablation U- Pb geochronological data for xenotime and monazite from the Highway Breccia which yield ages of ~1530 Ma and ~1490 Ma respectively.

2.3. Field Relationships and Petrology

Highway system is characterised by dominantly mono-lithologic quartz crackle breccias containing angular clasts that are cemented in a metal-rich matrix (Figure 3a-e). They contain chalcedonic siliceous vuggy veins, with characteristic epithermal textures, e.g., comb textured microcrystalline quartz, crustiform-colloform banded quartz and lattice textures (Figure 3d & f). The distinctive epithermal lattice texture in Figure 3f is interpreted to have formed by quartz replacing earlier deposited calcite in a hydrothermal system. Such textures are characteristic of low sulphidation, epithermal environments where boiling carbonate rich hydrothermal fluids culminated in the release of CO_2. Field and petrographic observations indicate that the breccias, the epithermal textured siliceous veins and the carbothermal veins and carbothermal lithologies (Figure 3g) were deposited from carbonate-rich hot spring (travertine), some contain chalcedonic clasts (Figure 3h).

In thin section, the epithermal lithologies exhibit very distinctive textures (Figures 4 and 5). Examples include interlayered quartz comb texture in a carbonate matrix (Figures 4a and 7 b), a vein of epithermal quartz cutting propylitically altered gabbro (Figures 4c & d) , a fine grained siliceous brecciated unit cemented by veins of epithermal quartz (Figures 4e and 7 f), cross cutting

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carbothermal and siliceous epithermal veins that indicate the impact of several hydrothermal events in the area (Figures 4g & h) and siliceous comb textures (Figure 4i). Also shown is a reflected light is an image of a Highway epithermal lithology containing multiple grains of gold.

Photomicrographs of typical Highway carbonate- and quartz-bearing carbothermal lithologies are show in Figure 5. Xenomorphic and idiomorphic grains of quartz typically exhibit slightly embayed grain boundaries that reflects effect of acid (HF?) corrosion (Figures 5 a to d and 5 g, h & i). Some of the carbothermal units are tourmaline-bearing (Figures 5 e & f) indicating that boron was an important component in the hydrothermal fluids.

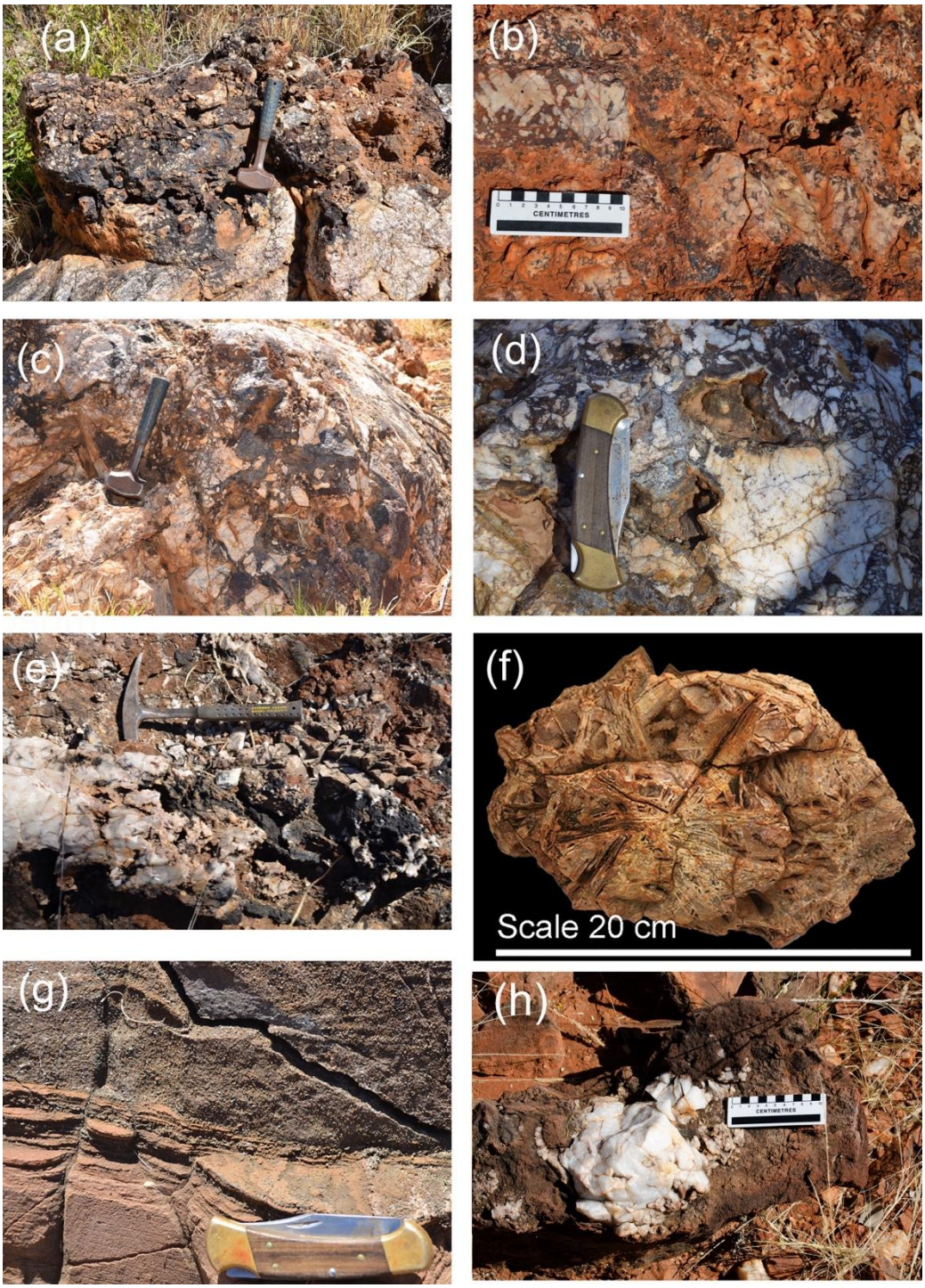

Figure 3. Highway Epithermal lithological variation (a to e & h) siliceous breccias showing embayed clasts of milky quartz, (f) rhombohedral carbonate pseudomorphed by epithermal silica, (g) carbothermal sinter and (h) embayed siliceous clast in a carbothermal matrix.

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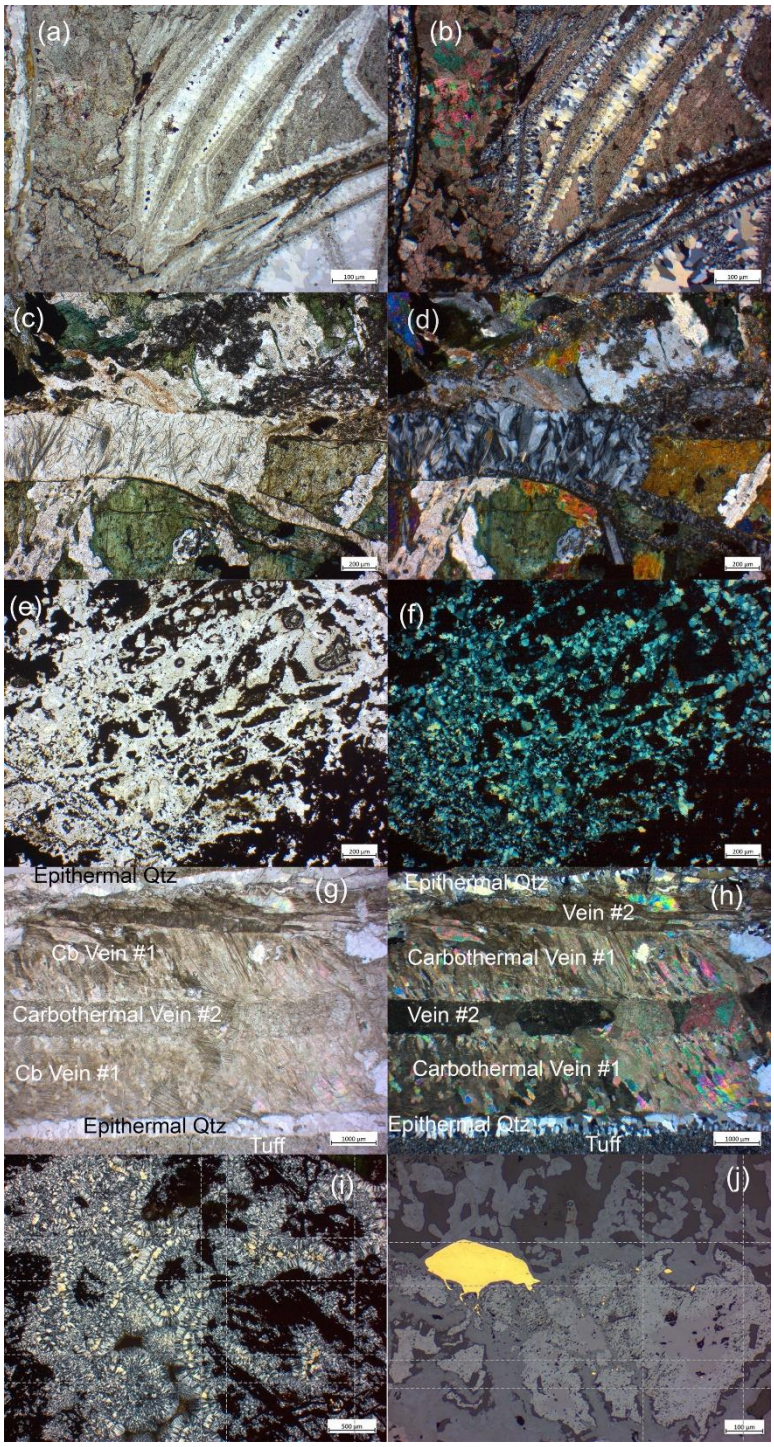

Figure 4. Photomicrographs of epithermal textures, (a & b) interlayered quartz comb texture in a carbonate matrix in plane and cross polarised light, (c & d) vein of epithermal quartz cutting gabbro that exhibits propylitic alteration in plane and cross polarised light, (e & f) fine grained siliceous brecciated unit cemented by veins of epithermal quartz in plane and cross polarised light, (g & h) carbothermal and siliceous epithermal veins showing evidence of several hydrothermal events in plane and cross polarised light,(i) classic epithermal siliceous comb textures in cross polarised light, (j) reflected light image showing grains of gold in Highway epithermal lithology.

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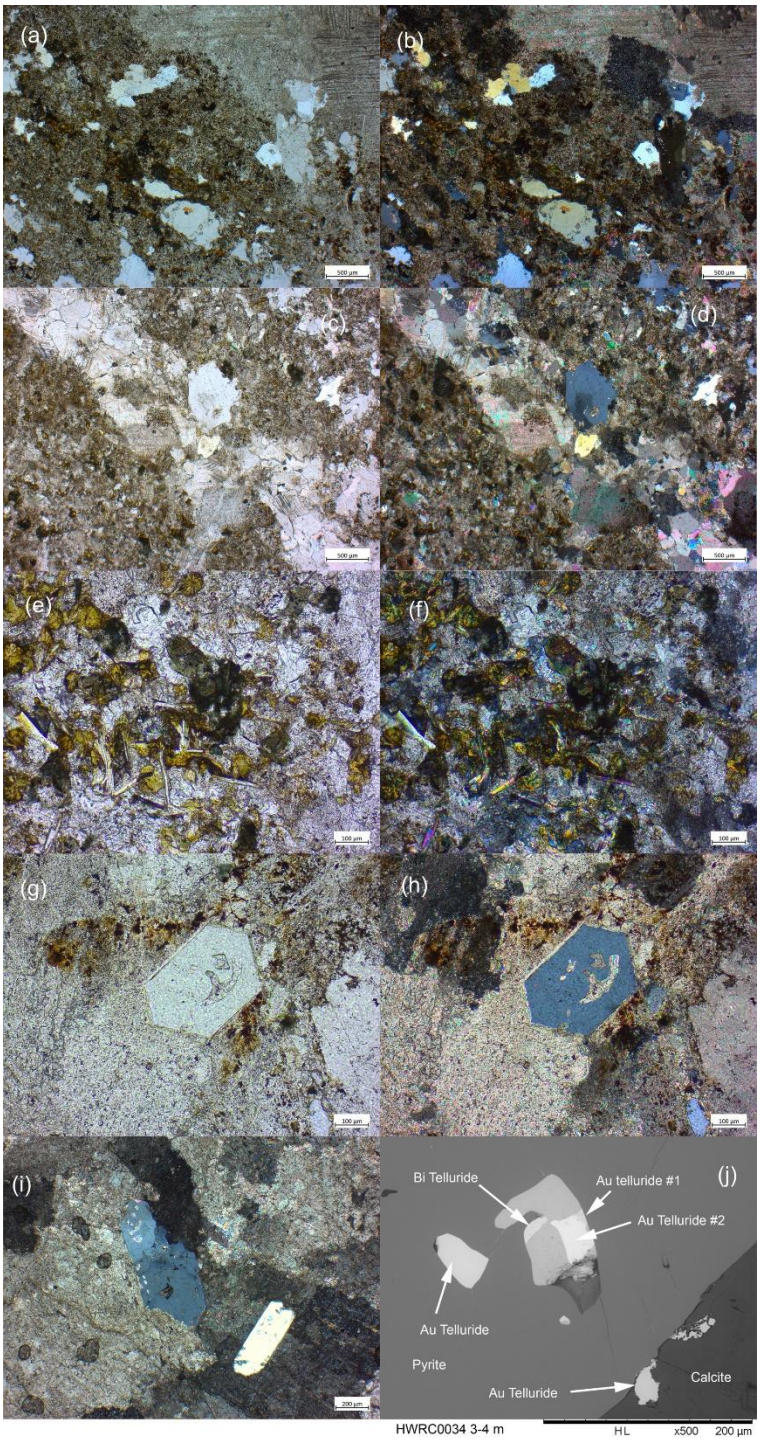

Figure 5. Photomicrographs of typical Highway epithermal lithologies in plane and cross polarised light, (a–d) and (g-i) carbothermal lithologies with xenomorphic and idiomorphic grains of quartz that show embayed grain boundaries caused by acid (HF?) corrosion, (e & f) tourmaline-bearing carbothermal lithology which shows the role of B in the fluid phase, (j) Back scattered electron image of Au and Bi tellurides in pyrite.

Boron is a significant component of porphyry-related systems that also host epithermal deposits, and its occurrence is common vector indicating the presence of a deeper porphyry copper deposit. Figure 5j shows a back scattered electron image pyrite containing inclusions of gold and bismuth telluride. Gold and bismuth tellurides are key components of some epithermal gold deposits, forming from hydrothermal fluids and often associated with alkaline or sub-alkaline magmatism [64].

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3.4. Duck Creek Mafic Intrusive Complex

Tholeiitic ferrogabbros, gabbronorites, leucogabbros and anorthosites are medium to fine grained, with intergranular to subophitic textures. A=bsence of deformational fabrics clearly indicates that the intrusions were emplace after thermotectonism in the NWMP (Figure 6).

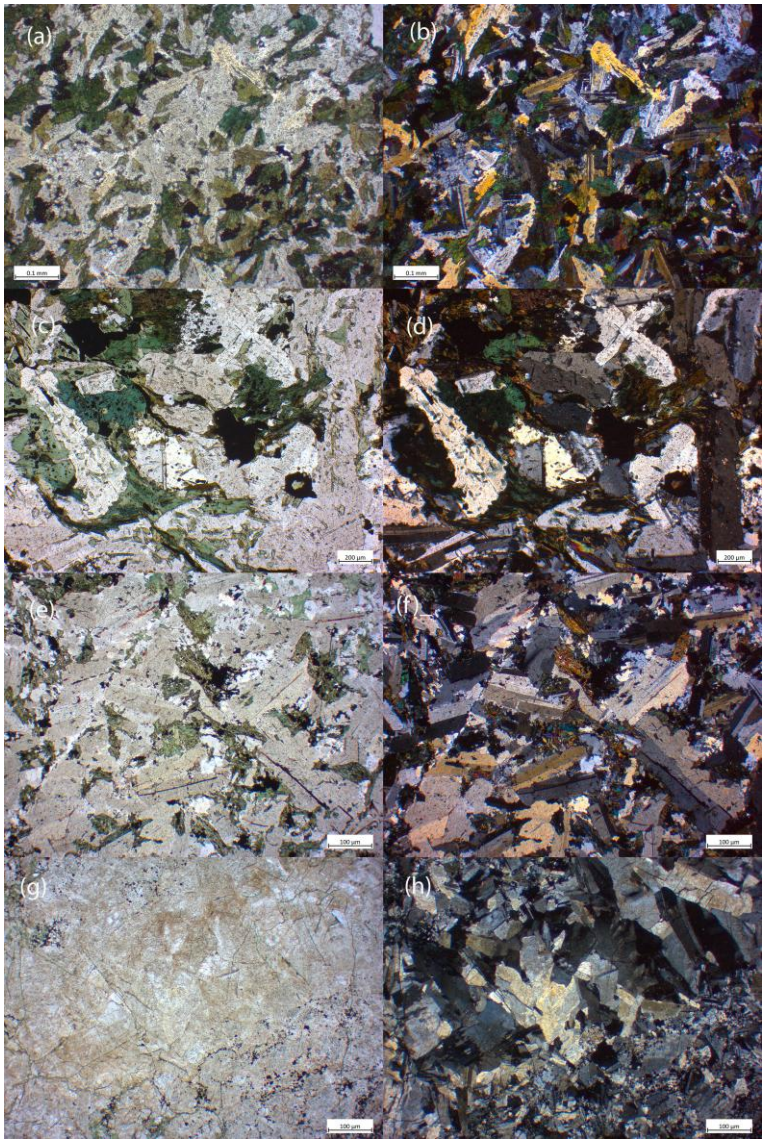

Figure 6. Photomicrographs of gabbro (a - d), leucogabbro (e & f) and anorthosite (g & h) from the tholeiitic ferro-mafic unit that intrudes the Bulonga Volcanics in plane and cross polarised light showing intergranular and sub-ophitic textures. These are clearly undeformed and are therefore interpreted to be post tectonic and therefore post-Isan orogeny. .

The fine grain size of these samples is interpreted to indicate that they represent the upper chilled margin of a much larger deeper intrusive complex. This is supported by the fact that they crop out as roof pendants within the Marraba Volcanics (Figure 2). The intrusion contains lithologies that resembles lithologies in the plume generated Nain Plutonic Suite, Labrador, Canada [65]. The lack of seismic anisotropy in the deep crust below the eastern margin of the North Australian Craton [48,49] shows the extensive nature of this mafic intrusive complex.

2.5. Alteration in the Highway - Duck Creek Mineral System

Marraba volcanics host lithologies, and members of the later ferrogabbro to monzogabbro suite all show the effect of penecontemporaneous pervasive alteration (Figure 7).

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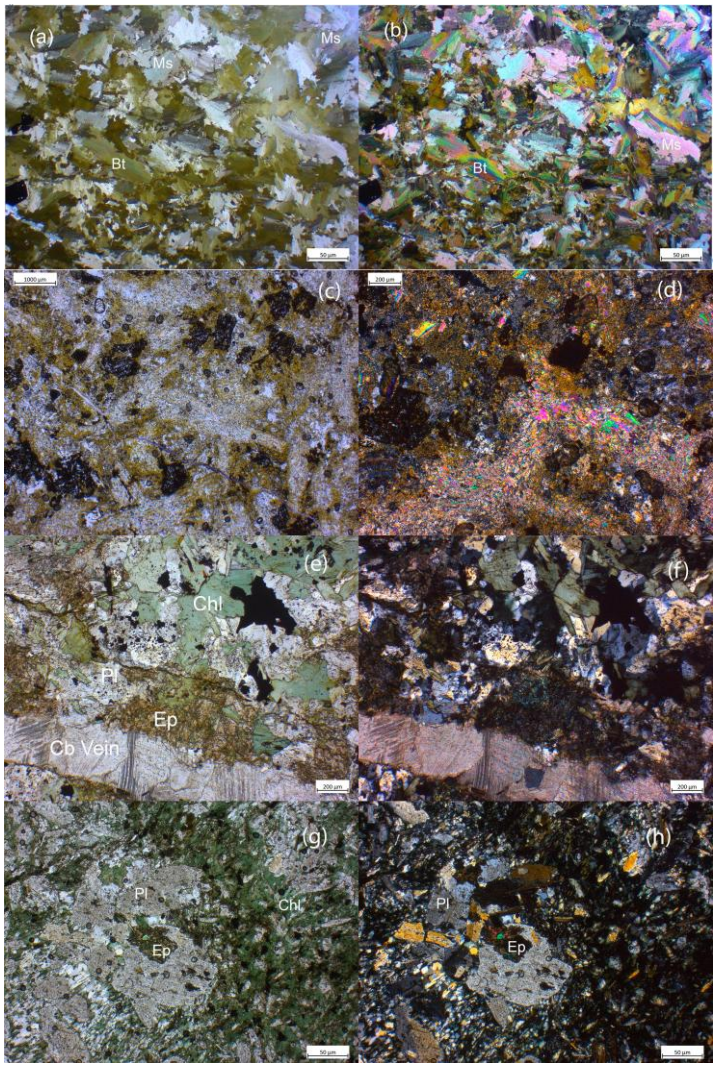

Figure 7. Photomicrographs in in plane and cross polarised light showing (a & b) phyllic alteration and (c – h) propylitic alteration.

The alteration assemblages overprint metamorphic microstructures in the Marraba volcanics and igneous textures in the ferrogabbro to monzogabbro suite. Typical alteration facies include:

(1) Phyllic alteration, is shown by the presence of white mica (sericite) and quartz together with chlorite, carbonate and tourmaline (Figures 7a & b). This alteration is caused by hydrolysis due to interaction between high temperature, moderately acidic fluids with host protoliths. The presence of carbonate and tourmaline indicate that the hydrothermal fluids were CO_2 and boron bearing. Tourmaline is a common phase in both phyllic and propylitic alteration haloes associate with epithermal and porphyry Cu-Au mineral systems [66]. Carbonate is also a common phase in phyllic alteration haloes and indicates that the alteration fluids have near neutral pH's [67].

(2) Propylitic alteration is shown by the presence of chlorite, epidote, carbonates (calcite), albite, and pyrite, often replacing original ferromagnesian minerals like biotite and amphibole (Figures 7c to h). It is also a common feature along the margins of porphyry copper and associated epithermal precious metal deposits.

(3) Potassic alteration is shown by the presence of secondary biotite, K-feldspar (orthoclase), and sometimes muscovite or sericite. It forms by reaction between protoliths of different compositions with hot (420-500°C) metal-rich fluids. Potassic alteration is commonly associated with sulphide mineralization, near the core of porphyry copper-gold systems.

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3. Highway - Duck Creek Mineralisation

Figure 8 shows reflected light and back scattered electron images examples of pyrite and Au telluride (Fig. 8a), scheelite and wolframite (Fig. 8b, c and f), gold mineralisation (Fig. 8d & e) titanomagnetite and pyrite (Fig. 8g), titanomagnetite (Fig. 8h & i) in Highway epithermal lithologies.

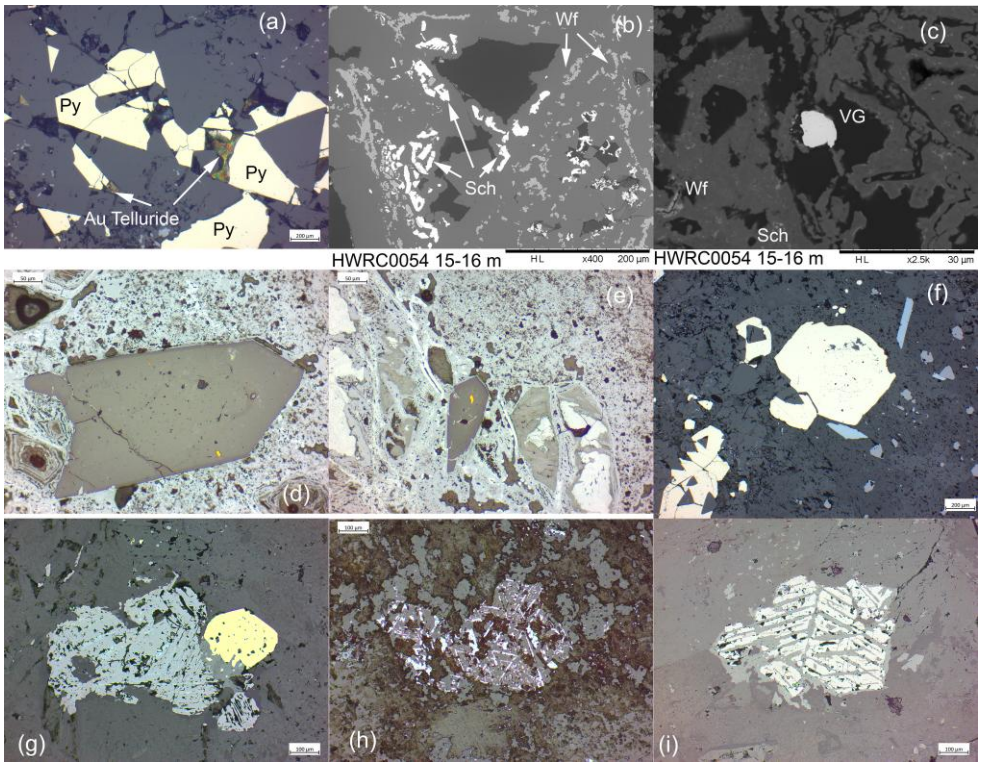

Figure 8. A selection of photomicrographs in reflected light and back scattered electron images showing (a) pyrite and Au telluride, (b) scheelite and wolframite, (c) scheelite, wolframite and gold, (d & e) gold in quartz crystals, (f) pyrite and wolframite, (g) titanomagnetite and pyrite, (h & i) titanomagnetite in Highway epithermal lithologies.

Rare earth elements in the Highway epithermal system occur in monazite and xenotime. As these phases occur in cross cutting veins (Figure 9) this mineralisation was clearly of hydrothermal origin.

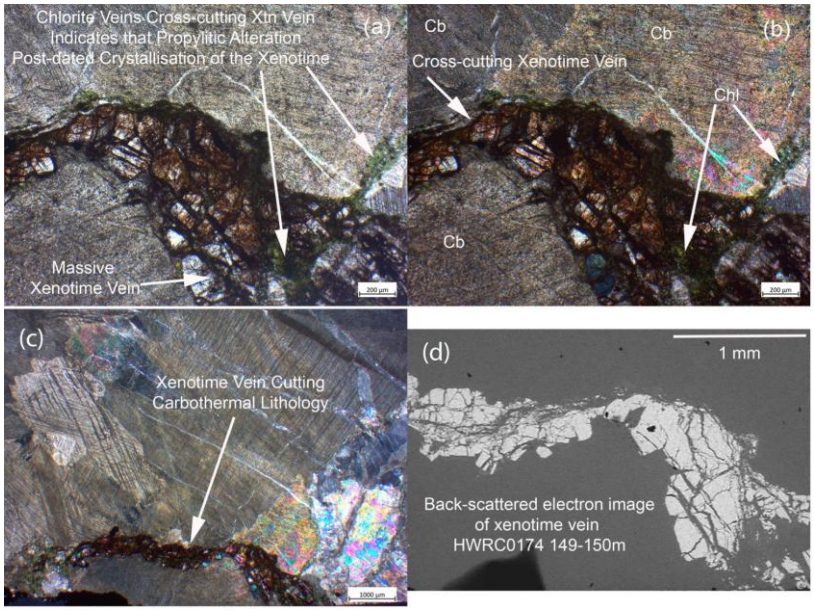

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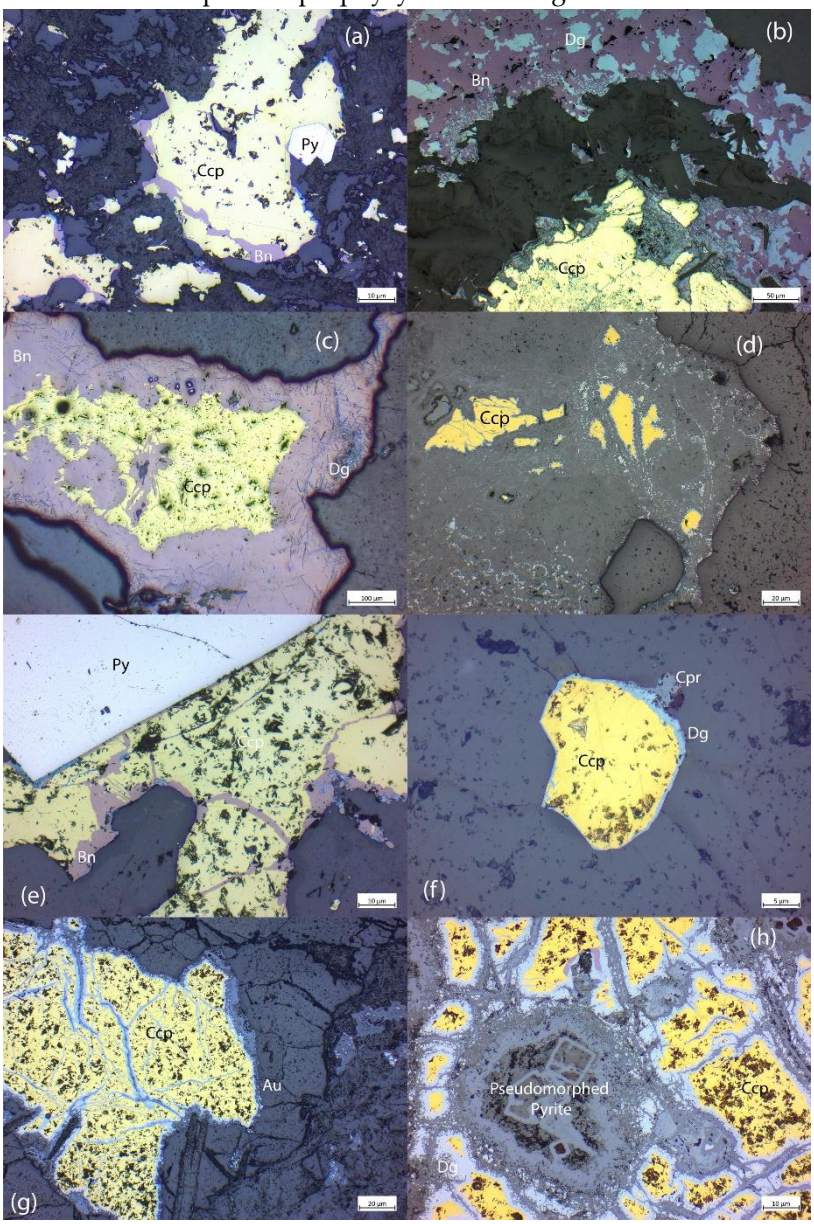
Figure 9. Carbothermal lithology showing details of a cross-cutting vein of xenotime in (a) plane polarised light, (b & c) cross polarised light and (d) back scattered electron image.

Textures and paragenetic relationships of Cu sulphides in the deeper Duck Creek mineral system are shown in Figure 10. Samples are mainly from diamond drill core from the New Dollar deposit, which is located between Columbiad and Horseshoe (Figure 11).

Sulphide paragenetic relationship show early formed crystallisation of euhedral pyrite and then subsequent crystallisation of chalcopyrite (Figure 10a & e). In some samples the early formed pyrite is brecciated and cut by narrow veins of chalcopyrite. Chalcopyrite typically exhibits coronas of bornite which is interpreted to have formed by oxidation (Figure 11b & c). Supergene alteration of chalcopyrite is shown by formation of digenite and covellite (Figure 11d, f & g) (h) pyrite pseudomorphs surrounded by chalcopyrite with rims of digenite New Dollar (19m).

The presence of early euhedral pyrite and later chalcopyrite is a common paragenetic feature of many Cu-Au porphyry systems e.g., Cadia in N.S.W. The presence of early formed euhedral pyrite in porphyries is interpreted to indicate that it crystallised from a hydrothermal fluid [68]. Bornite overgrowths on chalcopyrite are interpreted to have formed by hydrothermal alteration (Figures 10b-h). Bornite shows the pervasive effect of supergene replacement by digenite, covellite and chalcocite. The presence of abundant interstitial chalcopyrite in the ferro-gabbroic to monzogabbroic suite indicates a close relationship to the porphyry mineralising event.



 doi:10.20944/preprints202601.1172.v1

Figure 10. Photomicrographs of sulphides from the Duck Creek porphyry system in reflected light: (a) sulphide paragenetic relationship at Meteor (125m) showing euhedral pyrite, overgrown by chalcopyrite which is oxidised to bornite, (b & c) chalcopyrite rimmed by digenite, bornite and covellite at New Dollar (85.3m and 140m), (d) chalcopyrite with narrow rims of digenite New Dollar (140m), (e) euhedral crystal of pyrite with an overgrowth of chalcopyrite and rim of bornite produced by hydrothermal alteration, New Dollar (150m), (f & g) chalcopyrite with coronas of digenite New Dollar (19m), (h) pyrite pseudomorphs surrounded by chalcopyrite with rims of digenite New Dollar (19m).

4. Materials and Methods

4.1. Materials

Samples from the Highway and Duck Creek were collected by KDC and by DW and geological staff employed by Transition Resources between 2018 and 2025.

4.2. Methods

The polished thin sections were prepared in the Queensland University of Technology lapidary facility. Petrographic studies were undertaken by KDC using a ZEISS Axio-scope 5 Polarizing Reflected/Transmitted Light Petrographic Microscope in the School of the Environment, at the University of Queensland. Mineral identifications were confirmed using a Hitachi TM3030 Scanning Electron Microscope (SEM) equipped with a Bruker Energy Dispersive Spectrometer (EDS) also in the School of the Environment at the University of Queensland.

Analytical data reported and discussed in this paper, were obtained using XRF, fusion ICPOES, fusion ICPMS and fire assay provided by Intertek, Perth and Townsville. Certified litho-geochemical standards were analysed with all sample submissions and a full QA/QC assessment for accuracy was undertaken.

The data includes representative major and trace element analyses of lithological groups identified in rock chip and drill core from the Highway epithermal system (Table 1). These were selected from a data set of > 33,287 analyses of rock chips and both RC and DD drilling at Highway.

More than 17,915 rock chip, RC and diamond drill core samples from the Duck Creek area were also reviewed. Data reported in Table 2 are average analyses of 1801 individual analyses from historic mines in the Duck Creek area. This data was obtained with support from a Queensland Department of Mines, through a Cooperative Exploration Initiative (CEI) grant to Transition Resources. Data from Mount Cobalt deposit are from Philpott [69] and Collerson [42]. Analyses of 1m composite samples from a 896 m long diamond core drilled by Chinalco at the Elaine Dorothy deposit near Mary Kathleen (MKED0036) are from Collerson [46].

No generative artificial intelligence (GenAI) assistance was used in the preparation of this manuscript.

Table 1. Partial major element analyses for variably altered lithologies from the Highway epithermal system.

	#1	**#2**	**#3**	**#4**	**#5**	**#6**	**#7**	**#8**	**#9**	**#10**	**#11**	**#12**
Number of Analyses	154	783	426	888	12	64	1045	722	47	435	15	163
SiO_2	51.71	52.86	53.49	50.51	59.84	58.37	60.69	61.91	57.35	58.02	73.87	70.66
TiO_2	0.40	0.81	0.64	1.55	0.47	0.67	0.59	0.49	0.44	0.23	0.30	0.26
Al_2O_3	12.74	11.89	12.39	11.63	13.49	12.19	13.72	13.20	9.46	5.28	4.75	7.44
Fe_2O_3	16.99	13.91	10.74	16.42	14.48	9.64	9.21	7.58	8.87	7.77	12.21	12.32
MnO	0.13	0.17	0.13	0.17	0.18	0.13	0.08	0.12	0.26	0.38	0.69	0.22
MgO	4.40	4.73	4.00	4.55	2.96	4.76	3.81	3.58	3.32	2.94	1.52	1.77

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CaO	2.12	6.50	4.65	6.02	1.25	6.55	3.48	5.71	13.97	20.46	1.33	2.11
Na_2O	1.01	1.62	1.40	2.01	0.84	1.28	2.24	2.10	0.94	0.40	0.65	0.33
K_2O	2.23	2.27	2.37	1.38	3.18	3.25	3.01	2.15	2.24	1.37	1.56	1.73
P_2O_5	0.27	0.24	0.19	0.26	0.31	0.16	0.17	0.16	0.15	0.15	0.12	0.16
LOI	8.00	5.00	10.00	5.50	3.00	3.00	3.00	3.00	3.00	3.00	3.00	3.00
Total	100.0	100.0	100.0	100.0	100.0	100.0	100.0	100.0	100.0	100.0	100.0	100.0

Li	18.3	13.7	15.1	13.2	10.8	5.9	16.0	13.0	10.5	7.0	14.2	8.6
Sc	35	33	24	38	36	24	18	17	19	14	23	25
V	359	322	208	356	378	193	142	127	172	108	151	203
Cr	74	64	66	49	78	86	67	57	51	28	48	55
Co	126	73	57	78	75	40	34	28	60	237	229	170
Ni	55	47	42	41	50	48	42	37	37	38	44	43
Cu	53	94	73	413	30	39	35	29	34	86	38	22
Zn	26	21	28	37	22	26	25	12	31	12	10	21
Ga	22	21	19	22	26	19	20	20	16	11	9	14
Ge	1.33	1.27	1.44	0.73	1.14	3.88	1.74	1.28	1.63	0.85	1.30	1.17
As	4.8	5.6	3.8	7.0	2.9	4.0	2.7	2.1	2.9	6.5	7.3	4.7
Se	1.97	1.98	2.37	1.53	0.94	2.06	1.57	1.85	1.58	7.33	9.05	2.94
Rb	100	101	101	72	133	143	122	93	94	57	81	76
Sr	34	28	25	62	21	43	31	28	37	33	15	19
Y	47	38	35	50	46	32	31	29	35	72	244	112
Zr	168	170	201	161	185	174	214	224	160	109	65	112
Nb	13.3	15.0	12.9	11.7	13.5	11.6	13.9	13.9	12.6	12.1	5.8	12.7
Mo	2.25	4.19	3.01	2.73	3.83	3.26	2.69	3.13	3.09	13.07	5.31	4.63
Ag	0.22	0.17	0.29	0.32	0.07	0.45	0.35	0.15	0.37	0.45	0.34	0.28
Sn	3.91	4.13	3.92	2.87	6.57	3.13	4.00	4.34	3.92	4.35	2.13	4.59
Sb	0.96	0.53	1.66	0.73	0.76	5.32	1.50	0.70	3.40	0.83	0.53	0.58
Te	19.20	3.06	4.79	5.04	1.03	8.41	5.94	2.03	4.81	13.28	25.15	28.14
Ba	300	361	362	201	503	337	547	301	348	208	160	294
La	38.9	36.5	43.1	28.5	40.5	32.4	47.5	51.7	47.7	178.1	43.8	57.7
Ce	78.2	76.7	94.6	65.3	81.8	69.3	97.1	102.3	94.3	309.7	92.6	112.3
Pr	9.6	9.2	11.1	8.1	10.0	8.3	11.1	11.8	11.0	38.6	11.6	13.3
Nd	38.6	36.6	41.7	33.1	40.0	31.8	41.3	44.1	41.6	146.0	50.9	54.1
Sm	8.76	8.16	8.96	8.36	8.86	6.60	7.98	8.59	9.39	32.26	18.39	13.72
Eu	2.63	2.28	2.34	2.47	2.55	1.56	1.71	1.95	2.48	9.10	8.38	4.77
Gd	9.96	8.61	8.47	10.10	9.69	6.33	7.10	7.71	8.54	27.95	37.66	19.57
Tb	1.62	1.38	1.33	1.77	1.55	0.99	1.05	1.14	1.28	3.74	8.32	3.72
Dy	10.09	8.45	7.99	11.80	9.54	5.94	6.20	6.58	7.38	19.69	57.51	24.91
Ho	1.98	1.68	1.52	2.39	1.86	1.19	1.21	1.26	1.40	3.45	11.73	4.97
Er	5.50	4.64	4.24	6.50	5.35	3.40	3.44	3.48	3.83	8.92	30.50	13.29
Tm	0.78	0.65	0.59	0.90	0.79	0.49	0.49	0.50	0.53	1.22	4.13	1.87

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Yb	4.87	3.98	3.76	5.50	4.97	3.11	3.14	3.12	3.26	7.42	23.81	11.04
Lu	0.69	0.59	0.54	0.77	0.74	0.46	0.48	0.48	0.48	0.93	2.83	1.40
Hf	4.57	4.48	5.43	4.46	5.14	4.72	5.79	5.94	4.29	2.83	1.81	3.02
Ta	0.85	0.86	0.97	0.70	0.97	0.95	1.10	1.08	0.90	0.82	0.33	0.70
W	135	169	211	25	172	25	92	147	70	227	715	615
Au g/t	1.03	0.37	0.19	0.23	0.12	0.03	0.16	0.25	0.22	3.46	6.14	3.28
Tl	0.69	0.24	2.45	0.24	0.31	9.78	2.59	0.77	5.10	1.31	0.18	0.20
Pb	2.27	2.28	2.40	3.03	1.23	2.11	2.12	1.87	2.23	5.34	4.52	3.21
Bi	11.71	1.60	1.11	1.23	0.32	2.66	1.34	0.72	2.01	11.55	21.88	24.08
Th	6.79	6.33	12.35	5.37	7.78	17.26	15.99	14.99	11.68	5.99	4.72	5.24
U	4.46	3.94	4.74	2.12	5.90	8.31	6.07	5.81	6.42	10.37	7.06	5.17
S wt%	0.15	0.37	0.22	0.29	N.D.	0.06	0.13	0.24	0.22	1.35	1.01	0.47
Y/Ho	23.88	22.91	22.82	20.78	24.96	26.69	25.30	22.75	24.98	20.89	20.79	22.50
Zr/Hf	36.84	38.08	36.98	36.08	35.92	36.93	36.90	37.79	37.20	38.59	35.83	37.06
Nb/Ta	15.69	17.38	13.30	16.79	13.93	12.22	12.57	12.93	13.99	14.63	17.44	18.18
Nb/Y	0.28	0.39	0.37	0.24	0.29	0.37	0.45	0.49	0.36	0.17	0.02	0.11
La/Yb	7.99	9.17	11.45	5.18	8.15	10.41	15.11	16.56	14.63	24.00	1.84	5.22
Molar Cu/Au	159	786	1198	5571	784	4051	680	357	475	77	19	21

#1-Alterd pyroxenite; #2-Microgabbro; #3-Gabbro; #4-Ferrogabbro; #5-Glimmerite; #6- Qtz Monzonite; #7-Andesite;.

#8-Andesitic Tuff; #9-Carbothermal Veins; #10-Carbothemal sheet; #11-Highway Breccia; #12-Quartz Breccia.

N.D.-not derermined.

Table 2. Major Element Analyses and Trace Element Data for Lithologies from the Duck Creek Area with Comparative data from Mount Cobalt and Mary Kathleen.

	Columbiad	**Comet**	**Dulce**	**Easter Gift .**	**Eva**	**Forget-Me-Not**	**Hideaway**	**Horseshoe**
No. Analyses	n=130	n=44	n=61	n=4	n=13	n=157	n=66	n=481

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SIO_2	62.96	64.16	59.16	69.09	63.87	62.90	63.20	62.27
TiO_2	0.90	1.48	0.74	0.68	0.43	0.64	1.24	0.93
Al_2O_3	10.33	10.30	12.55	10.94	6.96	10.84	9.42	10.56
Fe_2O_3	16.39	15.58	15.60	11.51	20.51	15.05	18.65	15.34
MnO	0.09	0.32	0.13	0.09	0.11	0.08	0.12	0.12
MgO	4.76	2.30	5.83	3.78	3.43	6.11	2.32	5.45
CaO	1.46	3.28	2.86	0.43	3.31	1.49	2.75	2.60
Na_2O	2.14	1.28	2.03	2.52	0.74	1.74	1.07	1.83
K_2O	0.79	1.16	1.00	0.84	0.57	1.05	1.13	0.78
P_2O_5	0.18	0.14	0.10	0.12	0.08	0.11	0.11	0.13
SUM	100	100	100	100	100	100	100	100
Fe/(Fe+Mg)	0.63	0.77	0.57	0.61	0.75	0.55	0.80	0.59
Na_2O+K_2O	2.93	2.44	3.03	3.36	1.30	2.79	2.20	2.61
Trace Elements ppm								
Li	10.8	10.6	17.6	13.1	10.1	21.0	10.1	16.3
Be	1.65	1.26	0.94	0.85	0.66	1.10	1.00	1.05
Sc	21	30	25	19	18	24	24	26
V	224	329	216	213	274	215	352	244
Cr	52	54	92	43	70	81	68	55
Co	208	97	159	100	204	114	59	164
Ni	64	46	93	45	108	85	102	67
Cu	704	2544	5789	11596	19414	4645	1727	6382
Cu wt.%	0.07	0.25	0.58	1.16	1.94	0.46	0.17	0.64
Zn	28	27	37	22	29	36	41	34
Ga	20	15	22	14	14	21	15	19
Rb	37	50	47	31	26	74	54	30
Sr	30	58	50	16	24	36	60	40
Y	23	29	19	36	14	15	23	20
Zr	103	120	103	126	35	68	102	85
Nb	7.0	8.2	6.5	4.4	1.8	4.4	8.1	5.5
Mo	1.99	0.64	2.96	0.45	1.38	1.56	1.11	2.89
Pd	2.92	3.24	7.05	2.83	20.93	6.42	7.42	8.89
Ag	0.15	0.34	0.35	0.33	0.98	0.44	0.90	0.40
Sn	1.81	1.84	1.43	1.80	2.46	1.71	1.71	1.74
Sb	26.3	0.63	0.37	0.40	0.50	0.43	2.39	0.38
Cs	1.45	1.52	1.46	0.79	1.05	2.76	1.28	1.23
Ba	176	283	239	118	179	217	191	177
La	19.5	20.2	19.9	13.2	6.2	10.7	16.6	14.2
Ce	39.2	46.2	41.7	27.3	13.0	22.1	36.0	30.1
Pr	4.6	5.2	5.0	3.6	1.8	2.7	4.3	3.7
Nd	18.3	21.5	19.6	15.6	7.7	11.1	17.4	15.3
Sm	4.02	5.00	4.05	4.41	1.95	2.56	4.08	3.50

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Eu	1.03	1.46	0.90	1.15	0.48	0.62	1.08	0.83
Gd	4.26	5.55	3.82	5.50	2.20	2.71	4.31	3.65
Tb	0.70	0.85	0.57	0.91	0.36	0.44	0.68	0.59
Dy	4.30	5.43	3.62	6.11	2.40	2.90	4.33	3.82
Ho	0.87	1.11	0.72	1.27	0.51	0.60	0.88	0.78
Er	2.55	3.57	2.13	4.15	1.64	1.78	2.66	2.29
Tm	0.36	0.52	0.31	0.63	0.24	0.27	0.38	0.34
Yb	2.30	3.47	2.04	4.25	1.59	1.65	2.45	2.09
Lu	0.34	0.56	0.31	0.66	0.25	0.25	0.37	0.33
Hf	3.03	3.18	2.83	3.42	0.95	1.98	2.77	2.55
Ta	0.53	0.60	0.48	0.34	0.12	0.32	0.59	0.42
W	3.31	12.78	1.91	4.38	5.30	3.52	1.69	3.83
Pt	3.66	3.44	6.14	1.40	12.68	6.46	3.94	7.41
Au ppm	0.02	0.12	0.04	0.04	2.36	0.07	0.07	0.06
Tl	0.16	0.20	0.20	0.10	0.12	0.29	0.24	0.13
Pb	4.12	4.24	2.16	19.33	8.24	1.85	40.20	2.02
Bi	0.85	1.00	1.19	0.41	3.85	1.33	4.19	0.68
Th	5.49	6.14	6.66	5.89	1.10	3.39	5.53	3.62
U	2.27	1.97	2.13	4.56	3.44	2.37	1.55	1.99
TREE ppm	102	121	105	89	40	60	96	81
logAg/Au	0.91	0.46	0.98	0.90	-0.38	0.78	1.12	0.82
Molar Cu/Au	122967	66077	495563	866068	25471	195163	78261	326323
Y/Ho	26.29	26.48	26.65	28.42	27.87	25.52	26.09	25.35
Zr/Hf	33.90	37.63	36.46	36.87	36.44	34.48	36.73	33.53
Nb/Ta	13.41	13.84	13.50	13.02	14.52	13.84	13.62	13.12
Nb/Y	0.31	0.28	0.34	0.12	0.13	0.29	0.35	0.28
Nb+Y	29.95	37.70	25.59	40.46	16.09	19.84	31.12	25.26

Table 2. Major Element Analyses and Trace Element Data for Lithologies from the Duck Creek Area with Comparative data from Mount Cobalt and Mary Kathleen (Continued).

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	Junction	Micawber	Mountain Maid	New Dollar	Ready Rhino	Success	Mount Cobalt	Mary Kathleen*
Elements wt%	n=52	n=67	n=216	n=168	n=145	n=196	n=3	n=640
SIO_2	55.62	60.80	64.04	61.79	61.31	65.01	46.33	46.74
TiO_2	0.94	0.64	1.06	0.86	0.86	1.18	1.51	4.78
Al_2O_3	14.01	12.44	9.94	10.80	11.61	10.08	13.45	0.18
Fe_2O_3	15.37	13.91	15.53	15.70	14.38	16.98	17.74	19.02
MnO	0.13	0.11	0.10	0.11	0.11	0.16	0.16	0.32
MgO	6.50	5.66	4.28	6.08	6.69	3.21	9.20	3.99
CaO	3.07	2.65	1.85	1.69	2.04	1.34	5.05	23.65
Na_2O	3.00	2.30	1.71	1.43	1.53	0.98	1.88	0.65
K_2O	1.20	1.38	1.29	1.43	1.33	0.93	4.55	0.26
P_2O_5	0.17	0.11	0.20	0.12	0.15	0.11	0.12	0.42
SUM	100	100	100	100	100	100	100	100
Fe/(Fe+Mg)	0.54	0.55	0.65	0.57	0.52	0.73	0.49	0.85
Na_2O+K_2O	4.21	3.68	3.00	2.86	2.86	1.92	4.68	0.90
Li	19.7	17.2	14.3	25.9	29.1	15.7	26.7	6.89
Be	1.48	1.10	1.63	1.18	1.25	1.31		1.71
Sc	32	26	23	29	31	27	52	3.34
V	250	201	246	263	265	329	573	79.71
Cr	75	63	54	74	66	66	113	9.18
Co	60	124	156	80	76	208	148	104
Ni	67	101	62	70	80	67	147	203
Cu	187	2197	2932	4796	2774	4820	184	2039
Cu wt.%	0.02	0.22	0.29	0.48	0.28	0.48	0.02	0.20
Zn	42	38	33	32	38	25	34	23.44
Ga	22	21	19	19	19	17	36	15.45
Rb	55	79	73	65	61	43	423	11.16
Sr	69	51	39	29	48	33	59	18.90
Y	25	18	25	18	22	23	32	38.78
Zr	125	97	117	75	85	118	84	46.29
Nb	8.2	6.0	8.1	4.4	5.2	7.2	7.7	9.70
Mo	1.51	1.45	2.07	1.29	1.33	0.92	1.50	1.83
Pd	9.13	7.24	2.79	7.07	8.22	4.44	1.80	
Ag	0.18	0.16	0.32	0.29	0.28	0.42	1.67	0.11
Sn	2.12	1.40	1.82	1.33	1.43	2.39	0.08	45.72
Sb	0.57	0.33	0.33	0.40	0.48	0.53	0.43	0.33
Cs	2.89	4.06	2.66	3.39	2.68	1.18	13.78	0.36
Ba	293	292	217	189	234	200	461	52.6

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La	34.1	17.5	20.4	13.6	14.2	22.9	152.6	159.5
Ce	66.6	36.4	43.1	27.4	29.4	46.0	270.4	303.3
Pr	7.7	4.4	5.2	3.3	3.7	5.5	30.9	29.5
Nd	29.7	17.4	20.8	13.7	15.4	21.9	119.4	101.24
Sm	5.85	3.71	4.58	3.18	3.67	4.83	25.89	16.39
Eu	1.63	0.82	1.16	0.82	0.94	1.35	2.65	3.49
Gd	5.38	3.68	4.80	3.44	3.97	4.78	19.38	11.69
Tb	0.79	0.55	0.79	0.55	0.63	0.73	2.22	1.48
Dy	4.83	3.47	4.79	3.50	4.07	4.51	9.58	7.28
Ho	0.95	0.68	0.97	0.71	0.82	0.89	1.44	1.28
Er	2.75	2.04	2.84	2.10	2.43	2.62	3.35	3.22
Tm	0.40	0.29	0.41	0.30	0.34	0.37	0.43	0.43
Yb	2.54	1.95	2.63	1.93	2.24	2.46	2.39	2.59
Lu	0.38	0.31	0.39	0.29	0.34	0.37	0.31	0.39
Hf	3.47	2.68	3.25	2.13	2.37	3.20	2.30	1.09
Ta	0.61	0.45	0.58	0.31	0.36	0.54	0.23	0.26
W	3.34	2.04	3.27	3.58	4.26	3.10	1.00	13.76
Pt	8.43	5.52	2.76	6.66	7.52	3.97	1.50	
Au ppm	0.00	0.03	0.03	0.03	0.02	0.15	0.01	0.04
Tl	0.23	0.28	0.25	0.19	0.23	0.18	0.92	0.05
Pb	1.92	1.96	2.19	2.42	2.13	4.26		9.80
Bi	0.25	0.58	3.94	1.26	0.93	2.16	0.06	2.81
Th	8.03	6.52	5.79	2.37	3.75	6.02	1.22	55.81
U	2.43	2.23	2.57	2.06	1.93	2.17	0.86	17.95
TREE ppm	164	93	113	75	82	119	641	642
logAg/Au	1.62	0.74	1.06	0.94	1.06	0.45	2.44	0.41
Molar	133087	230183	333822	438564	345827	99015	94882	
Cu/Au								143049
Y/Ho	26.19	26.69	26.36	25.83	26.45	25.89	22.03	30.39
Zr/Hf	36.14	36.17	36.06	35.38	36.09	36.87	36.52	42.38
Nb/Ta	13.30	13.52	14.01	14.08	14.38	13.38	32.86	37.20
Nb/Y	0.33	0.33	0.32	0.24	0.24	0.31	0.24	0.25
Nb+Y	33.09	24.30	33.57	22.69	26.90	30.21	39.47	48.48

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5. Results

Figure 11 shows the areal extent and Cu, Au, Pd, Pt, Ni, Co and Sc anomalism in different deposits listed in Table 2, from in Highway-Duck Creek area. As this elemental anomalism extends for a strike length of more than 18.4 km, it is clear that a very large mineral system lies west and at depth below the Highway epithermal domain.

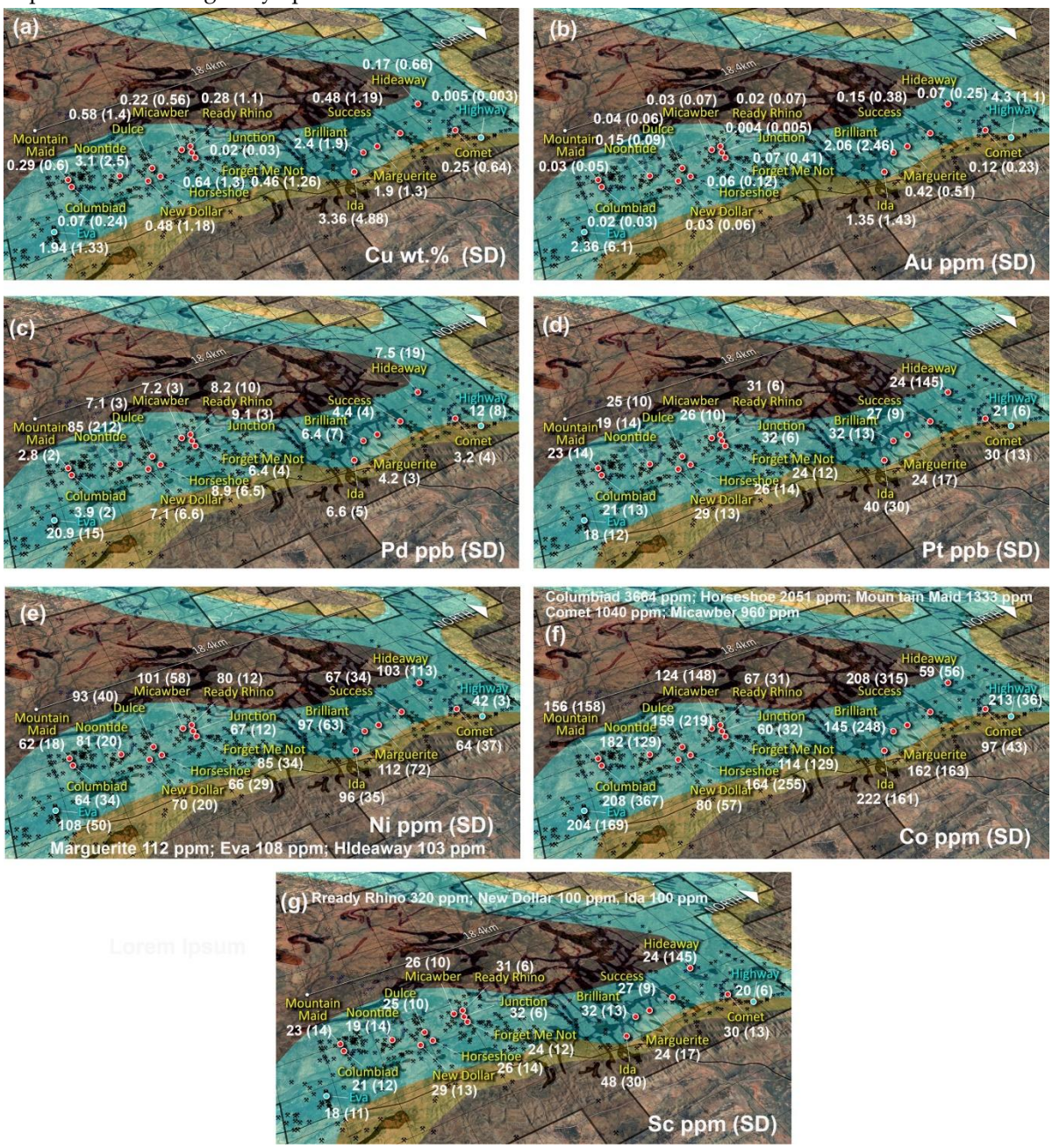

Figure 11. Maps showing the average and (SD) concentrations of (a) Cu wt.%, (b) Au (ppm), (c) Pd (ppb), (d) Pt (ppb), (e) Ni (ppm), (f) Co (ppm) and Sc (ppm) in selected prospects (red dots) along an 18.4 km strike length within the Duck Creek porphyry system (pale green ornamentation) and the overlying Highway epithermal system (pale yellow ornamentation). Shown by the small dark dots are the myriad of small historic supergene mineralised deposits.

The region is currently under active resource definition by Transition Resources. The Highway epithermal mineralisation exhibits significant concentrations of precious and critical metals (Table 1) e.g., Au intersections of up to 11m @ 9.58 g/t, 7m @ 9.02 g/t Au and 2m @ 41.2 g/t Au. The system is also heavy rare earth element rich with up to 157 ppm Dy_2O_3 and 23 ppm Tb_4O_7. Other key critical

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elements in the Highway system include WO_3 and Sc_2O_3 that range up to 1793 ppm and 64 ppm respectively.

The current global gold resource for the extreme northern segment of the Highway epithermal corridor Highway a depth of 125m is 1.2 Mt @ 3.1 g/t Au and 1600 ppm WO_3, plus significant HREEs, Co and Sc. Although limited resource drilling has been undertaken in the Duck Creek porphyry system, assays from a small segment of the exploration target indicate an inferred JORC resource of 5.44Mt @ 1.45% Cu, 0.11 g/t Au and 158 ppm Co. The current global resource for Duck Creek porphyry is estimated to contain 6.5 Mt @1.57%, Cu 0.12 g/t Au.

5.1. Metallogenic Affinity

5.1.1. K_2O-SiO_2 and Th-Co Discrimination Plots

CEI data for the Duck Creek deposits and the mean compositions of the Highway lithologies are plotted in the K_2O and SiO_2 discrimination projection of Peccerillo and Taylor [70] to understand the petrological affinity of the igneous source responsible for the formation of the Highway-Duck Creek mineral system. The Duck Creek data show a broad range of compositions extending from shoshonite to low K calc alkaline (Figure 12a).

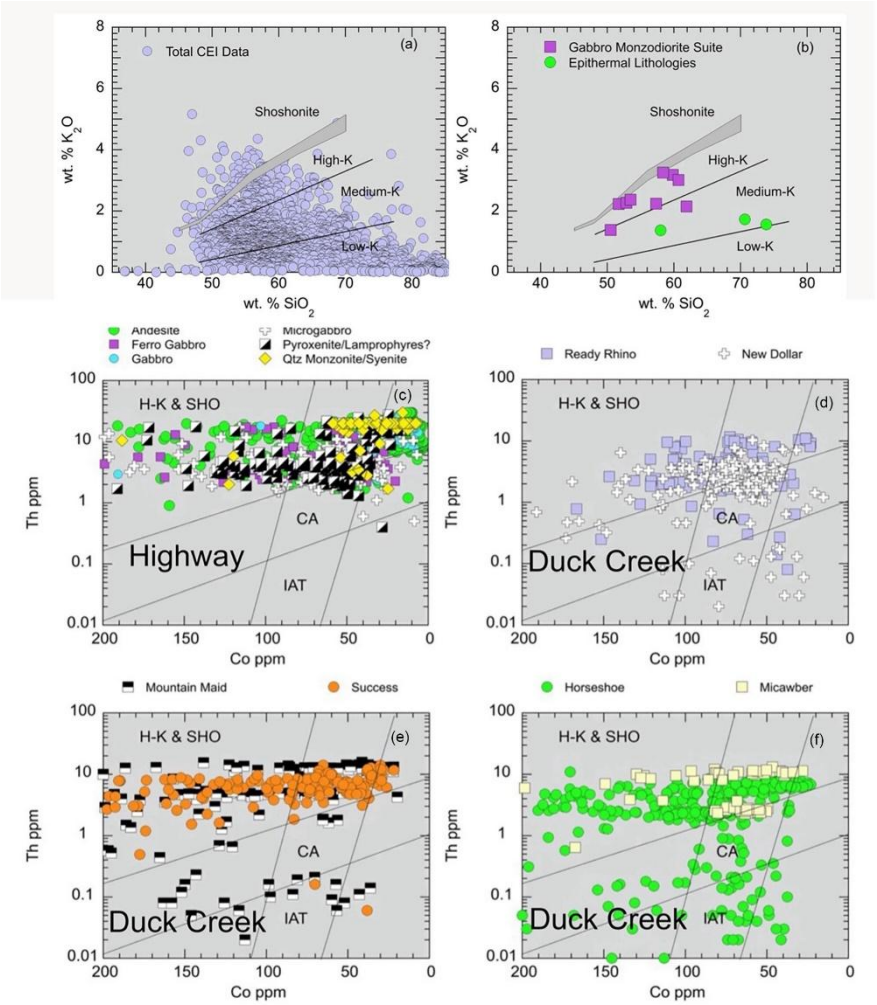

Figure 12. K_2O vs SiO_2 discrimination plot for calc alkaline, high potassium and shoshonite igneous suites after Peccerillo and Taylor [70] showing (a) data for Duck Creek deposits (CEI data set), and (b) average compositions of gabbro-monzodiorite suite lithologies and epithermal lithologies from the Highway epithermal system (Table 1).

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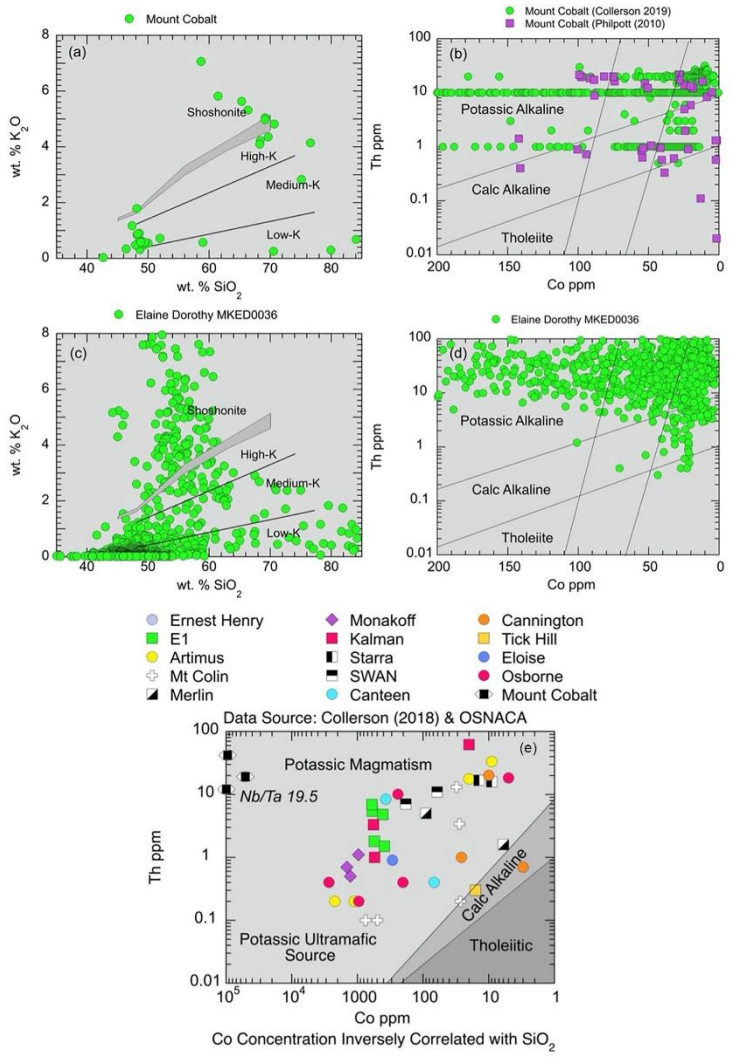

Figure 13. K_2O vs SiO_2 and Th vs Co plots comparing data for selected mineral deposits in the eastern NWMP. (a & b) Mount Cobalt, (c & d) Mary Kathleen-Elaine Dorothy and (e) deposits classified as "IOCG" systems in the eastern NWMP and southern NWMP. The majority of "IOCG" deposits plot within the potassic field, using Th vs Co systematics that extend from ultramafic Mt Cobalt to felsic compositions. Mount Cobalt has a plume-like Nb/Ta ratio of 19.5 . Data sources [42, 45, 69, 82].

The mean data for the different Highway lithology assemblages in Figure 12b, is more coherent and falls within the field from high K to medium K. The spectrum of variation shown by K_2O is interpreted to largely be caused by loss of K_2O during the phyllic and propylitic alteration event that permeated the system following crystallisation of the primary igneous lithologies [72]. In addition, as these lithologies show phyllic or propylitic alteration, it is likely that all have either lost or gained Si.

To circumvent the effect of loss or gain of K during post-crystallisation alteration, Hastie et al. [73] suggested using Th as a proxy for K and Co as a proxy for Si and proposed a discrimination projection that was more robust to alteration. Figures 12c to f shows Th and Co data for Highway and Duck Creek lithologies. They indicate that the dominant metal source at Highway and Duck Creek was an alkaline (high K) igneous suit[42,45,69,82e, although samples from some Duck Creek deposits, e.g., Ready Rhino, New Dollar, Mountain Maid, and Horseshoe, exhibit a secondary calc alkaline to tholeiitic petrologic vector.

These divergent trends are interpreted to indicate the presence two petrological/mineralising systems in the Duck Creek area. First, a subduction related porphyry mineralising event with a magmatic arc source, that formed during subduction and back arc extension. This was followed by a

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second metallogenic event, after collision and closure of an ocean basin east of the North Australian Craton. It is likely that a palaeo-slab tear was accreted to North Australian Craton lithosphere at this time thereby providing a window that facilitated the ingress of a plume magmas into the mantle wedge and subcontinental lithosphere. The post-tectonic tholeiitic to alkaline suite thus provided the metal source for formation of a later porphyry system with high-level epithermal carapace. In this regard, the identification of a potassic magmatic source has significant implications for the metal endowment, because some of the largest porphyry systems are hosted by K-rich magmas [74]. As continental collision zone porphyry deposits are typically hosted by potassic and ultrapotassic magmas [7,54,75] evidence for presence of such a magmatic source for the Highway Duck Creek system provides additional support for the continental collision zone porphyry model.

Identification of a potassic magmatic source for the Highway - Duck Creek mineral system also has important regional implications. For example, this potassic signature is also seen at Mount Cobalt (Figures 13 a & b) and at the Elaine Dorothy deposit near Mary Kathleen (Figures 13 c & d). This signature is also prominent in many of the "IOCG" deposits in the Cloncurry area (Figure 13e).

5.1.2. Multi-element Geochemical Variation

Primitive mantle-normalized multi-element plots, where elements are arranged in order of increasing compatibility are an effective way to compare geochemical data for different protoliths. The primitive mantle values used for element normalization are either from McDonough and Sun [76] or Lyubetskaya and Korenaga [77]. Analyses of the Duck Creek deposits and the different Highway lithologies are compared in Figures 14a and b. The two systems exhibit broadly uniform shaped patterns, with positive spikes in W, U, Pb, P and Li and negative spikes in Ba, Th, Nb, Ta and Sr. The pronounced depletion in Sr is a common feature seen at both Duck Creek and at Highway. It could reflect the fact that early fractionation of a Ca-Sr component (possibly a carbonatite) has occurred during evolution of the source magma. Alternatively, the widespread Sr depletion seen at Highway and Duck Creek could be alteration induced [78–80]. Highway also shows a pronounced negative Ti spike, which is absent at Duck Creek suggests the fractionation of a Ti-bearing mineral into the gabbro – monzogabbro intrusion that underlies the Mitakoodi anticline.

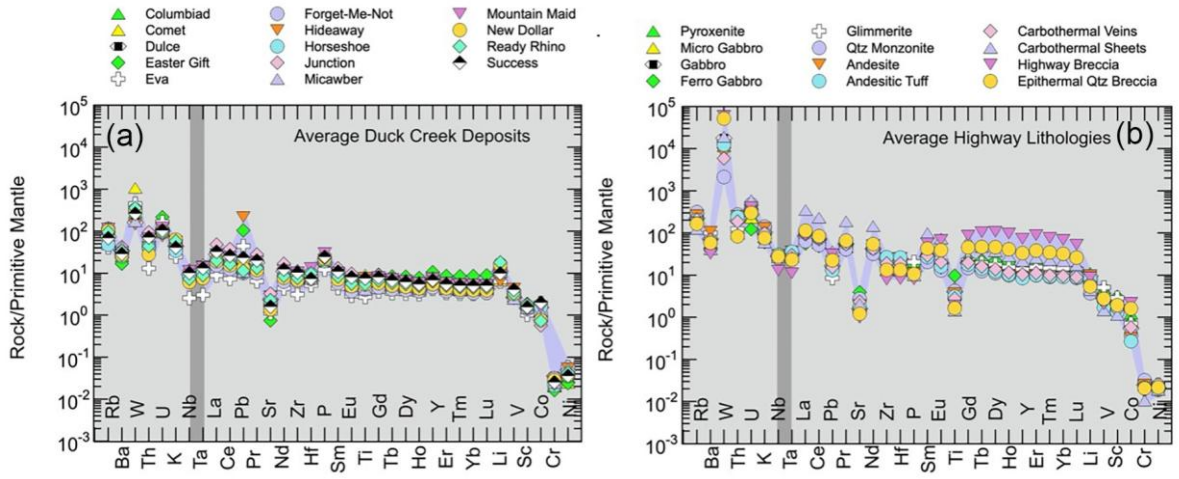

Figure 14. Primitive mantle-normalised multi-element variation diagrams showing (a) average compositions for a selection of Duck Creek deposits and showing (b) average compositions of Highway lithologies. The patterns are remarkably similar, which supports the interpretation that they are genetically related systems. Furthermore, lithologies from both systems show distinctive negative spikes in Sr, This, is interpreted to have been caused by Sr loss from the entire mineral system during pervasive hydrothermal alteration.

In addition to the significant level of W enrichment seen in all Highway lithologies, they are also significantly enriched (50 to 100x PM levels) in the rare earth elements (REEs). These include the high value heavy rare earth elements, Dysprosium (Dy) and Terbium (Tb) that are hosted by xenotime

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and monazite. For example, average compositions are 8.32 ppm Tb and 57.51 ppm Dy breccia (n= 15), 3.74 ppm Tb and 19.69 ppm Dy carbothermal (n=435) l and 3.72 ppm Tb and 24.91 ppm Dy epithermal quartz breccia (n=163). These lithologies are also significantly enriched in Au, averaging 6.14 g/t (breccia), 3.46g/t (carbothermal sheets) and 3.28 g/t (quartz breccia).

Although, the level of W enrichment in all Highway lithologies is significantly greater than at Duck Creek, the overall similarity in pattern, with relatively consistent immobile-incompatible element ratios, results in parallel profiles at Duck Creek and Highway. This supports the model that they are related systems and that the metals in the Highway epithermal system were delivered by fluids derived from the deeper Duck Creek system.

5.1.3. Multi-element Geochemical Variation

To discriminate between different mineral deposit styles, with specific reference to the IOCG class of deposit, Brauhart and Groves [34] proposed the use of multielement plots normalised to average crustal abundance [81]. This projection highlights the granitophile La and U enrichments seen in both IOCG and some porphyry systems (Figure 15).

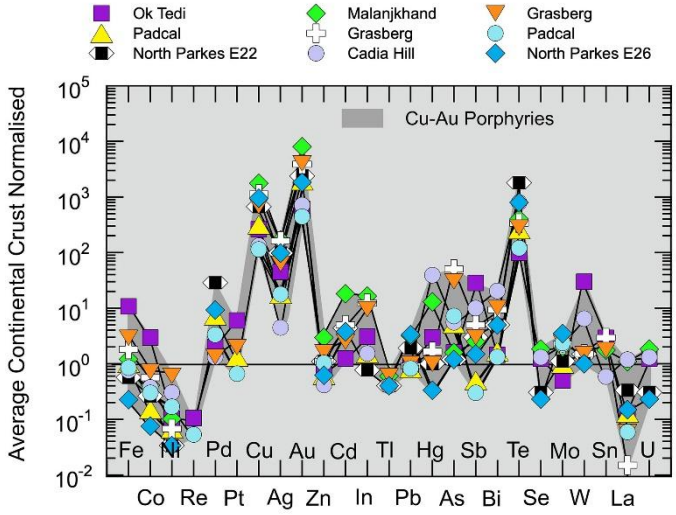

Figure 15. Ore and pathfinder element plots showing data for Cu-Au porphyries from the OSNACA database [82] normalised to average crustal abundances from Rudnick and Gao [81].

The projection also highlights the enrichment in 'ultramafic' suite of elements, such as the transition elements (Co-Ni) and the platinum group metals (Pd-Pt). In these plots, the presence of spikes in these elements in a mineral system geochemical profile, indicates the role of a mantle continental collision zone porphyry system rather than a crustal component in the metallogenic process.

The principal focus of the Brauhart and Groves [34] publication was to compare chemical variability of Australian intracratonic deposits, classified as IOCG and ISCG systems. However, in some locations, e.g., Chile [32] and the Gawler [31] Cu-Au porphyry deposits occur close to IOCG systems. Cu-Au porphyry deposits also occur in cratonic settings [7] especially, in collisional cratonic environments, like the eastern margin of the North Australian Craton [52]. Thus, to provide, a more representative platform for comparing between IOCG and Cu-Au porphyry systems, crustal abundance multi-element normalised plots for representative Cu-Au porphyry deposits, using data from the OSNACA database [82] are shown, in Figure 15. Porphyry systems exhibit a downward trend from Fe to Ni, but they also show positive spikes in Pd, Cu, Au, Te and W.

Crust normalised data for the Highway mineral system lithologies are shown in Figures 16 a & b. The Highway patterns are remarkably similar, to the Cu-Au porphyry field with prominent spikes in Au, Te and W. It is also enriched in Co, La and U, relative to Cu-Au porphyries that are hosted by K-rich igneous lithologies. The depletion at Highway reflects the fact that Cu has largely been

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sequestered by sulphide precipitation into the underlying Duck Creek porphyry system. Such depletion in Cu is a typical feature of epithermal mineralisation above porphyry systems [83].

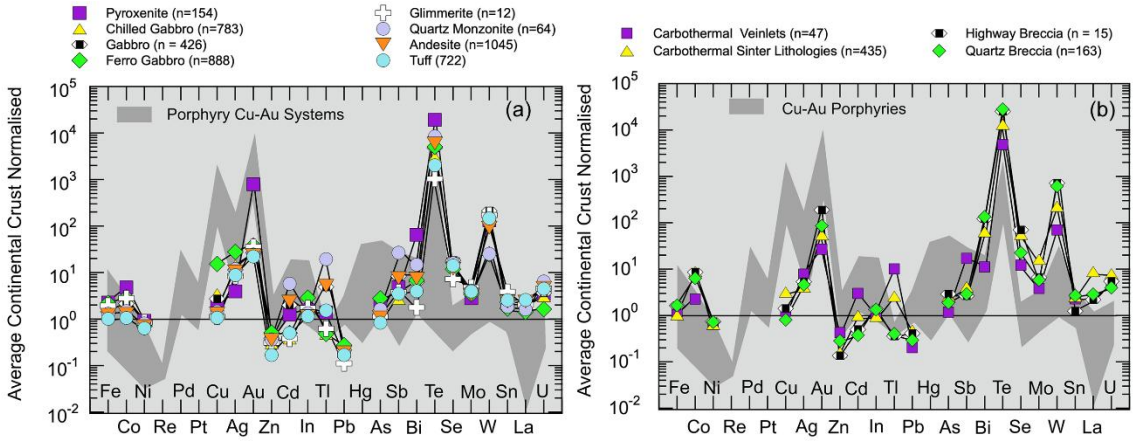

Figure 16. Ore and pathfinder element plots normalised to average crustal abundance showing data for average Highway lithologies (n=numbers of analyses). The Highway patterns are remarkably similar, to the Cu-Au porphyry field, except that Highway epithermal field is Cu depleted, because Cu has largely been sequestered into the underlying Duck Creek porphyry system. Such depletion in Cu is a typical feature of epithermal mineralisation above porphyry systems. The Highway system also exhibits prominent spikes in Te and W. It is also enriched in Co, La and U, relative to Cu-Au porphyries that are hosted by K-rich igneous lithologies.

Crustal abundance multi-element normalised plots for several Duck Creek deposits, i.e., Success, Mountain Maid, Horseshoe, Dulce, Eva, Furnace, Glory Hole, Comet and Daisey are shown in Figure 17.

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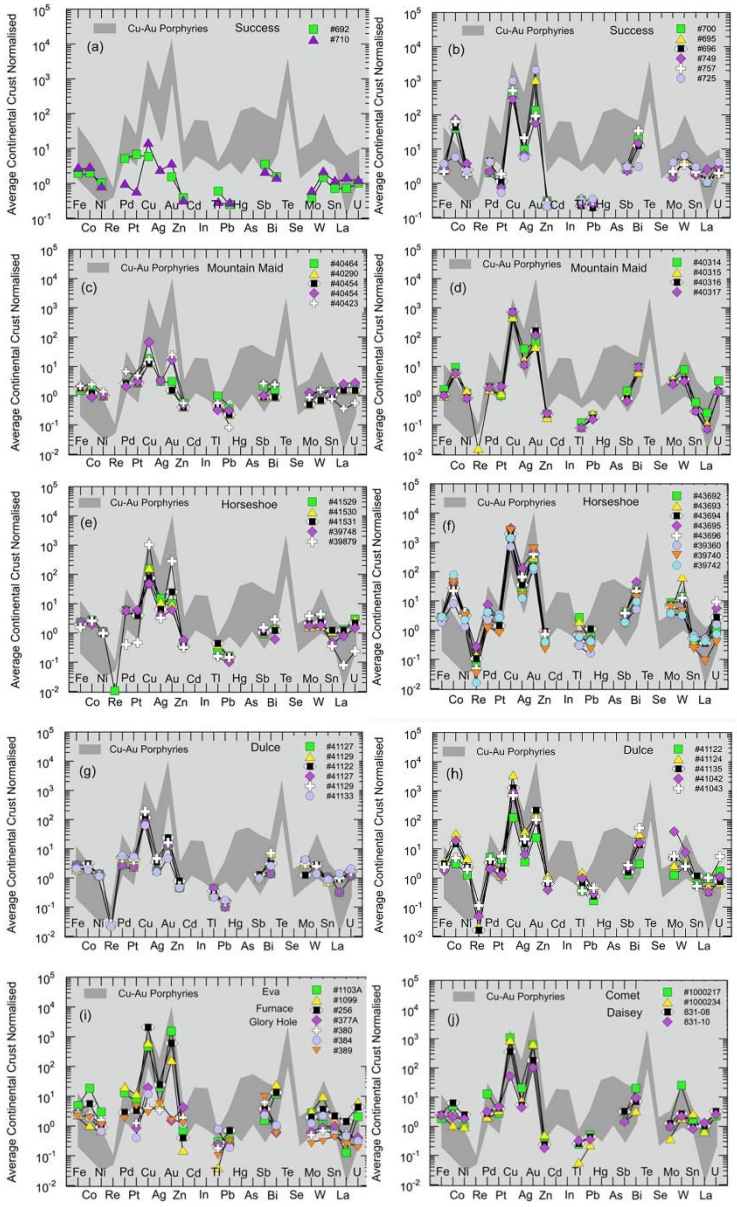

Figure 17. Ore and pathfinder element plots normalised to average crustal abundance for several the Duck Creek deposits, Success, Mountain Maid, Horseshoe, Dulce, Eva, Furnace, Glory Hole, Comet and Daisy. Success (a & b), Mountain Maid (c & d), Horseshoe (e & f) and Dulce (g & h) have enrichment patterns similar, to the Cu-Au porphyry field. A subgroup of Co-rich samples indicates the possible involvement of two metal sources a subduction generated component similar to the porphyry field and an ultramafic to mafic component most likely derived from the post-tectonic tholeiitic to alkalic plume system.

Samples exhibit topologies that either resemble porphyries with a downward trend from Fe to Ni and the positive spikes in Cu, Au, Te and W, or they resemble the Highway sub-pattern with positive spikes in Co as well as in Cu, Au, Te and W. For example, samples from Success (a & b), Mountain Maid (c & d), Horseshoe (e & f), Dulce (g & h) as well as at Eva, Furnace, Glory Hole, Comet and Daisy (i & j). The presence of these different patterns suggests the possible involvement of two metal sources in the metallogenic evolution of the Duck Creek system, viz., a subduction generated component, generating patterns like the porphyry field, and an ultramafic to mafic component most likely associated with the post-tectonic tholeiitic to alkalic plume component.

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5.1.4. Metal Source Constraints from Te Cu Ni Systematics

Post- subduction geodynamic settings are characterised by significant mantle-to-crust fluxes of metals and volatiles. Te is a tracer of this flux, and of the magmatic and hydrothermal processes that drive the flux. Post-subduction magmato-hydrothermal systems are Te enriched, and it exhibits both incompatible and chalcophile behaviour. At magmatic temperatures (> 1000 °C), Te has chalcophile behaviour like Au, Cu and the PGEs. Thus, at magmatic temperatures, Te enters sulphide melts, which fractionate and forming Pt–Pd-telluride melts at ~900 °C and then Pt–Pd-telluride minerals at <400 °C [84]. However, under hydrothermal conditions, at ~300 °C Te is mobilised as chloride complexes [85], as polytellurides in S- and CO_2-rich fluids [86] and as telluride–bismuthide melts [87]. Thus, Te is an ideal tracer of both magmatic and hydrothermal processes providing important information regarding mantle-to-crust fluxes of metals showing spatial and genetic links between deep magmatic Ni–Cu–PGE–Au–Te mineral systems and shallower, alkali-enriched porphyry–epithermal Cu–Au-(PGE-Te) deposits [83].

As a result of this behaviour, Holwell et al., [83] demonstrated that Ni, Cu and Te systematics are effective vectors to resolve petrogenetic process and metallogenic sources. Figures 18 & 19 show Cu, Ni and Te systematics for representative samples from Highway and from several Duck Creek deposits. These figures also show compositional fields of porphyry, epithermal, carbonatite, magmatic Cu-Ni-PGE and IOCG mineralisation.

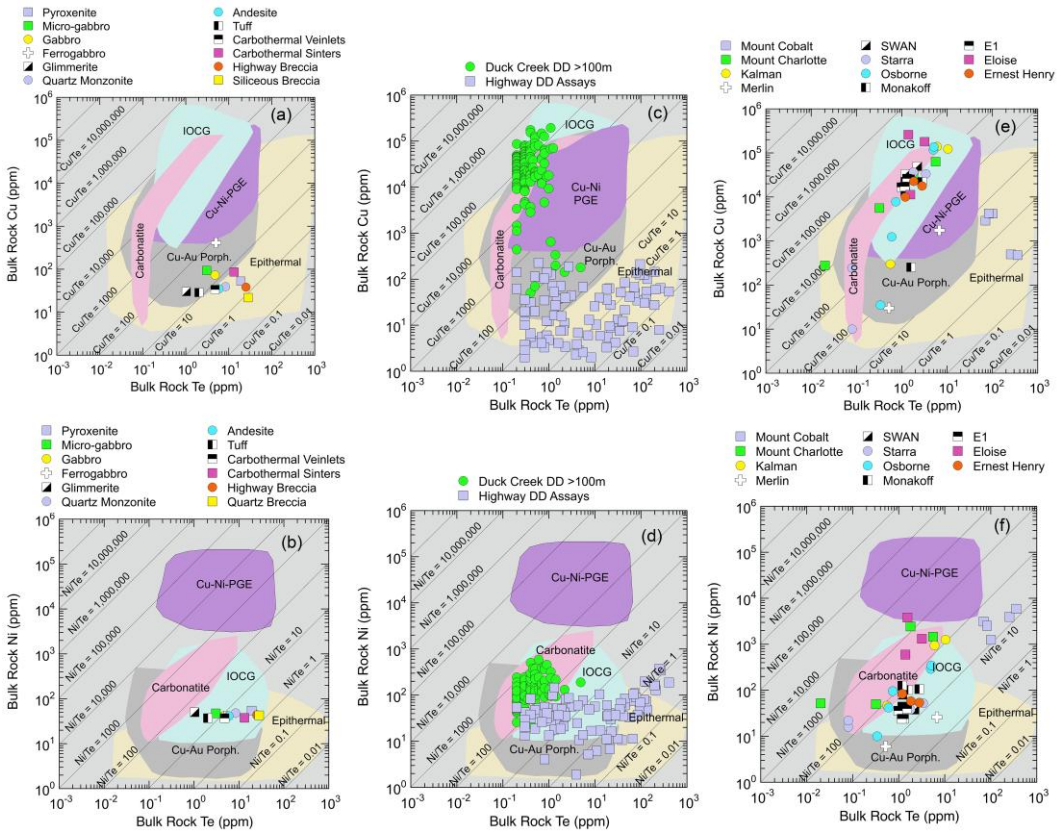

Figure 18. Cu vs Te and Ni vs Te covariation showing that the Highway samples fall in the epithermal field. Comparative fields are from [82]. The Cu-Au porphyry field has been enhanced with data for the Skouries Cu-Au porphyry [121].

Figure 18 a & b shows the compositional variability of individual Highway lithologies in Cu-Te and Ni-Te space. They lie within the epithermal field and in Cu-Te space within the Cu-Au porphyry field. Analyses of diamond drill hole samples from Duck Creek at a depth of >100m (Figures 18 c & d) show an even more prominent compositional difference between the volatile dominated Highway system and sulphide dominated Duck Creek systems. Figure 18 e & f show data for NWMP deposits that are classified as IOCG systems [82].

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The projection also highlights the enrichment in 'ultramafic' suite of elements, such as the transition elements (Co-Ni) and the platinum group metals (Pd-Pt). In these plots, the presence of spikes in these elements in a mineral system geochemical profile, indicates the role of a mantle-derived rather than a crustal component in the metallogenic process.

The principal focus of the Brauhart and Groves [34] publication was to compare chemical variability of Australian intracratonic deposits, classified as IOCG and ISCG systems. However, in some locations, e.g., Chile [32] and the Gawler [31] Cu-Au porphyry deposits occur close to IOCG systems. Cu-Au porphyry deposits also occur in cratonic settings [7] especially, in collisional cratonic environments, like the eastern margin of the North Australian Craton [52]. Thus, to provide, a more representative platform for comparing between IOCG and Cu-Au porphyry systems, crustal abundance multi-element normalised plots for representative Cu-Au porphyry deposits, using data from the OSNACA database [82] are shown, is shown in Figure 15. Porphyry systems exhibit a downward trend from Fe to Ni. But they also show positive spikes in Pd, Cu, Au, Te and W.

Covariation between Ni/Te and Cu/Te (Figure 19) showing a clear discrimination between epithermal and porphyry systems, with the Te-rich Highway lithologies plotting in the epithermal field along a Cu/Ni ratio vector of ~1 and Success, Hideaway, Chinaman and Eva showing the effect of Cu enrichment and plotting within the porphyry field.

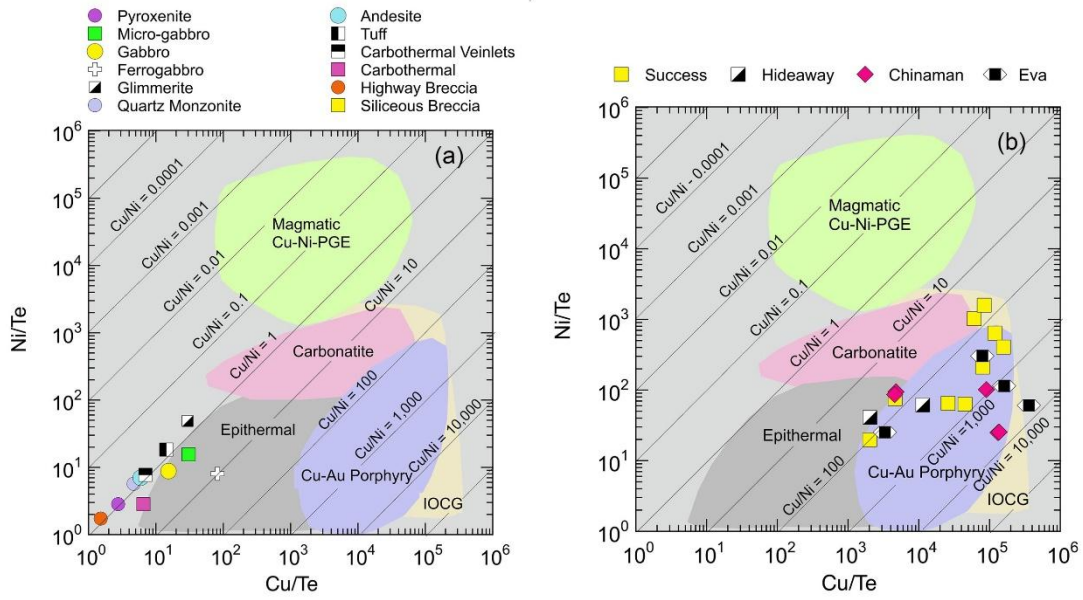

Figure 19. Plot of Ni/Te and Cu/Te showing that the Highway samples fall in the epithermal field. Comparative fields are from the Brauhart et al., (2017). The Cu-Au porphyry field has been enhanced with data for the Skouries Cu-Au porphyry [121].

5.1.5. Chalcophile Element and Highly Siderophile Element Systematics

Chalcophile element fertility or the chalcophile metal abundance and the availability of sulphur during source magma evolution, is a critical factor required to enable the formation of porphyry Cu ± Au deposits. Some mineralized porphyry systems are known to contain hydrothermal Pd and Pt [14], showing that hydrothermal fluids can mobilize Pd and Pt and in rare cases, porphyry Cu deposits can contain up to 2400 ppb Pd (e.g., Skouries Cu porphyries, Greece; [13]). Park et al., [88] have shown that the platinum group elements (PGEs) can be used as a tracer for chalcophile element fertility during magma evolution

Recent studies on PG) geochemistry of subvolcanic rocks associated with porphyry Cu ± Au deposits [89–91] have suggested that the relative timing of fluid and sulfide saturation plays an important role in the formation of porphyry Cu ± Au deposits. The PGEs are used in preference to Cu and Au for two reasons. First, their partition coefficients in immiscible sulfide melts are one to two orders of magnitude higher than those of Cu and Au respectively [92]. Second, because they are

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appreciably less affected by alteration than Cu and Au, they are more likely to preserve original igneous geochemistry [93,94]. The relative abundances of Cu and Pd porphyry systems are shown in Figures 20 and 21. Figure 20 shows covariation between Cu against Pd in several major porphyries Cu and Au-Cu deposits. Although Cu varies by several orders-of-magnitude in Bingham Canyon and Cadia Hill it does not correlate with increase in Pd as might be expected if Pd was an important component in the hydrothermal fluid. This indicates that the abundance of Pd is controlled by the chemistry of the porphyry host magma.

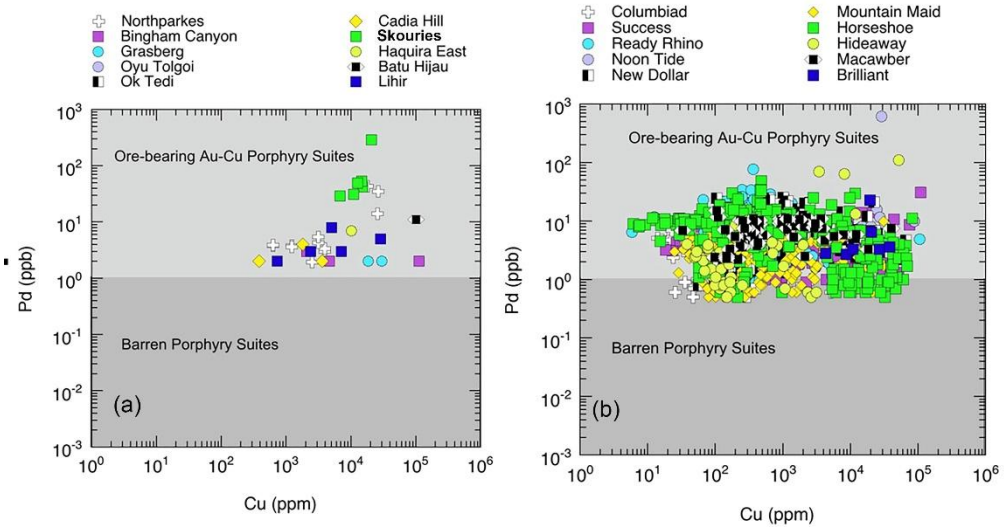

Figure 20. Covariation between Cu against Pd in (a) several major porphyries Cu and Au-Cu deposits and (b) deposits from the Duck Creek area. The similarity in Pd – Cu systematics is striking showing that the Duck Creek system is Pd-rich with compositions that resemble the Skouries Cu-Au porphyry.

Samples from the Duck Creek system all plot with the Cu-Au porphyry field In Figure 20b. Many of these are significantly enriched in Pd, in fact, several have similar Pd contents to the PGE rich Skouries Cu-Au (Pt, Pd, Te) porphyry [87].

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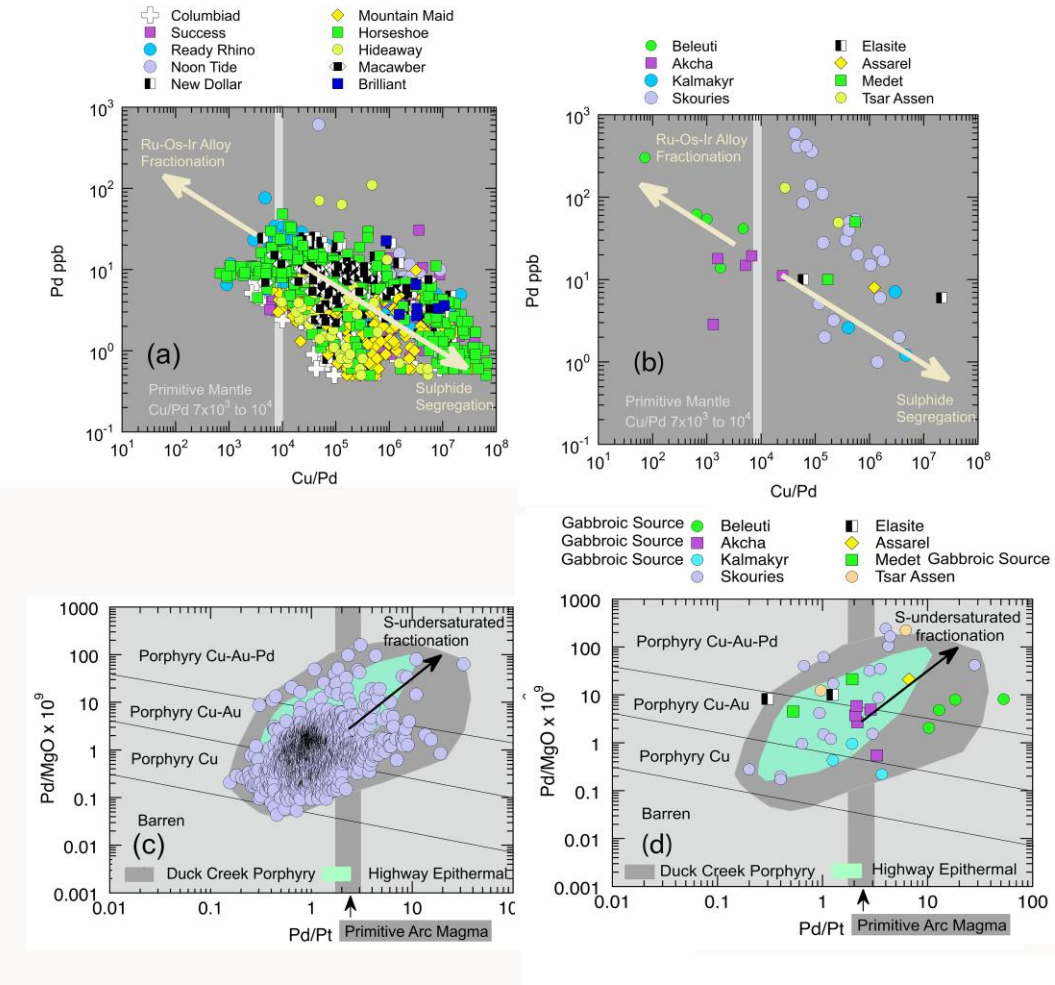

Figure 21. (a & b) Pd and Pd/Cu systematics showing that the Duck Creek Porphyry system lies on a similar sulphide segregation fractionation vector to that exhibited by Skouries and by several other gabbro-hosted porphyry systems e.g., Beleuti, Akcha, Kalmakyr, Elasite, Assarel, Medet and Tsar Assen [95], (c & d) show that Highway, Duck Creek and the gabbro-hosted porphyry systems (Figure 21d) define similar trends indicating metallogenesis involved S undersaturated fractionation.

This is also shown in Figure 21, which compares Pd with Cu/Pd and Pd/MgO with Pd/Pt ratios. Figures 21a & b shows that the Duck Creek Porphyry system lies on a sulphide segregation fractionation vector that is identical to the Pd vs Cu/Pd systematics exhibited by Skouries and by several other gabbro-hosted porphyry systems e.g., Beleuti, Akcha, Kalmakyr, Elasite, Assarel, Medet and Tsar Assen [96] that all occur in collisional orogenic environments. Importantly, the Highway and Duck Creek lithologies (Figure 21c) and the gabbro-hosted porphyry systems (Figure 21d) all define a trend indicating that metallogenesis involved S undersaturated fractionation.

The remarkable similarity between Cu, MgO and Pd systematics of the Duck Creek deposits with porphyry mineral systems such as Cadia, Bingham Canyon, Grasberg and Northparkes, Beleuti, Akcha, Kalmakyr, Elasite, Assarel, Medet and Tsar Assen provides vector support for the model that the Duck Creek area is likely underlain by a significant porphyry Cu-Au mineral system. . The similarity between the Duck Creek system and PGE-Cu systematics of other mafic-hosted porphyries is remarkable and has a significant impact on confirming the validity of the mineral system interpretation.

5.1.6. Deducing Redox Conditions using Scandium and Vanadium Systematics

In low oxygen fugacity magmas, V^{3+} behaves like Sc^{3+} and substitutes for Fe in amphibole, pyroxene. Thus, the V/Sc ratios of reduced magmas, remains constant at around 7 [96]. However,

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in hydrous oxidized porphyry magmas, V behaves as an incompatible element and is enriched in these melts. Scandium on the other hand substitutes into clinopyroxene and hornblende [97]. Hence, in oxidized and hydrous melts, crystallization of hornblende and clinopyroxene lowers Sc in the melt, causing an increase in V/Sc ratios. Although hornblende fractionation is commonly interpreted to play a role in V/Sc fractionation trends [97,98], clinopyroxene has a higher partition coefficient for Sc [99] and thus formation of clinopyroxene-rich cumulate lithologies would be expected to be Sc-rich and be characterised by low V/Sc ratios. The pyroxene-bearing lithologies at Mount Cobalt, with 52 ppm Sc and 575 ppm V and a V/Sc ratio of 11 confirm this interpretation. Thus, Sc and V systematics in the Highway-Duck Creek mineral system provide useful information regarding the redox history of the system where fertile porphyry Cu magmas show a trend of increasing V/Sc with increasing SiO_2 [98].

According to Loucks [98] the effect of increasing water content in a melt promotes the crystallization of hornblende well ahead of magnetite. The Sc concentration in porphyry Cu-forming magmas typically range between 3 and 8 ppm yielding V/Sc ratios of 10 to >15. V-Sc data for Duck Creek and Highway are shown in Figures 22a, c & d. These show a fractionation trend typical of reduced mafic and ultramafic magmas caused by olivine fractionation in the ferrogabbroic intrusion that hosts the Duck Creek mineral system. However, at low Sc concentrations, Duck Creek lithologies, and to some extent Highway lithologies both develop a trend to very high V/Sc ratios, like those seen in Cu-Au porphyries (Figure 22e). Thus, the Duck Creek mineral system has clearly evolved from highly reduced source magmas to strongly oxidised porphyry magmas, similar to Grasberg, Northparkes and Cadia. Importantly, several of the so-called IOCG deposits in the eastern Mount Isa Block (Figure 22f), including Elaine Dorothy at Mary Kathleen (Figure 22b, also define a mixing-trends that indicate evolution from reduced primary source magmas to strongly oxidised evolved magmas.

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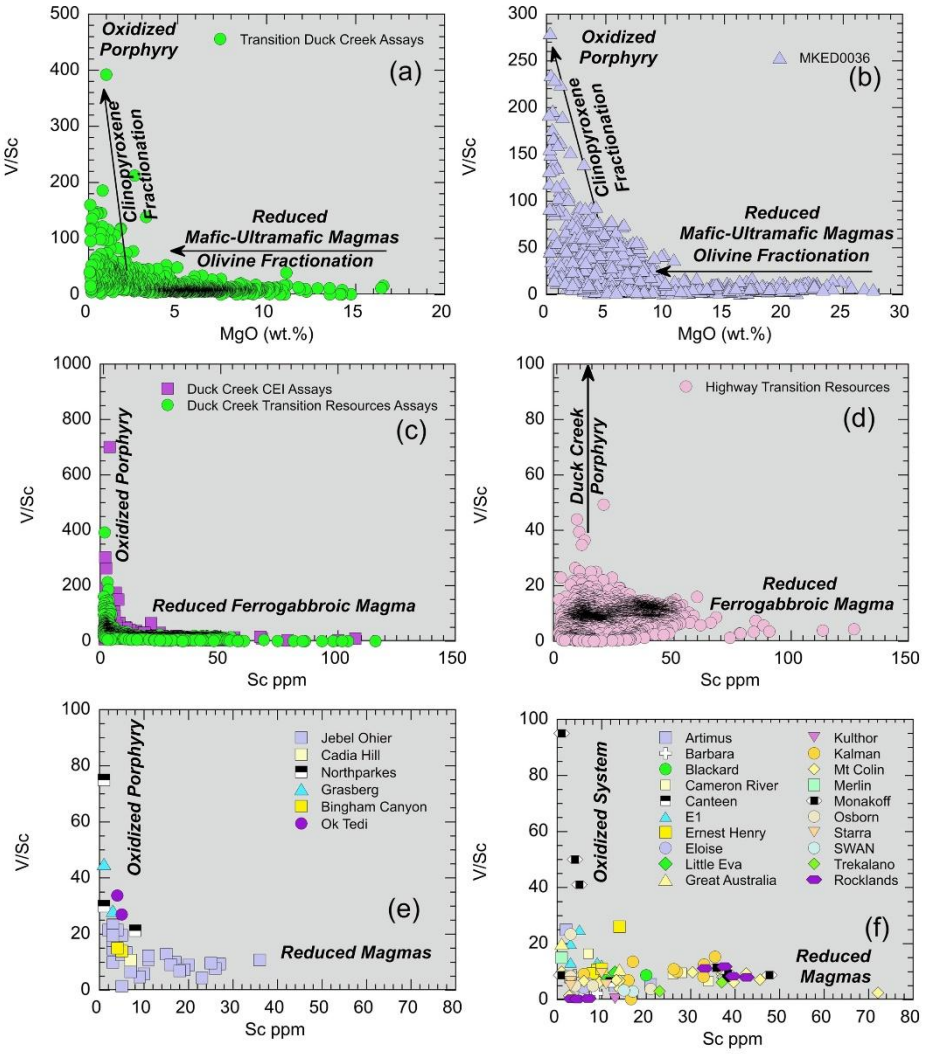

Figure 22. (a,c &d) V-Sc systematics of Duck Creek and Highway showing a fractionation trend typical of reduced mafic and ultramafic magmas caused by olivine fractionation in the ferrogabbroic source intrusion that hosts the Duck Creek mineral system. At low Sc concentrations, Duck Creek lithologies, and to some extent Highway lithologies evolve to high V/Sc ratios, (e) shows V-Sc systematics in Cu-Au porphyries. (f) and (b) V-Sc systematics of so-called IOCG deposits in the eastern Mount Isa Block including Elaine Dorothy.

5.1.7. Trace Element Vectors Indicating Tectonic Setting and Role of Sulphate Fractionation

Multi element ratio Nb/Y versus Nb+Y plots for samples from the Highway-Duck Creek mineral system are shown in Figure 23a, b & c. Also shown in Figure 23c are data from Elaine Dorothy. Figure 23d shows data for Jebel Ophier (Saudi Arabia), Bingham Canyon, Porgera and Northparkes Cu-Au porphyry deposits and Figure 23e shows data for several of the so-called IOCG deposits in the eastern Mount Isa Block. This plot, initially developed by Whelan and Hildebrand [100], allows discrimination between tectonic settings, i.e., Slab Tear, Magmatic Arc and Mantle Plume. In Figures 23a -c samples from Highway and Duck Creek mineral systems plot in the slab tear and magmatic arc fields (Slab Window) with a subtle vector towards the mantle plume field. Both Highway and Duck Creek systems define a vector that extend into the right-hand bottom quadrant of the diagram with very low Nb/Y and high Nb+Y values caused by enrichment in Y caused by sulphate – fluid fractionation that involved REE-SO_2- complexing [101,102].

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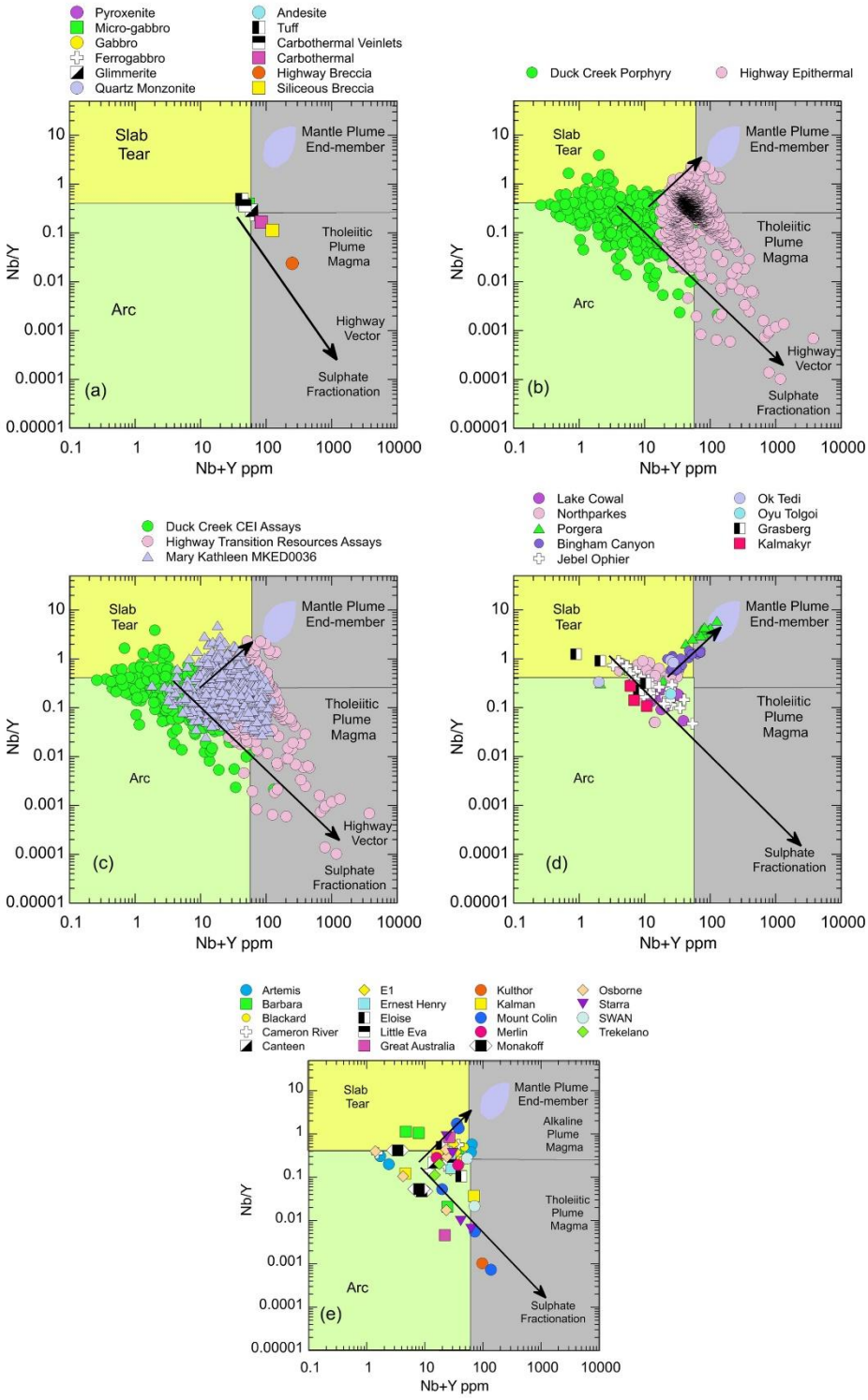

Figure 23. Multi element plots showing Nb/Y versus Nb+Y systematics showing (a) Highway lithologies, (b) Highway and Duck Creek data, (c) comparison between Highway-Duck Creek and Elaine Dorothy deep MKED0036 core, (d) comparative data for selected Cu-Au porphyries, (e) Cloncurry "IOCG" systems. All show the impact of sulphate fractionation leading to an increase in Y (HREE) concentrations.

The Highway-Duck Creek deposits are interpreted to have evolved via magmatic–hydrothermal processes from a deep porphyry to a shallow epithermal environment. The system exhibits a definitive phyllic, propylitic and potassic alteration halo. The metallogenic process initially involved crystallisation of reduced potassic ultramafic to intermediate (pyroxenites-gabbros-monzogabbros) magmas that evolved to oxidized magmas saturated with metal-rich aqueous fluids. The final stage of the mineralization event was controlled by sulphate reduction that was likely

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initiated by magnetite crystallization. This caused a decrease in pH and also in the oxidation potential of sulphate, resulting in substantial fractionation of Y from Nb and leading to the enrichment in the HREEs seen at Highway (Figure 23). In this regard it is significant that the most REE enriched Highway samples have the highest S contents (Table 1).

5.1.8. Variation in Molar Cu/Au with Depth

Molar Cu/Au ratios and bulk metal contents, specifically Au grades in porphyry deposits are interpreted to be controlled by magmatic chemistries [103–105] and by fluid phases separation into brine and vapor resulting in Cu-Au fractionation into a vapor phase and possibly, to partial separation of the two metals [103–107]. However, ore grades are ultimately controlled by precipitation efficiency of Cu-sulphides and native Au during cooling [108].

A positive correlation between Cu/Au ratio with depth has been demonstrated [109–111]. However, magma source is not the sole factor determining the bulk metal ratio of the deposits, and according to Murakami et al., [112] the systematic variation of molar Au/Cu ratios of porphyry-style ore deposits with depth is related to the density evolution of cooling magmatic hydrothermal fluids. This is because fluid pressure controls the extent of fluid phase separation into brine and vapor, causing fractionation of ore-forming components. In addition, fluid density evolution together with temperature also affects the differential solubility and selective precipitation of ore minerals. Thus, molar Cu/Au ratios are controlled initially by magma source chemistries and subsequently by depth where the physical-chemical evolution of the ore forming hydrothermal fluids occurred.

Figure 24 shows that molar Cu/Au ratios for samples, obtained by reverse circulation drilling, versus depth exhibit an increase in molar Cu/Au ratios from ~10 to 100,000 in the Highway area (east), through "Comway", to between 200,000 and 10,000,000 in Duck Creek (west) system. This supports the exploration model that the Highway system is a shallow Au dominated epithermal system and Duck Creek is a deeper Cu-Au bearing porphyry. The "Comway" deposits lie between Highway and Duck Creek at an intermediate crustal depth, with molar Cu/Au ratios between 60,000 and ~120,000, typical of alkaline porphyry systems, to a deeper Cu dominated porphyry systems at Duck Creek with ratios up to ~2,000,000. Note a near surface sample from Success with a molar Cu/Au of ~80,000,000 is Cu enriched due to supergene alteration.

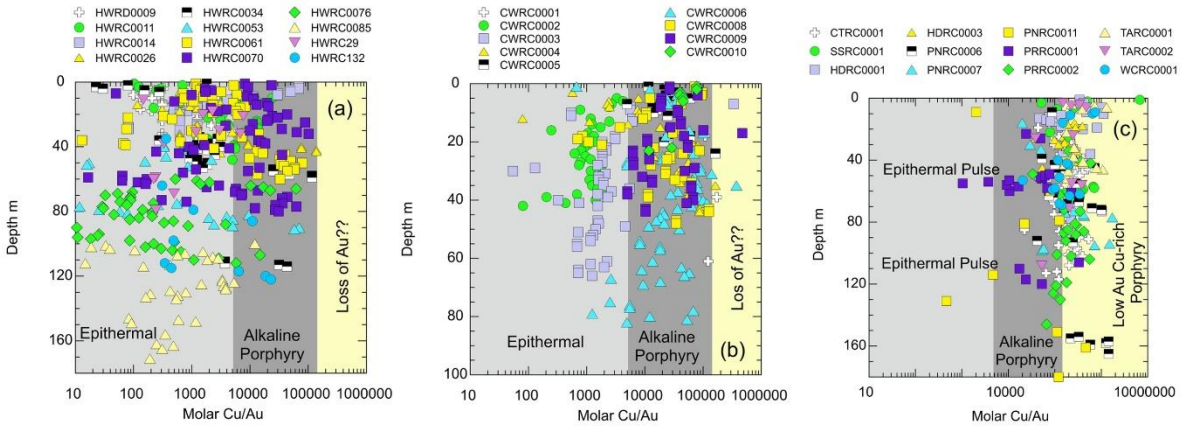

Figure 24. Plots showing molar Cu/Au ratio versus depth in a crustal transect from shallow at Highway to deep at Duck Creek. The data show an increase molar Cu/Au with depth from East (Highway) to west (Duck Creek). This provides further support to the exploration model that the Highway system is a shallow Au dominated epithermal system and Duck Creek is a deeper Cu-Au bearing porphyry. Comway, lies between Highway and Duck Creek at an intermediate crustal depth, with molar Cu/Au ratios between 60,000 and ~120,000, typical of alkaline porphyry systems, to a deeper Cu dominated porphyry systems at Duck Creek with ratios up to ~2,000,000. Note a near surface sample from Success with a molar Cu/Au of ~80,000,000 is Cu enriched due to supergene alteration.

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5.1.9. Nb/Ta Ratios and the Role of a Mantle Component in NWMP Mineralisation

The Nb and Ta are "twin elements", with identical charge (+5) and almost identical ionic size do not generally fractionate during the partial melting. The Nb/Ta ratio of depleted mantle and the continental crust are assumed to be geochemically complementary reservoirs. However, the Nb/Ta ratio of the continental crust ranges from 10-12 [113,114], and the Nb/Ta ratio in the depleted (upper) mantle is 15.5 [115]. These are significantly less than chondritic value of 19 [116]. Thus, mixtures of depleted upper mantle and continental crust do not produce the chondritic Nb/Ta ratios and balanced the primitive mantle [115]. This indicated that another geochemical reservoir with superchondritic Nb/Ta is required to balance the low Nb/Ta observed in the continental crust. Kamber and Collerson [115] showed that fractionation of the chemical twins Nb and Ta can occur by dehydration during subduction of oceanic lithosphere. Melting associated with this process produced arc lavas and continental crust with low Nb/Ta ratio leaving an eclogitic residue with superchondritic Nb/Ta ratio. Kamber et al., [114] have shown that weathered continental crust represented by alluvial sediments derived from Precambrian to Tertiary crust have an average Nb/Ta ratio of 9.76 indicating that the crustal low Nb/Ta ratio is extremely robust value.

The Nb/Ta ratio is useful vector to discriminate between metallogenic crustal and mantle sources as Nb and Ta do not fractionate during crustal partial melting. This is particularly true for lithologies associated with deep sourced mantle plumes that originate from the Earth's primitive lower mantle, e.g., kimberlites with Nb/Ta of 17.9 to 23 [11] and oceanic carbonatites with Nb/Ta ranging from 32 to 474 [117,118].Nb/Ta ratios of >~17 exhibited by ocean island basalt (both EM1 "enriched mantle" and "HIMU" endmembers) are also interpreted to reflect their lower mantle primitive source compositions [11]. Alternative models for the location of the high Nb/Ta reservoir include the subcontinental lithospheric mantle [44] possibly containing melts derived from subducted carbonate [119], or in Earth's core [120].

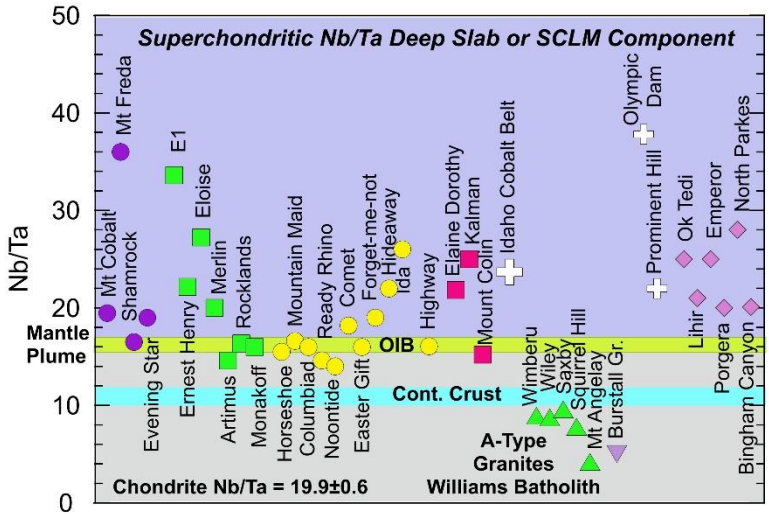

Figure 25. Nb/Ta ratio variability in NWMP deposits and Williams Batholith postulated sources. Also shown for comparison average Nb/Ta ratios of porphyry deposits, Olympic Dam and Prominent Hill IOCG deposits and REE rich Co ores from the Idaho Cobalt Belt. Sources of data: Highway (Tabel 1); Duck Creek deposits (Table 2); NWMP deposits - Shamrock, Evening Star, Mt Freda (Collerson 2019), Ernest Henry, Artimus, E1, Monakoff, Eloise, Rocklands, Merlin (OSNACA data [82]; Elaine Dorothy [46]; Wimberu Granite [63], Squirrel Hill [122]; Mt. Angelay granite [123]; Saxby and Wiley granite [124]; IOCGs and Cu-Au Porphyries (OSNACA data, [82]); Mount Cobalt Idaho Cobalt Belt [125]) Chondrite Nb/Ta [116]; Ocean Island Basalt – Mantle Plume [115], Continental Crust [113,114].

Many of the Cloncurry area "IOCG" deposits e.g., E1, Monakoff, Merlin, Eloise etc, and the Highway-Duck Creek mineral system exhibit superchondritic Nb/Ta ratios (Figure 25). The "IOCG" deposits are spatially associated with the Williams and Naraku batholith (1545–1490 Ma) a suite of

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intrusions interpreted to be the primary source of metals in these deposits. However, this is unlikely, as Williams Batholith intrusions, e.g., Wimberu Granite, Saxby Granite, Wiley Granite, Mount Anglesay Granite and the Squirrel Hill Granite all exhibit subchondritic Nb/Ta ratios <10. Whereas the "IOCG" deposits all have significantly higher Nb/Ta ratios, well above the average Nb/Ta ratio (Nb/Ta 10 -12) of continental crust [114].

The high Nb/Ta ratio of these deposits shows that metals in this system were clearly not derived from a low Nb/Ta crustal source. The Sc, Ni, Co, Pd and Pt tenor of many of these deposits supports the interpretation that mantle-derived melts, like the potassic ultramafic – intermediate pyroxenitic-gabbroic-monzogabbroic intrusions identified at Highway and Duck Creek. Several other domains characterized by elevated Nb/Ta ratios are present in the NWMP e.g., at Mount Cobalt (average Nb/Ta ~25), Elaine Dorothy (average Nb/Ta ~48.6) and Lawler (average Nb/Ta ~54) and it is likely that metals derived from such domains played a major role in post collisional NWMP mineralisation. For example, calcium-rich alkaline lithologies from a deep diamond drill hole MKED0036 between ~250 and 850m, at Mary Kathleen. In these lithologies, CaO ranges from ~20 to 39 wt%, Cu contents are up 1.15 wt %, and Ni, Co, U, Th and REEs range up to 890 ppm, 64 ppm, U 400 ppm, Th 2000 ppm and 1% respectively.

From the almost ubiquitous high Nb/Ta ratio character of deposits in the eastern NWMP (Figure 25), it is likely that this previously unrecognised Ca-rich magmatic system played role in Mesoproterozoic post collisional metallogenesis in the area. These Ca-rich intrusions preserve igneous textures and are clearly post-tectonic and hence post collisional. They are interpreted to represent part of an alkaline suite that is likely genetically associated with the tholeiitic ferrogabbros and monzogabbros, seen in the Duck Creek system. As lithologies with elevated Nb/Ta ratios extend, from Mary Kathleen to Mount Cobalt and Northeast to Ernest Henry, this high Ca alkaline system is interpreted to underly a large area in the eastern Mount Isa block.

This alkaline and tholeiite system is interpreted to have been emplaced through a palaeo-slab window that accreted to the eastern margin of the North Australian craton following Numil Terrane collision from the east (Figure 26). This collisional environment provided the geodynamic setting, the metal source and the plumbing system for formation of post collisional mineralisation in the Highway-Duck Creek area that exhibits all the hallmarks of a Cu-Au porphyry – epithermal mineral system. There are two major metal sources for this mineral system, one associated with tholeiitic ferrogabbros the other was a high Ca alkaline system with high Nb/Ta ratios.

A possible deep porphyry magnetic target has been imaged at a depth of ~ 1 km below the northern portion Highway epithermal system using sub-audio magnetics. If this magnetic anomaly extends for the entire length of the Highway epithermal system viz ~20 km (Figure 1), then based on the epithermal-porphyry model proposed in this paper, the Duck Creek Cu-Au porphyry clearly represents a major new Cu system discovery in the NWMP.

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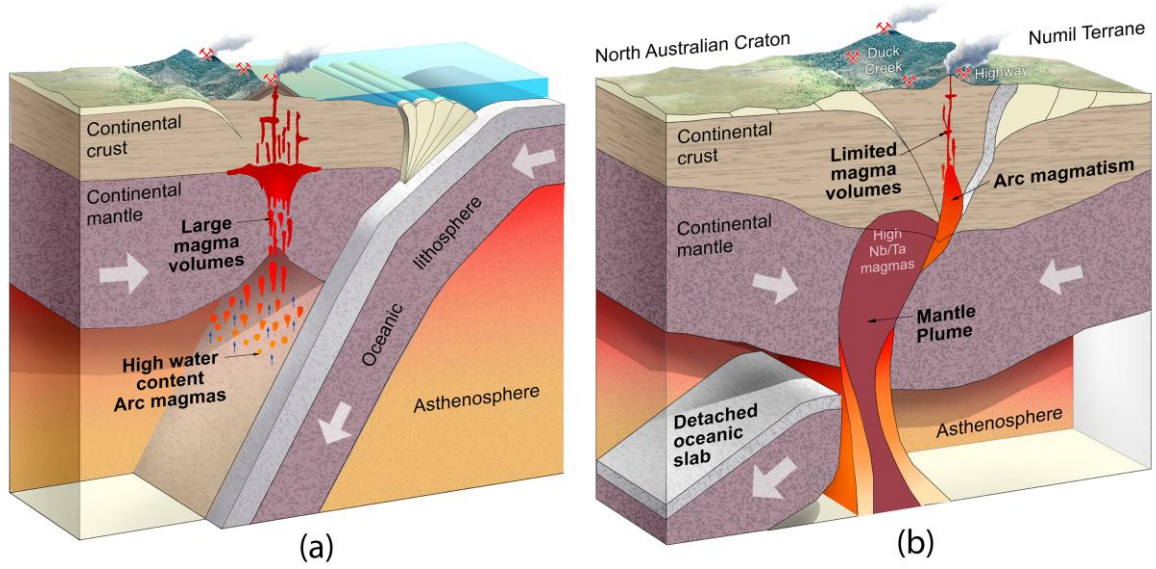

Figure 26. : Schematic geodynamic cartoon comparing subduction generated Cu-Au porphyry systems, where metals are derived from subducting oceanic crust, with post-tectonic collisional porphyry systems, where a slab window has allowed passage of a deep mantle upwelling that interacts with high Nb/Ta slab (eclogite) located in the transition zone and upper lower mantle. The superchondritic Nb/Ta slab and the upwelling plume are additional metal sources for post-collisional porphyry that explain the presence of superchondritic Nb/Ta ratios as well as elevated Pt and Pd chemistries. .

6. Summary and Conclusions

Lithological, geochemical and geodynamic models for the eastern margin of the north Australian craton all support the interpretation that the Highway-Duck Creek area hosts a post-subduction collisional Cu-Au epithermal - porphyry mineral system. The Cu-poor epithermal system at Highway overlies Cu-Au porphyry mineralisation. Timing of mineralisation based on U-Pb, LA ICPMS geochronology of xenotime and monazite is between ~1490 and 1530 Ma. The epithermal Au-Te-W-REE mineralisation is hosted by siliceous breccias that show classic epithermal comb textures. Euhedral and subhedral crystals of quartz in carbothermal lithologies exhibit embayed grain boundaries that indicate acidic corrosion by F or B-rich fluids as evidenced by the presence of tourmaline. Epithermal lithologies have high Au contents and low molar Cu/Au ratios of 6.16 g/t and 19 (epithermal breccia) and 3.46 g/t 475 (carbothermal lithologies) typical of epithermal systems. However, Highway underlying mafic intrusion and Marraba volcanic host rocks also have low, molar Cu/Au ratios e.g., ferrogabbro (5571), quartz monzonite (4051) and andesite (680), with average Au concentrations ranging from 0.03 to 0.35 g/t. Gold occurs as native Au and as Au telluride. Bismuth is also enriched (24 ppm), occurring as Bi telluride. The occurrence of Au and Bi tellurides are typical of epithermal systems. Tungsten is hosted by wolframite and scheelite. REEs occur in xenotime and monazite and thus the system is heavy REE enriched. This heavy REE enrichment was caused by sulphate fractionation, at a late stage of the fluid evolution of the system.

The two systems exhibit broadly uniform shaped geochemical patterns, with positive spikes in W, U, Pb, P and Li and negative spikes in Ba, Th, Nb, Ta and Sr. The pronounced depletion in Sr, a common feature seen at both Duck Creek and Highway, is interpreted to have been caused by breakdown of Ca-Sr bearing minerals during pervasive phyllic, argillic, propylitic and potassic alteration associate with the porphyry mineralisation. Highway also shows a pronounced negative Ti spike, which is not seen at Duck Creek suggesting the fractionation of a Ti-bearing mineral into the gabbro – monzogabbro intrusion which was the source of the epithermal fluids. Although, the level of W enrichment in all Highway lithologies is significantly greater than at Duck Creek, the patterns are broadly similar with relatively consistent immobile-incompatible element ratios. These

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parallel profiles at Duck Creek and Highway support the view that they are related systems and that metals in the Highway epithermal system were delivered by fluids derived from the deeper Duck Creek post-tectonic potassic intrusive suite.

The Cu-Au-Ni-Co-Pt-Pd-Sc rich porphyry is hosted by gabbro and norite an association similar too other collisional orogen hosted Cu-Au porphyry deposits, e.g., Beleuti, Akcha, Kalmakyr, Elasite, Assarel, Medet and Tsar Assen, (Li et al., 2020). The Duck Creek porphyry sulphides are characterised by an assemblage of early crystallised pyrite, followed by chalcopyrite and bornite formed by oxidation of the chalcopyrite. Cu sulphides also show the effect of supergene oxidation alteration producing rims of covellite, chalcocite and digenite.

Redox conditions deduced from V/Sc systematics show the Duck Creek porphyry system contains both oxidized (typical of porphyries) and reduced source lithologies. Covariation plots of Ni/Te and Cu/Te show that Highway epithermal lithologies and Duck Creek porphyries lithologies clearly plot within the definitive fields of comparative global mineral system data sets. Chalcophile and highly siderophile element (Cu, MgO and Pd) systematics show that there is a remarkable similarity between the compositions of Duck Creek samples with similar data from porphyry mineral systems, such as Cadia, Northparkes, Bingham Canyon, Grasberg, Skouries, Kalmakyr, Elaisite, Assarel and Medet.

Geochemical vectors indicate three metal sources in the Duck Creek porphyry system, a tholeiitic ferrogabbro, potassic ultramafic to mafic and a Ca-rich alkaline source. The latter two sources impart the high Nb/Ta ratios characteristic of many deposits. The alkaline and tholeiite systems are interpreted to be genetically associated in upwelling thermochemical plume that penetrated a palaeo-slab window that had accreted to the lithospheric mantle below the North Australian craton following collision from the east Numil Terrane. This collisional environment provided the geodynamic setting, the metal source and the plumbing system for formation of post collisional mineralisation in the Highway-Duck Creek area with all the hallmarks of a Cu-Au porphyry – epithermal mineral system.

Author Contributions: Conceptualization, KDC & DW; Methodology, KDC & DW); Formal analysis KDC; Investigation, KDC: Resources, DW.; Data curation, DW.; Writing—original draft, KDC; Writing—review & editing, KDC & DW; Visualization, KDC & DW.; Supervision, KDC & DW; Project administration, DW. All authors have read and agreed to the published version of the manuscript.

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Conflicts of Interest: The authors declare no conflicts of interest. .

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